

Biotite chemistry as a monitor of magma fertility and mineralisation potential: Results from the Devonian granitoids of New Brunswick

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Abstract: There are over 150 granitoid intrusions in the New Brunswick; however, all the mineralised intrusions formed in relation to Acadian and Neocadian orogenic phases of Appalachian accretion. These granitoids range in age from 423 to 360 Ma, and include examples of pre-, syn-, late-, and post-tectonic emplacement with affinities ranging from primitive to highly evolved A-, S-, and I-types granitoids along with their hybrid varieties. Many of these are spatially and temporally related to specific styles of mineralization, producing deposits of Sn, Ta, Li, Sb, W, Mo, Cu, and Au, as well as other base-metals and U.

Igneous biotite crystallises over a wide range of conditions and reacts very sensitively to physio-chemical conditions like halogen and oxygen fugacities, pressure, temperature and chemical composition of the magmas. This sensitivity makes biotite a suitable mineral for identifying the petrogenetic processes, mineralization and alteration of the host granitic rocks. The following features make biotite a valuable probe of magma composition: i) It is the most important reservoir of any excess aluminium in granites that do not contain modal garnet, cordierite, or the Al₂SiO₅ polymorphs; therefore, it directly reflects the peraluminosity of the host magma in such rocks; ii) it is the most readily available indicator of oxidation state; and iii) it can provide information about the F and Cl composition of the magma.

Previous studies have shown that biotite, and to lesser extent hornblende and magnetite, continuously equilibrates with host liquids. Consequently, a core-to-rim study of these minerals and their compositional zoning can provide a record of magma evolution so that the origin and evolution of granitoids can be documented.

The aim of this study is to calculate fluoride and chloride activity of aqueous fluids based on measuring F and Cl contents in the minerals containing hydroxyl and halogens, using a combination of electron microprobe and Laser Ablation ICP-MS. These data will be combined and compared with whole-rock trace element geochemistry. The results are expected to help constrain crystallisation conditions, volatile exsolution, and fluorine-chlorine activity of fluids associated with these intrusions, and also to examine the degree of subsolidus re-equilibration using various geothermobarometry techniques. By linking these results to the various styles/types of granitoids and their associated mineralisation it is hoped to establish biotite composition as a robust indicator of an intrusions ore potential.

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Biotite Chemistry as a Monitor of Magma Fertility and Mineralisation Potential: Results From the Devonian Granitoids of New Brunswick



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ABSTRACT

There are over 150 granitoid intrusions in the New Brunswick, however, all the mineralised intrusions formed in relation to Acadian and Neo-Acadian orogenic phases of Appalachian accretion. These granitoids range in age from 423 to 360 Ma, and include examples of pre-, syn-, late-, and post-tectonic emplacement with affinities ranging from primitive to highly evolved A-, S-, and I-types granitoids along with their hybrid varieties. Many of these are spatially and temporally related to specific styles of mineralization, producing deposits of Sn, Ta, Ti, Sb, W, Mo, Cu, and Au, as well as other base-metals and U.

Igneous biotite crystallises over a wide range of conditions and reacts very sensitively to physio-chemical conditions like halogen and oxygen fugacities, pressure, temperature and chemical composition of the magmas. This sensitivity makes biotite a suitable mineral for identifying the petrogenetic processes, mineralization and alteration of the host granitic rocks. The following features make biotite a valuable probe of magma composition: 1) It is the most important reservoir of any excess aluminium in granites that do not contain modal garnet, cordierite, or the Al₂SiO₅ polymorphs; therefore, it directly reflects the peraluminosity of the host magma in such rocks; 2) Biotite is a reliable indicator of oxidation state; and 3) It can provide information about the F and Cl composition of the magma.

Previous studies have shown that biotite, and to lesser extent hornblende and magnetite, continuously equilibrates with host liquids. Consequently, a core-to-rim study of these minerals and their compositional zoning can provide a record of magma evolution so that the origin and evolution of granitoids can be documented.

The aim of this study is to calculate fluoride and chloride activity of aqueous fluids based on measuring F and Cl contents in the minerals containing hydroxy and halogens, using a combination of electron microprobe and Laser Ablation ICP-MS. These data will be combined and compared with whole-rock trace element geochemistry. They are evaluated against the conditions of crystallization of the granitoids. The results are compared with various geothermometry techniques. By linking these results to the various styles/types of granitoids and their associated mineralisation it is hoped to establish biotite composition as a robust indicator of an intrusions ore potential.

INTRODUCTION

Mica group minerals are among the most common mafic minerals of intermediate to felsic igneous rocks; they can also be found in most of metamorphic and some of sedimentary rocks. Among them, biotite is particularly characteristic of the intermediate rocks of calc. alkaline affinities and occurs in a wide range of rocks of hybrid origin (Deer et al. 1992).

Mineralized and barren rocks are characterized by different chemical variations in biotite. For instance, mineralized biotite is characterized by lower Mg and Ti contents relative to biotite from barren rocks; they also have higher amount of Al compared to biotite from barren phases (Goomer et al. 2010).

Biotites in the granitoid of New Brunswick area usually altered to chlorite and to some extent muscovite along their cleavages. The alteration products are chlorite, pyrite, pyrrhotite, and sulfide minerals (e.g., pyrrhotite, pyrite) as mineral inclusions. Their colour also varies from green to brown and reddish brown reflecting differences in their redox conditions (cf. Lalonde and Bernard 1993).

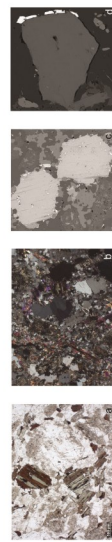


Fig. 1. Photomicrographs and SEM-BSE images of some New Brunswick biotite-bearing assemblages illustrating the mineralogy and texture. (a) Partly altered brownish biotite to chlorite along their cleavages, Lake George granitoidite, sample 1647. (b) Pyrrhotite (Prt) and pyrite (Py) inclusions in biotite from the Lake George granitoidite, sample 1647. (c) Pyrrhotite (Prt) and pyrite (Py) inclusions in biotite from the Lake George granitoidite, sample 1190.

TECTONIC SETTING

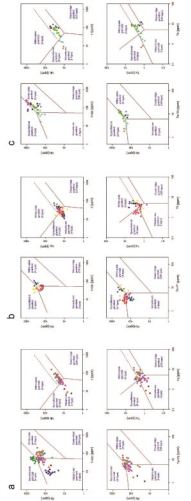


Fig. 2. Tectonometric discrimination diagrams of New Brunswick Devonian granitoids data from Whalen (1987), McFarlane and Kerr (1993), and McFarlane and Kerr (1996). (a) McFarlane and Kerr (1993) and (b) McFarlane and Kerr (1996) diagrams. (c) McFarlane and Kerr (1996) diagram. (d) McFarlane and Kerr (1996) diagram. (e) McFarlane and Kerr (1996) diagram. (f) McFarlane and Kerr (1996) diagram. (g) McFarlane and Kerr (1996) diagram. (h) McFarlane and Kerr (1996) diagram. (i) McFarlane and Kerr (1996) diagram. (j) McFarlane and Kerr (1996) diagram. (k) McFarlane and Kerr (1996) diagram. (l) McFarlane and Kerr (1996) diagram. (m) McFarlane and Kerr (1996) diagram. (n) McFarlane and Kerr (1996) diagram. (o) McFarlane and Kerr (1996) diagram. (p) McFarlane and Kerr (1996) diagram. (q) McFarlane and Kerr (1996) diagram. (r) McFarlane and Kerr (1996) diagram. (s) McFarlane and Kerr (1996) diagram. (t) McFarlane and Kerr (1996) diagram. (u) McFarlane and Kerr (1996) diagram. (v) McFarlane and Kerr (1996) diagram. (w) McFarlane and Kerr (1996) diagram. (x) McFarlane and Kerr (1996) diagram. (y) McFarlane and Kerr (1996) diagram. (z) McFarlane and Kerr (1996) diagram.

FUTURE STUDIES

(a) Calculating zircon saturation temperature, (b) calculating apatite saturation temperature, (c) studying zoning in apatites, (d) apatite, and (e) zircon and trying to correlate their zoning in order to better understanding the history of magma evolution, (f) calculation of the trace element distribution and partition coefficient of elements between melt and crystals, (g) calculation of the trace element distribution and partition coefficient of elements between melt and crystals (that are traced by SEM-BSE and ChromSENCL), (h) analyzing the studied granites using laser from rim to rim, and calculating trace element distribution between whole rock and biotite grains, (i) Advancing the LIS method done at the University of New Brunswick.

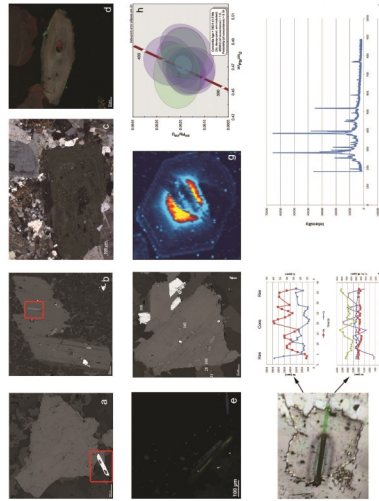


Fig. 3. Geological maps of New Brunswick showing the location of various granitoid intrusions. (a) Geological map of New Brunswick showing the location of various granitoid intrusions. (b) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (c) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (d) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (e) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (f) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (g) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (h) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (i) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (j) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (k) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (l) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (m) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (n) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (o) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (p) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (q) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (r) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (s) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (t) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (u) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (v) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (w) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (x) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (y) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions. (z) Geological map of the Canadian and adjacent New England Appalachians showing the location of various granitoid intrusions.

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