

DESCRIPTIVE NOTES

The map area (NTS 87-H12) lies within the Minto Inlet, a ~300 km long by 100–150 km wide belt of gently folded sedimentary and igneous rocks of early Neoproterozoic (late Tonian–early Cryogenian) age. The Neoproterozoic sedimentary strata belong to the Shaler Supergroup, an approximately 4 km thick succession of shallow marine carbonate rocks and evaporate rocks with interbedded terrigenous rocks that were mainly deposited in a shallow intracratonic epicritic sea referred to as the Hudson Basin (Rainbird et al., 1994; Rainbird et al., 1995a; Thompson and Topp, 1992; Young, 1981). The basin is considered to have formed within the supercontinent Rodinia and exposures of similar rocks, in what are now the Mackenzie Mountains of the northern Cordillera, suggest that it extended for more than 1000 km to the southwest (Long et al., 2008; Rainbird et al., 1995a). The sedimentary succession is intercalated with mafic sills of the ca. 720 Ma Franklin gneiss event (Heaman et al., 1962). The silt area of variable thickness up to 100 m, but most are 20–40 m thick. In many cases, individual sills extend for 20 km or more along strike with little significant change in thickness. Sills constitute anywhere from 10 to 50 per cent of the stratigraphic section. Sills of similar type and age also occur in the Copegmine Formation, Brook Inlet and Duke of York Inlet to the south (Rainbird et al., 1996b; Steinhilb et al., 2004) and, in general, geochemically similar intrusions and volcanic rocks associated with the Franklin event extend from Greenland to the western Yukon (Denyszyn et al., 2009; Heaman et al., 1992; Macdonald et al., 1994). The Shaler Supergroup in Minto Inlet is capped by a succession of food-basalt flows and interflow sedimentary rocks (Nukaskiak Fm.) more than 1 km thick, which are the calcareous equivalent of the Silesites (Burgar, 1979; Jefferson et al., 1995). Rare north-west-trending dykes are interpreted to have intruded along siph-magmatic normal faults, to feed sills and possibly the food basalt (Bédard et al., 2012). Three magma populations are identified in the lavas, which have composites in the different sill subtypes. The oldest sills and corresponding basal lavas are enriched in incompatible trace elements and may have differentiated bases. Younger diabasic sills correspond to the major sheet-flow units of the lava succession. Basal strata of the Shaler Supergroup (Duke Group) are exposed east of Minto Inlet, near Haldy Bay. Here they unconformably overlie Palaeoproterozoic Palaeoproterozoic sedimentary rocks, which, in turn, unconformably overlie Archean granitic rocks (Campbell, 1981; Rainbird et al., 1994). The irregular edge of Minto Inlet is defined by an erosional unconformity that separates the Neoproterozoic rocks from Lower Cambrian sandstones and siltstones that pass upward into a thick succession of mainly dolomitic carbonate rocks, ranging in age from Cambrian to Devonian (Thompson and Topp, 1992). Structurally, the Minto Inlet is relatively simple, composed of the open, north-trending Haldy Island Syncline and a smaller Vialkye Bay anticline to the northwest. Beds typically dip no more than 10° and there is generally no penetrative cleavage or other significant outcrop-scale fabric. The origin of the folding is unknown but it occurred after deposition of the early Neoproterozoic rocks and before uplift, erosion and deposition of overlying lower Cambrian siliclastic rocks, which are not related. All rocks are dissected by east-northeast to east-trending faults that form a north and graben system with up to 200 m of stratigraphic separation on individual faults. The zone of faulting is about 100 km wide and stretches from the head of Minto Inlet in the west to Wynniatt Bay in the east and is spectacularly imaged as prominent lineaments on recently published aeromagnetic maps (e.g. Kiss and Oneschuk 2010).

NTS 87-H12 is underlain by stratigraphic units from the Jago Bay, Minto Inlet and Wynniatt Formations of the Shaler Supergroup. Together with diabase sills, the Neoproterozoic strata comprise the gently south-east-dipping northern limb of the Haldy Island Syncline. The outcrop pattern of bedrock units on this map sheet is strongly influenced by evenly spaced, west and graben, with variable along-strike separation and are true akin to strike slays. Protruding, generally north-south rotated dikes of various fales, an exemplified along the Qiyukluak River (UTM 50 9697 E, 7644450N). There are no radiometric dated sections in the Shaler Supergroup. The Jago Bay Formation consists of massive, laminated white and orange-brown siliceous sandstone, intercalated with reddish, oxidatively alteration that is interpreted to reflect chemical weathering and development of incipient rippen. This would likely have occurred in the ca. 200 Ma interval between intrusion of the Franklin diabase and deposition of the unconformably overlying Cambrian sandstones. Rocks of the Jago Bay Formation are mainly dolomitic, deposited in shallow subtidal to intertidal environments and are best characterized by their yellowish-weathering. Stratigraphically overlying strata of the Minto Inlet Formation are mainly coarsely, crumbly weathering, laminated white and orange-brown siliceous sandstone, intercalated with reddish, oxidatively alteration and nodular gypsum. The dark gray limestone member which outcrops close to Adakialuk Lake, should be distinguished from the overlying Wynniatt Formation. Strata of the lower carbonate member of the Wynniatt Formation outcrop sporadically in the south-east part of the map area. A topographically prominent, apparently cross-cutting, irregular intrusion trends north-northwest from the northern shore of Minto Inlet on the west side of the map area (all referred to as King Inlet). This feature may represent a broad magmatic feeder because it cuts across strata and parts of the body appear to feed laterally into more typical, continuous sills. In the northern half of the map area, the Cambrian sedimentary rocks are generally exposed in the lower class unit and overlying thin dolomite unit, which features distinctive orange-weathering. An excellent section is exposed in the time Qiyukluak River (UTM 50 9697 E, 7644450N). This section illustrates the sedimentary sequence of Cambrian strata and occurred at the same time as folding of the Shaler Supergroup, but reactivation during Cambrian time is suggested by thickness variations between adjacent fault panels. The sub-Cambrian unconformity and Cambrian strata are offset across faults indicating significant post-Cambrian, probably normal, tectonics.

Abstract

Le feuillet NTS 87-H12 expose des roches de la Shaler Supergroup, les formations de Minto Inlet et les roches de la Shaler Supergroup. Ensemble avec les sillons diabasiques, ces roches constituent le membre nord-est de la Sienne. Les roches sont composées de bancs de sable à grains fins et de grès, avec des argiles et des grès à grains fins. Les roches sont principalement constituées de sables à grains fins et de grès à grains fins, avec des argiles et des grès à grains fins. Les roches sont principalement constituées de sables à grains fins et de grès à grains fins, avec des argiles et des grès à grains fins. Les roches sont principalement constituées de sables à grains fins et de grès à grains fins, avec des argiles et des grès à grains fins.

Résumé

Le feuillet NTS 87-H12 expose des roches de la Shaler Supergroup, les formations de Minto Inlet et les roches de la Shaler Supergroup. Ensemble avec les sillons diabasiques, ces roches constituent le membre nord-est de la Sienne. Les roches sont composées de bancs de sable à grains fins et de grès, avec des argiles et des grès à grains fins. Les roches sont principalement constituées de sables à grains fins et de grès à grains fins, avec des argiles et des grès à grains fins. Les roches sont principalement constituées de sables à grains fins et de grès à grains fins, avec des argiles et des grès à grains fins.

#7015	#1016	#1015	#7014
#1015	CGM 103	CGM 82	#7014
#1014	CGM 59		#7013
#1013	CGM 104	CGM 155	#7012

Cover illustration
Postglacial raised beaches beneath thick Franklin diabase sill near Qiyukluak, Northwest Territories. Photograph by R.H. Rainbird, 2012/08.

National Topographic System reference and index to adjoining Geological Survey of Canada maps

Catalogue No. M183-182-2014E-PDF
ISBN 978-1-100-23174-7
doi:10.4069/290460

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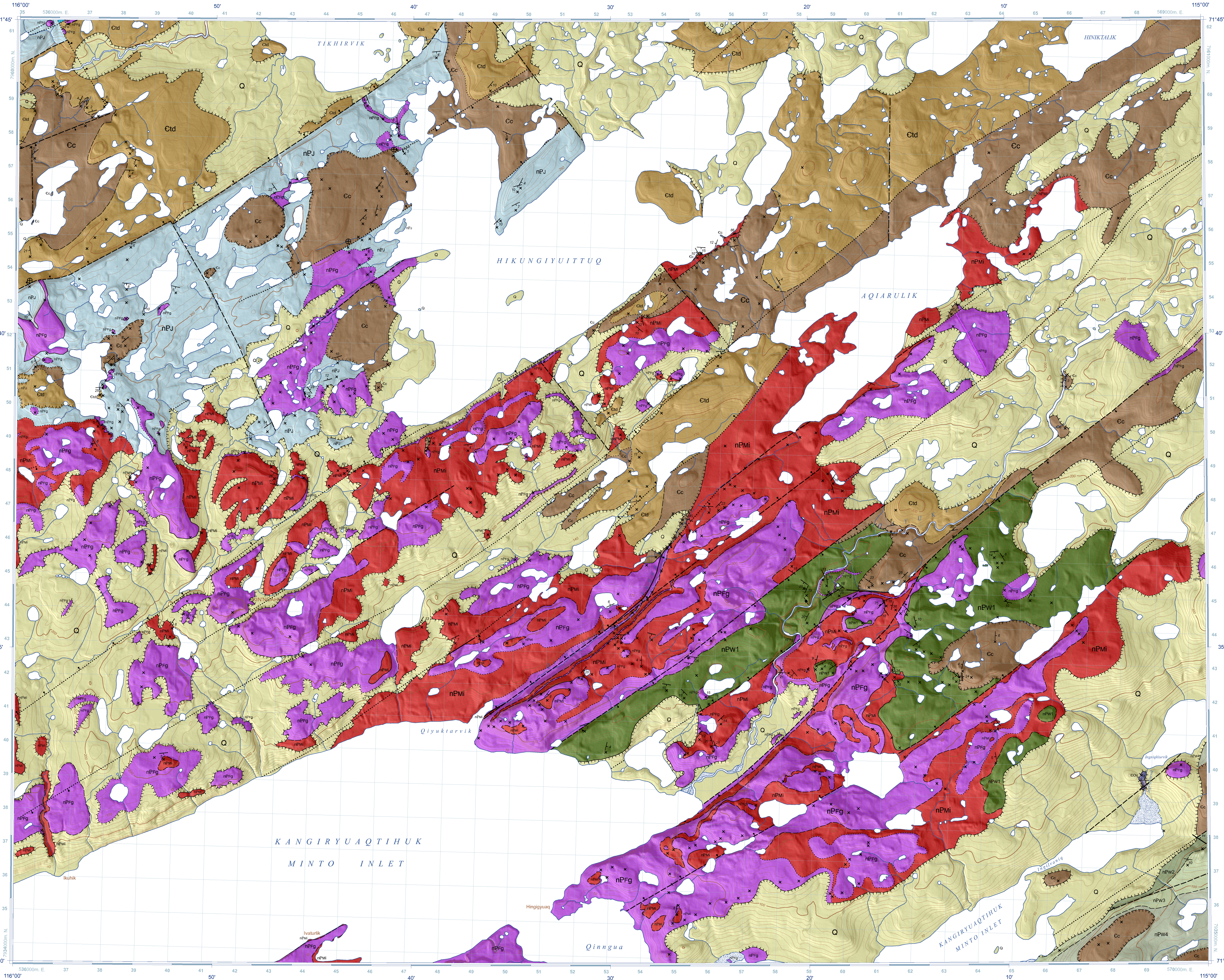
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CANADIAN GEOSCIENCE MAP 82

GEOLOGY

KANGIRYUAQTIHUK / MINTO INLET

Victoria Island, Northwest Territories
1:50 000



CENOZOIC

Q Quaternary sediments.

CAMBRO-ORDOVICIAN

C/O Victoria Island Formation: Light grey to almost white weathering, fine to coarse crystalline fabric; dolomitic dolomite that is widespread throughout northwestern Victoria Island. Locally preserved primary structures include horizontal, cross-bedded calcarenites, stromatolitic dolomites, brownish dolomitic siltstones, micritic lamination, and intertuffaceous conglomerate. Fossils are rare but include indistinct gastropods. Silicification is widespread in the upper half of the unit. Chert occurs as prominent white-weathering beds and nodules, 5 to 60 cm thick, and composed of microcrystalline chert or silicified stromatolites. Some vugs contain crystalline quartz. Depositional environment is interpreted as a broad, shallow carbonate platform.

CAMBRIAN

C/C Tan dolomite unit: Light brown dolomitic to dolarenite. Stromatolite beds are locally well developed and together with metre-scale cross-stratification suggest a shallow marine setting. Although fossils were recovered from this unit, the lower contact is gradational with mudstones that contain Early Cambrian trilobites. Thickness is 30–45 m.

LOWER CAMBRIAN

C/C Clastic unit: Red-brown to orange weathering fine to coarse grained quartz arenite and mudstone at the base of the Palaeozoic succession. Sedimentary structures are lamination, wavy and curved rippled, and 10 cm to 2 m thick cross-stratified beds. Recrystallization surfaces and porosity with rounded tops indicate an influence by tidal currents. Depositional environment is considered to be shallow marine. Mudstone contains Lower Cambrian trilobites. Distribution and thickness are variable; thickness ranges from 0 to 90 m.

NEOPROTEROZOIC

n/P Framkin intrusions: Typically massive, laterally extensive, diabasic sills with columnar jointing ($5\text{--}15\text{ m}$ thick, rarely up to 100 m). Some sills are composite with internal intrusive contacts. Two types: 1) An older, more primitive type is commonly layered, with microdiabasic lower and upper border zones and olivine-enriched basalt cumulate (olivine gabbro or feldspathic wellite) that may be capped by a 10–20 m felspathic pyroxenite (magnesian) horizon containing abundant olivine. 2) Youner, less primitive type is commonly layered, with microdiabasic lower and upper border zones and olivine-enriched basalt cumulate (olivine gabbro or feldspathic wellite) that may be capped by a 10–20 m felspathic pyroxenite (magnesian) horizon containing abundant olivine. 2) Youner, less primitive type is commonly layered, with microdiabasic lower and upper border zones and olivine-enriched basalt cumulate (olivine gabbro or feldspathic wellite) that may be capped by a 10–20 m felspathic pyroxenite (magnesian) horizon containing abundant olivine.

n/PW1 Wynniatt Formation (n/PW1–n/PW4): Base characterized by distinctive nodular, black calcarenite shale, overlain by thin, rhythmically bedded and normally graded, quartz-sandy calcarenite. Upper, metre-scale alternations of stromatolitic dolomites and cross-bedded intertidal granitoids. Local hornington cross-bedded quartz arenite and microbially laminated lime mudstone. Chert is common. Approximately 300 m thick.

n/PW3 Stromatolitic carbonate member: Stromatolitic dolomite with built-ups that have local sigmoidal and/or vertical, main building contains oncos up to 20 cm. Interbedded intratidal granitoids with rips and scours. Multibedded and multibanded structure. Parallel or marginally laminated dolomite with mudcracks, and leopore structures. Sharp, erosive upper contact. Approximately 160 m thick.

n/PW2 Black shale member: Dark grey parallel-laminated siltstone and silty mudstone with discontinuity to continuous beds of ripple-logged quartz arenite common near top. Structures include flute and gutter casts, ball and pillow structures, channels and fill structures, and climbing ripples in siltstone. Carbonate nodules and pyrite are present throughout. Up to approximately 200 m thick.

n/PW1 Lower carbonate member: Lower parallel-laminated dolomite and calcillite with mudcracks, leopore structures and chert, minor carbonate mudstone. Middle hummocky cross-stratified dolomite and calcillite with scours. Rip-ups and molar-tooth structure, stromatolitic, intratidal dolomites. Upper stromatolitic dolomite, intratidal dolomites and carbonaceous siltstone capped by regionally extensive orange-weathering stromatolitic dolomites, up to 10 m thick. Approximately 120 m thick.

n/MI Minto Inlet Formations: Four informal members in ascending stratigraphic order: maroon-green gypsiferous siltstone-calcillite; dark grey limestone; buff and grey argillite, reef gypsiferous siltstone and buff to grey calcillite. Chickiwere, nodular argillite and cross-cutting subaqueous ventails common in upper evaporite as are up to 2 m thick beds of crystalline gypsum. Carbonate lithoclasts include dark grey to buff-grey laminated to thin-bedded dolomite with molar-tooth and fine-grained dolarenite with hummocky cross-stratification. Approximately 200 m thick.

n/PJ Rippenak Point Group (n/RP): Lower predominantly wavy-bedded calcillite with black shale partings, molar-tooth structure, minor beds of cross-bedded quartz-sandy calcarenite and stromatolitic limestone. Upper is mainly laminated calcillite with subordinate granitoids, calcareous sandstone, and gypsum near upper gradational contact with Minto Inlet Formation. Approximately 200 m thick.

n/RP Boon Inlet Formation: Rhythmically laminated to finely bedded, dark-grey with dolomite-calcillite and light-grey dolomitic-calcillite with common molar-tooth structure, gutters, scours, tool marks, convolute bedding and other loading features. Interbedded with trough cross-stratified to hummocky cross-stratified dolarenite/calcarenite, including tabular siltite arenite (granitoid) in the type area, near Boon Inlet; the upper part of the formation is characterized by a thick, tabular, stromatolitic reef complex composed of strongly dolomitic, digitate stromatolites interbedded with dolomitic granitoids. Only the upper half (<math><250\text{ m}</math>) of the formation is exposed in the map area.

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Station location
x Ground observation

Planar structure
Fault
Jointing
Inclined
Vertical
Bedding
Inclined
Inclined
Cross-bed forest

Linear structure
Sedimentary structure
Sub-arcuate or slump
Fold
Fold axis
Anticline with plunge

CONTACT: depositional-depositional-conformable or intrusive
Defined
Approximate
Inferred

CONTACT: depositional-unconformable
Defined
Approximate
Inferred

Fault, generic, steep dip
Defined
Approximate
Inferred

Fault, normal, upright
Defined
Approximate
Inferred

Structural lineament
Defined
Approximate
Inferred

Recommended citation
Rainbird, R.H., Bédard, J.H., Dewing, K., and Haldy, T., 2014. Geology, Kangiryuaqtihuk / Minto Inlet, Victoria Island, Northwest Territories. Geological Survey of Canada, Canadian Geoscience Map 82 (preliminary), scale: 1:50,000. doi:10.4069/290460

Canadian Geoscience Maps

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Geomatics by É. Grand
Cartography by N. O'Hare

Authors: R.H. Rainbird, J.H. Bédard, K. Dewing, and T. Haldy
Geology by R.H. Rainbird, J.H. Bédard, K. Dewing, and T. Haldy, 2011
Geomatics by É. Grand
Cartography by N. O'Hare

Initiative of the Geological Survey of Canada, conducted under the auspices of the Victoria Island PGE/Basé Métales project, as part of Natural Resources Canada's Geo-mapping for Energy and Minerals (GEM) program.
Map projection Universal Transverse Mercator, zone 11, North American Datum 1983

CANADIAN GEOSCIENCE MAP 82
GEOLOGY
KANGIRYUAQTIHUK / MINTO INLET
Victoria Island, Northwest Territories
1:50 000
1 0 1 2 3 4 km

Base map at the scale of 1:50 000 from Natural Resources Canada, with modifications. Elevations in metres above mean sea level.
Shaded relief image derived from the digital elevation model supplied by GeoBase.
Illumination: azimuth 225°, altitude 45°, vertical factor 1x
Proximity to the North Magnetic Pole causes the magnetic compass to be erratic in this area.
Magnetic declination 2014, 19°35'E, decreasing 55' annually.

The Geological Survey of Canada welcomes corrections or additional information from users.
Data may include additional observations not portrayed on this map. See documentation accompanying the data.
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Preliminary publications in this series have not been scientifically edited.

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GEOLOGY
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