### Introduction

As the exploration for porphyry systems becomes more challenging, there exists a need for exploration methods that can detect deposits under cover. This study investigates the physical and chemical characteristics of tourmaline in stream sediment samples around the Casino calc-alkaline porphyry Cu-Au-Mo deposit in the unglaciated terrain of west-central Yukon, Canada. Tourmaline is known to occur in porphyry deposits worldwide (Fig. 1) but its application as an indicator mineral has not yet been utilized.

The Casino deposit, owned by Western Copper and Gold, is one of the largest undeveloped porphyry Cu-Au-Mo deposits in Canada. The deposit is hosted in Late Cretaceous quartz monzonite and associated breccias and is known to contain dispersed tourmaline throughout the deposit. Bulk (10-14 kg) coarse-grained stream sediment samples were collected in creeks around the deposit, nearby by porphyry Cu occurrences, and background areas (Fig. 2).

Tourmaline was recovered in the <2 mm heavy (>3.2 specific gravity (SG)) and mid-density (2.8–3.2 SG) fractions of stream sediments in the study area (Fig. 2). The Casino deposit has a tourmaline anomaly downstream of the deposit, but also in some surrounding streams that do not contain known porphyry mineralization. Tourmaline anomalies in stream sediments can not be used on their own to identify prospective drainages for porphyry mineralization because other rocks may contain tourmaline. Additional characteristics of the tourmaline grains need to be investigated to identify prospective porphyry grains. The purpose of this study is to provide practical tools which can be applied during routine exploration for porphyry deposits in both glaciated and unglaciated terranes.

### **Porphyry Tourmaline Worldwide**



squares represent deposits that do not contain tourmaline. Note how the occurrences of tourmaline correspond to the major porphyry belts and some of the most well-endowed, mineralized porphyries worldwide



### The Casino Deposit



### Bedrock Geology

PALEOCENE TO LOWER EOCENE RHYOLITE CREEK: rhyolite and dacite RHYOLITE CREEK: andesite

LATE CRETACEOUS

PROSPECTOR MOUNTAIN SUITE: syenite CASINO SUITE: quartz-feldspar porphyry MID-CRETACEOUS

- WHITEHORSE SUITE: granodiorite, diorite WHITEHORSE SUITE: guartz monzonite, granite, leucogranite MOUNT NANSEN: andesite to dacite flows
- UPPER CRETACEOUS CARMACKS: augite-olivine basalt and breccia CARMACKS: andesite, porphyry
- shale, tuff, coal LATE TRIASSIC TO EARLY JURASSIC

MINTO SUITE: granodiorite, gneissic schlieren **NEOPROTEROZOIC AND PALEOZOIC** LATE TRIASSIC

STIKINE SUITE: gabbroic Hbl orthogneiss

MIDDLE TO LATE PERMIAN

- SULPHUR CREEK SUITE: granite, <sup>-</sup> metaporphyry
- KLONDIKE SCHIST: qtz-ms-chl schist **MISSISSIPPIAN** SIMPSON RANGE SUITE: metagranodiorite
- <sup>J</sup>metadiorite, metatonalite **DEVONIAN, MISSISSIPPIAN, AND(?) OLDER** FINLAYSON: intermediate to mafic volcanic
- and volcaniclastic rocks FINLAYSON: carbonaceous metasedimentary rocks. metachert
- FINLAYSON: ultramafic rocks, serpentinite;
- LATE DEVONIAN TO MISSISSIPPIAN MT BAKER SUITE: gneissic granodiorite, diorite, monzogranite, gabbro, minor pyroxenite CARMACKS: sandstone, pebble conglomerate, **ORDOVICIAN TO LOWER DEVONIAN** 
  - SCOTTIE CREEK: quartzite, Qtz-Ms-Bt±Grt SNOWCAP: quartzite, psammite, pelite, marble; minor greenstone and amphibolite
  - SNOWCAP: marble



# DISCOVERING A PORPHYRY DEPOSIT USING TOURMALINE: A CASE STUDY FROM YUKON

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# **Tourmaline Morphology**



bridge of tourmaline from the Casino deposit. a) image of hypogene tourmaline from drill core (93-177) at the Casino deposit showing the typical Fur) grains with accessory pyrite (Py). b) clusters of tourmaline crystals from stream sediment sample 1023 that are typical of tourmaline source from a porphyry deposit. Note the variable color of the grains from blue to brown as well as the attached minerals (jarosite-Jrs, quartz-Qz). c) individual tourmaline grains from sample 1023 that are typical grains sourced from background rocks.

Tourmaline morphology is a significant indicator of grains derived from porphyry systems (Beckett-Brown et al. 2023a, 2023b). These clustered crystals (Fig. 4a, b), also described as radiating aggregates of prismatic grains, reflect conditions of formation. This morphology reflects rapid crystallization as a result of pressure release during porphyry formation. Tourmaline in other environments forms more commonly as individual isolated grains that commonly contain inclusions of other minerals (Fig. 4c).

## **Trace Element Discrimination**



Figure 5: Tourmaline trace-element classification diagram. Dashed lines represent the typical field where tourmaline analyses from the respective environments listed plot. The porphyry Cu tourmaline field (thick dashed lines) represents a 99th percentile region of porphyry tourmaline analyses drawn in a straight line. This line coincides with the 10% marker on the ternary. The additional dashed line in the porphyry field represents the 95th percentile of the data at the 3% interval on the ternary scale. Grains plotting in this field should be considered most prospective, with samples plotting in the 99th percentile less so. Overlap between metamorphic (blue dots) and orogenic samples (brown dots) and potentially a reflect misidentification of environment of formation. Published data used for comparison are referenced in Beckett-Brown et al. 2023b. Porphyry Cu samples represent data from 7 deposits, granitic samples are from 9 different localities, metamorphic samples are from 6 different localities, orogenic Au samples are from 25 different localities, pegmatitic samples are from 9 different localities, W ±Sn ±Cu samples are from 8 different localities, VMS samples are from 3 different localities.

A ternary plot of Sr/Pb–Zn/Cu–Ga reveals that tourmaline associated with mineralized porphyry systems have trace element compositions that predominantly occur along the Sr/Pb-Ga join that reflect the chemistry of porphyry fluids (Beckett-Brown et al. 2023b). Porphyry tourmaline contain high Sr/Pb (avg: 297), variable Ga (avg: 58 ppm), and low Zn/Cu (avg: 4.8) values (Fig. 5). Conversely, the tourmaline associated with orogenic Au deposits clusters along the Sr/Pb-Zn/Cu join, and those having a granitic or granitic pegmatite affiliation plot along the Zn/Cu-Ga join (Fig. 5).

Tourmaline can exhibit high degrees of chemical variability that exists at the grain-scale for individual crystals but the application of trace-element ratios allows for any chemical zone to be analyzed and their distinct porphyry signature (*i.e.*, high Sr/Pb, low Zn/Cu and variable Ga) is still retained (Fig. 5). However, some porphyry-related tourmaline plot outside the porphyry tourmaline field (to the left of the Sr/Pb – Ga join) due to relative enrichments in Zn. These tourmaline samples are from a late-stage dike within the porphyry and may explain the Zn enrichment, similar to the zonation of Zn seen in porphyry systems. Based on this assumption, samples of tourmaline with high Zn could potentially be considered less prospective (*i.e.*, typically have concomitant decreases in Cu) because they occur distal to the porphyry center or are entirely unrelated to the mineralizing process.

Figure 2: a) Map of the Yukon Territory, Canada showing the location of the Casino deposit. b) Map of the Casino deposit area showing the local bedrock geology Yukon Geological Survey, 2015), location of mineral occurrences (Yukon Geological Survey MDI), and stream sediment samples (red tars). Sample numbers are listed in eside each sample site. leaend presented below

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### **Granitic and Pegmatitc Tourmaline**



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