

Using biotite composition of the Devonian Lake George granodiorite, New Brunswick, as a case study for W-Mo-Au-Sb mineralized magmatic hydrothermal systems

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Abstract: The granodioritic Lake George polymetallic deposit is located approximately 35 km southwest of Fredericton, New Brunswick. This intrusion is a metaluminous to weakly peraluminous, calc-alkaline body that shows an evolved I-type, volcanic arc affinity. Based on an Early Devonian age determination, (412 \pm 5/-4 Ma, zircon U-Pb) it is related to the Hakshaw granite phase of the Pokiok Batholith. The Lake George stock is cut by Hibbard stibnite-quartz veins and quartz-scheelite-molybdenite veinlets that contain significant amounts of gold. Fresh biotites from this intrusion were analyzed at both the core and rim by electron microprobe, and along rim to rim transects by laser ablation ICP-MS at the University of New Brunswick to build an understanding of the halogens responsible for mineralization and trace element distribution within this deposit.

Biotites of this intrusion are reddish brown in colour (indicative of reduced I-type source) and mostly altered to chlorite. They usually contain apatite, zircon, titanite, ilmenite, rutile, and sulphide minerals as mineral inclusions. Temperature was calculated with the Ti-In-biotite geothermometer, in which the results showed a variation between 583 and 745°C.

Two depths of emplacement were determined based on the Al in hornblende and biotite geobarometries confirming the porphyritic texture of this intrusion (4.3 and 1.5 km, respectively). Forming amphiboles at that depth clearly indicates a high water content of the source magma; in addition, hydroxyl is the most dominant component of the hydroxyl site (Average 1.89 wt%) in biotites. The limited range of IV(F/Cl) values of the Lake George biotites suggested that they all equilibrated with one fluid.

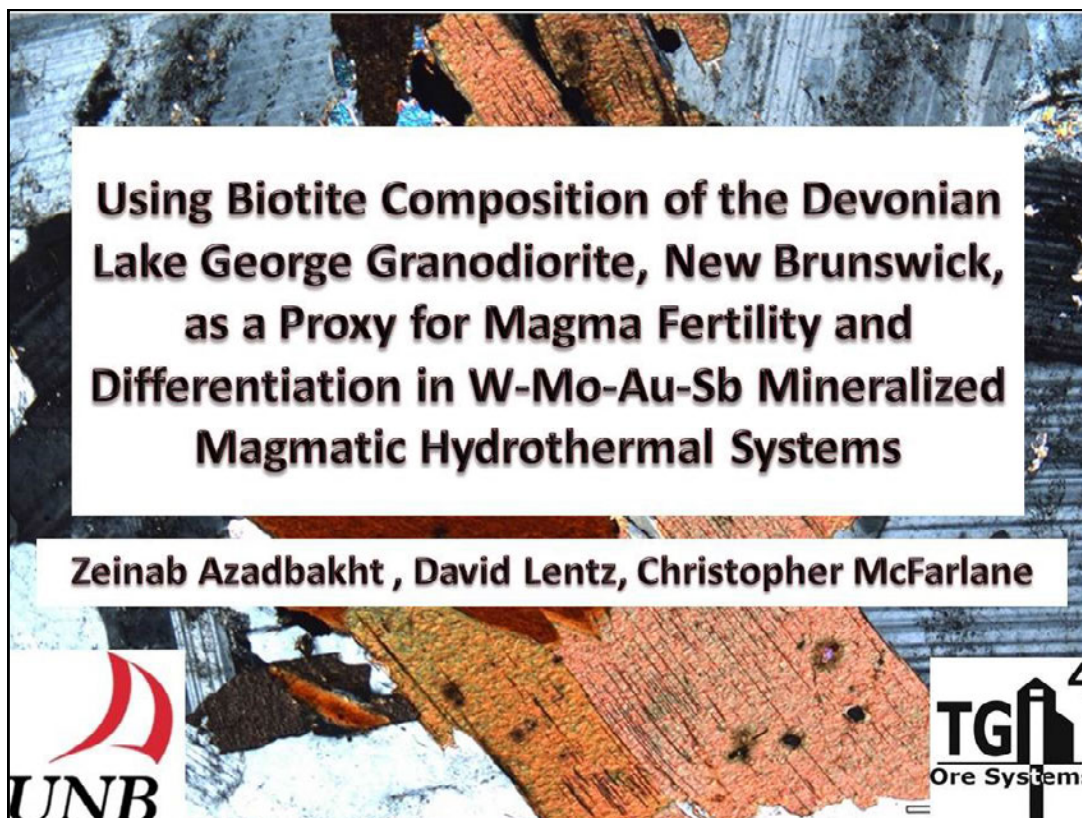
Even though there is no noticeable difference in major elements from core to rims, evidence of magma evolution is recorded by biotite grains by their trace elements. For instance, Cu, Rb, Cr, K, Mo, Sn, Cs and W increase from core to rims, whereas Ba, Ni, Mn, and Li act inversely. Sb has a negligible variation from core to rim. Interestingly, the partition coefficient (biotite/whole-rock) is significantly small for Sb, W, and Mo (main associated mineralization) with 0.06, 0.28 and 0.13 in pure magmatic biotites and increase to 0.77, 1.93 and 0.20 in more altered biotites reflecting enrichment of these elements towards the late stage fluid.

Based on these observations, the concept of using mica composition to help identify fertile Acadian magma systems was proved; this method may be a useful tool to indicate the difference between barren and mineralized granophile-element rich systems.

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Using Biotite Composition of the Devonian Lake George Granodiorite, New Brunswick, as a Proxy for Magma Fertility and Differentiation in W-Mo-Au-Sb Mineralized Magmatic Hydrothermal Systems

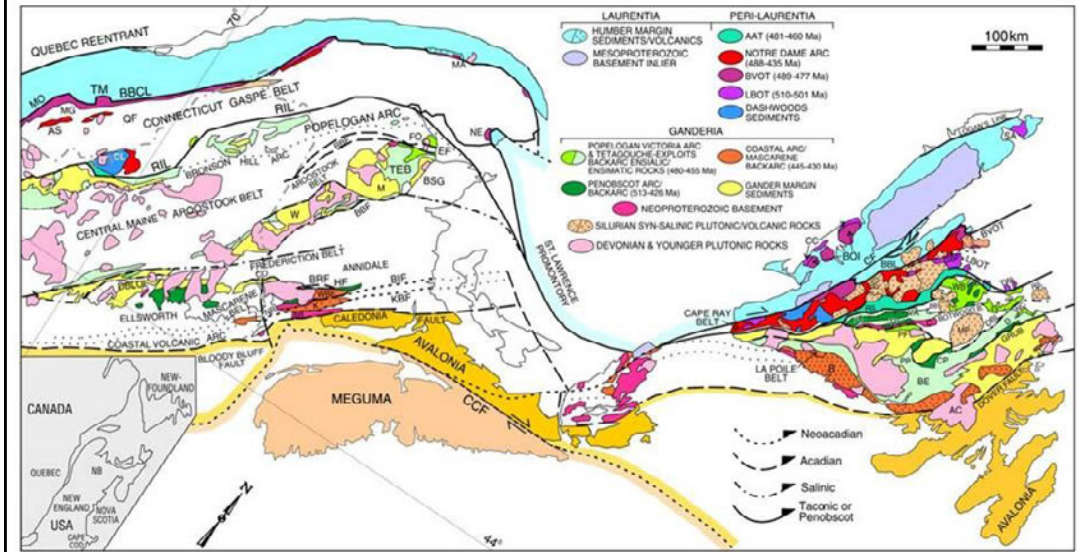
Zeinab Azadbakht , David Lentz, Christopher McFarlane



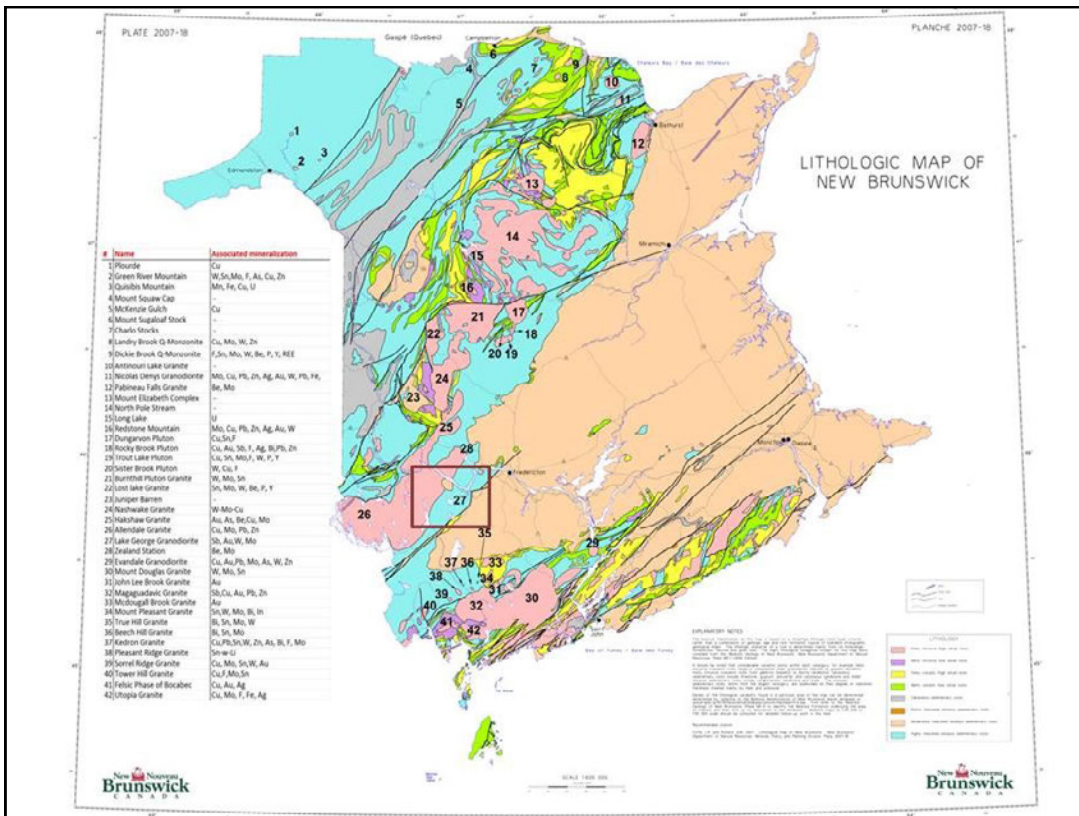
Project Goals

- Document detailed chemistry of coexisting phyllosilicate mineral phases to reveal distinctive element signatures that may be specific to fertile granitoids.
- Provide a geochemical fingerprint for more-specific mineralization styles and prospectivity in the Acadian Plutonic Complex.

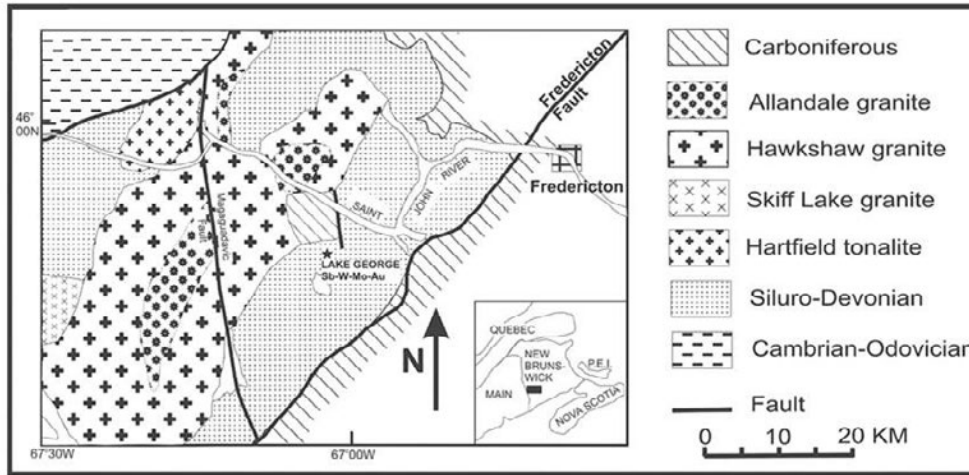
Geological and Tectonic Setting



Van Staal et al. (2009)

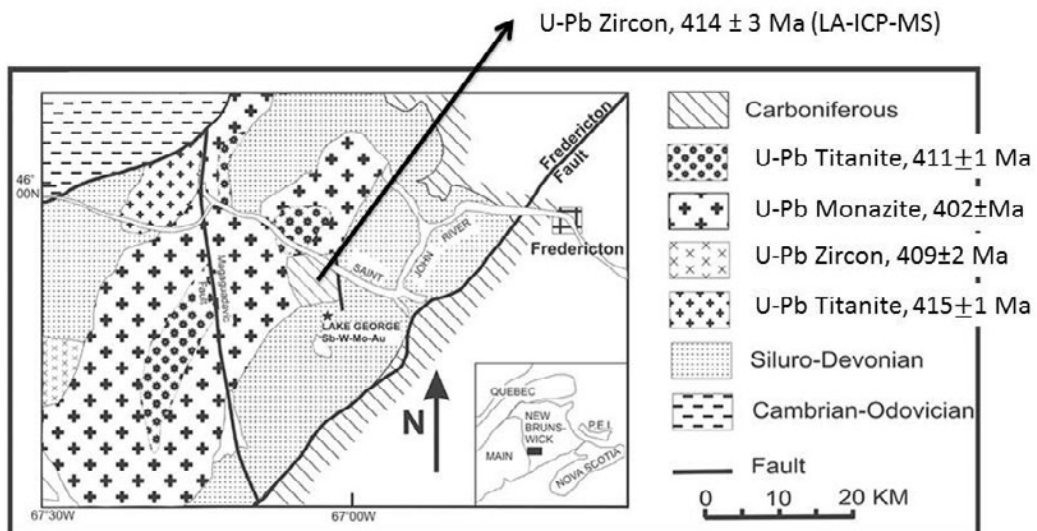


Geological Setting



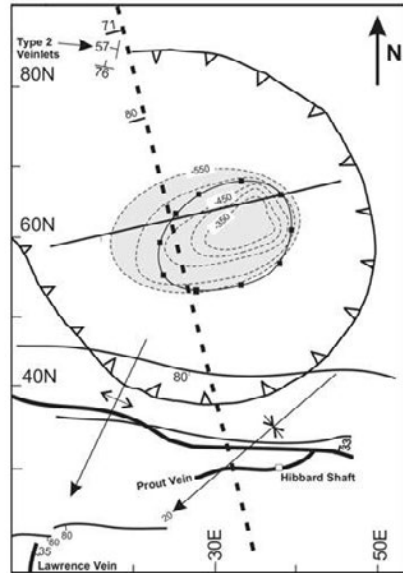
Yang et al. (2002)

Geological Setting



Yang et al. (2002)

Geological Setting



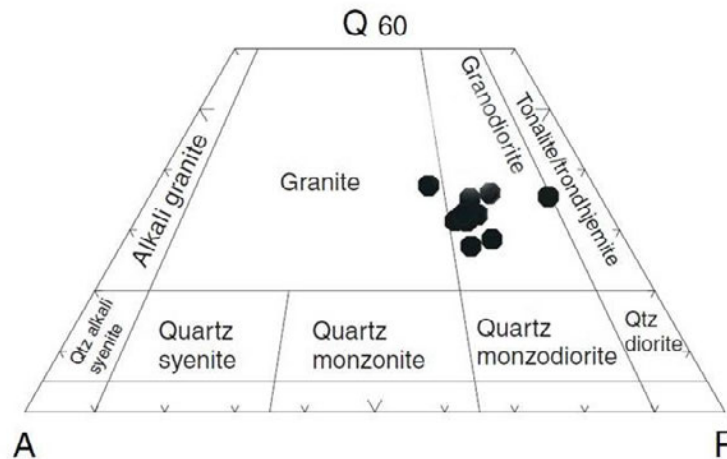
LEGEND

- Stibnite-quartz vein
- - - Quartz-feldspar porphyry dyke
- Lamprophyric dyke
- Silurian metasedimentary rocks
- - - 350 Structural contour of the roof of the granodiorite stock (metres below surface)
- ∇ Biotite isograd
- Mottled hornfels isograd
- ↔ 20 Fold axis

0 500 m

Leonard et al. 2005

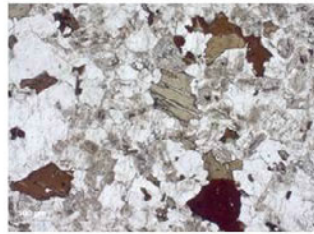
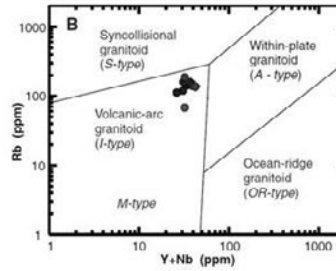
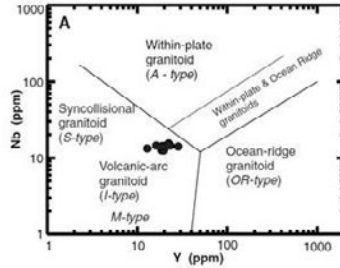
Wholerock Geochemistry



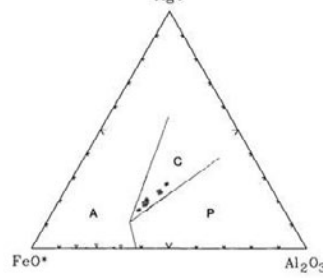
Quartz (Q)-Alkali feldspar (A)- Plagioclase (P) mesonormative classification diagram(after Le Maitre et al., 1989) (Yang et al.(2002)

Tectonic Setting

Yang et al.
(2002)

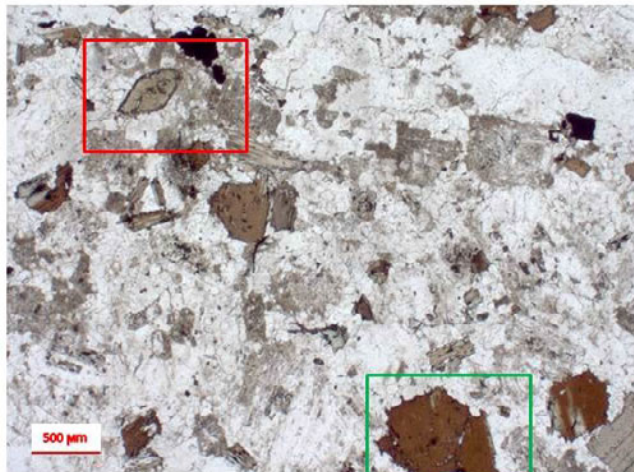


Lalonde and Bernard (1993)



Abdel-Rahman (1994)

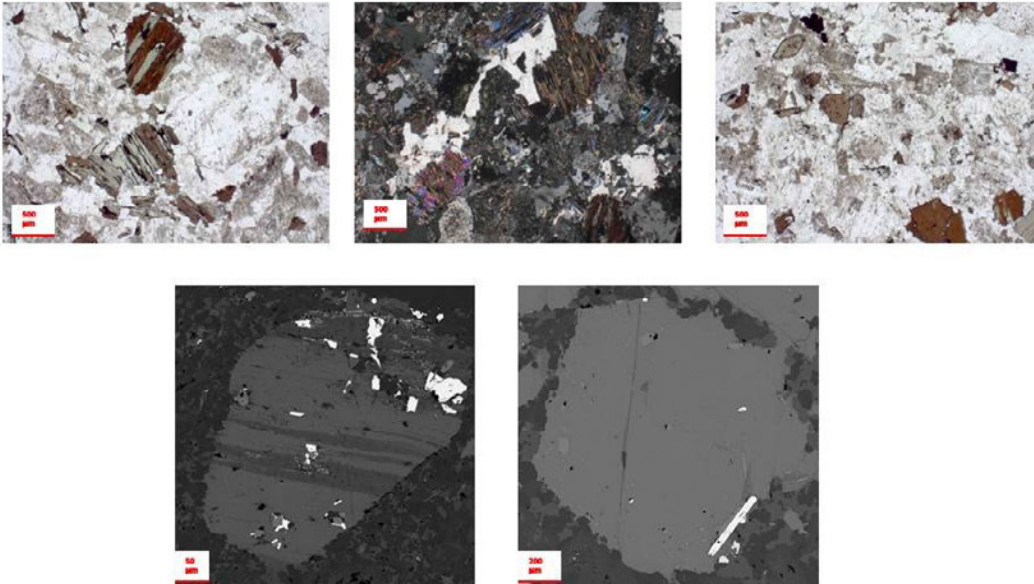
Geobarometry



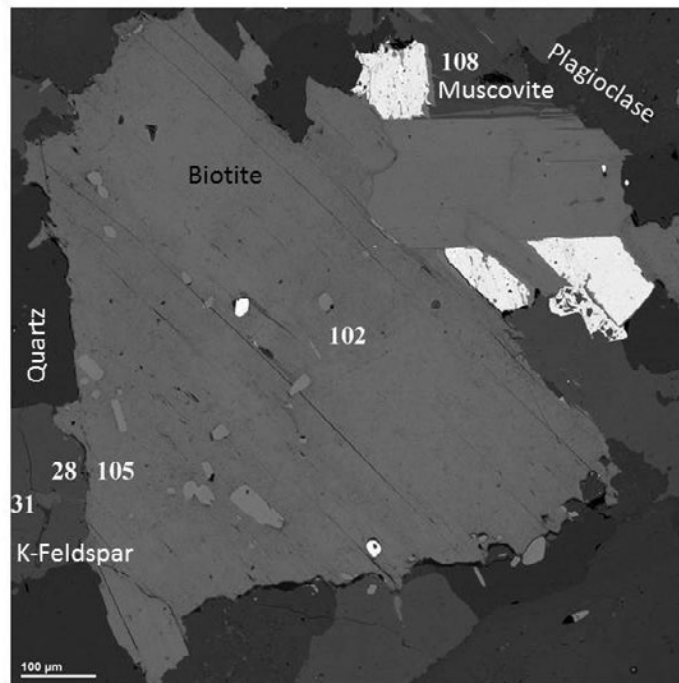
(~3.4 - 4.6 kb)
Done by Yang et al.(2002)

1~2 kb
Uchida et al.(2007)

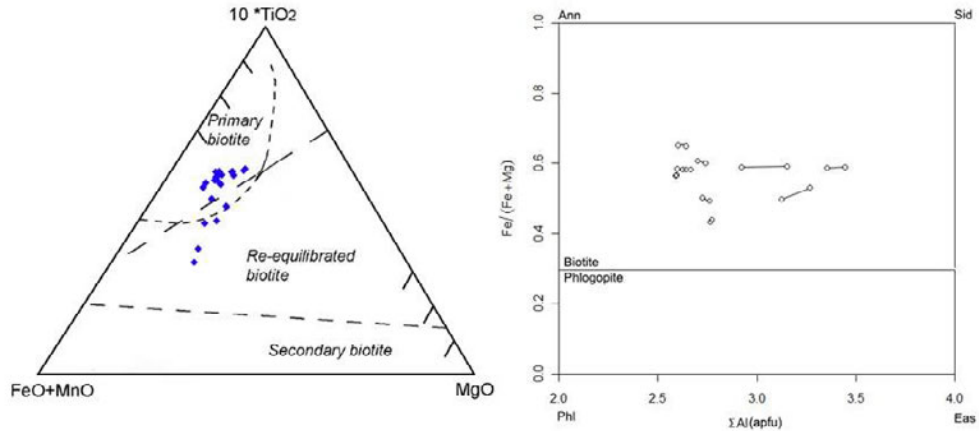
Petrography



Methodology

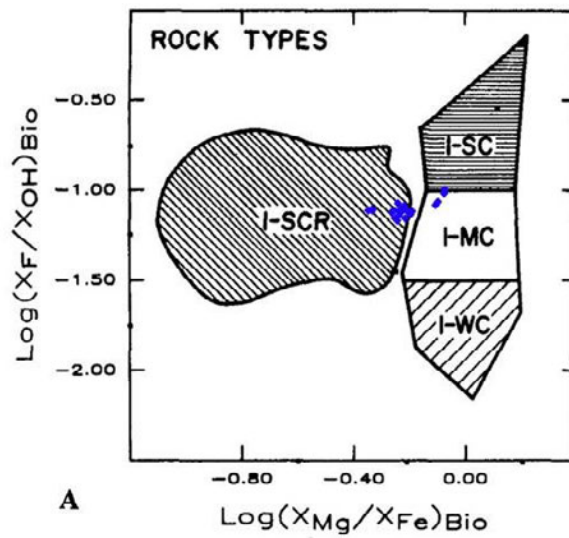


Biotite Classification



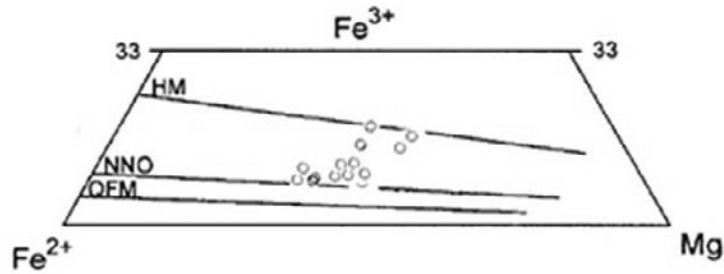
10xTiO₂ – (FeO + MnO) – MgO ternary plot of biotite composition from the Lake George granodiorite; the discrimination field boundaries are from Nachit et al. (2005).

Biotite Classification



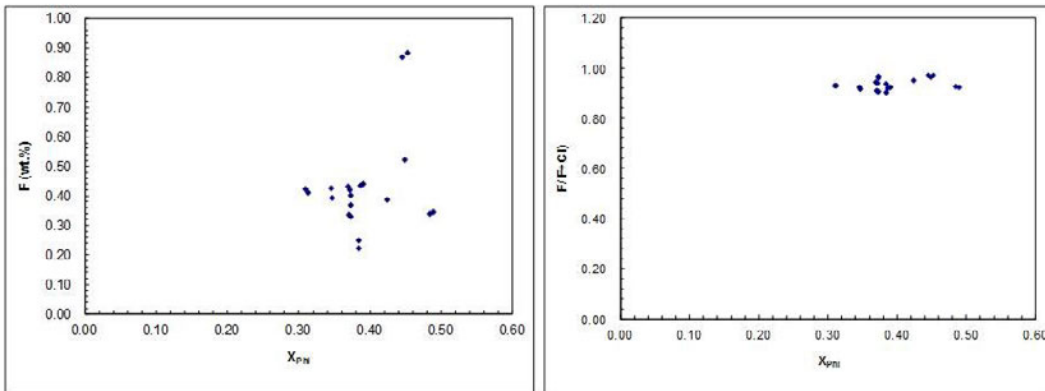
Plot of $\log(X_{Mg}/X_{Fe})$ and $\log(X_F/X_{OH})$ of the magmatic biotite I-SC, strongly contaminated I-type; I-SCR, strongly contaminated and reduced I-type; I-WC, weakly contaminated I-type; after Ague and Brimhill (1987).

Oxygen Fugacity



Mica+
Yavuz (2003 a and b)

Biotite Halogen Study



Composition of the biotite plotted with respect to F Wt% vs. X_{pht} for granodiorite samples of Lake George. $X_{pht} = \text{Mg}/\text{Sum octahedral site}$

Biotite Halogen Study

Munoz 1984

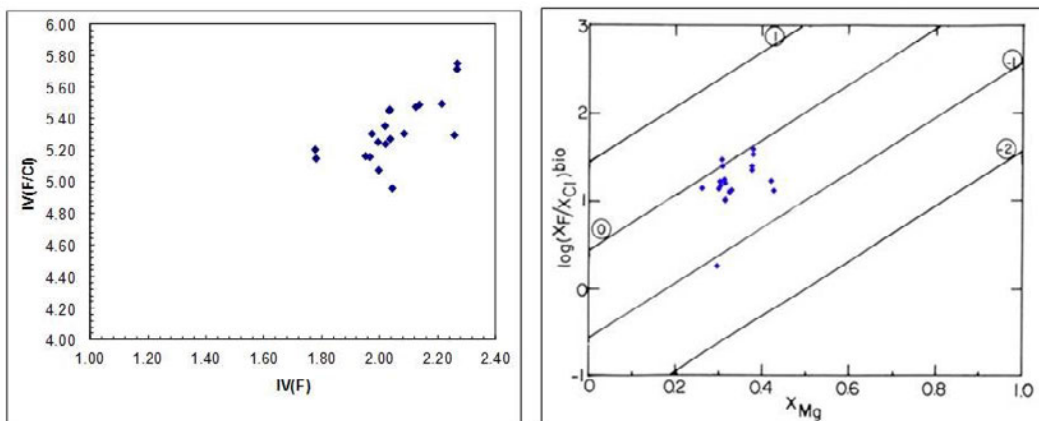
$$IV(F) = 1.25 X_{\text{phl}} + 0.42 X_{\text{ann}} + 0.20 X_{\text{sid}} - \log(X_F/X_{\text{OH}})$$

$$X_{\text{sid}} = [(3-\text{Si}/\text{Al})/1.75] (1 - X_{\text{phl}})$$

$$X_{\text{ann}} = 1 - (X_{\text{phl}} + X_{\text{sid}});$$

X_F and X_{OH} = mole fractions of F and OH respectively on the hydroxyl site

Biotite Halogen Study



Plot showing values of IV(F) vs. IV(F/Cl) of biotite grains from granodiorite of Lake George

Magmatic Hydrothermal Fluid

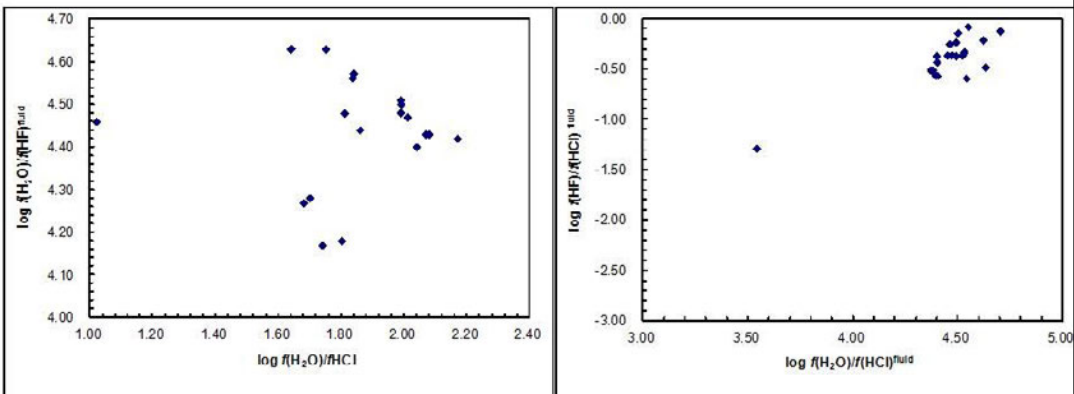
$$\begin{aligned} \log(f_{\text{H}_2\text{O}}/f_{\text{HF}})_{\text{fluid}} \\ = 1000/T[2.37 + 1.1(X_{\text{Mg}})_{\text{biotite}}] + 0.43 \\ - \log(X_{\text{F}}/X_{\text{OH}})_{\text{biotite}}, \end{aligned}$$

$$\begin{aligned} \log(f_{\text{H}_2\text{O}}/f_{\text{HCl}})_{\text{fluid}} \\ = 1000/T[1.15 + 0.55(X_{\text{Mg}})_{\text{biotite}}] + 0.68 \\ - \log(X_{\text{Cl}}/X_{\text{OH}})_{\text{biotite}}, \end{aligned}$$

$$\begin{aligned} \log(f_{\text{HF}}/f_{\text{HCl}})_{\text{fluid}} \\ = -1000/T[1.22 + 1.65(X_{\text{Mg}})_{\text{biotite}}] \\ + 0.25 + \log(X_{\text{F}}/X_{\text{Cl}})_{\text{biotite}}, \end{aligned}$$

Munoz 1992

Magmatic Hydrothermal Fluid

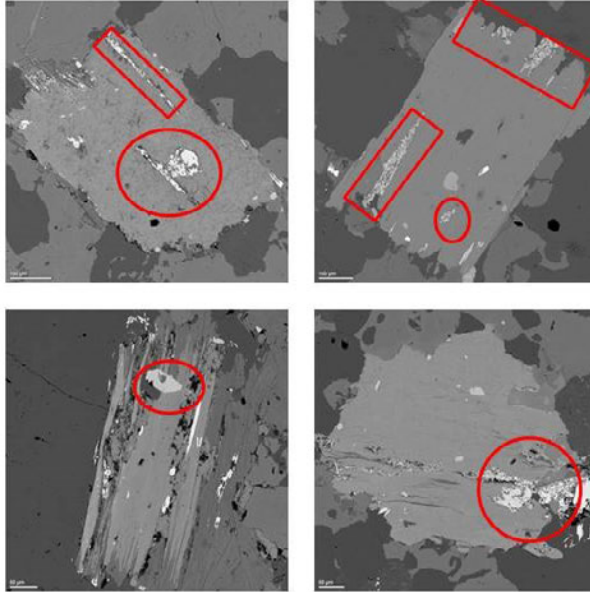


Ti in Biotite Thermometry

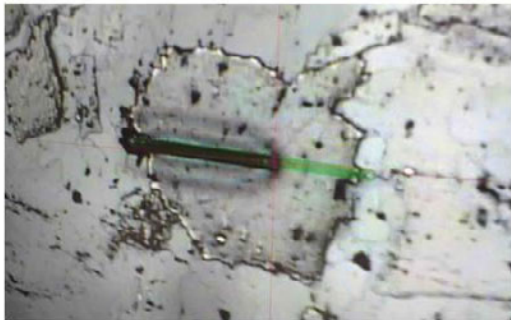
#	Core	Rim
1351	725	725
1190	744	746
1886	719	718
1741	726	729
1291	735	729
1647	720	717
80-33-A	710	711
612	748	746
1567	718	716
80-33-B	719	719

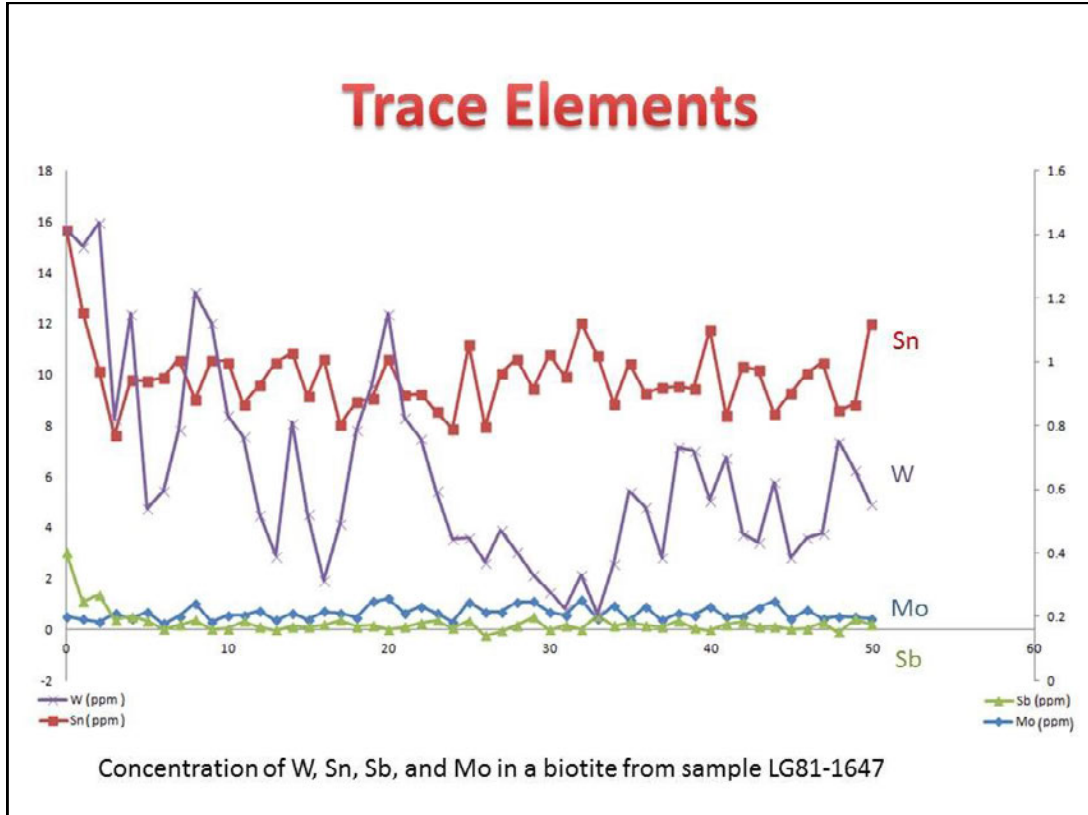
Henry et al. 2005

Biotite grains containing secondary rutile show a lower calculated temperature following the Henry method.



Trace Elements



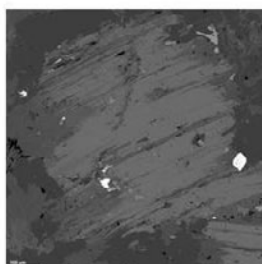
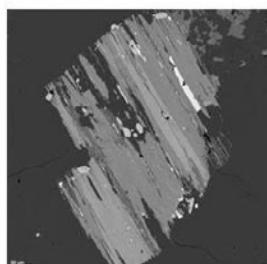


Trace Elements

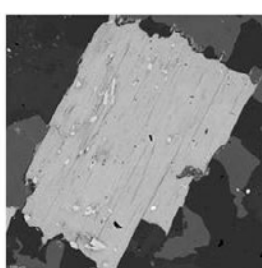
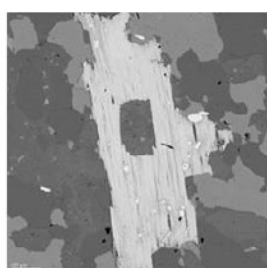
(ppm)	80-33A	1291	1351	1647	1741	1886
Li	257.42	265.23	275.35	321.72	281.47	392.44
P	2.93	23.86	14.04	72.06	19.43	2174.51
Ca	292.38	452.92	347.68	266.09	732.63	6141.64
Ti	24877.94	31702.44	27927.93	27990.00	28495.14	27782.75
V	385.32	640.43	454.60	523.67	486.95	544.28
Cr	47.51	130.86	60.52	69.05	70.14	63.88
Mn	6634.01	2263.85	4287.13	4652.31	3940.02	4649.05
Co	44.53	31.43	58.46	49.38	62.84	45.65
Ni	44.84	46.29	61.08	48.67	62.42	40.49
Cu	3.66	4.99	1.15	0.55	3.81	12.50
Zn	531.73	391.19	455.21	476.70	449.32	484.71
Ga	50.72	63.22	46.02	49.68	45.97	50.80
Rb	802.79	843.97	733.01	763.71	755.35	746.56
Sr	1.73	5.65	1.59	1.70	2.32	5.09
Mu	0.97	0.34	0.74	0.35	0.62	0.32
Sn	63.94	18.77	8.02	10.57	8.21	12.11
Sb	0.77	2.36	0.27	0.07	0.33	-0.09
Cs	90.91	58.39	17.94	26.41	20.55	42.73
Ba	1304.40	1439.19	1885.19	2053.30	1466.99	2468.81
W	1.07	1.29	0.78	0.76	0.59	0.84

Yellow: Highest concentration
Orange: lowest concentration

Trace Elements



1291- Sb-W



80-33A Sn-Mo

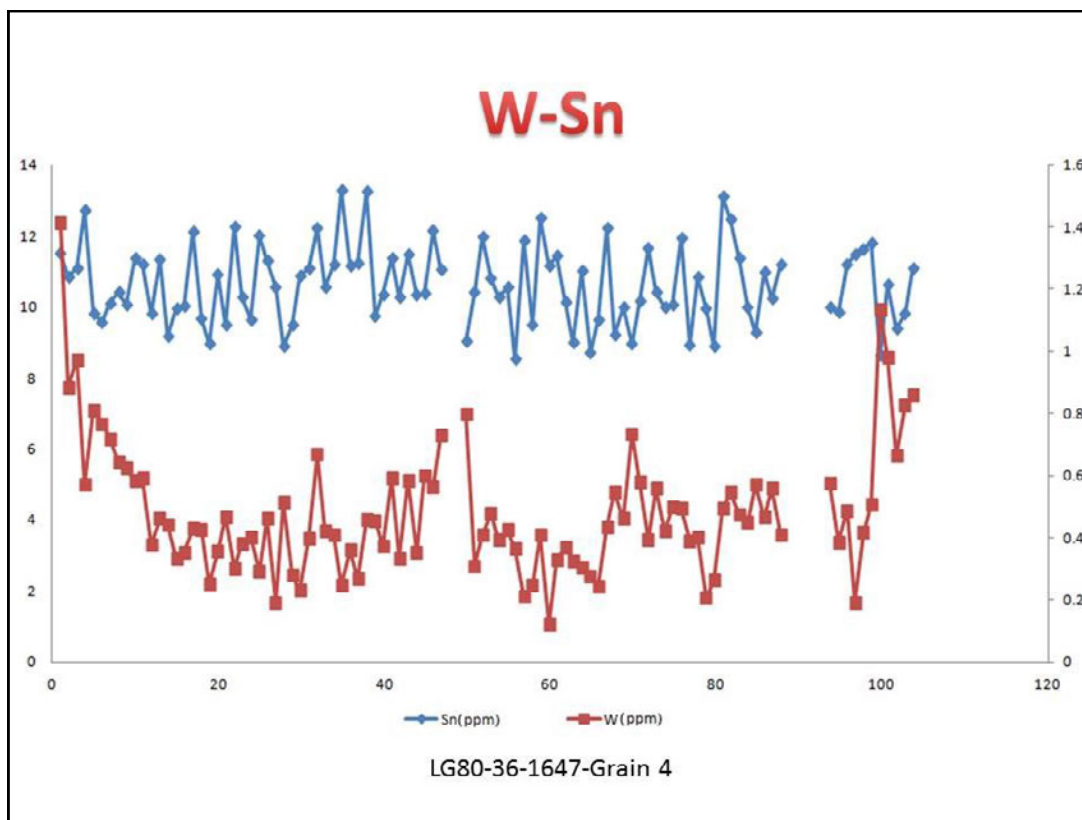
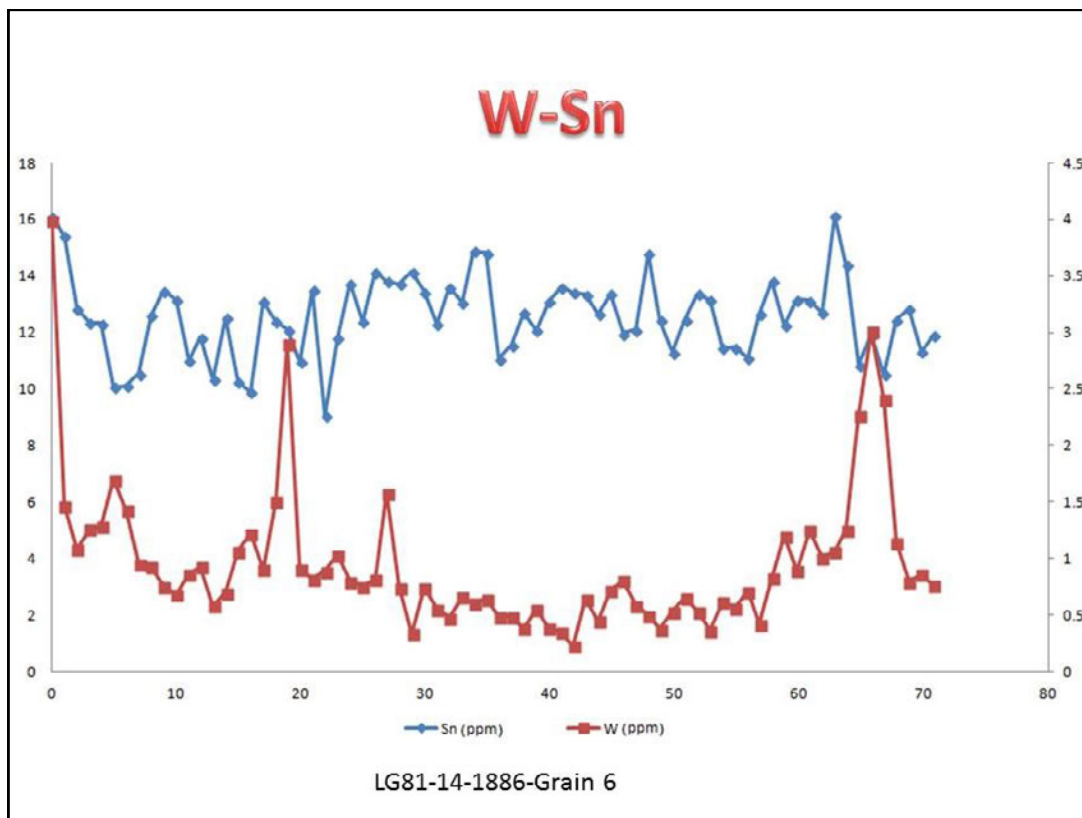
Trace Elements

	Mo	Sn	Sb	W
1351	6	<1	4.8	<1
1741	3	<1	4.3	8
1647	7	<1	6	6
1291	6	1	6.5	<1
1886	2	<1	5.6	5

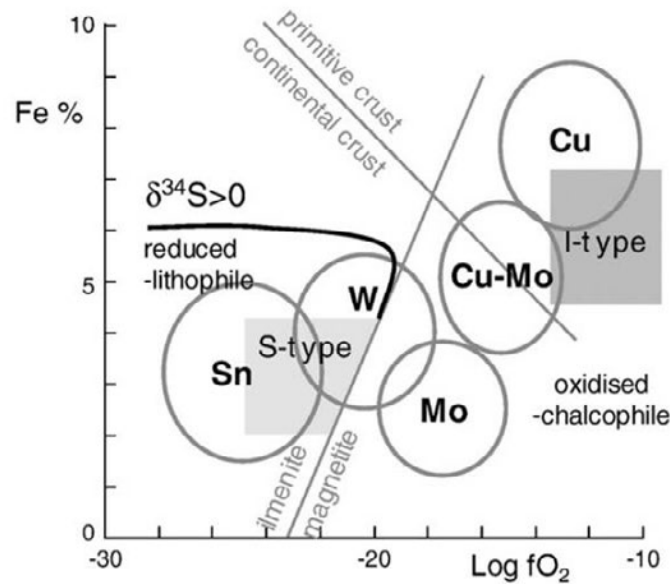
Whole rock trace element Yang et al. (2002)

Averages of Mo, Sn, Sb and W in biotites of the Lake George granodiorite (Right)

	Mo	Sn	Sb	W
1351-6	0.64	9.91	0.26	0.64
1351-7	0.75	9.31	0.22	0.67
1351-8-1	0.71	6.89	0.33	0.92
1741-2	0.64	8.66	0.47	0.64
1741-4	0.54	8.05	0.75	0.69
1741-5	0.65	8.01	0.4	0.59
80-33-A-2	1.00	66.18	0.76	1.30
80-33-A-3	1.12	62.56	0.66	1.06
80-33-A-5	0.74	62.82	1.16	0.92
1647-3	0.33	10.82	0.06	0.56
1647-4	0.44	10.62	0.01	0.48
1647-5	0.29	10.14	0.16	1.54
1291-2	0.29	20.78	0.32	0.64
1291-7	0.39	16.37	3.59	2.24
1291-8	0.34	17.18	11.14	2.91
1886-4	0.33	11.70	-0.2	0.77
1886-6	0.29	12.56	0.11	0.95



What I think



Conclusion

- Biotite chemistry reflects whole rock geochemistry
- fH_2O is higher than fHF and $fHCl$
- Disequilibrium partitioning of halogens and high metal concentrations in biotite, good reflection of contamination?
- Internal zoning of trace element good record of physio-chemical changes during biotite crystallization
- Proof of concept that mica chemistry may aid in identification of fertile Acadian magma systems