IMPLICATIONS FOR VMS PROPECTIVITY IN THE FLIN FLON BLOCK FROM SULPHIDE STUDIES AND MAPPING OF PALEO-HYDROTHERMAL CORRIDORS

INTRODUCTION

Mineralogical, chemical and isotopic studies of the sulphide orebodies and hangingwall strata to the five Flin Flon block VMS deposits are used to constrain ore deposition and develop a model for detecting buried mineralization. Here we report an extensive suite of major and trace elements for the precious metal-bearing deposits, define key mineral and chemical vectors, their host minerals and illustrate the remobilization of elements within hydrothermal corridors in the hanging-wall mafic volcanic strata.



Figure 1: General geology and sample location map of the Flin Flon mining camp odified from Simard et al., 2010). The distribution of mafic tuffaceous sedimentary rocks in the Flin Flon and Louis Formations and the five Fe-Ti-P basalt occurrences in the Hidden Formation are shown. Sulphide deposits are situated proximal to felsic volcanic centres near the top of the Flin Flon Formation and buried by the Hidden Formation basaltic lavas (contact thick red line). Generalized vertically projected outlines of the principal VMS bodies (Callinan, 777 and Main Flin Flon) are indicated by dashed lines, A-A' South Main to Sipple corridor.

Geochemical classification readily identified Fe-Ti-P basaltic andesite in the lower metamorphic grade rocks south of the Railway fault (Fig 2). Fe-Ti-P basaltic andesites are discontinuous over a 6 km strike length shown by the 1920 Flin Flon Pit and Phantom occurrences and dykes in the Reservoir and South Main localities (Fig. 1). South of the RWF stilpnomelane is the key Fe-rich mineral in these basaltic andesites whereas, to the north, the presence of asicular tschermakitic horneblende (Fig 10) readily distinguishes the 1920 unit in the field.

SULPHIDE STUDIES OF THE FLIN FLON BLOCK VMS DEPOSITS

Table 1: Major an	d trace	e elen	nent co	ompo	sitions	of su	phide	s fron	n five I	Flin Fl	on blo	ock VI	//S de	posite	6
	FE.T	CU	ZN	РЬ	Au	Ag	S.T	SIO2	Al ₂ O ₃	MgO	CaO	Na ₂ O	K ₂ O	CO2.T	REE.T
	%	%	%	%	ppb	ppm	%	%	%	%	%	%	%	%	ppm
Western FF block															
Flin Flon (13)	32.42	5.47	10.85	0.80	27287	62	39.45	4.82	0.51	0.71	1.09	0.03	0.02	1.77	8.76
777 (10)	26.90	4.71	15.22	0.57	4064	102	34.06	8.38	1.96	0.75	1.36	0.40	0.26	1.51	10.99
Callinan (10)	31.22	6.14	9.87	0.20	5377	113	36.49	7.24	1.38	0.89	2.47	0.25	0.21	2.25	8.19
Eastern FF block															
Schist Lake (11)	28.58	2.84	14.91	0.81	3501	118	37.99	7.91	0.58	0.36	0.55	0.01	0.01	0.95	2.40
Mandy (13)	26.84	16.21	12.70	0.32	3625	55	34.48	2.33	0.87	0.50	0.37	0.00	0.00	0.58	8.70
	As	Ba	Bi	Cd	60	Cr	60	На	Tn	Mo	Ni	Sh	Se	Sn	Te
	nnm	nnm	nnm	nnm	nnm	nnm	nnm	nnm	nnm	nnm	nnm	nnm		nnm	
Western FF block	PP	PP'''	PP···	PP	PPm	PP'''	PPm	PP'''	PP'''	PPm	25.	PP'''	PP'''	PP'''	PP
	2407	1.4								40	1/	110	124	11	ЛЛ
Flin Flon (13)	2400	14	4	300	243	44	12	8	6	13	16	110	164	11	
777 (10)	2486 481	24	4	300 372	243 195	44 37	12 21	8 4	6 2	13 11	16	158	90	16	20
Fin Flon (13) 777 (10) Callinan (10)	2486 481 4355	24 618	4 2 2	300 372 260	243 195 432	44 37 36	12 21 8	8 4 7	6 2 3	13 11 21	16 18 6	110 158 54	90 102	16 6	20 44
Fin Flon (13) 777 (10) Callinan (10) Eastern FF block	2486 481 4355	24 618	4 2 2	300 372 260	243 195 432	44 37 36	12 21 8	8 4 7	6 2 3	13 11 21	16 18 6	158 54	90 102	16 6	20 44
Fin Flon (13) 777 (10) Callinan (10) Eastern FF block Schist Lake (11)	2486 481 4355 5140	24 618 4	4 2 2 1	300 372 260 309	243 195 432 273	44 37 36 39	12 21 8 22	8 4 7 13	6 2 3 2	13 11 21 4	16 18 6	110 158 54 285	90 102 145	16 6 15	20 44 20

Figure Tectonize massive sulfide ore type from the Schist and Mand VMS deposits, eastern F Flon block.



00GIA-533B. McKinnon ~1920. SP1134-SCHL.

Chalcopyrite-rich stringer ore chlorite-chalcopyrite-pyrite, minor biotite porphyroblasts, Kfeldspar and trace pyrrhotite - magnetite - cubanite sphalerite and biertherine. Sample 07AV-FF54.

B) Massive chalcopyrite - pyrrhotite ore consisting of quartz pyrrhotite - chalcopyrite, minor magnetite, pyrite and trace cubanite - sphaleritesericite - carbonate and Fe-rich minerals berthierine and stilpnomelane. Sample 07AVFF63. C) Massive pyrite-sphalerite banded ore with minor chalcopyrite and pyrrhotite intergrown with sphalerite

interstitially to pyrite. Ore contains trace arsenopyrite magnetite - carbonate - stilpnomelane and chlorite. Sample 06AV-251 D) Fine grained massive pyrite-sphaleritechalcopyrite ore

from the 777 deposit, sample 06AV-254. Typical mineral assemblages include massive and granular pyrite, minor sphalerite - pyrrhotite - chalcopyrite and trace cubanite galena-tetrahedrite-arsenopyrite-antimony and stilpnomelane sericite or chlorite E) Fragmental ore, altered quartz phyric rhyolite fragments

with semi-massive chalcopyrite/sphalerite in matrix, from 1600L- #5lens, Callinan Cu-Zn-(Pb, Au) deposit. Sample 00GIA-CALL-531.2A2

F) Semi-massive pyrite - chalcopyrite ore hosted in biotite rich, rhyolite breccia. Fragmental ore sample 06AV-504. G) Semi-massive pyrite - sphalerite - chalcopyrite ore from the

Callinan deposit, sample 06AV- 239, drillhole 4Q- 43, 2413 ft. H) Hand sample of massive sulphide from discovery outcrop (Figure 7A), Flin Flon Cu-Zn- (Pb, Au) deposit collected by A.T. McKinnon in 1920. Chalcopyrite - pyrite/Mg-chlorite fragmental ore with trace tellurides, and pink siliceous, fault gouge. Sample ooGIA-FL-10088 3 (Archive no. GSC-EA-18.3).



Figure 3: Typical metamorphosed ore types from the Flin Flon-777-Callinan orebodies, western Flin Flon block.

ORE MINERALOGY AND MINERAL CHEMISTRY

The Flin Flon massive sulphide ore has not been studied in over 60 years in a comprehensive mineralogical study (Table 2). Applied mineralogy studies were undertaken at the Callinan deposit in the 1980's (Healy and Petruk, 1990).

Detailed sulphide mineralogy of over 50 ore samples representative of the Flin Flon-777 Callinan deposits and 25 samples from the Schist and Mandy deposits was completed The metamorphosed mineral assemblage of the Flin Flon VMS system contains over comprising sphalerite, chalcopyrite, magnetit and pyrite and the trace minerals bein dominantly antimony-bearing (Table 3). The Schist and Mandy deposit ore samples are severely deformed mylonites with S-C fabrics (Fig. 4).



Important trace element associations for vectoring

Au-bearing minerals Native gold (Au,Ag) Electrum (Ag,Au,Sb,Hg) Sylvanite (Au,Ag)₂ Te₄ Allargentum (Ag,Au)5 Sb

Ag-bearing minerals Ag-chalcopyrite Ag-CuFeS₂ Freibergite (Ag,Cu,Fe)₁₂Sb₄S Acanthite Ag₂S

Dyscrasite Ag₃Sb Hessite Ag2Te

Co-As-bearing minerals Arsenopyrite Fe(Co)AsS

Cobaltite CoAsS Mattagamite Co₂Te

Antimony (analyzed by ICP-MS) is a key exploration vector for these Flin Flon VMS deposits



Figure 5: Sphalerite compositions of five Flin Flon block VMS deposits and the Lakeview showing. Mineral chemistry data from 120 analyses from this study including 68 analyses from the Flin Flor main deposit and 33 analyses from the Schist deposit from Bristol (1974).

GOLD IN THE 5 FLIN FLON BLOCK VMS DEPOSITS



Figure 6: Global VMS deposits showing the Au-rich VMS deposits of the Flin Flon domain based on tonnage (Flin Flon-777,) or average gold grade (Vamp, Photo Lake), Lalor not shown although preliminary calculations suggest Au> 1 Moz (modified after Mercier-Langevin et al., 2011).

Mineral chemistry of gold-electrum in five VMS deposits

Gold was found in the Flin Flon depositon the 390mL pyrite-pyrrhotite ore, as well as different orebodies in the 777 deposit (Augo, Ag_{10}) with little to no mercury or antimony impurities. Electrum in the Schist and Mandy deposits are intermediate in Ag:Au ratios and also have few deleterious trace elements whereas Sb-Hg bearing electrum (Ag_{75} , Au_{20}) is common in the Callinan and 777 deposits.

Mineralogic repository of mercury in the Flin Flon ores: Separate Hg-minerals in the ores have not been found however gold and electrum contain 0.9 - 3.9% Hg. Mercury in sphalerite is very low in the five VMS deposits (<0.1 % Hg).





A) Tectonized massive ore composed of distinct bands of granular pyrite-sphalerite, chalcopyrite-sphalerite and chloritized wallrock, Mandy Cu-Zn-(Au) deposit. Sample B) Example of tectonized massive chalcopyrite-pyrite ore with minor wisps of folded phalerite Mandy Cu-Zn-(Au) deposit.. GSC archives sample EA73-11 collected by A.T. C) Example of mylonitized massive sulphide. massive banded Cu ore with porphyroclasts and discontinuous streaks of sphalerite defining the tectonic fabric Mandy Cu-Zn-(Au) deposit. GSC archives sample EA73-2 collected by A.T. McKinnon ~1920. D) Stringer ore containing chalcopyrite and sphalerite in a chloritized felsic breccia from then Mandy Cu-Zn-(Au) deposit 00GIA-533C E) Strongly deformed massive fine-grained pyrite with discontinuous wispy sphalerite bands cut by late quartz-chalcopyrite veins Schist Lake Cu-Zn-(Pb) deposit. Sample F) Massive milled thinly-banded sphalerite/pyrite ore with trace chalcopyrite / pyrite / galena Schist Lake Cu-Zn-(Pb) deposit. Sheared and transposed, Sample SP1138-SCHL. G) Massive pyrite-sphalerite ore with quartzose augens and chalcopyrite in the pressure shadows, Schist Lake Cu-Zn-(Pb) deposit. Sample 06AV-534. H) Tectonically banded sphalerite-pyrite ore in sharp contact with massive pyrite ore, Schist Lake Cu-Zn-(Pb) deposit. Sample 06AV-536.

<u>Sb- bearing minerals</u>
Tetrahedrite (Cu,Fe) ₁₂ Sb ₄ S4 ₁₃ Native antimony Sb Valentinite Sb ₂ O ₃ Gudmundite FeSbS Ullmannite NiSbS Breithauptite NiSb Boulangerite Pb ₅ Sb ₄ S ₁₁
Pb-bearing tellerides and selenides

Altaite PbTe Rucklidgeite (Bi,Pb)₃Te₄ Clausthalite PbSe



Table 4: Tetrahedrite-freibergite compositions in Flin Flon block VMS deposit Callinan Mandy Schis n=13 n=3 n=8 n=14 6.27 5.60 5.9 6.58 1.43 1.32 0.4 16.56 11.3 26.06 27.64 0.20 0.0 23.23 23.43 23.96 23.63

Tetrahedrite has long been known as a minor component of the Flin Flon ore assemblage (Brownell and Kinkell, 1935), identified in the Callinan ores (Healy and Petruk, 1990) and now the 777, Schist and Mandy deposits (Tables 2-4). This antimony rich (21.2-32.9 wt%) mineral found in all five deposits is generally Ag-rich (1.7-18.7 wt. %) and probably accounts for most of the Sb and Ag in the ore. Tetrahedrite hosts Zn (1-3 wt%) in most deposits, Ni (2.28 wt %) at 777 and rare arsenic (<4.6 wt%) at the Mandy deposit.

The sphalerites from the five Flin Flon block deposits display a broad range of Fe contents whose average deposit compositions form two distinct clusters reflecting the deposits in the western FF block and eastern FF block. The Lakeview showing sphalerite compositions compare favourably with the western FF block orebodies and not Schist and Mandy (Fig. 5).

HIGHLIGHTS OF MINERAL-CHEMICAL AND OXYGEN ISOTOPE **MAPPING OF PALEO-HYDROTHERMAL CORRIDORS**





of the Flin Flon area.

() Distribution of Fe and Mg-rich chlorite in the Flin Flon main camp determined by microprobe analyses. Fe-rich chlorite is present in the mafic rocks above the three deposits associated with mafic tuffs roximal to Fe-Ti-P basalt including from north to south, the 1920 unit and the Reservoir. South Main and Phantom Lake "icelandite" occurrences. Fe-rich chlorite occurs in the Millrock member at South Main and Phantom Lake.

B) Chlorite mineral abundances were determined by Reitveld analyses of x-ray diffraction patterns of volcanic and altered rocks in

the hanging wall sequences. Abundant chlorite is, present in the mafic rocks south of the open pit (F), locally abundant in southwestern exposure of the Hidden Formation. Approximate positions of previously mapped metamorphic isograds are shown M1- hornblendein, M2- actinolite-out, M3- epidote + chlorite-out (Digel and Gordon, 1995).

C) Distribution of biotite in the Flin Flon camp. The sporadic and irregular occurrence of biotite throughout the area emphasizes the complexities of hydrothermal gradients and drawing metamorphic biotite-in isograds in VMS hosting terranes. The abundance of biotite in the hanging wall strata is generally < 5 wt % throughout the Flin Flon camp with anomalous abundances and Fe-rich mineral compositions restricted to the hanging wall above the 3 Flin Flon deposits Anomalously high amounts of biotite (10-30 wt. %) occurs locally above the Flin Flon main mine and locally anomalous (10-15 wt. %) annite above the Callinan and Flin Flon deposits and south near Phantom Lake (Ames et al., 2011).

D) Histogram of (Fe/Fe+Mg) in biotite and chlorite showing similar trends between the tw phyllosilicate minerals in the Flin Flon camp. Flin Flon chlorite compositions are compared to the ominant chlorite compositions in the Paleoproterozoic Kristenberg area of the

Skellefte district and the Archean Noranda camp (Hannington et al., 2003).

Figure 7: Textural features of gold-bearing assemblages in the three Western Flin Flon block VMS deposits A) Gold intergrown with sylvanite (sy) interstitial to coarse pyrite (py) with pyrrhotite (po) and altered pyrite (upper right). Flin Flon main deposit 390 level, sample FL D87C (3a), long edge 0.25 mm. Note the stronger yellow

tint in the colour of gold compared to electrum in B and C reflecting the higher Ag content in electrum. B) Electrum (el) and allargentum (al) in pyrrhotite intergrown with magnetite at the 777 deposit. Sample 06AV-254C (5), long edge ca. 0.15 mm.

C) Electrum (el) in galena (gn) which are interstitial to coarse pyrite porphyroblasts (py) in sphalerite (sp) Sample (06AV-248A), 777 deposit.

D) Two gold grains at either side of pyrite (py) in chalcopyrite intergrown with sphalerite (py), hessite (he) and rucklidgeite ru=(Bi,Pb)3Te4). Sample TSEV 534.3B, long edge 0.12 mm 777

E) Electrum (el) intergrown unidentified Ag-Au-Sb mineral (arrow), willyamite (wi), pyrite (py), pyrrhotite chalcopyrite (cpy). Long edge of photo 0.20 mm. reflected light. ooGIA-531.1A1, Callinan deposit.

F) Pyrrhotite (po) mixed with gold that occurs as turguoise tarnished inclusions veining pyrite (py) in chalcopyrite (cpy). reflected light. Sample ooGIA-531.1A **Callinan deposit**

G) Acanthite (ac) grey with iridescent chalcopyrite (vellow). Acicular gangue mineral is stilpnomelane. Long edge of Sample ooGIA-531.1A1, Callinan deposit.

H) Electrum (golden) with willyamite (light blue) interstitial to pyrite (cream) intergrown with sphalerite (medium grey) and gangue (dark grey) in chalcopyrite. Long edge of photo is 0.26 mm (reflected light). Sample ooGIA-531.1A1, Callinan deposit.

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distinguishes the 1920 unit.

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Figure 8: Chorite and biotite distribution and mineral chemistry of volcanic sedimentary rocks and alteration in the hanging wall rocks



Mapping syn-volcanic paleo-hydrothermal corridors

Epidote-guartz, silicification, chlorite ±magnetit

Phyllosilicates (biotite, chlorite, stilpnomelane)

Volcanology - synvolcanic structures, dyke swarms,

sulfide, actinolite-epidote, feldspar-quartz

neral chemistry and abundance important

Epidote of clinozoisite composition:

subsidence structures, permeability corridors.

Fe-poor Al-enriched

Useful vectors for Flin Flon VMS camp:

Sb, Hg, δ^{18} O geochemistry

/drothermal alteration

Geochemical tools

Figure 10: A) Regional distribution of the Fe-rich assemblages with ferro-tschermakite, stilpnomelane he low temperature post-metamorphi alteration mineral corrensite. Similar to biotite and :hlorite the amphibole chemistry (Fe/Fe+Mg) is mor sensitive to upflow zones than abundance. Two main zones are marked by Fe- rich amphibole 1) extends from the Millrock Formation tuffaceous sediments i the footwall to the South Main-Sipple Hill corrido above in hanging wall basalt, epidosite an interpillow material and in the 2) Fe-rich amphibole is common along the TC paleofissure (Figure 1).

B) Amphibole mineral compositions in the western Flin Flon block

South of the RWF, ~20% stilpnomelane is the key Fe-rich mineral whereas to the north, the presence of ascicular tschermakitic horneblende readily



Figure 9: Epidote mineral compositions and distribution stratigraphically above the Flin Flon VMS deposits. A) Bimodal plot of Fe +3 versus Al (vi) in epidote in the hanging-wall strata above the Flin Flon-T-C VMS deposits. Generally, the Fe-Al substitution in epidote is a function of temperature and redox conditions with Fe-poor, Al-rich clinozoisite compositions indicative of higher temperature and reduced conditions and Fe-rich, Al-poor epidote indicative of low temperature, oxidizing conditions (Gillis and Thompson, 1994) 3) Distribution of epidote-clinozoisite based on microprobe analyses and reitveld modelling of XRD. Epidote

of clinozoisite composition distinguish zones of anomalous magmatic activity in the FTC dyke swarm, synvolcanic fault in the Phantom area as well as highlighting peperitic tuffaceous rocks at the base of the Hidden and Louis Formations.

) Amount of clinozoisite in the Flin Flon block. Fe-poor clinozoisite is most abundant in the high heat flow corridor north of the Railway Fault. The relative amount of clinozoisite is similar in the cauldron south of RWF, out of the cauldron and at the base of the Louis Formation (35-40%) however epidote dominates above the proposed cauldron wall in the South Main-Sipple Hill corridor (~ 20%).



Figure 11: Simplified geologic map of principal hanging wall lithologies, with locations and generalized whole-rock oxygen isotope compositions of samples from both the hanging wall and footwall with respect to the mine horizon (heavy line). Inset histogram plots illustrate the distributions of whole-rock d18O values for felsic and mafic rocks from the hanging wall (low- and high-temperature and "fresh" rock labels are discussed in the text). The approximate location of slightly generalized, vertically projected outlines of the principal VMS bodies (Callinan, 777 and Main Flin Flon) are indicated by dashed lines, in addition to the location of an area suggested for possible future exploration.

Large hydrothermal system at predominantly low-temperatures (Zn 250°C) punctuated by higher temperature short-lived pulses. In the hanging wall, the largest zone of high-temperature alteration is situated 1100 m above the 777 deposit.

A sulphide ore geochemical database of the Flin Flon belt deposits and mineralogical databases of the metamorphic assemblages of the Hidden and Louis Formations may be found in companion publications (Ames et al., 2011; Jonasson et al., 2009).

Detailed discussion of the study is in a paper entitled "Hanging wall vectoring for buried VMS deposits, Paleoproterozoic Flin Flon Mining Camp, Manitoba, Canada" to be published in a coming special issue of Economic Geology.

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