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CANADA DEPARTMENT OF MINES

Hon. W. A. Gordon, Minister; Charles Camsell, Deputy Minister

GEOLOGICAL SURVEY

W. H. COLLINS, DIRECTOR

Summary Report, 1930, Part B

CONTENTS

	PAGE
MILK RIVER AREA AND THE RED COULÉE OIL FIELD, ALBERTA: C. S. EVANS	1
SOME CLAY DEPOSITS OF WILLOWBUNCH AREA, SASKATCHEWAN: F. H. McLEARN	31
GROUND WATER RESOURCES OF MOOSE JAW, SASKATCHEWAN: W. A. JOHNSTON AND R. T. D. WICKENDEN	50
Interglacial Deposits in Southern Saskatchewan: R. T. D. Wickenden	
Variations in Thickness of Cretaceous Formations in Southern Manitoba: R. T. D. Wickenden.	72
DEEP BORINGS IN THE PRAIRIE PROVINCES: W. A. JOHNSTON	
OTHER FIELD WORK	
INDEX	99



OTTAWA
F. A. ACLAND
PRINTER TO THE KING'S MOST EXCELLENT MAJESTY
1931

No. 2289

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Variations in Thickness of Cretaceous Formations in Southern Manitoba: R. T. D. Wickenden.	72
DEEP Borings in the Prairie Provinces: W. A. Johnston	74
OTHER FIELD WORK	98
Types	99



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Note. Part B of the Summary Report formerly included reports relating to the provinces of Manitoba, Saskatchewan, and Alberta, and to the part of the North West Territories lying north of these provinces. It now contains only reports that deal with the southern and western parts of this region underlain chiefly by Palæozoic and later formations. Part C is a new part comprising reports that relate to the northern and eastern portions of the same region, which are underlain chiefly by Precambrian formations. What has hitherto been called Part C is now Part D. It relates to the provinces of Ontario, Quebec, New Brunswick, Prince Edward Island, and Nova Scotia, and to the part of the North West Territories lying north of these provinces and east of Hudson bay.

MILK RIVER AREA AND THE RED COULÉE OIL FIELD, ALBERTA

By C. S. Evans

CONTENTS

P	PAGE
Introduction	1
General geology	2
Structure	
Well logs	25
•	
Illustrations	
Figure 1. Milk River area	2
2. Structural contour map of part of Milk River area	
2 Structural contour man of Red Coulée field	92

INTRODUCTION

The area covered by this report lies in southern Alberta, just north of the International Boundary, and includes townships 1, 2, and 3, in ranges 13 to 19 inclusive, west of the 4th meridian.

Milk river flows eastward through the middle of the area, and, with its south branch, is the only permanent stream.

The eastern part of Milk River ridge lies north of Milk river in the western half of the area, and its upper surface is 500 to 600 feet above the elevation of the river. The northeast slopes of the ridge are fairly steep down to the prairie around Warner, 200 feet below Milk river, whereas the slopes to the south are long and gentle down to the valley of Milk river. A low prolongation of Milk River ridge extends eastward from the town of Milk River into range 15. South of Milk river, in ranges 17 and 18, there is a northeast-southwest trending ridge whose highest parts reach an elevation of about 4,300 feet, about 100 feet higher than the highest part of Milk River ridge within the area.

The remainder of the area is flat or rolling prairie, into which Milk river and its seasonal tributaries, as well as no longer existing streams, have eroded steep-walled valleys and coulées along which most of the outcrops of the area occur. Many of the coulées cannot be seen from the prairie surface except at their very edges.

There is abundant morainic and glacial outwash material over the area, in which boulders and rock fragments from the Canadian Shield

predominate. No glacial material was recognized as being derived from the Rocky mountains. Most of township 1, range 18, is typical pothole country with many small lakes.

Springs are fairly common in the western part of the area, along Milk river and Red coulée, and in an unnamed coulée in the south of township 2, range 17. These springs come from Pale, Foremost, Pakowki, and Milk River strata. East of the town of Milk River only two springs of potable water were found. Alkali lakes are present in Verdigris coulée and in the neighbourhood of Red coulée and these dry up during the summer. Red creek has not been known to dry up completely, but in the summer water is present only in isolated pools along its course.

A good, gravelled highway runs south from Warner, through Milk River to Coutts at the border. A secondary road runs west from Milk River to the Cardston district, and another runs eastward from Milk River. There are dirt roads along the road allowances in the farming districts, and in summer cars may be taken over the prairie surface, so that all parts of the area are readily accessible in dry weather.

A branch line of the Canadian Pacific railway runs south from Lethbridge through Milk River to Coutts.

The field work on which this report is based was done in the summer of 1930. Township plans were used as base maps. Traverses were run by stadia and plane-table. Temporary bench marks were set by stadia and transit. Elevations were carried on all known stratigraphic horizons exposed. Messrs. J. F. Caley, B. E. Souch, and John Souch, acted as assistants.

The writer wishes to acknowledge the co-operation of Mr. C. C. Ross, Supervisory Mining Engineer, and his staff at Calgary in providing ready access to logs and samples from the wells.

The writer is indebted to Alberta Pacific Consolidated Oils, Limited, Dalco Oil and Gas Company, Limited, Dixie Oil Company, Limited, Celtic Oils, Limited, Commonwealth Petroleum, Limited, Southern Alberta Exploration Company, Limited, Vanalta Oils, Limited, and Northwest Company, Limited, for permission to publish partial or complete logs of wells.

The writer wishes to thank Mr. S. E. Slipper, of the Canadian Western Natural Gas, Light, Heat, and Power Company, Limited, for records of sections of the Pakowki from shallow wells along Milk river, without which no full section of this formation could have been given.

GENERAL GEOLOGY

The area is underlain by Upper Cretaceous strata with dips measured in a few tens of feet a mile, so that the strata appear to be horizontal. The area is mantled by morainic and outwash material and soil, and the outcrops are confined to steeper slopes, often at springs, and along the sides of steep valleys and coulées.

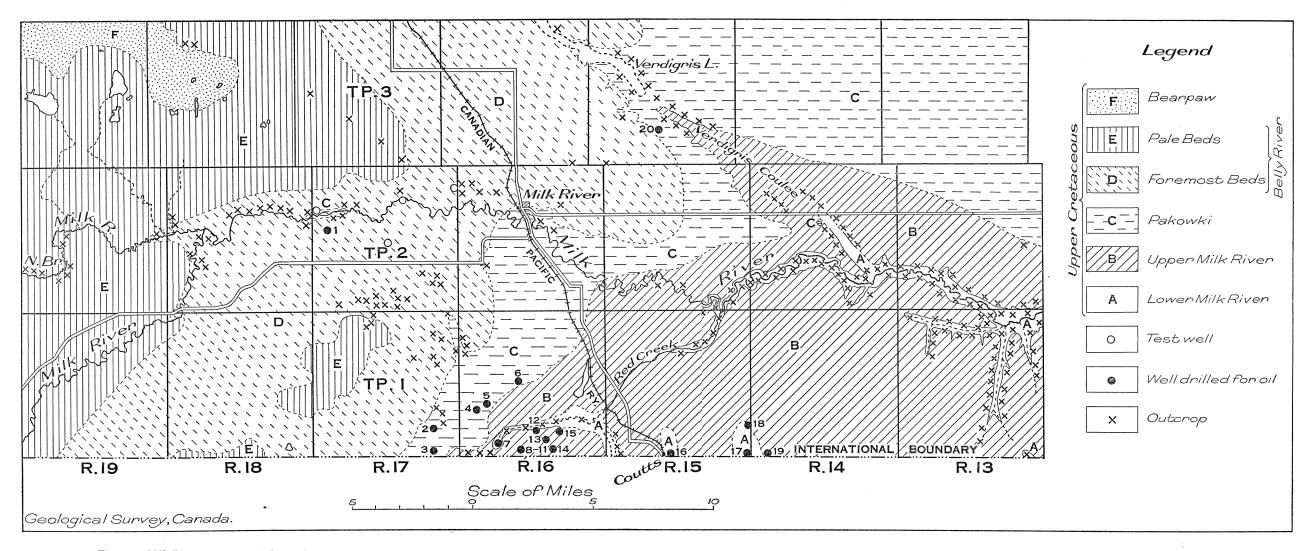


Figure 1. Milk River area, west of 4th meridian, Alberta. Oil wells: 1, Dalco No. 2; 2, Holgar; 3, Ko-Top No. 1; 4, Dixie No. 1; 5, Celtic No. 1; 6, Capitol No. 1; 7, Imperial Red Coulée No. 1; 8, Taylor No. 1; 9, Vanalta No. 2; 10, Vanalta No. 1; 11, Alberta Pacific Consolidated Red Coulée No. 1; 12, Devonshire No. 1; 13, Commonwealth Red Coulée No. 1; 14, Southern Alberta Exploration No. 1; 15, Dalco No. 1; 16, Urban; 17, Coutts-Sweet Grass; 18, Lethbridge Oils; 19, Border Oils; 20, Commonwealth Milk River No. 1.

3в

Table of Formations

	Bearpaw		No exposures found. Thought to be present on top of Milk River ridge		
		Pale Beds	Light grey, greenish grey shale and sandstone. Concretionary layers		
Upper	Belly River	Foremost Beds	Grey sandstones and shales, with darl grey and brown lignitic shales. Som lignite. Some impure bentonite Concretionary beds with Unios and Corbula. Heavy bed of medium sandstone at base		
Cretaceous	Pakowki		Grey and greenish grey sandstone and shale. Black chert pebble bed at base		
	Milk River	Upper Milk River	Fine to medium, light grey sandstone. Grey, greenish, and lignitic sandy shales. A little lignite		
	MIIK ILIVE	Lower Milk River	Heavily bedded, medium-grained, light grey sandstone. Shaly at base		
	Alberta shales		Mostly dark shales. Sandy at top and base. Bentonite, glauconite		
Lower Cretaceous	Kootenay (?)		Grey, light grey, green, and red shales. Fine and medium-grained, grey and greenish grey sandstone. Traces of coal. Some bituminous shale		
Upper Jurassic	Ellis		Grey, greenish grey, shale and cal- careous shale, with quartz and black chert grains. Sandy and limy at base. Glauconitic beds		
	Pre-Ellis		Limestone, sandy limestone, bitum- inous shale, gypsum		

Note. Strata below the Milk River are not exposed within the area.

PRE-ELLIS STRATA

Pre-Ellis strata are reached by three of the deeper wells. Summaries of parts of the logs of these wells are given below.

Commonwealth Petroleum, Limited, Milk River Well No. 1; L.S.D. 8, Sec. 9, Tp. 3, Range 15, W. 4th Mer., Elevation 3,163 Feet

Drilling depth

2,750 2,750–3,570 3,570–3,670 3,670–3,730 Base of Ellis White limestone

Missing

Dark bituminous limestone

3,730-3,750 Shaly, dark limestone, a little glauconite Sandy, light grey limestone—50 per cent very fine quartz sand 3,750–3,830 3,830–3,840

Cement?

Richly bituminous shale

3,840-3,870 3,870-4,735 Light grey and grey limestone and dolomitic limestone. Much-gypsum at top and some gypsum throughout. Samples are quite brownish, mostly due to iron rust, presumably from bit

Northwest Company, Limited, Erickson Coulée Well No. 1; L.S.D. 13, Sec. 8, Tp. 1, Range 12, W. 4th Mer., Elevation (?)

Drilling depth

2,500 2,500-3,360 Base of Ellis

White limestone, cherty from 3,330-3,360

3,360-3,450 Grey and dark grey limestone

3,450-3,530 40-50 per cent soluble in HCl, remainder very fine quartz sand

3,530-3,540 Mostly highly bituminous dark shale

3,540-3,550 Mostly white limestone, some bituminous matter

3,550-3,590 Mostly calcareous, greenish grey shale; pyrite and some fragments of anhydrite

3,590-3,650 Mostly gypsum and anhydrite

Urban Oils, Limited, Well No. 1; L.S.D. 2, Sec. 4, Tp. 1, Range 15, W. 4th Mer., Elevation about 3,450 Feet

Drilling depth

Base of Ellis

2,570 2,570-2,580 Mostly white limestone

2,580-2,611 30 per cent white limestone, 30 per cent grey, calcareous, and sandy shale, 30 per cent white, calcareous, and sandy shale; some glauconite

2,611-3,310 Mostly white limestone. Some samples missing and some represented by cement

3,310-3,390 Grey limestone

3,390-3,418 Dark grey limestone

The Erickson Coulée and Commonwealth Milk River well logs show a fairly close similarity in these deeper strata; the succession from top to bottom, in both, being:

- 1. White limestone
- 2. Dark limestone (bituminous in Commonwealth Milk River well)
- 3. Sandy, light grey limestone (50 per cent very fine quartz sand)
- 4. Richly bituminous shale
- . Gypsum and anhydrite with limestone

The depths of these zones below the base of the Ellis are:

	Commonwealth Milk River well	Erickson Coulée well
1	0 to 920 920 to 980 1,000 to 1,080 1,090 to 1,120 1,120 to –	0 to 860 860 to 950 950 to 1,030 1,030 to 1,050 1,090 to —

E. S. Perry¹ in writing on the Kevin-Sunburst field, states that there the Ellis is underlain by the Mississippian (Madison) limestone, 1,075 feet thick, and reports that fossils of Upper Devonian character were found in green shale that occurs 15 feet below black oil-shale. He also reports that Devonian fossils were blown from the Potlatch well by gas struck about 600 feet below the base of the Madison. He names the Devonian the Potlatch anhydrite formation and places its upper limit at the top of the bituminous shales and the base at the top of a light grey limestone about 940 feet below. The present writer knows of no fossil evidence indicating the age of the pre-Ellis strata in southern Alberta.

Applying this interpretation of Perry's to the Alberta field, the sandy limestones, zone 3, just above the dark bituminous shale would belong to the Mississippian and the underlying strata to the Devonian. Possibly the sandy character of the limestones of zone 3 might be regarded as a reflection of some change of conditions and thus indicate the contact of the Mississippian and Devonian strata. A more practical consideration is that the bituminous shale is readily spotted in the well samples and its top serves as a handy arbitrary horizon at which to place the contact.

Oil in commercial amounts was not obtained below the Mississippian limestone from any of the nine deep wells of the Kevin-Sunburst field² drilled to below this horizon, and no production has been encountered up to the present from the same strata in southern Alberta.

ELLIS

In the Kevin-Sunburst field, 15 miles to the south of Red Coulée field, 200 to 300 feet of shaly or limy sandstone and shaly limestone, lie above the Mississippian (Madison) limestone and are correlated with the Ellis formation of Upper Jurassic age. In the wells of the Red Coulée field, marine strata that occur between the top of the Mississippian limestone and the base of the continental deposits of Kootenay (?) age are also termed Ellis by the oil men, and this correlation, although not proved, is probably correct.

¹Perry, E. S.: State Bureau of Mines and Geology, Butte, Montana, Mem. No. 1, 2nd Edit., Sept. 1929.
²Perry, E. S.: Op. cit.

The Ellis of the Red Coulée field has a thickness of from 60 to 180 feet and consists of:

Upper part—Grey, brownish grey, and light greenish grey, slightly calcareous shales with disseminated pyrite and grains of quartz and black and red chert. Some beds hold glauconite.

Lower part—Light greenish grey, highly calcareous shales with disseminated pyrite and grains of quartz and red and black chert. Glauconite present but less prevalent than in upper part. Fine, shaly sandstone at base in one well.

Since the Ellis (and, or, Fernie) strata are regarded as favourable source rocks of oil, detailed descriptions of them as revealed in samples from the wells are given.

Commonwealth Petroleum Limited, Red Coulée Well No. 1; L.S.D. 10, Sec. 3, Tp. 1, Range 16, W. 4th Mer., Elevation 3,527 Feet

Drilling depth

2,514 Shaly sandstone and sandy shale; abundant carbonaceous remains
2,514-2,553 Medium-grained, quartz and black chert sandstone with slightly
calcareous matrix. Some siliceous grey shale. At base some
black chert grains, 5 mm. diameter

Few small fragments of coal and several microspores

Top of Ellis

2,553-2,572 Mostly light grey, sandy shale. Some medium to coarse chert and quartz sandstone. Much pyrite
2,572-2,576 70 per cent light greenish grey, calcareous 'pencil' shale

2,572-2,576
70 per cent light greenish grey, calcareous 'pencil' shale
15 per cent loose chert grains to 1.5 mm. diameter
10 per cent pyrite and marcasite

2,576-2,580 90 per cent light greenish grey, calcareous 'pencil' shale containing scattered sand grains. Pyrite

2,580-2,590 95 per cent light greenish grey, calcareous 'pencil' shale; disseminated pyrite and marcasite. Minor glauconite and medium-grained chert and quartz sandstone

2,590-2,598

90 per cent light greenish grey, calcareous 'pencil' shale; disseminated pyrite. Several thin bands of highly calcareous, fine-grained sandstone. A little medium-grained quartz and black chert sandstone. Minor glauconite. Several small, tubular structures in limestone.

2,598-2,603 90 per cent light greenish grey, finely sandy and calcareous shale, with disseminated pyrite. Minor glauconite, quartz and black chert sandstone, and highly calcareous fine sandstone

2,603-2,614 85 per cent of shale as above with 10 per cent fine, white, highly calcareous sandstone. Minor chert and pyrite

Base of Ellis

2,614 White, siliceous, dolomitic limestone

Celtic Oils, Limited, Red Coulée Well No. 1; L.S.D. 4, Sec. 17, Tp. 1, Range 16, W. 4th Mer., Elevation 3,531 Feet

Drilling depth

2,530-2,540 Light grey, dark grey, green, and red shales, with abundant carbonaceous remains

2,540-2,550 Missing

2,550-2,580 Medium-grained, quartz and black chert sandstone Coal fragments

Top of Ellis

Drilling depth

2,580-2,590 90 per cent slightly greenish grey, calcareous 'pencil' shale 10 per cent pyrite and marcasite

2,590-2,600 As above

2,600–2,600 As above, with some marcasite balls and a few grains of black chert to 7 mm. diameter

2,610-2,620 As above, with a few fragments of red shale

2,620-2,630 70 per cent slightly greenish grey, calcareous shale. 20 per cent pyrite

2,630-2,650 35 per cent chert grains, some 1 cm. diameter. A little red shale 35 per cent shale as above; 60 per cent fine, light greenish, highly calcareous sandstone. Scattered larger quartz and chert grains. Pyrite and a little glauconite

2,650-2,660 Much as above with 75 per cent of fine sandstone

Base of Ellis

Below 2.660 White limestone

Dalco Oil and Gas Company, Limited, Red Coulée Well No. 1; L.S.D. 4, Sec. 11, Tp. 1, Range 16, W. 4th Mer., Elevation 3,508 Feet

Drilling depth

2,400-2,440 Mostly red shale. A little fine sandstone

2,440-2,450 50 per cent red shale, 40 per cent grey, sandy shale. A little medium sandstone

2,450-2,460 65 per cent sandy grey shale and shaly sandstone, 30 per cent red shale.

Minor quartz and red chert grains

2,460-2,490 55 per cent grey, limy shale, 45 per cent red shale. Carbonaceous remains

2,490-2,510 70 per cent limy, light grey shale, disseminated pyrite 30 per cent red shale. Minor quartz and red chert grains 2,510-2,520 70 per cent light greenish grey, calcareous shale

2,510-2,520 70 per cent light greenish grey, calcareous shale
20 per cent light grey, sandy, and calcareous shale
10 per cent red shale
A few carbonaceous remains

Top? of Ellis

2,520-2,530 50 per cent light grey, sandy, and calcareous shale 20 per cent light greenish grey shale with disseminated pyrite 25 per cent red shale Several specks of glauconite

2,530-2,540 65 per cent light greenish grey, sandy, and calcareous shale 30 per cent red shale A little fine white sandstone

A few specks of glauconite

2,540-2,550 60 per cent light grey, calcareous shale with disseminated pyrite 25 per cent red shale

15 per cent light grey, sandy shale 2,550-2,570 40 per cent soluble in hot HCl Of remainder, 60 per cent grey shale

25 per cent red shale

10 per cent quartz grains
A little pyrite, white chert, and a few coal fragments

Base of Ellis

2,570-2,590 93 per cent soluble in hot HCl, remainder white chert, a little pyrite, and several grains of glauconite

Northwest Company, Limited, Erickson Coulée Well No. 1; L.S.D. 13, Sec. 8, Tp. 1, Range 12, W. 4th Mer., Elevation (?)

Drilling depth

2,300-2,310 Dark grey, green, red shale. Coal fragments

2,310-2,320 Medium-grained, quartz and black chert sandstone, slightly calcareous. Carbonaceous fragments (resembles Vanalta oil sand)

2,320-2,370 Missing

	Top of Ellis, probably at about 2,360 feet
Drilling dep	th
2,370-2,380	85 per cent grey, sandy, slightly calcareous shale 10 per cent glauconite Some pyrite
2,380-2,390	
2,390-2,400	60 per cent grey and light grey shale 35 per cent glauconite Some quartz grains and pyrite
2,400-2,420 2,420-2,440	Mostly light grey shale, some glauconite and pyrite Mostly light grey, highly calcareous shale A little glauconite and pyrite Pyrite cast of small gasteropod A little light grey shale showing mottling to red
2,440-2,450 2,450-2,470	Missing Mostly light grey, highly calcareous shale Some pyrite and a little calcareous quartz and chert sandstone
2,470-2,490	A few fragments of red shale Mostly greenish grey, shaly limestone with disseminated pyrite Some glauconite
2,490-2,500	As above, with a little, fine, calcareous sandstone
	Base of Ellis
2,500-2,520	65 per cent soluble in hot HCl, remainder white chert, fine white sandstone, and light greenish shale
2,520-2,530	60 per cent soluble in hot HCl, remainder 90 per cent spongy chert
Urban Oils, Lin	nited, Well No. 1; L.S.D. 2, Sec. 4, Tp. 1, Range 15, W. 4th Mer., Elevation about 3,450 Feet
Drilling dep	
	Mostly light grey and dark grey, sandy shale, with minor red shale and green shale. Carbonaceous remains
	Top of Ellis
2,390-2,400	40 per cent brown, sandy shale 30 per cent fine, white sandstone Misor diversity and light grey shale
2,400-2,410	Minor glauconite, red, green, and light grey shale Mostly dark brown, slightly calcareous shale, also greenish grey and dark grey, calcareous shales, with disseminated sand grains and red chert grains Some fragments with high per cent of glauconite
2,410-2,420	A little red shale Equal amounts of brownish grey, greenish grey, and dark grey shales as above
2,420-2,430	Some glauconite 50 per cent greenish grey, calcareous shale 30 per cent brownish grey, calcareous shale Some red shale
2,430-2,450	Minor glauconite, chert, quartz grains, and pyrite Mostly light greenish grey, calcareous shale Abundant pyrite Minor glauconite, and quartz and rod short grains

Mostly light greenish grey, light brownish grey, and brown shale with scattered sand grains

Some red shale and minor red chert grains and pyrite

Mostly grey, calcareous shale and light greenish grey shale with disseminated pyrite

Minor red chert grains and red shale Calcareous, grey shale and light greenish grey shale with dissemin-

Calcareous, grey shale and light greenish grey shale with disseminated pyrite
Mostly very finely sandy, and shaly limestone and calcareous shale.
Disseminated pyrite and minor red and black chert grains
Mostly brownish grey, sandy, and calcareous shale
Minor quartz and red chert grains
65 per cent grey, sandy, and light greenish grey, calcareous shale
25 per cent limestone
Minor glauconite

2,450-2,470 2,470-2,480

2,480-2,490

2,490-2,500

2,500-2,520 2,520-2,525 2,525-2,530 As above, but no glauconite

Minor glauconite

Drilling depth

2,530-2,535 80 per cent white, shaly limestone 10 per cent grey, calcareous shale Minor quartz, glauconite, and pyrite 40 per cent white limestone

2,535-2,545 30 per cent light grey, calcareous shale

15 per cent quartz grains Minor glauconite and pyrite

30 per cent white limestone; 65 per cent light greenish grey and brown-2,545-2,550ish grey, sandy, and calcareous shale with disseminated pyrite Minor glauconite

2,550-2,565 75 per cent cherty and siliceous limestone

20 per cent grey, calcareous shale Minor glauconite and pyrite

2,565-2,568

Mostly shaly, dark limestone, some white limestone Abundant black chert grains. Some pyrite 30 per cent white limestone, 60 per cent grey, calcareous shale. 2,568-2,570Minor glauconite, and white and black chert and quartz grains

Base of Ellis

2,570-2,580 Mostly white limestone

Commonwealth Petroleum, Limited, Milk River Well No. 1; L.S.D. 8, Sec. 9, Tp. 3, Range 15, W. 4th Mer., Elevation 3,163 Feet

Drilling depth

2,560-2,570 30 per cent brown, sandy shale 45 per cent shaly, fine sandstone

15 per cent white, sandy shale Minor red and black chert and quartz grains

Few carbonaceous remains

2,570-2,580 Mostly white, slightly calcareous, fine sandstone Some brown, sandy shale

Minor red and black chert and quartz grains; pyrite
Mostly medium-grained quartz and black and brown chert grains
(this resembles the Vanalta oil sand) 2,580-2,600

Top of Ellis

2,600-2,160 Mostly brownish grey, sandy shale Some glauconite

2,610-2,620 60 per cent brownish grey shale, 30 per cent ironstone Abundant glauconite

Some bentonitic clay 2,620-2,630

Mostly grey shale Some fragments nearly all glauconite in a black shale matrix

Brownish grey, sandy, and calcareous shale Some glauconite and pyrite Light greenish grey, highly calcareous shale 2,630-2,640

2,640-2,650 A little glauconite

Pyrite cast of small gasteropod; few marcasite balls Mostly light greenish grey, highly calcareous shale 2,650-2,670

Minor glauconite, ironstone, black chert, and fine white sandstone Mostly light greenish grey, highly calcareous shale A few fragments of green shale 2.670-2.680

Abundant pyrite Mostly light greenish grey, calcareous 'pencil' shale 2,680-2,700 Some brownish shale

Much pyrite and marcasite Some glauconite

Pyrite cast of small gasteropod

As above, with a little white, medium-grained sandstone 2,700-2,710 2,710-2,730

Light greenish grey, calcareous, and sandy shale with much disseminated pyrite

As above, some white limestone, some chert, and quartz grains 2,730-2,740 A few fragments of coal A little reddish shale

Base of Ellis

2,740-2,750 Practically all white limestone

Of the logs given above, those of the three eastern wells, the Commonwealth Petroleum, Limited, Milk River well No. 1, Urban Oils, Limited, well No. 1, and Northwest Company, Limited, Erickson Coulée well No. 1, agree fairly closely in thickness and general character of the sediments, and show some similarity in the detail of the succession of the different types. In each of these wells glauconite occurs at or near the base, above which it is absent until near the top of the section where it occurs much more abundantly than at the base. The quartz grains and the red and black chert grains that are scattered through the shales and limy shales are small, few being larger than 1 mm. in diameter. Comparing the logs of the three wells mentioned above with those of the Commonwealth Petroleum, Limited, Red Coulée well No. 1, and the Celtic Oils, Limited, Red Coulée well No. 1, in the western part of the area, certain marked differences are noted. In the eastern wells the thicknesses are greater (from 140 to 180 feet) than in the western wells (60 to 80 feet) and though a small amount of glauconite occurs at the base of the section of the western wells, as in the case of the eastern wells, no higher, more abundant glauconite is present. In the western wells the quartz and the red and black chert grains are more numerous and larger (a few reaching a diameter of 1 cm.) than in the eastern wells.

The presence in the west of the more numerous and larger quartz and chert grains renders it probable that the Ellis strata at the site of the western wells were laid down nearer shore than those to the east; the absence of the upper glauconite horizon and the markedly decreased thickness, in the western wells, probably indicate that either the sea receded from the western area before the time of deposition of the upper glauconite or that such deposits in the west were later eroded. In all five wells, beds of sand occur above the Ellis, and the sand is very similar in all of the wells, so that it appears that if erosion were responsible for the missing beds to the west, such erosion must have taken place before the deposition of this sand.

In the case of Dalco Oil and Gas Company well No. 1, in the western part of the area, it will be noted that red shale (very rare in all other logs) is very abundant, and that there is no sand that may be considered to be the Vanalta oil sand, that is, the sand that occurs just above the Ellis in all the other logs. The sand that does appear higher up in the Dalco No. 1 well is of a type wholly different from that of the Vanalta sand of the other wells. The 50 feet of sediments assigned to the Ellis contain material (aside from the red shale) very similar to that found at the base of the Ellis in all the other wells, though glauconite is quite minor in amount. The succeeding shale has a different aspect and compares closely with that found in the other wells in the Lower Cretaceous (Kootenay?). It appears as if erosion took place locally during or after the deposition of the Vanalta sand, and probably produced an erosion channel filled later with shale of continental deposition, much of which is now red shale.

The writer is of the opinion, then, that in this area erosion took place at the end of Ellis deposition, as might be expected during a recession of the sea. Definite evidence of erosion at the end of Ellis deposition in southern Montana is given by Knappen and Moulton.¹

¹Knappen, R. S., and Moulton, G. F.: U.S. Geol. Surv., Bull. 822A, pp. 9, 20 (1930).

Below is given a representative list of general descriptions of the Ellis in Montana.

Locality	Formation	Thickness	Character of strata
		Feet	
Kevin-Sunburst, 1 Mont	Ellis	200-300	Light brown to grey (locally yellowish), shaly or limy sandstone in upper part. Dove-coloured to black, shaly limestones in lower. Locally a pebble conglomerate at base. Marine fossils abundant. Sandstone members present. Great unconformity at base
Great Falls coal field ² , Mont	Ellis	80-120	Dove-coloured limestone, overlain by reddish brown, coarse-grained sandstone, coarsely conglomeratic at base Unconformably on Madison and Quadrant for- mations
Carbon county ³ , Mont	Sundance	to 475	50-100 —sandstone and sandy clay 200—olive-brown clay with subordinate sandy layers 175—very regularly interbedded, olive-brown clay and shale, pink and rose-coloured clay— thin limestone and a few thin beds of gypsum Bounded by erosional unconformities between Chugwater (Triassic?) below and Morrison (Cre taceous?) above
Birch River area,4 Mont	Ellis	240–310	Black to grey, calcareous shale with a few, thin, irregular beds of limestone and sandstone Disconformably on Mississippian limestone
South of Bearpaw mountains, Mont	Ellis	310	Buff and grey, petroliferous and fossiliferous limy shale and thin beds of limestone, with 75 feet of blue to black, thin-bedded, petro- liferous and fossiliferous limestone at base
Lewistown coal fields, Mont	Ellis	250	148 feet partly concealed—mostly grey sand- stone 80 feet concealed 10 feet fossiliferous limestone and shaly lime- stone 4 red sandy shale 20 pure white gypsum 92 feet limestone and gypsiferous shale On Madison limestone without apparent uncon- formity
Sweet Grass hills', Mont.		201	40 feet green, soft shales with dark stains 50 feet dark blue, calcareous shale with Bele- mnites densus 45 feet dark blue, impure limestone containing Gryphaea calceola 66 feet probably dark greenish, sandy shales, concealed

¹Perry, E. S.: State Bureau of Mines and Geology, Montana, Mem. No. 1, 2nd Edit., Table 1 (1929).

²Fisher, C. A.: U.S. Geol. Surv., Bull. 356, pp. 23, 27 (1909).

³Knappen, R. S., and Moulton, G. F.: U.S. Geol. Surv., Bull. 822-A, pp. 17-20 (1930).

⁴Stebinger, E.: U.S. Geol. Surv., Bull. 691-E, p. 155 (1918).

⁵Reeves, F.: U.S. Geol. Surv., Bull. 751-C, p. 75 (1924).

⁶Calvert, W. R.: U.S. Geol. Surv., Bull. 390, p. 20 (1909).

⁷Slipper, S. E.: Geol. Surv., Canada, Mem. 93, p. 66 (1917).

Fossils have been collected from the Ellis of most of these sections and in every case were reported as being of Upper Jurassic age. rather marked variation in thickness and type must thus be attributed to local conditions of deposition in the Upper Jurassic sea. Thus, probably, in places and at times, conditions were favourable for the formation of source rocks of oil, whereas at other places and times decidedly unfavourable.

A detailed section of the Ellis from the southern Bearpaw mountains. that seems to be a type distinctly favourable for oil, is given below:

50 feet brownish shale with characteristic Ellis fossils

20 feet igneous sill

15 feet dark shales and shaly limestones, petroliferous and characteristic Ellis fossils

20 feet drab, limy shale

10 feet thin-bedded, shaly, drab limestone. Belemnites 40 feet limy shale to black limestone, weathers in grey, barren surfaces. Characteristic Ellis fossils

20 feet igneous sill

50 feet grey, limy shales, petroliferous 5 feet dark blue, marly limestone, weathers light yellow. Petroliferous

30 feet limy, brown shale

6 feet limy, brown shale and grey, hard limestone full of Gryphaea calceola
6 feet limy brown shale, weathers yellow-grey
6 inches black to bluish grey, thinly laminated limestone with finely crenulated partings, probably
of organic origin. Very petroliferous and on fresh fracture is brownish black. Algal layer
70 feet blue to black limestone, thin bedded, petroliferous, and fossiliferous. Beds \(\frac{1}{2}\) to 6 feet.
Upper part very colitic. Top surface has a reddish tinge with ripple-marks in places

6 inches brown, sandy clay

o inches brown, sainty clay
2 feet soft, limy clay to argillaceous limestone
2 inches argillaceous, irregularly platy limestone in sharp basal contact with—

21 inches breccia or conglomerate of limestone and chert

None of the six wells furnishing the detailed sections of the Ellis given in this report has produced oil. The Ellis in them when compared with the section from the southern Bearpaw mountains, given above, does not appear to be a favourable source rock. In the case of the other wells now producing oil in the Red Coulée field, the oil comes from the Vanalta sand and, therefore, the underlying Ellis was not penetrated. It may be that at the sites of the producing wells local conditions were favourable for the deposition of source rocks during the time of accumulation of the Ellis strata; on the other hand the oil may have come from other strata.

Lower Cretaceous (Kootenay) Strata

The strata assigned to the Lower Cretaceous consist of grey, brownish grey, green, and red shales with thin (few inches to a few feet) sandstone beds, and a heavy sandstone bed at the base. No exposures of these strata occur within the area. They lie above the Ellis and vary in thickness from 580 feet in the western part of the area (range 16) to 490 feet in the eastern part (range 12). In general the upper part is predominantly shaly, and it is here that grey and green shales predominate, though some red shale is present in most of the well logs. Sandstone is greater in amount and present in thicker beds in the lower part where it is associated with red shales, though grey, green, and brown shales also occur.

The contact with the overlying Alberta shales is placed, in the logs, at the point where green shale or, in some instances, red shale occurs below drab-grey shales or grey sandstone. In some logs the grey shales and sand-

¹Howe, M. A., and Goldman, M.: Am. Jour. Sci., vol. 10, p. 314 (1925).

stones, above this colour change, are fairly carbonaceous, whereas in other well logs little carbonaceous material is found and they contain small amounts of glauconite. Although this contact may not actually mark the upper limit of the Lower Cretaceous or of the Lower Cretaceous continental deposits, it is a useful horizon, for it occurs in all the Red Coulée wells at a uniform distance below the horizon of the Colorado water sand (See later). The base of the Lower Cretaceous is placed in most of the wells at the base of a medium-grained sand (the Vanalta oil sand) lying just above the limy shales or shales containing glauconite that the writer has assigned to the Ellis. This sand is placed in the Lower Cretaceous rather than in the Ellis since it generally contains carbonaceous fragments.

The name Kootenai is applied to strata similar to the above in the Kevin-Sunburst field and this term (or rather Kootenay) is in general

use by the oil men in the Red Coulée field.

Bituminous shales were found in five of the logs in the Lower Cretaceous. They occur in the Celtic Oils, Limited, well, between 2,250 and 2,270 feet drilling depths, where this 20 feet is practically all bituminous shale. This same band is represented in Dalco Oil and Gas Company well No. 1; Southern Alberta Exploration Company well No. 1, and Urban Oils well No. 1 at about the same horizon as in the Celtic well, that is, 200 to 240 feet below the top of the Lower Cretaceous. In all but the Celtic well this band appears to be in several thin beds separated by other sediments. A small amount of bituminous shale occurs about 400 feet below the top of the Lower Cretaceous in the Capital Red Coulée well, but this horizon was not recognized in any other well.

The bituminous shales, though not regarded as source rocks for oil, are at least related to source rocks in origin, and their occurrence in the Lower Cretaceous (Kootenay?) indicates that these strata should not be

disregarded in considering the source of oil in this area.

ALBERTA SHALES

The Alberta shales are not exposed in the area, but in the wells a thick (1,750 feet) series of drab and dark grey shales and grey sandstones lies on the green and red shales of the Lower Cretaceous. These strata are generally termed Colorado, but since the writer collected from the upper part of these strata in Buckley coulée, Montana, specimens of Baculites cf. ovatus, of large size, that resemble very much those found in the lower part of the Montana strata to the northwest in the foothills, it may be that some Montana strata are present, and for this reason, the term Alberta shale is used in this report.

In the wells, the Alberta shales appear as a thick zone of fairly dark shales. Close inspection of drill samples shows, however, that the top 200 feet and the bottom 600 feet are distinctly greyer than the part between. This grey hue is due to the presence of sandstone and sandy shale fragments in the samples, and since no single 10-foot sample is wholly of sandstone fragments, probably none of the sandstone beds is as thick as 10 feet. Although the upper and lower sandy zones are present in most of the wells, so that a general correlation may be made of these zones, yet

¹Hume, G. S.: Geol. Surv., Canada, Sum. Rept. 1929, pt. B, p. 6.

only one sandstone bed seems to persist at about the same horizon throughout the field. This bed, termed the "Colorado water sand," carries water in some of the wells. The section between the upper and lower sandy zones is chiefly dark shale with occasionally sandy shales, limy shales, ironstone bands, and bentonitic beds. No single, highly bentonitic bed can be recognized as a horizon marker. Bentonitic beds are chiefly confined to three fairly thick, poorly defined zones, as follows:

At 250–550 feet above the base (in the lower sandy zone) At 700–800 feet above the base At 1,050 to 1,250 feet above the base

The contact of the Alberta shales with the underlying Lower Cretaceous has been described on a preceding page. The contact with the overlying Lower Milk River sandstone is very poorly defined, for between the Lower Milk River massive sandstone above, and the Alberta dark, marine shale below, there is a gradational zone, the lower part of which is of shales with but few thin bands of sandstone, whereas in the upper part thin sandstones predominate. In Buckley coulée, Montana, just south of the border, good exposures of the Lower Milk River-Alberta shales contact were seen. In the eastern part of the coulée the contact is as described above, and in one section, at about 50 feet below the base of the massive sandstone of the Lower Milk River a 1-foot bed of sandstone occurs in a zone that is predominantly shale. This bed can be traced westward for 4 miles, in which direction it gradually thickens, until at a point 3 miles west, where its base is last exposed, it has a thickness of 50 feet. At this point the top of the sandstone is about 25 feet below the base of the massive Milk River sandstone, a little over a mile farther west this interval has decreased to about 19 feet. The sandstone of this wedge-shaped body is in every way comparable with that of the massive Lower Milk River sandstone, but marine (and probably Montana) fossils were found in an ironstone band just above it. It thus appears as if this sandstone wedge interfingers eastward with the shale, and the shale and sandstone of definitely known marine origin, and westward becomes one with the Lower Milk River sandstone. No fossils and only a little coaly material are found in the Lower Milk River sandstone, so that so far as known it may be of marine shore, brackish, or continental deposition.

In an attempt to use the top of the Alberta shales as an horizon marker the writer has tried placing the contact at several places; at the base of the section showing nothing but sandstone; where sand and shale are equal in amount; and where sand is first found in quite minor amounts. However, no single method gives a consistent thickness for the Alberta shales, nor gives an horizon that accords with the structure as indicated by several deeper horizons which check with one another. An indication of the difficulties to be expected in trying to use the Lower Milk River-Alberta shales contact as an horizon is furnished by the sandstone wedge occurring in Buckley coulée, as described above. It seems likely that farther west the top of this wedge joins the base of the massive sandstone above, so that in the west the contact would most likely be placed from 50 to 75 feet lower than in the east.

Foraminifera have been found in the Alberta shales and one horizon has been established by their use (See later).

MILK RIVER SANDSTONE

The Milk River sandstone is the oldest formation exposed within the area. It is readily divisible into two parts; the Lower Milk River (Virgelle of oil men) and the Upper Milk River. The Lower Milk River consists of thick (70 to 100 feet), massive, fine to medium sandstone, composed chiefly of quartz grains (with minor black and brown mica) rather loosely cemented with a calcareous and shaly matrix. At its base, the massive sandstone gives way to alternations of sand and shale, and in the well logs the base is usually placed where sand first becomes subordinate to dark shale. The base is not exposed within the area covered by this report, but was examined in Buckley coulée just to the south, in Montana. Towards the top of the Lower Milk River, the sandstones are markedly crossbedded and there are many large (up to 10 feet diameter) ovoid concretions of sandstone cemented with ferruginous carbonate, which characteristically form the cappings of ledges and pillars or "hoodoos."

The contact with the Upper Milk River is placed where white or yellowish sand, either loose or firmly cemented into a hard ledge, is overlain by dark grey or brown, sandy shale. Where exposed the contact is usually sharp, but along the rivers and coulées the contact is commonly concealed, for it generally occurs on a narrow terrace covered by surface wash from higher exposures of Upper Milk River that are supported by several strong sandstone beds. The contact may be found, however, with a little digging. This contact is the only definite horizon exposed at the surface, and, accordingly, levels were run along it in an

attempt to make a structural contour map.

The Lower Milk River maintains its massive character in all but one outcrop. On the north bank of Milk river, in the west part of sec. 16, tp. 2, range 14, the massive sandstone passes laterally into sandstone and shale, which continue for about half a mile, beyond which the massive character is resumed.

The Upper Milk River consists, for the most part, of shales and sandy shales with several thick (8 to 20 feet) beds of sandstone. The shales and sandy shales are grey, dark grey (usually lignitic), yellowish and greenish grey. Weathered exposures are characterized by light grey, yellowish grey, black, purplish grey, and brownish bands.

The succession varies both in lithology and thickness from place to

place, but in general is as follows:

Pebble bed at base of Pakowki

20-30 feet of grey and yellowish shales and sandy shales; some bentonite and selenite

8-24 feet of coarse and medium-grained, yellowish sandstone 45 feet of shales, carbonaceous and sandy, with thin sandstone beds and beds and layers of nodules of iron carbonate, and greenish sandstone cemented with iron carbonate. These layers and nodules weather dark brown

8-22 feet of medium and coarse-grained, yellowish sandstone 45-50 feet of grey, greenish grey, and dark grey, sandy shales, with 2 to 4 thin lignitic

White or yellowish sand or sandstone, top of Lower Milk River

The two massive sandstone layers were used locally to determine the position of the contact with the Lower Milk River where this horizon could not be more directly determined.

The contact with the overlying Pakowki is placed at a bed of hard, grey, calcareous sandstone that contains black chert pebbles, averaging

31470 - 2

1 cm. in diameter. Locally the pebbles are very numerous, at other places

they are quite scarce.

A measured section in Buckley coulée gave the thickness of the Lower Milk River (Virgelle) as 167 feet. No single complete section of the Upper Milk River was found, but in Verdigris coulée, in the southwest part of sec. 22, tp. 2, range 14, the base and top are exposed within half a mile of one another, and the thickness, estimated from elevations and attitude of local horizons, is here 158 feet. These two measurements give a total thickness of 325 feet.

Only two wells in the area afford complete sections of the Milk River, unobscured by surface cavings. In the Commonwealth Petroleum Company Milk River well in sec. 9, tp. 3, range 15, the Upper Milk River has a thickness of 180 feet and the Lower Milk River a thickness of 240 feet, a total of 420 feet. In Dalco Oil and Gas Company No. 2 well, in sec. 19, tp. 2, range 17, the position of the contact between Upper and Lower Milk River is somewhat doubtful, but the total thickness is 340 feet. The Upper Milk River is probably 150 feet thick and the Lower, 190 feet. The thickness indicated by the Commonwealth well seems excessive, but since a thick wedge of sandstone occurs beneath the normal, massive Lower Milk River (Virgelle) in Buckley coulée, it is not at all improbable that similar conditions exist in the locality of the Commonwealth Milk River well.

The outcrops of Milk River sandstone in Alberta are continuous with those of the Eagle sandstone of the Sweet Grass arch of Montana, where the Lower Milk River is called the Virgelle¹. The fauna is described as Montana in age. No fossils were found in the Lower Milk River in the Alberta field, though Montana fossils were found just above the sandstone wedge near the Milk River-Alberta shale contact in Buckley coulée, Montana. The Lower Milk River is probably a marine shore deposit. The Upper Milk River contains thin (1 inch to 6 inches) beds of lignite and has yielded a small, deciduous flora. It is of continental origin and is continuous with the upper Eagle of the Sweet Grass arch.

PAKOWKI

Within the area the Pakowki consists of shales and sandstones containing marine fossils. The Pakowki lies above the Upper Milk River and below the Foremost beds. The base is placed at a thin, hard band of calcareous sandstone containing rounded black chert pebbles. The top was seen in only one small area on Milk river, in sec 25, tp. 2, range 18, and sec. 30, tp. 2, range 17, and there the contact with the Foremost was placed at an horizon where a massive sandstone of the basal Foremost passes into alternations of dark grey shale and thin sandstone beds in which are Pteria nebrascana. Williams² notes that the upper contact is marked by the occurrence of "thin lenses of 2 to 3 inches thick consisting of 'cones' of calcite starting at the top and bottom and interlocking near the centre". The writer found a bed of this type at the top of the Pakowki in sec. 30, tp. 2, range 17, on the south bank of Milk river, though just to the west on the north bank it was not present. Similar cone-in-cone concretionary layers were seen in Upper and Lower Milk River sandstones and in the Foremost, as well as within the Pakowki.

¹Stebinger, E.: U.S. Geol. Surv., Bull. 641-C, p. 58 (1916). ²Williams, M. Y., and Dyer, W. S.: Geol. Surv., Canada, Mem. 163, p. 14 (1930).

No good sections of the Pakowki occur within the area, but Dr. S. E. Slipper, of the Canadian Western Natural Gas, Light, Heat, and Power Company, Limited, very kindly provided the writer with logs of test wells drilled along Milk river, through the Pakowki, and allowed him to examine the core samples. A typical section is given below:

1	Orilli	ing deptl	n	
Ft.		Ft.		
0		10		Overburden
10		39	6	Mostly fine to medium-grained, grey and greenish grey sandstone, with a few thin shale bands
				(Top of Pakowki)
39	6	58		Interbedded, dark grey shale and fine to medium-grained, grey sandstone
58		64	6	Dark grey shale with a few layers of fine- grained sandstone
64	6	64	9	Hard ironstone, cone-in-cone structure
64	9	91	6	Dark grey shale
91	6	92 -		Dark grey, sandy shale
92		94		Greenish grey, shaly sandstone
94		94	3	Hard ironstone
94	3	110		Greenish, fine sandstone
110	-	118	6	Light grey, fine-grained sandstone
118	6	120	-	Dark grey, sandy shale
120	-	129.		Dark grey shale
129		131		Greenish grey sandstone
131		137		Dark grey, sandy shale
137		137	6	Greenish, fine-grained sandstone
137	6	139	-	Dark grey, sandy shale
139		143		Dark grey shale
143		146		Greenish grey, fine-grained sandstone
146		147	6	Dark grey, fine-grained, shaly sandstone
147	6	150	1	Greenish to grey, fine-grained sandstone
		~		(Fragment of Baculite at this horizon in
				another log)
150		166		Drab to dark grey, very shaly sandstone
166		170		Dark grey, sandy shale
170		191	6	Dark grey shale
191	6	192		Bentonite
192		214	-	Dark grey shale
214		222	1	Dark grey, sandy shale, black chert pebble
	-		29	(Base of Pakowki)
222		224		Light grey sandstone
224		227		Grey shale and sandy shale
224			6 4	OTES SHARE ARE SALES SHARE
		39.	n.b	

Small collections of typical Pierre fossils were obtained from the Pakowki. Among the species, identified by L. S. Russell, are the following:

	Other occurrence
Inoceramus subdepressus Meek and Hayden?	Claggett
Pteria nebrascana (Evans and Shumard)	Uppermost Alberta shale to Fox
	Hills
Cardium speciosum Meek and Hayden?	
Callista deweyi (Meek and Hayden)	Bearpaw to Cannonball
Mactra formosa Meek and Hayden	
Panopaea occidentalis Meek and Hayden	Claggett
Baculites compressus Say	Uppermost Alberta shale to Bear-
	paw
Baculites cf. anceps Lamarck	

31470-21

BELLY RIVER

Lying above the Pakowki and below the Bearpaw is a series of freshwater shales and sandstones. Williams and Dyer¹ review the nomenclature of these beds and propose "to follow Dawson's original intention and to define the Belly River as being composed of the continental beds lying between the Pakowki and the Bearpaw". This proposal is followed in this report. In the disturbed belt to the northwest, the name is commonly applied to the continental series lying above the Alberta shale and below Bearpaw or Edmonton. It is obvious, of course, that this greater series of continental beds includes more than the Belly River as defined above.

Williams and Dyer² accept the division of the Belly River of southern Alberta, into two parts, Foremost below and Pale beds above. They state "although the two members are more or less alike and the boundary between them is by no means definite, they are sufficiently different to form desirable mapping units. The outstanding difference . . . is a difference in colour, the Pale beds being very much lighter than the Foremost The Pale beds include massive strata of hard, yellow, and grey sandstone which are rare in the Foremost Coal seams are common in the Foremost but occur only near the top of the Pale beds."

The Foremost and lower Pale beds are intermittently exposed along a watercourse on the ridge in secs. 32 and 33, tp. 1, range 17. At about 300 feet above the base the strata quite definitely have the characteristics of the Pale beds, and at about 240 feet above the base they are fairly definitely Foremost. Between these horizons are exposed several sandstone layers that could not be assigned definitely. The same difficulty was met with in connexion with exposures along Milk river near the junction of the south branch. Up to a certain exposure in going west there is little doubt but that the strata belong to the Foremost, but just west of this, beds outcrop that can not be placed definitely, whereas still farther west they are certainly Pale beds. The nature of the exposures makes it impossible to say whether the change from Foremost to Pale beds is gradational or abrupt, or whether it does or does not take place at a fairly uniform horizon.

A partial section (114 feet) of the Foremost was obtained by trenching in sec. 2, tp. 2, range 17. It showed alternations of drab grey shale, brown lignitic shale, and shaly sandstone with a few thin seams of lignite and several concretionary ironstone layers, one of which held Unios and Corbula perundata. At the base just below a thin seam of lignite, is a 4-foot bed of impure, white to yellowish bentonite. Similar bentonitic beds were found in separate outcrops westward for about 2 miles, and were used as a local horizon.

In sec. 25, tp. 2, range 18, about 30 feet of yellowish grey, medium-grained sandstone forms the base of the Foremost. A heavy bed of similar sandstone was found to occur near the base of the Foremost in sec. 35, tp. 1, range 17, and again in sec. 12, tp. 2, range 17, where, however, the total thickness of the sandstone is not exposed.

Good exposures of part of the Pale beds occur along Milk river in tp. 2, range 19. They consist of light grey and light greenish grey sandstones, in beds from 1 foot to 20 feet thick, alternating with grey and greenish shales, and occasional 1-foot ironstone concretionary layers.

¹Williams, M. Y., and Dyer, W. S.: Geol. Surv., Canada, Mem. 163, pp. 16, 17 (1930). ²Op. cit., p. 17.

BEARPAW

No exposures of Bearpaw were found, but it seems likely that it occurs under the drift on the highest parts of Milk River ridge within the area.

STRUCTURE

In an attempt to determine the structure of the area, the elevations of all exposed horizons were determined. Of these horizons, that at the top of the Lower Milk River is much more widely exposed. Other well-established horizons include the base and the top of the Pakowi. Locally developed horizons were used where available to give an indication of

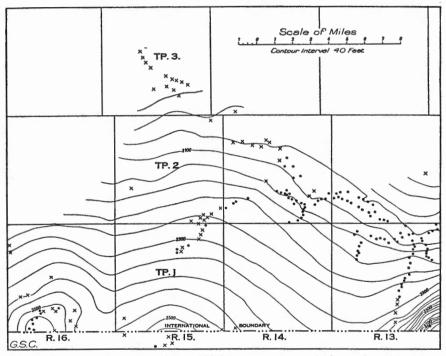


Figure 2. Structural contour map of part of Milk River area, west of 4th meridian, Alberta; contours are approximate and represent surface of top of Lower Milk River; outcrops where the elevation of top of Lower Milk River was directly determined are indicated by solid circular dots, outcrops where elevation of same datum plane was indirectly determined are indicated by crosses.

local dip. Exposures of recognizable horizons were found to be fairly continuous on parts of Red creek, along Milk river east of the town of Milk River, and in Verdigris, Van Cleeve, and Police coulées, but over most of the remaining parts of the area exposures are widely separated and at many of them no established horizon could be recognized. As a result of the general conditions found to obtain, an interpretation of the structure of only a part of the field could be made. This part of the area is mainly confined to townships 1 and 2, ranges 13, 14, 15, and 16, represented

in outline on the accompanying figure (Figure 2). The structural contours of this figure represent the attitude of the top of the Lower Milk River and depend on measured elevations of outcrops of this and other horizons. From an inspection of the figure it is evident that the outcrops permitting determination of the elevation of the datum horizon are so distributed that the structure can be only approximated. It is realized that if the outcrops were more uniformly distributed over the whole area, the structures would probably markedly differ from those depicted. Furthermore, wells in the Red Coulée field seem to indicate that the top of the Lower Milk River may not follow in detail the deeper structures. The present writer, therefore, is inclined to believe that the top of the Lower Milk River may be used only to indicate the broad, general structure.

The evidence from the wells in the Red Coulée field that seems to indicate that the structure of the top of the Lower Milk River is different from that of lower horizons is as follows. The elevation of the top of the Lower Milk River is 50 feet higher in the Commonwealth Petroleum Company Red Coulée well No. 1 than in the Southern Alberta Exploration Company well and some other nearby wells, whereas elevations of three deeper horizons, as determined in all the wells, agree in showing that the top of the Lower Milk River in the Commonwealth well should be 30 feet lower than it is in the Southern Alberta Exploration Company well if the structure on this horizon is the same as those of the lower horizons. The apparent vertical discrepancy is 80 feet, possibly part of this may be due to the fact that it has not been possible to locate definitely the top of the Lower Milk River in any of the wells being compared except the Commonwealth well. It does not seem likely, however, that the assumed positions of the horizon can be in error by an amount so great as 80 feet.

In attempting to determine the general structure of part of the area represented in Figure 2 and the structure of the Red Coulée field as represented in Figure 3, use was made of the various horizons described in the next paragraphs.

Top of Pakowki. This horizon was found outeropping in one locality and was placed where dark shale and medium-grained sandstone became about equal in amount, just below a massive sandstone at the base of the Foremost, and just above sandy beds containing Pteria nebrascana. Arenaceous foraminifera were found at the top of the Pakowki in samples from Dalco Oil and Gas Company No. 2 well.

Base of Pakowki. The base of the Pakowki is marked in Verdigris coulée by limy sandstone beds containing rounded black chert pebbles of 1 cm. average diameter. Numerous pebbles occur in the bed at one place, whereas a few hundred feet distant they are quite scattered. This probably explains why samples from this horizon from certain wells that have penetrated the Pakowki show no chert pebbles.

Top of Lower Milk River. This contact where exposed on Milk river and in the various coulées is fairly sharp. The writer placed it at the contact of sandy and carbonaceous shale above with sandstone below. The sandstone in places is very loosely consolidated and at other places is cemented into a hard ledge.

Bullopora-Clavulina Horizon. R. T. D. Wickenden of the Geological Survey has found a foraminiferal zone in the upper part of the Alberta shales of the Red Coulée field. The horizon is marked by the highest occurrence of two foraminifera: a species of Bullopora occurring 20 to 30 feet above a species of Clavulina. This horizon has been found in five wells. It lies about 400 feet below the top of the Alberta shale.

Colorado Water Sand. In many of the wells of the Red Coulée field and the Montana Border field, a water horizon is found 65 to 75 feet above the top of the Lower Cretaceous. Since the wells were drilled by standard cable, chip samples only are available, so that the thickness of this sand cannot be determined. The sand present in the samples representing 10-foot intervals varies from 30 to 75 per cent of the sample and in some wells is present in several 10-foot samples. It is probable that at this horizon there are several thin beds of sand. In spite of the sand not being clearly defined in the logs, this horizon as determined in wells where water was encountered, or in the other wells by the most sandy 10-foot sample at this general horizon, occurs at very uniform distances above the top of the Lower Cretaceous and below the foraminiferal horizon in the Alberta shale.

The Colorado water sand consists of fine to medium-grained quartz sand with a few scattered black chert grains and pieces of black mica. The cement is sparse and is clayey and limy.

Three analyses of the water at this horizon, made by W. P. Campbell, Assistant Chemist of the Supervisory Mining Engineer's office, Calgary, are given below.

			Par	ts per m	illion			
Well	Total solids	K and, or, Na	Са	Mg	Cl	SO ₄	Bicarb.	Carbo- nate
Dalco 1	13,064 13,336 14,074	4,896	68 44 318	31 23 203	6,932 6,567 6,901	44 21 33	1,120 1,912 2,085	0 0 0

Top of Lower Cretaceous. The contact of Lower Cretaceous and Alberta shale is placed at the place where greenish sand or shale or red shale, or both, occur below grey sand and drab shale.

Vanalta Gas Sand. In most of the wells a fine white sandstone occurs with its base 47 to 60 feet above the top of the Vanalta oil sand. This sandstone seems to be less than 10 feet thick, since from 10 to 70 per cent of the 10-foot samples in which it occurs consist of shale. It is probable that there are several beds of this sandstone, since it is present as fragments in two or three of the 10-foot samples. Small flows of gas were encountered at this horizon in several of the wells.

Vanalta Oil Sand. The Vanalta oil sand is the horizon from which all commercial production of oil has been obtained in the Red Coulée field. Here it consists of from 45 to more than 59 feet of fine to mediumgrained, loosely cemented, quartz and black chert sand. In two wells the black chert grains are quite coarse towards the base. Much pyrite

occurs in the sand. In five wells the sand is continuous and in two others, 4- and 5-foot bands of sandy shale occur within it. It could not be located in the Dalco No. 1 well and was represented only by scattered sand, of the type noted, in shale in the Urban well.

In the Border oil field in Montana, just south and continuous with the Red Coulée field of Alberta, the basal part of the Lower Cretaceous

is thought to be as follows:1

	Feet
Gas sand	10-45
Variegated shale	40-80
Vanalta sand	5-15
Grev, green, pyritic shale	10-15
Cosmos sandstone	30-40
Ellis formation	

"The gas sand is commonly regarded as the counterpart of the Sunburst sand in the Kevin-Sunburst field, but it probably occupies a slightly higher stratigraphic horizon. The variegated shale is regarded as the westward expansion of the 'yellow shale' horizon of the Kevin-Sunburst field. The Vanalta and Cosmos sandstones occupy the horizon of the 'laminated sandstones' of the Kevin-Sunburst field."

Ellis Water Sand. Water was encountered in Dalco No. 1, Celtic No. 1, and in the Commonwealth Red Coulée No. 1 wells close to the base of the Ellis, in highly calcareous, fine sandstones. The drilling report of the Dalco No. 1 well gives this as a "sulphur water".

Top of the Limestone. This is placed at the top of the white or white and cherty limestone at the summit of the pre-Ellis strata.

STRUCTURE OF THE RED COULÉE FIELD

Four of the horizons listed above were employed in an endeavour to determine the structure at depth in the Red Coulée field.

- (1) Vanalta Oil Sand. This horizon has been located definitely in seven wells. In one well it does not seem to be present, in three others it has not been reached, in the case of another well the samples are missing, and in one other the horizon is doubtful, as samples from the well are not available.
- (2) Top of Lower Cretaceous. This horizon is well established in nine wells; three others are not deep enough to reach it, and in one it is doubtful.
- (3) Colorado Water Sand. This horizon is present in eleven of the wells; two others are not deep enough to reach it.
- (4) Bullopora-Clavulina Horizon. This horizon has been found in all the wells in which it has been looked for, five in number in this field.

The Colorado water sand has been chosen as a datum plane to interpret the structure of the Red Coulée field, and the results are shown in Figure 3. The Colorado water sand was used in preference to either the Vanalta sand or the top of the Lower Cretaceous because it is known in more wells and is a bed deposited under marine conditions, and, furthermore, because the elevation of this horizon in most of the wells just south in Montana is also known.

¹Personal communication, Dr. C. E. Erdmann, 1931.

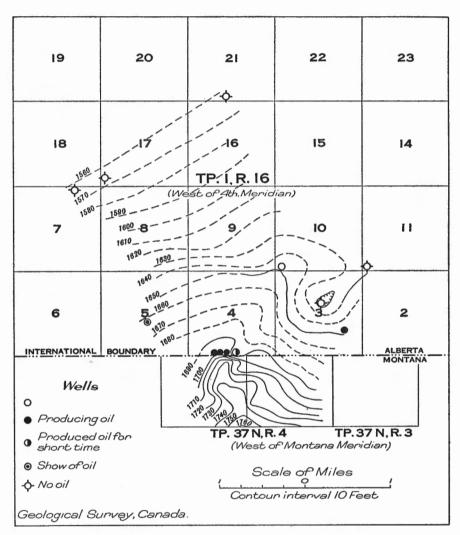


Figure 3. Structural contour map of Red Coulée field, west of 4th meridian, Alberta, contours represent surface of Colorado water sand.

The three wells, Vanalta No. 1 and No. 2, and Taylor No. 1, are close together in the south of sec. 4, tp. 1, range 16, W. 4th mer., and lie on the crest of a narrow, plunging arch and close to the foot of a fairly steep terrace slope. Vanalta No. 4, drilled since the close of field work and reported as a producer, lies about 1,000 feet due north of Taylor No. 1 and, apparently, on the same "flat" as the three wells first mentioned. Southern Alberta Exploration well No. 1, the only other producer, is in sec. 3, tp. 1, range 16, and since it is the most easterly well the details of the structure in its vicinity are unknown. However, between this well and the Commonwealth Red Coulée well No. 1 to the northwest, the structural contours must bend sharply northeast, and it is not improbable that the Southern Alberta Exploration Company well lies west of a small Devonshire well No. 1 is farther northwest, in the southwest corner of sec. 10, tp. 1, and appears to be on a "nose" northwest of a depression reached by Commonwealth Red Coulée well No. 1. The structural elevation is about the same as at the Southern Alberta Exploration The Devonshire well is a possible producer. Should it prove dry the search for other producers would be confined to short distances north and west of the Vanalta No. 1 and No. 2, and Taylor No. 1, wells and to the east and south of the Southern Alberta Exploration well No. 1.

STRUCTURE IN THE VICINITY OF DALCO OIL AND GAS COMPANY WELL $\operatorname{No.} 2$

The Dalco No. 2 well is in L.S.D. 7, sec. 19, tp. 2, range 17, W. 4th mer., just south of a small "window" of Pakowki within Foremost outcrops along Milk river. Along the river the attitude of the top of the Pakowki indicates an easterly dip of about 20 feet to the mile. One mile to the west of the Pakowki outcrop local horizons in the Foremost along the river indicate a westerly dip of about 60 feet to the mile. in secs. 3 and 4, tp. 2, range 17, about 4 miles south of the river and to the east of the Pakowki outcrop, show an easterly dip of about 15 feet a mile. The regional dip, as indicated by the elevation of the Pale beds-Foremost contact, located just south of the south boundary of tp. 2, range 17, and again near the north boundary of the township, is northward at the rate of about 45 feet to the mile. It thus appears that the Dalco well No. 2 is on the crest of an arch, but it is not known whether there is closure to the south of the well or whether the strata rise gradually and continuously in that direction. If the arch rises gradually to the south there is a possibility that a structure might be found close to the border, near the boundary of ranges 17 and 18. Any well drilled there should be regarded as exploratory.

WELL LOGS

Vanalta Oils, Limited, Red Coulée Well No. 1; L.S.D. 3, Sec. 4, Tp. 1, Range 16, W. 4th Mer., Elevation 3,571 Feet

Drilling depth	Horizons	Lithology
0→ 190 190→ 220	220, base of Lower Milk River	Missing Fine-grained, grey sandstone and shaly sandstone
220- 250 250-1,480		Dark grey shale; some shaly sandstone Dark grey shale, with sandy shale at 350-360, 390-420, and 1,150-1,170. Some bentonite at 740-750, 780-790, 820-830, and 860-870
1,480-1,490		50 per cent light grey, fine to medium-grained sand- stone, remainder shale
1,490-1,520 1,520-1,540 1,540-1,550		Dark shale and white bentonite, 50 per cent fine-grained, light greenish grey sandstone, remainder shale and bentonite
1,550-1,610		Dark grey and brownish grey shale, small amount of bentonite
1,610-1,940		Dark grey, brownish grey, and grey shale and sandy shale with sandstone beds as follows: 1,650-1,660-50 per cent fine grey sandstone 1,710-1,720-70 per cent fine grey sandstone, some glauconite 1,850-1,860-50 per cent fine to medium, grey sandstone
	1875, Colorado water sand	1,870-1,900—40 per cent fine to medium, white sand- stone Some bentonite at 1,680-1,710, 1,820-1,840, 1,860-1,870, and 1,920-1,930
1,940-1,980	1940, top of Lower Creta-	Grey shale, some green and red shale; few coal frag- ments, some bentonite
1,980-1,990		Grey and green shale and greenish grey, fine sand- stone
1,990-2,030 2,030-2,060		Grey and green and some red shale 70 per cent fine, greenish grey sandstone; remainder grey and green shale
2,060-2,080 2,080-2,090 2,090-2,110		Grey, green, and red shale 60 per cent fine, greenish grey sandstone Grey and greenish shale, some light greenish grey sandstone
2,110-2,120 2,120-2,180 2,180-2,290		80 per cent fine, light grey sandstone Green and brownish grey shale Predominantly fine, brownish grey sandstone and medium, light grey sandstone. Some grey, green, red, and light grey shale
2,290-2,320 2,320-2,340 2,340-2,350		Light grey, green, and red shale; a few coal fragments As above, with a little sandstone 80 per cent fine to medium, light grey sandstone; 20 per cent red shale
2,350-2,380 2,380-2,400	Gas reported at 2,390 (gas sand)?	Light grey, green, red shale Missing
2,400–2,410 2,410–2,450 2,450–2,470 2,470–2,477		90 per cent fine, light grey sandstone Variegated shale and fine sandstone Fine-grained, light grey sandstone Missing

Alberta Pacific Consolidated Oils, Limited, Well; L.S.D. 2, Sec. 4, Tp. 1, Range 16, W. 4th Mer., Elevation 3,585 Feet

Drilling		
depth	Horizons	Lithology
0- 60	60, base (?) of Upper Milk River	Drift and medium, brownish grey sandstone; a few fragments of coal
60 110 110 240	240, base (?) of Lower Milk River	Drift and brownish grey, medium sandstone Fine to medium, light grey sandstone
240- 310 310- 480 480-1,200		Missing Grey, shaly, fine sandstone and sandy, dark shale Mostly dark grey shale, partly sandy
	610, top of Bullopora zone 640, top of Clavulina	
4 800 4 000	zone	W
1,200-1,600 1,600-1,670		Dark shale and sandy, bentonitic shale
1,670-1,690		Grev. shalv sandstone
1,690-1,750		Grey, shalp sandstone Grey shale and brownish grey, bentonitic and sandy
1 770 1 000		shale: a little glauconite at 1.740
1,750-1,880		Dark shale, a little sandstone; traces of bentonite at 1,750-1,760, 1,780-1,800, and 1,840-1,850; a little glauconite at 1,760-1,770 and 1,810-1,820
1,880-1,930		50 per cent fine white sandstone
1,930-1,975	sand (water) 1,975, top of Lower Cre- taceous	Mostly dark grey shale, a little white sandstone
1,975-2,030		Grey, brownish grey, and greenish shale
2,030-2,040		60 per cent medium, light grey sandstone; a few coal fragments
2,040-2,130		Grey, light grey, brownish grey, and green shale
2, 130-2, 140		Grey, light grey, brownish grey, and green shale Fine grey, pyritic sandstone
2,140-2,160		Missing
2, 160-2, 180 2, 180-2, 190		Fine to medium, quartz and black chert sandstone
2, 180-2, 190		Fine grav gandstone
2,200-2,210		Grev and green shale sandy
2,210-2,220		Missing
2, 220-2, 240		(trev and green shale a little sandstone
2, 240-2, 260		Missing
2,260-2,270	l <i></i>	Light greenish grev and brownish grev shale
2,270–2,300 2,300–2,310		60 per cent fine to medium, grey sandstone; remainder
2,310-2,320 2,320-2,340		Mostly red shale, some green and brownish grey shale 60 per cent fine to medium grey sandstone; red, green, and grey shale
2,340-2,400		Missing
2,400-2,410	2,400, gas sand	Mostly white sandstone, a little red shale
2,410-2,430		Red shale
2,430-2,440		Green and red shale, some light greenish sandstone
2,440-2,447		Brownish grey, red, and green shale
2,447-2,452	9 459 top of Worolds oil	Missing
2,452-2,462	2,452, top of Vanalta oil	Fine to medium, quartz and black chert sand
2,462-2,468		Grey and greenish grey, sandy shale
2,468-2,480		Fine, light grey sandstone and green shale
2,480-2,487		Fine quartz and black chert sandstone
2,487-2,493		Grey and greenish shale and fine, greenish sandstone
2,493-2,500 2,500-2,515	Oil	Fine to medium, quartz and black chert sandstone
2,000-2,010	2,010, Wallet	permits urresult

Southern Alberta Exploration Company, Limited, Well No. 1; L.S.D. 8, Sec. 3, Tp. 1, Range 16, W. 4th Mer., Elevation 3,502 Feet

Drilling depth	Horizons	Lithology
0- 10 10- 40 40- 70 70- 220 220- 260 260- 360 360- 390 390-1,130	River	Weathered, fine to medium, yellowish sandstone Medium, light grey sandstone Fine, light grey sandstone and shaly sandstone; few coal fragments Dark grey shale and grey, sandy shale Dark shale Fine, grey sandstone and grey, sandy shale Dark shale, with traces of bentonite at 590-610 and 720-730
1,130-1,140 1,140-1,180 1,180-1,240 1,240-1,250 1,250-1,310 1,310-1,380 1,380-1,420		Fine, grey sandstone and dark shale Dark shale Dark shale with some bentonite Fine, grey sandstone and dark shale Dark grey shale; a little sandstone and bentonite Dark shale Fine-grained, grey sandstone, medium, light grey
1,420-1,460 1,460-1,520 1,520-1,610		sandstone, and dark shale Dark shale Dark shale and light grey, bentonitic shale Shaly and bentonitic, fine, grey sandstone and grey shale
	water sand	Dark shale Grey, sandy shale and light grey, brownish grey, fine and medium sandstones, partly bentonitic and con- taining some glauconite
1,860-1,870 1,870-1,890 1,890-1,900 1,900-1,910 1,910-1,920		Dark shale Dark shale Dark shale and fine, grey sandstone
1,920-2,110		Grey, green, and red shale with trace of bentonitic material
2, 110–2, 150 2, 150–2, 230	Bituminous shale, 2,150- 2,170	Grey and green shale with 30 per cent fine to medium, light greenish sandstone; a little red shale Brownish grey shale, impure bentonitic material, and a little fine, greenish grey sandstone; some of the
2,230-2,240 2,240-2,260		brown shale is highly bituminous Missing Fine, grey sandstone; medium, light grey sandstone; grey shale
2,260-2,270 2,270-2,300		Red and light grey shale Fine, light grey sandstone; grey shale and a little red shale
2,300–2,340 2,340–2,350 2,350–2,360 2,360–2,390 2,390–2,410 2,410–2,440	2,340, gas sand (?)	Grey, light grey, green, red shale Fine, light grey sandstone Grey shale Missing Grey, light grey, and green shale Medium to coarse, quartz and black chert sand
	2,429; Oil at 2,429- 2,439	

Commonwealth Petroleum, Limited, Red Coulée Well No. 1; L.S.D. 10, Sec. 3, Tp. 1, Range 16, W. 4th Mer., Elevation 3,527 Feet

Drilling depth	Horizons	Lithology
0- 20		Drift; some yellowish sandstone
	River	77 11 11 11 11 11 11 11 11 11
20- 90		Yellowish, medium-grained sandstone
90 160 160 200		Fine to medium, light grey sandstone Fine, light grey, shaly sandstone
100- 200	River	rine, nght grey, shary sandstone
200- 670		Dark shale and sandy shale; little grey sandstone,
	zone	some beds limy
670-1,180		Dark shale
1,180-1,220		Dark, sandy shale and fine, grey sandstone
1,220-1,280		Dark shale, bentonitic
1,280-1,310		
1,310-1,370		Dark shale and sandy shale; some fine, grey sandstone Dark shale, some bentonite
1,370-1,450 $1,450-1,470$		Sandy shale and 40 per cent medium, quartz and black
1,400-1,410	1	chert sandstone
1,470-1,520	1	Dark shale, a little thin sandstone
1,520-1,530		Dark shale and bentonitic, medium sandstone
1,530-1,680		Dark shale, some bentonite
1,680-1,700		Dark shale and bentonitic, fine to medium sandstone
1,700-1,740		Dark shale
1,740-1,820		Dark shale and fine to medium sandstone with some
1 000 1 050		glauconite
1,820-1,850 1,850-1,890		Grey shale, some bentonite Grey shale and bentonitic, sandy shale
1,890-1,910		Dark shale and fine, grey sandstone
1,910-1,920	1 910 Colorado water	Medium-grained sandstone and dark shale, a little
1,010-1,020	sand	glauconite
1,920-1,980	1	Dark grey shale, some sandy and some bentonitic beds
1,980-2,020	1,980, top of Lower Cre-	Grey and greenish grey shale, some greenish, bento-
	taceous	_ nitic sandstone
2,020-2,100		Dark grey, light grey, green, and red shale
2,100-2,150		Fine to medium, grey and greenish grey sandstone
2,150-2,160		Grey, green, and red shale Grey, green, and red shale; minor grey and greenish
2,160-2,270		sandstone; little coal
2,270-2,280		Missing
2,280-2,290		Grev, green, and red shale
2,290-2,310		Grey and green shale, some red shale, 20 per cent white,
		medium sandstone
2,310-2,370		
2,370-2,390		Red shale
2,390-2,400		60 per cent medium, light grey sandstone
2,400-2,430 2,430-2,450	?Gas sand	Grey, greenish, red shale, some grey sandstone Medium, light grey sandstone
2,450-2,450	Gas sand	Brownish shale
2,495-2,505		Brownish shale and medium to coarse, grey sandstone
2,505-2,553	Vanalta oil sand, water	Medium to coarse, quartz and black chert sandstone;
	at 2 535	nyrite some coal fragments
2,553-2,610	2,553, top of Ellis	Light greenish grey shale and shaly limestone; pyrite;
	1	some calcareous and glauconitic sandy beds at base
2,160-2,614	2,614, water	White, sandy dolomite, 30 per cent sand
2,614-2,625	0.005 4612	Sandy dolomite and grey shale
2,625-2,732	2,029, top of limestone	White limestone; some grey shale 2,689 to 2,695
	<u> </u>	1

Celtic Oils, Limited, Red Coulée Well No. 1; L.S.D. 4, Sec. 17, Tp. 1, Range 16, W. 4th Mer., Elevation 3,531 Feet

Drilling depth	Horizons	Lithology
0- 110 110- 150		Mostly drift; some yellowish sandstone Light brownish sandstone; some traces of coal
150- 290	290, base of Lower Milk River	Fine to medium, light grey sandstone, quite shaly
290- 340	niver	200–220 and at base Dark grey shale with thin bands of grey sandstone
340- 350 350- 360		Missing
360- 370		Dark grey, sandy shale Missing
370- 410		Dark grey shale
410- 440		Dark shale ribboned with thin sandstone layers; thin ribbons of coal
440 460 460 470		Dark grey shale
470- 500		Dark grey shale
500- 750	l	Missing
750-1,100 1,100-1,140		Dark grey shale and fine grey sandstone
1,140-1,280		Dark grey shale, traces of bentonite at 1,170-1,180 and 1,230-1,240
1,280-1,340		Dark grey shale and sandy shale and fine, grey sand- stone
1,340-1,440 1,440-1,480		Dark grey shale, bentonitic 1,340-60 Dark grey shale with 30 per cent fine, grey, shaly sand-
1,480-1,490	 	stone Missing
1,490-1,560		Dark grey shale with bentonite
1,560-1,570 1,570-1,590		Fine, grey, shaly sandstone Dark grey shale
1,590-1,600		30 per cent medium, quartz and black chert sand- stone, remainder shale
1,600-1,680 1,680-1,930		Dark shale, bentonitic and slightly sandy Dark shale, sandy at 1,730-1,750, bentonitic at 1,770- 1,820 and 1,920-1,930, glauconitic at 1,890-1,900
1,930-2,040	sand	Dark shale; some glauconite at 1,930-50, 50 per cent
2,040-2,080	2,040, top of Lower Cre- taceous	Fine to medium, light grey and greenish grey sand- stone
2,080-2,240		Grey, brownish grey, and green shale, somewhat sandy
2,240-2,250 2,250-2,270	Bituminous shale	Missing Dark bituminous shale
2,270-2,290 2,290-2,300		Light grey sandstone, grey and greenish grey shale
2,300-2,330		Mostly fine, light grey sandstone
2,330-2,380		Grey, greenish grey, brownish grey shale, and some grey sandstone
2,380-2,420		Missing
2,420-2,540 2,540-2,550	2.540, top of Varalta	Dark grey, grey, brownish grey shale, some red shale Medium, quartz and black chert sandstone, with minor
2,550-2,560	sand; water at 2,543	grey, green, and red shale Fine, white sandstone and medium, quartz and black
		chert sandstone
2,560-2,570 2,570-2,580	Water at 2,573	Missing Fine to medium, quartz and black chert sandstone
2,580-2,630	2,580, top of Ellis	
2,630-2,650		Greenish grey shale, with 60 per cent of fine, light greenish sandstone; pyrite, black chert grains, and a little glauconite
2,650-2,660 2,660-2,670	2,660, top of limestone.	As above, with a little white limestone

Dalco Oil and Gas Company, Limited, Red Coulée Well No. 1; L.S.D. 4, Sec. 11, Tp. 1, Range 16, W. 4th Mer., Elevation 3,508 Feet

Drilling depth	Horizons	Lithology
0-220	220, base of Lower Milk River	Missing
220-1,550		
2,150-2,170 2,350-2,360		Some bituminous shale in fine, greenish sandstone
2,420-2,435		Reported as coarse sand, very little sand present in samples, which are mostly red shale; typical Van- alta sand not indicated in samples
	2,560, sulphur water 2,570, top of limestone	Looks like Ellis

Norm. An outline only is given of this well, since Lower Cretaceous appears in the samples at 1,550 feet in disagreement with the drilling records. The writer is of the opinion, however, that the samples are correct from 2,150 feet down, as the bituminous shales at that depth agree with those found in the Southern Alberta Exploration well No. 1 and the Celtic No. 1 well.

Dalco Oil and Gas Company, Limited, Red Coulée Well No. 2; L.S.D. 7, Sec. 19, Tp. 2, Range 17, W. 4th Mer., Elevation 3,630 Feet

Drilling depth	Horizons	Lithology
0-50 50- 120	120, base of Foremost	Missing Drift; grey sandstone, dark carbonaceous shale, a few coal fragments
120-290	290?, base of Pakowki	Grey and slightly greenish grey, fine sandstone; shaly sandstone and sandy, grey shale; foraminifera pre sent at 130-150 and 280-290
290- 310		Mostly cavings of drift; a little grey shale and grey sandstone
310- 340		Light grey, shaly sandstone, a little greenish sand stone, and white "mud"
340- 380		Mostly white, sandy "mud," some greenish and brown shale
380 400		Light grey and greenish grey shale, some brown shale
400 420		Fine, light grey, shaly sandstone
420- 430		Fine, light grey sandstone, brown shale, and coal frag
430- 440		Grey, sandy shale and shaly, medium sandstone
440- 460		Light grey, medium sandstone
460- 470		As above, with some brown shale
470- 480		Light grey, sandy shale and brown shale
4 80- 540		Light grey, fine and medium sandstone and shaly sand stone
540- 650	Lower Milk River	Light grey, fine to medium sandstone and shaly sand stone
650-1,890	1,040, top of Bullopora	Dark grey shale, some sandy shale
	1,070, top of Clavulina	
	1,900, water	
1,890-1,920		Dark shale and some light grey, medium sandstone a little bentonite
1,930-1,970		Dark shale, a little bentonite

SOME CLAY DEPOSITS OF WILLOWBUNCH AREA, SASKATCHEWAN

By F. H. McLearn

CONTENTS

	PAGE
Introduction	31
Stratigraphy	32
Horizons of coal seams	34
Clays	35

INTRODUCTION

In the field season of 1930 the west half of the Willowbunch sheet¹ was studied. In particular, areas in the vicinity of Willows, Willowbunch lake, Bonneau lake, Harptree, Big Muddy, and Rockglen were examined. Particular attention was devoted to stratigraphy, structure, oil possibilities, and the clay deposits. The results regarding oil possibilities are largely negative. In this preliminary report, therefore, on last season's activities, attention is confined chiefly to the clay deposits. A brief account of the stratigraphy is given, however, for a thorough understanding of it is a necessary prelude to all economic studies in southern Saskatchewan.

Very satisfactory field assistance was given by O. L. Backman, Edward Leith, and L. S. Russell. All instrument work, including the running of levels, was done by Backman. The clay tests were made in the Ceramics Division of the Mines Branch, H. Frechette, Chief of Division, and were reported on by J. G. Phillips. Acknowledgment is made to Professor W. G. Worcester for valuable advice on clay deposits.

By Lake of the Rivers valley is meant the valley containing Lake of the Rivers and Willowbunch lake. By Twelvemile Lake valley is meant the valley containing Twelvemile lake, Bonneau lake, Rivard lake, etc. The two valleys unite near sec. 6, tp. 4, range 25, W. 2nd mer., to form Big Muddy valley which extends eastward to Big Muddy lake and thence directly south to cross the International Boundary in sec. 3, tp. 1, range 22, W. 2nd mer.

The refractory and semi-refractory clays are the most important of the mineral resources of southern Saskatchewan. They are suited to the manufacture of a great variety of ceramic products, include some of the most refractory clays in Canada, and form the basis of a growing industry.

From an early date the Geological Survey and Mines Branch have directed attention to these clays and have endeavoured to encourage their utilization. In 1885, McConnell noted the "white band" near the bottom of the "Laramie." In 1895, Hoffmann examined a sample from sec. 28, tp. 12, range 24, W. 2nd mer. (Claybank), recorded its refractoriness and recommended its suitability for fairly refractory fire-brick and pottery. In 1906, Chambers observed the winter haulage of clay from the hills near Wood Mountain to Moose Jaw for manufacture into common brick and

¹Dept. of Interior, Top. Surv., Canada, sec. sheet No. 19. ²McConnell, R. G.: Geol. Surv., Canada, Ann. Rept. 1885, N. S., vol. I, pt. C, pp. 27, 28 (1886). ³Hoffmann, G. C.: Geol. Surv., Canada, Ann. Rept. 1892-3, N. S., vol. VI, pt. R., p. 91 (1895).

some fire-brick.¹ The next records, those of Ries and Keele, based on field work of 1910 and later years and published between 1912 and 1915, are more elaborate and include the results of actual sampling of deposits and laboratory tests.² Between 1914 and 1916 Rose published the results of some clay tests.³ Davis' valuable work, published in 1916, treated the clays in very considerable detail.⁴ These studies between 1910 and 1916 gave great encouragement to the commercial development of this important resource. The recent work of the Geological Survey has included primarily a study of the distribution of the deposits as controlled by geological conditions, a study of the origin of the deposits, and the location of some prospects.⁵

Worcester notes that the first private report on the clays, known to him, was by Professor Edward Orton, jun., in 1907, and was based on samples from Dirt hills, collected by Daniel Diver, a prospector. Worcester's own valuable investigations, made under provincial auspices, have included much sampling of deposits in the field, laboratory tests, and the devising of methods of treatment and utilization.

STRATIGRAPHY

The formations in the west half of the Willowbunch sheet, studied in 1930, are given in the following table, as well as some important stratigraphic data.

As the stratigraphy has been fully treated in previous reports.8 a description of it need not be repeated here. It is not known whether the Triceratops (Lance) or basal zone of the Ravenscrag extends as far east as Willowbunch area. Near Willows dinosaur bones were found on the surface of beds just above the Whitemud, but whether they came from the bottom of the Ravenscrag or from the glacial drift could not be proved. Owing to lack of persistent and reliable lithologic boundaries it is not possible to subdivide the Ravenscrag nor to apply to the Canadian sections the formational classification used in Montana. Divisions into grey and yellow or buff facies have locally considerable value, but it is not proved that they are of general application to all of southern Saskatchewan. It does not seem possible to make the beds above and below the Willowbunch member separate mapping units or formations, for the Willowbunch member is difficult to identify and trace; in places it passes laterally into beds difficult to recognize as typical of it and refractory and non-refractory beds do occur at other horizons in the Ravenscrag. Moreover, the beds over and below it are not markedly different. The Colgate member in Montana is possibly an equivalent of the Whitemud beds.

⁹See Thom, W. T., jun., and Dobbin, C. E.: Bull. G.S.A., vol. 35, pp. 481-506 (1924).

¹Chambers, R.: Geol. Surv., Canada, Sum. Rept. 1905, p. 74.

²Ries, Heinrich, and Keele, Joseph: Geol. Surv., Canada, Mem. 24E (1912); Mem. 25 (1913); Mem. 65 (1915);

Mem. 66 (1915).

³Rose, Bruce: Geol. Surv., Canada, Mem. 89, pp. 69-83 (1916).

⁴Davis, N. B.: Mines Branch, Dept. of Mines, Canada, "Rept. on Clay Resources of Southern Saskatchewan" 1918).

^{(1918). **}McLearn, F. H.: Geol. Surv., Canada, Sum. Rept. 1927, pt. B, pp. 21-43; 1928, pt. B, pp. 30-44; 1929, pt. B, pp. 48-63.

Worcester, W. G.: Trans., Can. Inst. Min. Met., 32, p. 256 (1929).

Worcester, W. G.: Trans., Can. Inst., Min. Met., 32, pp. 255-269 (1929). Also other papers, including unpublished reports to the Saskatchewan government.

McLearn, F. H.: Geol. Surv., Canada, Sum. Rept. 1927, pt. B, pp. 24-39; 1928, pt. B, pp. 30-42; 1929, pt. B, pp.

TABLE I

	100				T	Γ	Ä			215				
il seams	Names				BIG MUDDY		HARPTREE MINE	KEOGH	ST. VICTOR	ANXIETY BUTTE		WHITEMUD		
Some coal	Horizons													
Volcanic	ash													
Refractory,	clays	•												
Floras					Upper flora		Upper flora		Middle flore	(r. crassinolis)	Lower flore (F. ceratops) present ?	Whitemud flora (Nelumbites n.sp)		
Vertebrate	faunas	Merychippus							Leidyosuchus	Echmatemys	Triceratops present ?		Sauropterygian rib	
Invertebrate	faunas				V. trochiformis		V. trochiformis					Anodonta sp. l.	Anodonta sp.2.	
FORMATION		Wood Mountain		Absent	-3	Willowbunch	ŞE	scr	uə <i>ı</i>	Ra	Basal part present ?	Whitemud	Sandstone E	Bearpaw
AGE	ENE		OLIGOCENE	E LATER EOCENE (POSTPALEOCENE)	INE		CENE) EOCE				sı	CEO	LA ETA	ES CE
			10	YAAI							WESOZOIC			IW

The stratigraphic data shown in the table include faunas, floras, refractory and semi-refractory clay zones, a volcanic ash zone, and some coal seams. In the following tentative correlation of coal seams the assumption is made that refractory and semi-refractory, clay-bearing zones, identified as the Willowbunch member in the cliffs south of Big Muddy and Harptree, are exactly equivalent stratigraphically to the typical Willowbunch south of the town of Willowbunch.

HORIZONS OF COAL SEAMS

By Big Muddy seam is meant the thick one over the Willowbunch member south of Big Muddy lake. At a similar horizon, but not necessarily exactly the same seams nor exactly at the same horizon, are the thick seam over the Willowbunch zone south of Harptree and the Willowbunch seam which just overlies the Willowbunch clay member south of the town of Willowbunch.

By Harptree Mine seam is meant the one high in the Ravenscrag but below the Willowbunch member and the one mined south of Harptree in NW. ½ sec. 34, tp. 3, range 26, W. 2nd mer. In the same area, but at a lower horizon, is the Contact seam, which occurs at the contact between a grey below and a yellow or buff facies above. The beds for a few feet just over it are white, pale grey, and pale purplish weathering silts and clays and with the Contact seam form locally an important key horizon or zone. Where measured in SW. ½ sec. 3, tp. 4, range 26, W. 2nd mer., it is 1 foot 8 inches thick, but it is thicker at other localities.

It is probable that the Keogh seam south of Big Muddy lake is at a somewhat higher horizon than the Contact seam. The Keogh occurs near the top of a grey facies in the area south of Big Muddy lake and is the one worked on a small scale near the Keogh ranch buildings. Locally two seams of fair thickness occur over it.

West of Harptree area, and south and west of Gye, is a fairly thick seam worked in an adit and an open pit near the middle of sec. 26, tp. 4, range 27, W. 2nd mer., near valley-bottom level. This may be called the Gye seam. It is at the same or at a somewhat lower horizon than the Contact seam to the east. A seam which occurs a little above it in Gye area may be designated the Bonneau seam. A seam mined near valley-bottom level at Willowbunch is at an horizon similar to that of the Gye seam.

Yet, farther west near St. Victor the St. Victor seam lies just over the volcanic ash horizon and possibly at a little lower horizon than the Gye seam, but not much lower. It is mined east of St. Victor in the Brenner mine. Two or three seams of fair thickness occur over the St. Victor seam in the yellow or buff facies, west of St. Victor on the south side of the valley.

East of Twelvemile lake a small seam is present near the top of the grey facies and just under the horizon at which volcanic ash discontinuously occurs. This is the Krieg seam. It thins out eastwardly toward St. Victor, just as the slightly higher St. Victor seam does not extend west to Twelvemile lake. The Krieg seam has two small seams under it in the grey facies and at least two fair seams over it in the yellow or buff facies in the area east of Twelvemile lake.

The first seam in the Ravenscrag is poorly developed in Lake of the Rivers valley and Big Muddy valley. It is more distinct and thicker at Twelvemile lake, but is variable there. Far to the west, in East End area, the No. 1 or Ferris seam is more uniform and persistent. It is not necessarily the same seam nor exactly at the same horizon as the first seam in the east, however. In East End area seams occur over the No. 1 seam, including the Anxiety Butte seam, which occurs at the base of the cliff at Anxiety butte and has been mined there.

By Anchor seam is meant the one low in the Ravenscrag formation, about 55 feet over the No. 1 seam, on the south shore of Willowbunch lake and recently opened near the west boundary of sec. 10, tp. 6, range 27, W. 2nd mer. It is a useful horizon marker locally. The No. 1 seam is here 1 foot 8 inches, including some lignitic shale, but not including an 8-inch shale parting. It occurs just over the Whitemud beds and is the lowest seam in the area, as there is no coal below the Whitemud beds. As pointed out in previous reports there is no coal-bearing Estevan formation below these Whitemud beds.

Another seam that may have some stratigraphic value in Wood Mountain district is the Vogelberg seam in the buff facies of the Ravenscrag formation, more than 100 feet below the Willowbunch member. It is worked in the Six-mile mine in sec. 16, tp. 4, range 4, W. 3rd mer. In this district, the No. 1 seam, a little over the Whitemud beds, is a well-developed

and workable seam and a good horizon marker.

CLAYS

Refractory and semi-refractory clays occur at two horizons. The lower is that of the Whitemud formation in late Cretaceous age and the

higher is that of the Willowbunch member of Palæocene age.

The origin of the Whitemud clays and the geological conditions affecting their distribution have been considered by the writer in recent reports.² Only a few observations need be added now. The writer's theory of intra-regional origin includes the kaolinization of feldspathic sands after deposition and in Whitemud time. This part of the theory is also held by Worcester.³ C. S. Ross, describing the sands of the Colgate member of Montana, notes that "Clay minerals make up a considerable part of the interstitial material and most of these appear to have formed in place by the alteration of detrital rock grains."

Additional evidence supporting a non-marine origin of the Whitemud beds was obtained. At one locality the freshwater pelecypod *Anodonta* was found. At another very fine, paper-thin, bedded clays carry leaves of the "water lilies" *Nelumbo* and *Nelumbites*, which record a freshwater

aquatic environment.

Long, vertical, tap-like roots occur in some of the Whitemud beds. That these were roots penetrating downward and not buried stems or stalks can be proved, for they pierce fossil leaves lying prone along the bedding. This is considered to be evidence that the water-table, at times at least,

 ¹McLearn, F. H.: Geol. Surv., Canada, Sum. Rept. 1928, pt. B, p. 35; 1929, pt. B, p. 58.
 ²McLearn, F. H.: Trans. Roy. Soc., Canada, 3rd ser., vol. 22, p. CIX (1928); Geol. Surv., Canada, Sum. Rept. 1927, pt. B, pp. 25-39; 1928, pt. B, pp. 35-38; 1929, pt. B, pp. 51-55.
 ³Worcester, W. G.: Trans. Can. Inst. Min. Met., vol. 32, p. 260 (1929).
 ⁴Dobbin, C. E., and Reeside, John B., jun.: U.S. Geol. Surv., Prof. Paper 158 B, p. 25 (1929).

was at some appreciable depth below the surface. A periodic fall of the water-table to promote weathering was postulated¹, as part of the local or intraregional theory of origin of the Whitemud beds.

Although the Whitemud clays of the area studied in 1930, i.e., the west half of the Willowbunch sheet, vary from place to place, they maintain a good average quality and some of the best clays of Saskatchewan are found in this area and at this horizon. The original deposition and alteration all seem to have been largely favourable to them and they were little affected by the erosion at the end of Whitemud time which destroyed so much of the clays in other areas. The best clays occur at the tops of the sections.

Over most of the west half of the Willowbunch sheet the Whitemud formation is deeply buried. It occurs, however, on both sides of Lake of the Rivers valley from the south end of Lake of the Rivers to the west end of Willowbunch lake, where it is carried below lake-level in about sec. 9, tp. 6, range 27, W. 2nd mer., by the east dip. It reappears in a low, broad arch on both sides of Big Muddy valley between sec. 2, tp. 4, range 25, W. 2nd mer., and the west end of Big Muddy lake, where it is carried below lake-level by the east dip.

The clays of the Willowbunch member high in the Ravenscrag formation occur just under the highest uplands. They are widely distributed in the west half of the Willowbunch sheet. The upland south of the town of Willowbunch is high enough to carry them and they are exposed in the tributary coulées heading southward into this upland. They can be traced in a southeastward direction along the north border of the upland to a long tributary coulée west of Bonneau lake in sec. 4, tp. 5, range 27, W. 2nd mer. Exposures reappear on the south side of the main valley south of Harptree, where the upland is also high enough to carry the clays of this zone. The high upland in the vicinity of Buffalo gap and the highest part of the upland east of the Bengough (34) highway on the south side of Big Muddy valley carry them. They are also exposed in the northern edge of the high upland south of Big Muddy lake, on both sides of Big Muddy valley between the lake and the International Boundary, and on a small, high, upland remnant in about sec. 6, tp. 2, range 23, W. 2nd mer. A small, and the highest part of the ridge, near the International Boundary in about sec. 2, tp. 1, range 24, W. 2nd mer., appears to carry it. In Rockglen valley, north, south, and southeast of the town of Rockglen the highest ridges carry what appear to be clays of this zone.

Small sections of semi-refractory or refractory clays occur also at other horizons near that of the Willowbunch member. South of Harptree there is a thin zone about 40 or 50 feet over it. In the Rockglen there is such a zone 80 or 100 feet below the probable Willowbunch member. The occurrence of these additional zones makes the identification of the true Willowbunch member difficult and even tentative in places, e.g., at Rockglen and possibly south of Harptree.

The Willowbunch clays vary much from place to place, are mostly in thinner sections than those of the Whitemud, and on the whole are not the commercial equal of them. Locally, however, very good Willowbunch clays are found.

McLearn, F. H.: Geol, Surv., Canada, Sum. Rept. 1927, pt. B, p. 36.

The tests recorded below are not complete, but serve to show the location of some promising refractory clay deposits.

The following symbols are used in the tables:

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Sd = selected, not general, sample.
S = sticky
St = stiff
p = plastic
f.p. = fairly plastic
v.p. = very plastic
s.p. = slightly plastic
v.s. = very sticky
G = greasy
T = tough
p.c.e. = pyrometric cone equivalent
```

In the tables, where the grade has been determined, it is recorded as such, i.e., clay, sandy clay, etc. Elsewhere the rock term, based on field observation, is given, i.e., shale, etc. In the tables all beds are given in their natural stratigraphic order, i.e., the highest beds are at the top of the page.

WHITEMUD SECTIONS

Good exposures of Whitemud clays are found in the vicinity of Willows, and several quarries have been operated. An unworked deposit, containing good clays, occurs in SW. ½ sec. 7, tp. 8, range 28, W. 2nd mer. A section is as follows:

Thick	ness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0 1 1 2 3 1 0 3 2 1 1 0 0 2 0 1 0 0 0 1 0 0 0 0 0 0 0 0	2 0 3 0 0 6 9 0 0 6 10 10 6 8 8 6 6 8 6 6 8	W1189 W1186 W1186 W1185 W1184 W1183 W1190 W1192 W1193 W1194 W1195 W1196 W1197 W1198 W1199 W1200 W1201 W1202 W1203 W1204	Coal Purplish grey clay Purplish brown clay. Purplish brown clay. Pale mauve-grey clay, yellow stains. Pale mauve-grey clay Brownish clay. Brownish clay. Thin beds, pale mauve-green clay and silt. Brownish grey clay Brownish grey clay Light grey, silty clay Light brownish grey clay Light brownish grey clay Light brown clay. Grey clay Light grey, sandy clay. Brown, carbonaceous, silty clay. Light brownish grey, hard clay. Grey, silty clay. Brownish grey, hard clay. Grey, hard, silty clay. Dark brown, hard clay.	33 · 6 33 · 3 31 · 6 29 · 6 30 · 0 30 · 0 28 · 3 34 · 0 32 · 0 32 · 0 32 · 0 34 · 3 22 · 0 36 · 3 34 · 6 26 · 4 32 · 3 27 · 0	v.p. v.p. v.p. p p v.p. v.p. v.p. v.p.	32 32 29 30 28 31 27 28+ 31 29 28+ 26 20 25+ 15 20 14 14

Nos. W1189, W1188, and W1190 are very high-grade refractory. The others from W1197 up are second or low-grade refractory clays.

In SW. ½ sec. 3, tp. 8, range 29, W. 2nd mer., on the north side of a coulée northwest of Willows, the lower part of the Whitemud is exposed. The high zones with plastic clay have been eroded here, but are probably present farther back in the bank. Only the lower zone of sandy clays is present. The section exposed is as follows:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
2 2 2 5 5 4 0	8 0 2 8 0 0 8	W1205 W1206 W1207 W1208 W1209 W1210 W1211 W1212 W1213	Grey clay, boulder clay, etc. Light grey, sandy clay, grey clay lenses. Light brown clay. Light brown, sandy clay. Light grey, sandy clay. Light grey, very sandy clay. Light grey, sandy clay. Light brownish grey, very sandy clay. Light brownish grey, very sandy clay. Light grey, very sandy clay. Light grey, very sandy clay. Cay. Light grey, very sandy clay. Grey, very sandy clay.	24·0 29·3 20·0 22·0 22·3 25·0 19·0 20	p v.p. p s.p. s.p. s.p. s.p. s.p.	28 20 27 29 23 20 14 13

The above section is interesting for it shows decreased refractoriness below. Mechanical separation tests show that this is co-ordinated with less clay content below. Therefore, the kaolinization that took place in Whitemud time was dying out below.

In the SE. $\frac{1}{4}$ sec. 24, tp. 7, range 29, W. 2nd mer., on the north side of a small coulée to the west of the highway, a rather thin section of the Whitemud occurs:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
2 1 1 3 5 0 0	6 0 5 6 6 3 5	W1354 W1353	Pale brownish grey clay Pale grey, iron-stained clay Pale grey, silty clay. Pale grey, sandy clay White, sandy clay Clay ironstone. Brown, silty clay. Grey, silty clay.	27·3 27·3 25·6 24·3 34·6	v.p. v.p. v.p. f.p. p	28 26 26 29 29 29

The top of the section was probably not obtained here.

There is a good deposit of Whitemud clays in SW. ½ sec. 30, tp. 7, range 28, W. 2nd mer.

Thic	kness	Sample No.	Colour, grade	Tempering water,	Working properties	P.c.e.
$\mathbf{Ft.}$	Ins.	140.		per cent	properties	
0 0 0 0 1 1 1 4 2 1 0 1	11 2½ 9 3 3 11 6 8 10 9		Lignite Brownish grey, carbonaceous silt Lignite Brown clay. Grey clay. Grey clay. Grey clay, iron spots. Grey clay, iron spots. Brownish grey clay. Pale brownish grey clay. Brownish grey clay. Brownish grey clay.	34·0 32·3 28·3 30·0 28·3 30·6		
2 2 2 2 1	11 2 5 4	W1277 W1276 W1275 W1274	mostly of latter Grey, sandy clay. Grey clay. Light grey, sandy clay. Very fine sandstone and dark grey clay.	22·6 31·3 22·6	v.p., G p v.p. p	30 30+ 29 29
4	8	W1273	Light grey, sandy clay	22.0	s.p.	28+
2	6	W1272	Thin layers grey, very fine, sandy, and			10
2	0	W1271	Grey, sandy clay		v.p. p	19 23

The two top beds, samples W1269, W1268, are very high-grade refractory clays. The beds with samples W1267, W1266, W1265, W1280, and W1278 are high grade.

North of highway 13 and east of the Canadian Pacific Railway trestle, at end of a spur, in SE. ½ sec. 6, tp. 7, range 27, W. 2nd mer., is a deposit of Whitemud clay:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
5 2 3 3 1	0 9 0 6 0 5	W1219 W1218 W1217 W1216 W1215 W1214	Fissile green shale. Pale mauve-grey clay. Light mauve-grey clay. Light mauve-grey, silty clay. White, silty clay. Light grey to white clay, silty clay. White, sandy clay.	32·3 31·0 30·0 30·3 28·0 25·6	v.p. v.p. v.p. p	31 31 31 30 29 30

The clays of W1219, W1218, and W1217 are very high-grade refractory.

Just south of the above and just north of highway 13, east of trestle, is a small pit from which a test shipment has been made. Three feet of clay in the upper part of the exposed section gave the following results:

Thie Ft.	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
3	0	W1178	Light grey clay	30.0	v.p.	28+

This is a low-grade refractory clay.

In a coulée on the south side of Willowbunch lake, near its western end and about in sec. 7, tp. 6, range 27, W. 2nd mer., is a good deposit of the Whitemud. The section exposed just south of a fault on the east side of the coulée is as follows:

Thic	kness	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0	7 3		Lignite			
0	8		Lignite			
ő	4		Dark purplish grey shale			
2	2	W1307	Light brownish, mauve-grey clay	30.0	v.p.	30+
ĩ	4	W1306	Light brownish, mauve-grey clay,			
-	_		iron stains	32.0	v.p., S	30+
2	8	W1305	Light mauve to brownish grey clay	30.3	v.p.	30+
4	0	W1304	Grey clay	28.6	p, G	30+
1	0	W1303	Yellow clay-many minute ferrugin-			
			ous concretions		p	15+
0	10	W1302	Grey clay		v.p.	30+
0	10	W1301	Grey clay		v.p., G	28+
6	6	W1300	Brownish, grey, silty clay		v.p.	28+
1	4	W1299	Grey to white, sandy clay		p	30+
2	5	W1298	Grey to white, sandy clay		f.p.	28+
1	5	W1297	Grey to white, sandy clay		s.p.	28+
2	4	W1296	Grey to white, sandy clay	21.6	s.p.	28
	l	l	1	1		

Samples W1307, W1306, W1305, W1304, and W1302 are high-grade refractory clays. From W1301 down are low-grade refractory clays.

On the south side of Willowbunch lake and in approximately the SW. $\frac{1}{4}$ sec. 10, tp. 6, range 27, W. 2nd mer., at about lake-level, is the following section:

Thie	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
5		W1176Sd	Shale and ligniteGrey clay, minute ferruginous concretions	30.7	v.p.	32

The grey clay is a very high-grade refractory. The sample, however, is not representative of the entire 5 feet and the occurrence can only be regarded as a promising prospect for refractory clay.

On the south side of Big Muddy valley west of the Bengough highway (34), and near the centre of sec. 27, tp. 3, range 24, W. 2nd mer., is a good deposit of Whitemud clays.

Thic	kness	Sample	Colour, grade	Tempering water.	Working	P.c.e.
Ft.	Ins.	No.	Colour, Brade	per cent	properties	
0	3 4		Lignite and lignitic shale			
0	4 2 6		ShaleLignite.			
7	6	W1167	Brownish grey clay Dark, coaly shale	28.3	v.p.	32
4	6 0	W1166 W1165	Very dark, carbonaceous clay Dark brown, sandy clay, bands white silt		p s.p.	31 26
3	6	W1164 W1163	Light brownish, silty clay Brownish grey clay	29·3 33·3	p v.p.	20- 20
1 4 4	0 6	W1162 W1168 W1169	Light brown clay Light grey, sandy clay Light grey, sandy clay	22.6	v.p. s.p. s.p.	17 29 29
3 4	2 8	W1170 W1171	Grey, coarse, sandy clay Light grey, sandy clay	23.3	s.p.	29 28

Samples W1167 and W1166 are very high-grade refractory.

At the end of a spur in NE. $\frac{1}{4}$ sec. 24, tp. 3, range 25, W. 2nd mer., the top of a section of the Whitemud is exposed:

Thiel	Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0	5	W1465	Bentonite	61.6	T, v.s.	8
0	9		Brown, fissile shale			
Ŏ	4		Brown to black, carbonaceous shale, top of Whitemud			
1	2	W1464	Greyish brown clay	29.7	v.p.	31
1	5	W1463	Grey clay	29.0	p	31
1	9	W1462Sd	Light grey silt, admixed ferruginous			
			concretions and yellow clay not in- cluded in sample	28.6	p	31_

These are high-grade refractory clays.

Some good clay occurs near the top of a section on the south side of Big Muddy valley, in about the middle of sec. 25, tp. 3, range 25, W. 2nd mer.

Text	Thic	kness	Sample	Colour, grade	Tempering water,	Working	P.c.e.
1 W1461 Dark brown, carbonaceous clay, ferruginous S.p. 31	Ft.	Ins.	No.	557542, 52545		properties	
Tuginous S.p. 31			W1461				
0 6 W1460 Brown clay 28.6 p 31 2 5 W1459 Grey clay, some ferruginous concretionary lenses 25.0 p 29 3 8 W1458 Pale greyish green clay, ferruginous lenses 29.3 v.p. 30 1 5 W1457 Pale green, silty clay, ferruginous lenses 29.3 v.p. 29 0 9 W1456 Greyish green clay, ferruginous lenses 29.3 v.p. 29 1 5 W1455 Pale greenish grey clay, dark streaks 30.3 v.p. 29 1 5 W1455 Pale greenish grey clay, dark streaks 30.3 v.p. 29 1 5 W1455 Pale greenish grey clay, dark streaks 30.3 v.p. 29 1 7 W1454 Pale prewish grey, silty clay 32.6 v.p. 2 0 8 W1453 Grey, silty clay 32.6 v.p. 2 0 8 Prown, salty clay 2	1	1	W 1401		31.0	8.D.	31
tionary lenses	0	6	W1460				31
3 8 W1458	2	5	W1459			-	
1 5 W1457 Pale greyish green, silty clay, ferruginous 29.6 p. 26 p. 26 p. 27 28 29 29 29 29 29 29 29	_	_		tionary lenses		p	29+
1	3	8	W1458	Pale greyish green clay, ferruginous			
1	1	_	7371457	Pole gravish graen silty elev formus		v.p.	30
0 9 W1456 Greyish green clay, ferruginous lenses 29·3 v.p. 29 1 5 W1456 Pale greenish grey clay, dark streaks 30·3 v.p. 29 1 7 W1454 Pale brownish grey, silty clay 32·6 v.p. S 0 8 W1453 Grey, silty clay 34·0 p 29 0 8 Brown, carbonaceous and lignitic shale. 34·0 p 29 1 9 W1452 Brown, silty clay 32·3 p 27 0 6 W1450 Brownish grey clay 25·0 p 26 0 3 Fine, sandy clay 25·0 p 26 0 1 Hard, purplish silt 30<	1	9	W1497			_ n	26
1 5 W1455 Pale greenish grey clay, dark streaks 30·3 v.p. 29 1 7 W1454 Pale browfish grey, silty clay 32·6 v.p., S 26 0 8 W1453 Grey, silty clay 32·0 p 29 0 8 W1453 Brown, carbonaceous and lignitic shale p 29 1 9 W1452 Brown, silty clay 32·3 p 27 0 6 W1450 Brownish grey clay 25·0 p 26 0 3 Fine, sandy clay 25·0 p 26 0 1 Hard, purplish silt 0 7 W1451 Dark grey, sandy clay 27·3 p 36 0 1 Fine, grey, sandy clay 27·3 p 30 1 5 W1449 Brown, carbonaceous, lignitic, silty 1 5 W1448 Brownish clay 34·0 p 16 2 W1446 Very dark grey clay, films fine white	0	9	W1456				29
Tolerand Tolerand							29
0 8 W1453 Grey, silty clay 34.0 p 29 0 8 Brown, carbonaceous and lignitic shale 32.3 p 27 1 9 W1452 Brown, silty clay 25.0 p 26 0 6 W1450 Brownish grey clay 25.0 p 26 0 3 Fine, sandy clay 25.0 p 26 0 1 Hard, purplish silt 30 1 27.3 p 36 0 1 Purplish shale 30 27.3 p 36 30 </td <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>26</td>							26
Brown, carbonaceous and lignitic shale							29
1 9 W1452 Brown, silty clay. 32·3 p 27 0 6 W1450 Brownish grey clay 25·0 p 26 0 3 Fine, sandy clay 25·0 p 26 0 1 Hard, purplish silt 27·3 p 30 0 1 Purplish shale 27·3 p 30 0 1 Purplish shale 27·3 p 30 0 1 Fine, grey, sandy clay 27·3 p 30 0 4 Lignite and lignitic shale 20 30 30 30 0 8 W1449 Brown, carbonaceous, lignitic, silty clay 36·6 s.p. 17 1 5 W1448 Brownish clay 34·0 p 16 2 W1447 Very dark grey clay, films fine white sand 34·3 v.p. 17 6 2 W1446 Light grey, sandy clay 21·0 s.p. 25 0 5 Finely banded silt, carbonaceous films 34·3 v.p. 25 0 3 W1445 Brownish clay, fine ferruginous con-			111200			P	20
1 9 W1452 Brown, silty clay 32·3 p 27 0 6 W1450 Brownish grey clay 25·0 p 26 0 3 Fine, sandy clay 25·0 p 26 0 1 Hard, purplish silt 0 7 W1451 Dark grey, sandy clay 27·3 p 36 0 1 Purplish shale 0 4 Lignite and lignitic shale 0 8 W1449 Brown, carbonaceous, lignitic, silty clay 36·6 s.p. 17 1 5 W1448 Brownish clay 34·0 p 16 1 5 W1446 Very dark grey clay, films fine white sand 34·3 v.p. 17 6 2 W1446 Light grey, sandy clay 21·0 s.p. 25 0 5 Finely banded silt, carbonaceous films 0 3 W1445 Brownish clay, fine ferruginous con-	U	°					
0 6 W1450 Brownish grey clay 25.0 p 26 0 3 Fine, sandy clay	1	9	W1452	Brown, silty clay	32.3	ъ	27
0 3 Fine, sandy clay 0 1 Hard, purplish silt 0 7 W1451 Dark grey, sandy clay 0 1 Fine, grey, sandy clay 0 4 Lignite and lignitic shale 0 8 W1449 Brown, carbonaceous, lignitic, silty clay 1 5 W1448 Brownish clay 34.0 p 16 1 5 W1447 Very dark grey clay, films fine white sand <				Brownish grey clay	25.0	ά	26
0 1				Fine, sandy clay			
0 7 W1451 Dark grey, sandy clay 27·3 p 36 0 1 Purplish shale				Hard, purplish silt	l	l <i></i>	
0 1 Purplish shale. 0 1 Fine, grey, sandy clay. 0 4 Lignite and lignitic shale. 0 8 W1449 Brown, carbonaceous, lignitic, silty clay. 1 5 W1448 Brownish clay. 1 5 W1447 Very dark grey clay, films fine white sand. 2 W1446 Light grey, sandy clay. 3 4.0 p 16 3 W1447 Very dark grey clay, films fine white sand. 5 Sand. 6 2 W1446 Light grey, sandy clay. 7 Finely banded silt, carbonaceous films. 7 Brownish clay, fine ferruginous con-				Dark grev sandy clay	27.3	l n	30
0 1		l i		Purplish shale		l	
Description of the content of the	0	Ī		Fine, grev, sandy clay			
0 8 W1449 Brown, carbonaceous, lignitic, silty clay	Ŏ	4					
1 5 W1448 Brownish clay	0	8					
1 5 W1448 Brownish clay			.55	clay		s.p.	17
1 5 W1447 Very dark grey clay, films fine white sand	1	5	W1448	Brownish clay	34.0	p	16
6 2 W1446 Light grey, sandy clay	1	5	W1447	Very dark grey clay, films fine white		-	
6 2 W1446 Light grey, sandy clay				sand	34.3	v.p.	17
films	6	2	W1446	Light grey, sandy clay	21.0	s.p.	23
0 3 W1445 Brownish clay, fine ferruginous con-	0	5			}		
	0	3	W1445	Brownish clay, fine ferruginous con-	40.5		
		_		cretions		v.p.	14
0 5 Finely banded, coarse silt.							
5 8 W1444 Grey, sandy clay 23.0 short 18	5	1 8	W1444	Grey, sandy clay	23.0	short	15

The samples W1461, W1460, W1458, and W1451 are high-grade refractory. Samples W1447 and W1445 crack in drying. There is also some low-grade refractory clay in the section.

Near the west end of Big Muddy lake on the south side of the valley, the Whitemud outcrops in a few places. On a small butte in sec. 14, tp. 3, range 23, W. 2nd mer., sandy clay is overlain by pale greenish grey clay. Just to the southeast of this, at the end of a short spur, one foot of clay at the very top of the Whitemud shows. A sample gave the following results:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1	0	W1473	Light brown clay	32.6	р	32

This is a high-grade clay and the locality should be further investigated.

The top of the Whitemud also shows in NE. $\frac{1}{4}$ sec. 15, tp. 3, range 23, W. 2nd mer. Only a few feet of clay are exposed. A sample from 1 foot of section gives the following results:

Thie	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1	0	W1475	Greyish brown clay	27.3	р	31

WILLOWBUNCH SECTIONS

Sections of the clays of the Willowbunch member south of the town of Willowbunch were described in a recent report.¹ These clays are also exposed to the southeast of Willowbunch in a long tributary coulée west of Bonneau lake. In NE. ½ sec. 4, tp. 5, range 27, W. 2nd mer., high on the coulée side, is the following section of the Willowbunch member:

Thickness		Sample	Colour, grade	Tempering water.	Working	P.c.e.
Ft.	Ins.	No.		per cent	properties	1 .0.0.
1 3 1 1	0 8 9 0 6	W1324 W1323 W1322 W1321Sd	Grey and yellowish, sandy clay Pale grey to white, mottled, sandy clay Light grey clay, yellowish at top Brown, carbonaceous shale, coaly	$\begin{array}{c} 24 \cdot 3 \\ 32 \cdot 6 \end{array}$	s.p. p, G, T p v.p.	14 30 28+ 23
3 0 0 0 1	6 9 8 9 0	W1320 W1319 W1318	streaks. Light grey, sandy clay. Light brownish grey clay. Light brownish grey clay. Coal and brown, carbonaceous shale Light brown clay.	21.0	short v.p. v.p.	23 14 14

W1323 is a high-grade refractory clay.

¹McLearn, F. H.: Geol. Surv., Canada, Sum. Rept. 1929, pt. B, p. 61.

In the same valley, to the southeast of the above, in SW. $\frac{1}{4}$ sec. 3, tp. 5, range 27, W. 2nd mer., and about 120 feet stratigraphically below the Willowbunch member, is a zone which contains some very low-grade, semi-refractory clay:

Thickness Ft. Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0 8 0 2 4 5 0 9 0 10 0 6 1 1	W1330 W1329 W1328 W1327 W1326 W1325	Lignite and lignitic shale Brownish grey, mottled clay. Lignite, etc. Dark grey clay. Dark grey to black clay. Grey clay. Light brownish grey clay. Grey, somewhat mottled clay. Lignite. Brown shale.	33·0 34·0 35·6 31·6 24·0 30·0	v.p., S v.p. p, T v.p., T p v.p.	15 9 11 13 13 12

A good section of the Willowbunch member is exposed a little below the level of the high upland on the south side of Twelvemile Lake valley, south of Harptree, just west of the road to Coronach and in about NE. \(\frac{1}{4}\) sec. 33, tp. 3, range 26, W. 2nd mer.

Thic	kness	Sample	Colour, grade	Tempering water,	Working	P.c.e.
Ft.	Ins.	No.		per cent	properties	
3 2 3 1 5 1 0 1 0 2	6 10 0 8 1 11 4 2 2 8 4	W1257 W1256 W1255 W1254 W1253 W1252 W1251 W1250 W1249	Top section Grey and yellowish grey, silty clay. Pale grey to white, sandy clay Dirty grey clay. Pale purplish and yellowish clay Light grey, yellowish in places, clay Pale greenish grey clay. Grey clay. Dirty white clay, some fine iron spherules. Purplish brown, silty clay. Light grey to white, sandy clay as below, but yellow Light grey and green, film-banded	27.0	p p v.p. v.p. v.p. v.p.	20+ 27 23 28+ 23 29 18
1	0	W1248				
0	10	W1247	ClayDirty mauve-grey clayDark grey silt, light silt bands	31·6 32·6	p v.p.	16 15
ŏ	6	W1246	Light grey, silty clay	25.4	g.p	19
ĭ	ĭ	W1245	v.f. banded, bluish grey and light			
2	4	W1244	grey clay Light yellowish grey, sandy clay	27·6 22·6	p s.p.	18 23

W1251-W1257 might serve as low heat duty refractory clays.

The Willowbunch member was also sampled in NE. $\frac{1}{4}$ sec. 5, tp. 4, range 26, W. 2nd mer.

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
3 4 5 7	6 0 4 9	W1317 W1316 W1315 W1314	Very calcareous, light yellowish green clay	30·3 24·0	p p s.p. short	2 23 23 23 23

At this locality the Willowbunch clays are rather poor, and they are very variable in this vicinity. About 40 feet over the Willowbunch member at this place is a zone containing some semi-refractory clay and low-grade refractory clay:

Thic	Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1 1 0	7 11 10	W1311 W1310 W1308Sd	Brownish to purplish clay	20·3 26·6 29·0	p p, T, G v.p.	16 26 15

W1310 is a low-grade refractory clay.

In this area south of Harptree the Willowbunch member was also sampled in NW. ½ sec. 5, tp. 4, range 26, W. 2nd mer.:

Thic	kness	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
2 2 2 2 3 2 1	10 0 10 7 9	W1262 W1261 W1260 W1259 W1258 W1263	Top section Pale mauve-grey clay Grey clay Grey, partly mottled clay Light purplish and yellowish grey, silty clay Light brownish grey clay. Brown, hard, silty clay	21.6	v.p. v.p. p	27+ 26 19 23 28 20

Omitting W1260 and W1263, this section contains low-grade refractory clay.

Refractory and semi-refractory clays also occur in the vicinity of Buffalo gap. They appear to be at the horizon of the Willowbunch member. Outcrops are few, however, and the exact stratigraphic position is difficult to determine. On the east side of the road through the gap and in about the southern part of sec. 11, tp. 3, range 25, W. 2nd mer., under burnt shale, is some grey shale from which a sample was taken:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
•••••		W1443Sd	Light grey clay	24.6	р	28+

This is a prospect for low-grade refractory clay.

About $3\frac{1}{2}$ miles northwest of Buffalo gap, near the end of a northward extending spur, and in about sec. 17, tp. 3, range 25, W. 2nd mer., are fine sandstones, shales, etc., weathering pale yellowish, etc. High in the exposed section are the following beds:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
i		W1182Sd	Light greenish, silty clay Purplish clay	35.9	v.p.	33+

The above sample was taken from only about one foot of section, but the results are enough to show that refractory clay of very high grade does, locally, occur in the Willowbunch or near zones and that this is a prospect that invites further examination.

On the south side of Big Muddy valley, east of the Bengough (34) highway, and in sec. 13, tp. 3, range 24, W. 2nd mer., the following section was studied:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
3 4 4 1 3	0 0 7 0 5	W1441 W1440	Pale grey to white, silty clay. Light grey, silty clay. Light grey to white, sandy clay. Ferruginous, sandy clay. Hard grey clay.		v.p. v.p. f.p.	20+ 26 26

W1441 and W1440 are low-grade refractory clays.

A section of the Willowbunch member in Big Muddy valley south of the lake was described in a recent report.¹

¹McLearn, F. H.: Geol. Surv., Canada, Sum. Rept. 1929, pt. B, p. 61.

Near the International Boundary and in sec. 2, tp. 1, range 24, W. 2nd mer., about 10 to 15 feet of whitish silt is exposed that may be of the Willowbunch member. A sample was taken from near the middle of the section:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1	0	W1145Sd	White, silty clay	26.1	р	18

This, of course, is a poor clay.

North of Rockglen, in about the centre of sec. 11, tp. 3, range 30, W. 2nd mer., at the top of a cone-shaped hill, is a section of what is probably the Willowbunch member:

Thie	Sample		Colour, grade		Tempering water, per cent	Working properties	P.c.e.
0	2 2	W1425	Brownish grey silt	24.0	v.short	30+	
2	5	W1424	Light brown clay and silt	25.0	р	23	
3	5 6	W1423 W1422	Light brownish grey clay and sandy clay Pale green, silty clay	21·5 27·5	p p	20 23	

Near the above and at the same horizon, in about NW. $\frac{1}{4}$ sec. 11, tp. 3, range 30, W. 2nd mer., is the following section:

Thic	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0 0 0 0 2	5 2 5 6 8		Brown clay. Green shale	32.5	v.p., S	11 15 15

W1421 cracks on drying. The Willowbunch carries poor clay at this locality.

Better refractory clay occurs at a lower horizon, i.e., about 80 or 100 feet below the above described and apparent Willowbunch zone. This lower zone was sampled in SW. 4 sec. 11, tp. 3, range 30, W. 2nd mer.:

Thic.	Thickness Sample No.		Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1 2	0	W1434 W1433	Purplish brown clayGrey and brownish clay	29·0 26·0	p, T, St	29 15

What is probably the same lower zone occurs southwest of Rockglen, in about SW. ½ sec. 33, tp. 2, range 30, W. 2nd mer.:

Thic	kness	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
1 0 0 0 0	2 44 27 2	W1382 W1381 W1380 W1379	Dolomitic, grey, irregularly banded clay Lignitic layer. Brownish clay, pea-like white lenses Lignitic layer. Brown clay. Mottled brown and grey clay.	34·6 29·3 29·6	v.p. p	32 30 20

The clay of W1381 is very high-grade refractory. That of W1380 is high-grade refractory.

The higher and probable Willowbunch zone was sampled at the east end of a high ridge in the southern part of sec. 26, tp. 2, range 30, W. 2nd mer.:

Thickness		Colour grade		Tempering water.	Working	P.c.e.	
Ft.	Ins.	No.		per cent	properties	1.0.6.	
2 1 1 1 1 2	1 5 0 ¹ 0 ¹ 5 ¹	W1416 W1415 W1414 W1413	Yellow weathering silt	42.6 19.0 24.0 23.3	v.p.,T,S s.p. p s.p. s.p. v.p.	28 23 23 23 23 23 24 14	

¹Approximate.

W1416 cracks in drying.

A good section of the Ravenscrag is exposed on the east side of Rockglen valley, in SW. $\frac{1}{4}$ sec. 29, tp. 2, range 29, W. 2nd mer. At the top is what possibly may be the Willowbunch zone:

Thic Ft.	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
2 1 3	0 10 0	W1409 W1408 W1407	Dark grey to chocolate clay	28.0	p p p	15 17 15

49в About 10 feet below the above is the following section:

Thic.	kness Ins.	Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
0 0 2 1 1	2 7 0 0 0	W1403	Whitish, silty clay Mottled, dark grey and whitish clay Mottled, dark grey, silty clay Fine, yellowish grey silt Yellowish green, sandy clay	30.0	p p v.p. v. short	13 15 14 14

About 25 feet below the above is the following section:

Thie	Thickness Sample No.				Working properties	P.c.e.
4 1 1	4 1 0	W1399 W1397 W1398	Dark grey, calcareous clay	34.0	v.p. v.p. p	9 14 3

About 100 feet below the above is the following section:

Thickness Ft. lns.		Sample No.	Colour, grade	Tempering water, per cent	Working properties	P.c.e.
i	0	W1410	Coaly streak	27.3	р	14

GROUND WATER RESOURCES OF MOOSE JAW, SASKATCHEWAN

By W. A. Johnston and R. T. D. Wickenden

CONTENTS

	PAGE
Introduction	50
General character of the district	51
Precipitation	
General geology	54
Description of areas from which water is obtained	55
Illustrations	
Figure 4. Surface deposits of Moose Jaw area, Saskatchewan	52
5. Location of flowing water wells in Darmody-Mawer area, Saskatchewan,	63

INTRODUCTION

An investigation of the ground water resources of Moose Jaw area was made in the summer of 1930, with a view to determining what new sources of water supply for the city could be made available. In 1929 new developments in connexion with water supply were undertaken, as it was realized that, on account of the exceptionally dry year and the increased demand for water for industrial and other purposes, there was likely to be a shortage of water. A large gravel-walled well was put down at Sandy creek, one of the existing sources of water about 20 miles west of the city, and for a time it was thought that the well would yield nearly a million gallons a day and that an adequate supply at least for the immediate future had been obtained. But after prolonged tests the yield was found to be only about 450,000 gallons a day, and this additional supply, together with the supplies available from other sources, was scarcely sufficient to satisfy the present needs.

An area of about 2,500 square miles, extending in a belt 10 to 50 miles wide southeast through Moose Jaw and Regina nearly to Weyburn, is a clay plain in which supplies of ground water are limited and in places are The area is a very important one agriculturally and devoted almost exclusively to grain growing, so that the water supply problem has not been very acute, but if the farmers are forced by soil drifting and by other factors to undertake mixed farming the water supply problem will be a very important one. Diversion of water from South Saskatchewan river to supply the needs of towns and villages in the clay plain as well as the cities of Regina and Moose Jaw was long ago proposed, and was shown by elaborate surveys and investigations to be a feasible but costly scheme. The work, however, may not be undertaken for some time, partly because of the large expenditure involved, several million dollars, and partly because there is sufficient water in the vicinity of Regina to supply the city for a number of years, even allowing for a growth of the city equal to that of the past few years. The water resources of the clay plain were investigated to some extent as well as those in the vicinity of Moose Jaw.

It was found that these were limited in general, but that there is a large artesian water area lying between Moose Jaw and South Saskatchewan river. This area, described in a later part of the report, may prove to be of importance as a future source of supply not only for Moose Jaw but for the clay plain in general, as the water, although not of very good quality, can be used when mixed with a proportion of hard water and can be obtained at much less cost than that from South Saskatchewan river.

Acknowledgments are made to Mr. H. C. Ritchie, City Engineer, for co-operation in the present investigation, to Mr. L. McManamna, of the Layne Canadian Water Supply Company, Limited, for information regarding test drillings, to Mr. D. M. McNeely, Moose Jaw, and to Mr. Chris Olemen, Riverhurst, for information regarding deep wells drilled by them in the Mawer area.

GENERAL CHARACTER OF THE DISTRICT

TOPOGRAPHY AND DRAINAGE

Moose Jaw is located (See Figure 4) at the junction of Moosejaw creek, flowing northwest in a valley that is nearly 100 feet deep in its lower part and is narrow and steepsided, but has a flat bottom, and Thunder creek, flowing east in a wider valley that is flat bottomed in places and has gently sloping sides. Moosejaw creek is the larger, though its bed, as well as that of Thunder creek, is nearly dry in summer. Below the city it flows in the main valley northeast for 20 miles to Qu'Appelle river, draining east into the Assiniboine. Thunder Creek valley extends west and northwest for many miles into the hilly upland country near Riverhurst, on South Saskatchewan river, and is a broad, deep valley which, however, is nearly dry. The area north of Thunder Creek valley and lying between it and Qu'Appelle valley is an undulating plain, the highest part of which is near Qu'Appelle valley and is about 100 feet above Thunder Creek valley. There is very little drainage from this area into Thunder Creek vallev: most of the surface drainage is impounded by dams in the tributary valleys. The area south of Thunder creek is a gently sloping plain interrupted in places by low hills and ridges and extending for 15 to 20 miles to the abrupt rise known as the Missouri escarpment near Buttress, south of Moose Jaw and extending northwest to near Riverhurst. There is a rise of about 500 feet within 2 or 3 miles in places, and a more gradual slope in other places, to an upland that has many irregular hills and undrained hollows and slopes towards the southwest to the broad, shallow basins occupied by lakes Johnstone and Chaplin. These lakes, as well as much of the upland region, have no drainage to the sea. There are, however, broad valleys, of which Sandy creek, draining into Thunder creek, is the most important in the Moose Jaw region, which extend far back into the upland and receive part of their water supply from it. Other valleys, which drain into Thunder creek from the south, but do not enter the upland, are Forsythe creek, about 4 miles east of Sandy creek, and Snowdy Springs creek, entering about 4 miles west of Moose Jaw.

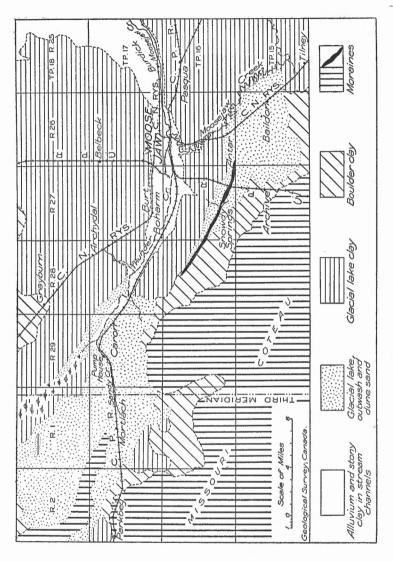


Figure 4. Surface deposits of Moose Jaw area, Saskatchewan.

PRECIPITATION

One of the chief conditions affecting ground water supply in Moose Jaw area is the comparatively small amount of the annual precipitation. The annual precipitation according to records of the Meteorological Service of Canada is shown in Table I.

TABLE I

Precipitation, Moose Jaw, Saskatchewan

Year	Rainfall	Snowfall	Total precipita- tion
			Inches
1910 1911 1912 1913 1913 1914 1915 1916 1917 1918 1919 1920 1921 1922 1922 1923 1924 1925 1926 1927 1928 1927 1928	11·33 10·84	24·3 36·8 14·2 21·7 56·9 59·0 56·5 48·9 49·1 32·9 45·9 30·1 36·0 32·5 40·3 37·2 21·7 51·5	12·59 16·25 14·03 14·94 14·55 13·70 16·53 14·70 14·60 14·41 17·65 19·81 16·54 16·54 16·54 14·88 12·70 10·10

The precipitation in 1929 was the smallest since 1908. It was somewhat greater in 1930 up to the end of October than in 1929, but was considerably below the normal. The precipitation at Sandy creek was practically the same as that of Moose Jaw, although the amount each month differed considerably at the two places. The mean annual precipitation is practically the same at Moose Jaw as at Regina, although it differs considerably in some years. It decreases somewhat towards the south and southwest, but that of Moose Jaw probably is nearly the average for the area extending west and southwest of the city to Missouri coteau. Table I shows that during the past 20 years only in 1929 did the precipitation fall below 12 inches. In the 20-year period previous to 1910, judging by partial records at Moose Jaw and at Regina, the average was 1 or 2 inches less than during the past 20 years, and was below 12 inches in several years. There were three consecutive dry years, 1892-3-4, and in 1897 and 1908 the precipitation was less than in 1929. The recurrence of exceptionally dry years at irregular intervals, and especially two or even three dry years in succession, is of particular importance in the case of the Moose Jaw water supply, as the water is derived mainly from surface sands, which are affected by drought to a much greater extent than a deeply buried reservoir; in times of drought much of the water in the sandy beds near the surface is lost by operation, whereas in deeply buried, water-bearing beds this is not the case. It is necessary, therefore, to provide for a supply that will be adequate, not only in normal years but during periods of low rainfall which may last from 1 to 3 years.

GENERAL GEOLOGY

The surface deposits of the area are the unconsolidated, chiefly glacial, deposits, which overlie the older consolidated or partly hardened bedrocks. They consist of wind-blown sand, alluvial sand, and clay deposited on the flood-plains of the present streams, stratified clay and sand deposited in ancient glacial lakes, and stony clay and sand deposited by the ancient glaciers, partly in the form of moraines and partly as outwash and boulder clay plains.

Wind blown sand, in the form of old dunes that have begun to drift somewhat in recent years, occurs in Sandy Creek area and south of Moose Jaw. Much of the alluvium in Thunder Creek valley above Moose Jaw is a heavy clay that is nearly impervious to water, so that the valley bottom does not form a good reservoir for ground water; the water brought into the valley by tributary streams appears to disappear largely by evaporation.

Stratified lake clay and sand deposited in an ancient lake, known as glacial Lake Regina,¹ formed as a result of damming of South Saskatchewan river by the retreating ice-sheet, occur in Moose Jaw district in places up to an elevation of about 1,950 feet above sea-level, and to a somewhat higher level near Mortlach and Parkbeg west of Moose Jaw, where an earlier glacial lake existed at a higher level. The sands occur in the higher parts near the shores of the ancient lake basin, as at Sandy Creek and south of Moose Jaw; much of the wind-blown sand has been formed from the lake sand. The clay occupies a large part of the area north of Thunder Creek valley, has been eroded away in the valley itself, and forms a belt a few miles wide along the south side of the valley. The sands readily absorb the rainfall and, therefore, form good sources of ground water, whereas the clay because of its extremely heavy character is nearly impervious to water. The distribution of the lake sands in the vicinity of Moose Jaw is shown on Figure 4.

Outwash sands and gravels occur in the area northwest of Mortlach and are favourable for the accumulation of ground water, but are not very

extensive and are underlain at shallow depths by clay.

A narrow, morainic ridge, composed in part of sand and gravel, occurs at Snowdy springs and extends northwest nearly to Sandy creek and southwest nearly to Moosejaw creek. It forms a sort of natural dam of the drainage of the sandy area to the southeast, so that most of the water from this area passes through Snowdy Springs valley, the only opening through the ridge. A great, broad, morainic belt occurs along the Missouri coteau or upland, but there is very little drainage from this moraine towards Moose Jaw and much of the water stored in the moraine has no free drainage and is saline.

¹Trans. Roy. Soc., Canada, sec. IV, 1930, pp. 41-49.

Boulder clay plains—much of the boulder clay contains few boulders, but is unstratified and usually contains at least a few small stones—extend over parts of the area north and south of Thunder Creek valley. Areas along Thunder Creek valley, particularly below Moose Jaw, are very bouldery. The boulders were derived from the drift and were concentrated in ancient stream channels.

The boulder clay plains are not very favourable for the storage of ground water as the material is not very permeable and the water usually is hard and saline, at least to some extent, except in places where sand and gravel deposits are included in the clay, and these deposits as a rule are not very large.

The thickness of the glacial drift varies greatly from place to place and there is no way of predicting what the thickness is at any locality, as the contours of the underlying bedrock differ greatly from those on the surface. The thickness in Thunder Creek valley at Moose Jaw is only 15 to 30 feet, at Sandy creek it may be about 80 feet, and at a few places is upwards of 300 feet.

The bedrock underlying the surface deposits in the vicinity of Moose Jaw is the Pierre (Bearpaw) shale of Upper Cretaceous age. As the shale is of considerable thickness and does not contain potable water in this region it is important in drilling to distinguish between it and the glacial drift. But this is not always easy, for much of the drift is derived from the shale and in places is nearly similar to the weathered upper part of the shale. It can be distinguished as a rule by the presence in it of pebbles and stones of limestone and granitic rocks that are foreign to the district and were transported by the ice-sheets for long distances from the northeast.

DESCRIPTION OF AREAS FROM WHICH WATER IS OBTAINED

The present water supply is derived from four sources: Snowdy springs, $7\frac{1}{2}$ miles southwest of the city; from a well and collecting gallery at Sandy creek, 20 miles west of the city; from a collecting gallery in Thunder Creek valley, just west of the city; and from a well on Forsythe creek, 2 miles south of Caron. The last source was developed during the present investigations and has yielded an additional supply of about 300,000 gallons a day.

Snowdy springs are on the southeast quarter of sec. 9, tp. 16, range 27. The springs are seepages from sand and gravel overlain by clay and extend for about 200 feet along the south side of a pond formed by a concrete dam placed across the outlet of the depression in which the springs occur. The spring pond is on the south side of a valley about 50 feet deep, draining a sandy area lying to the southeast. The valley is cut through a drift ridge below the spring pond and is dammed to form a large reservoir which, when full, is at the same level as the spring pond and adjoins it. The large reservoir becomes filled in the spring from melting snow and from the flow of the creek, which is fed in part from springs higher up the valley. In the latter part of the summer and in autumn there is practically no flow of the creek into the large pond or overflow from the small pond, so that the water in the reservoir becomes stagnant and unfit for use except after filtering and even then has an objectional odour, especially

at times when a considerable part of the water has been drawn off. There are small "secondary" springs in the valley bottom about 1,000 feet below the main dam. This water probably is seepage water coming from the sides of the valley or from higher up in the valley and may be affected by back wash from the filters. The springs in the small pond are the main source of supply, and derive their water from the sandy area lying to the southeast. This is shown by the fact that the water is not highly mineralized. The water found in the clay areas to the south is very hard and saline in places and it is evident that the spring water if it were derived from the hilly region to the south and southwest would likewise be saline. The sandy area supplying the springs consists in part of dune sand in irregular hills and in part of lake sand and silt. The extent of the sandy area drained by Snowdy Springs Creek valley is about 25 square miles. Though the area is not large it is favourable for the absorption of rain water, so that the springs are permanent. Furthermore, the morainic ridge crossing Snowdy Springs creek below the springs acts as a dam preventing escape of the underground water except at the springs.

Sandy creek joins Thunder Creek valley about 2 miles northwest of Caron and drains a large area to the southwest. The creek valley in its lower part passes through a sandy area which extends west nearly to Mortlach and from 2 to 4 miles northwest and southeast of the valley. The greater part of the area is underlain at the surface by old sand dunes: a part is lake sand. In the valley of the creek itself there is alluvial and lake sand interbedded with thin layers of silt and clay. The clay layers are nearly impervious and tend to prevent absorption of the surface water that accumulates in the spring, so that a part of this water is lost by runoff and by evaporation. Some water seeps into the valley from the sides. but the main flow of the creek, except during freshets and in wet years, comes from springs on the south side of the valley bottom 2 to 3 miles east of Mortlach. The springs occur in the northwest quarter of sec. 19, tp. 17, range 29, and in the north half of the adjoining section to the west. The water comes from sand and fine gravel which may overlie clay at a shallow depth. The creek above the springs has a very large drainage area extending west into the Coteau hilly country. During freshets considerable water may be derived from the upper part of the basin, but in dry years the creek bed above the springs contains little or no water. The springs, however, may be supplied by underground water coming from the southwest part of the basin and, therefore, may be permanent. The flow of the springs increases somewhat in the autumn but is said to be fairly uniform throughout the year, though the flow is much less in dry years than in wet years, apparently because the water comes mainly from the surface sand and not from any deep artesian basin.

Trouble has been experienced in the collecting gallery at Sandy creek by plugging of the holes with small stones, and with caving. The section at the lower end was relaid in 1930 and crushed stone placed on the outer side, apparently with better results. The ground, however, in which the pipe is laid consists of fine sand and there are thin clay layers interbedded with the sand, so that the water does not readily drain into the gallery from the sides. Though the pipe extends upstream for about half a mile to where it receives the main flow of the creek, the water obtained by pumping in exceptionally dry periods is little greater than the flow of the

creek as determined by weir measurements about 1 mile upstream. The sandy deposits near the lower end of the gallery have a thickness of 25 feet to 40 feet, as shown by test borings, and the material is reported to be somewhat coarser in the lowest 2 to 3 feet.

The new gravel-walled well at Sandy creek is 80 feet deep and derives its water from sand and gravel at depths of 51 to 79 feet. A layer of clay, probably boulder clay, separates the lower sand from the surface sand and clay. The water is artesian, but is lowered about 57 feet by pumping and recovery amounts to only about 1 inch a day. In 1912 two test wells had been put down, one near the site of the present No. 1 well and the other 300 feet downstream. These wells had a small flow and suggested that artesian conditions existed. The log of the new well from the surface down showed: sand and silt with layers of clay 25 feet, clay 11 feet, dirty sand 2 feet, clay 13 feet, sand 5 feet, clay and boulders 6 feet, sand and gravel 17 feet, and clay 1 foot. The marked draw down of water in the well when pumped, and slow recovery, show that there is no very large supply of water in the gravel buried beneath clay; probably the surface water does not readily find access to the gravel because of the wide extent of the overlying clay. Two other gravel-walled wells, one a mile upstream and the other 700 feet downstream from the pump house, were put down to the same sand and gravel beds and also showed a marked draw down of the water. The one well probably obtains all the water that can be drawn from these beds without exhausting the supply. The drawing of water from the gravel beds beneath the clay does not appear to affect the supply obtained from the flow of the creek, as this supply comes from springs issuing from the sand and gravel above the clay. A number of test borings in Sandy Creek valley both above and below No. 1 well were made by the Layne Canadian Water Supply Company, and the possibilities of lower, water-bearing beds appear to have been thoroughly tested.

Forsythe creek flows into Thunder Creek valley from the southwest about 2½ miles east of Caron and is a spring creek deriving its water locally from sandy beds overlying clay. The creek is only about 5 miles long and in dry seasons the upper 2 or 3 miles where the valley is shallow contain no water except in ponds formed by dams across the valley. springs, which are merely seepages from the sand, are in the lower part where the valley is about 40 feet deep. Test borings by the Lavne Canadian Water Supply Company showed that the depth of the sand in the valley bottom in the southeast quarter of sec. 13, tp. 17, range 29, is 30 to 55 feet and that the depth decreases down stream to about 12 feet $\frac{1}{4}$ mile to the east. The new well was placed at a point where the sand was deepest near the line between sections 12 and 13 and about 500 feet west of the east line of the sections. The spring water is derived from the sandy beds through which the creek flows and as these are only a few square miles in extent it is not to be expected that a supply greatly in excess of the flow of the creek, about 150,000 gallons a day, can be obtained from one or more wells. The springs, however, flowed at about this rate in August, 1930, after two exceptionally dry summers, and it is probable that the flow in wet years will be considerably increased.

Britannia Park collecting gallery is excavated in alluvial sand, gravel, and clay down to the shale or bedrock. A large well about 15 feet in diameter was also put down nearly to the bedrock in the valley flat beyond

the southern end of the gallery, but yielded only about 13,000 gallons a day. The gallery extends only a part of the distance across the flat, but the small yield from the large well indicates that production would not be greatly increased by extension of the gallery. Moreover, the character of the water renders it less desirable than that from other sources.

MINERAL CHARACTER OF THE WATER

The mineral character of the water from the various sources is shown in the following table. The analyses were made at the Provincial Laboratories, Regina, for the city of Moose Jaw. The results are shown in parts per million.

Water Analyses

	Forsythe Creek well	Snowdy springs	Sandy Creek gallery	No. 1 well, Sandy creek	Britannia Park gallery
Total solids Solids in solution. Solids in suspension. Ignited residue. Organic and volatile matter. Chlorides. Sulphates. Carbonates. Carbonates. Calcium. Magnesium. Sodium. Iron. Total hardness.		640 46 39 160 197 66 42 98	494 486 8 440 46 19 94 171 110 28 9 0.3	1,218 1,218 11,218 nil 1,106 112 79 363 282 50 19 348 nil	1,814 1,674 141 1,514 160 18 814 189 224 82 134 7

The city tap water is a mixture of these waters in various proportions, depending upon the amounts available from each source at different periods. In winter when the storage water from the large pond at Snowdy springs has been used the main supply is from Sandy creek. This water is comparatively soft and of good quality. The Britannia Park water is much harder than any of the other waters, it contains some iron, and there is some danger of pollution of the water because of nearness to the city and the sluggish character of the drainage in the valley. The analysis of the Forsythe Creek well showed that the total solids in the water are 41.0 grains a gallon, made up of the following constituents in the order of comparative quantity: calcium sulphate, magnesium sulphate, calcium carbonate, and calcium chloride. The water from No. 1 well at Sandy creek differs somewhat from the water derived from the surface sand. It contains a much higher proportion of sodium sulphate and sodium carbonate, but is a softer and, therefore, a more desirable water.

WATER REQUIREMENTS

The population of Moose Jaw is about 21,000. There are very few private water wells in the city, there is a large industrial consumption of water amounting to about one-third of the total, and in dry years such as

1929 and 1930 only a small supply of water from rainfall is caught in cisterns. Based on an allowance of 75 gallons a person a day, the average daily requirement should be about 1,575,000 gallons a day. Because of the importance of abundant water for industries and of the shortage of water from sources other than the city water supply, it is probable that 2,000,000 gallons of good water could be used to advantage, and, allowing for normal growth of the city, the amount should be increased to 3,000,000 within a few years.

The average daily consumption of water in 1929 was about 1,000,000 gallons and was somewhat greater in 1930. The average consumption was small partly because of shortage of water. The shortage of water in spite of the new source developed at Sandy creek was due to the exceptionally dry season; in a normal year there would have been an abundant supply of water. As it is necessary to provide for an ample water supply in dry years as well as in normal years, the amounts of water that can be obtained in dry years need only be considered. The yield of the present sources with the addition of the supply from the Forsythe well is approximately 1,300,000 gallons a day, of which nearly one-half is obtained from Sandy creek.

OTHER POSSIBLE SOURCES OF WATER SUPPLY

Surface Water Supplies

Moosejaw creek has a very large drainage area, but is dry in summer except in the lower part for 3 or 4 miles up from the junction with Thunder creek where one or more springs furnish a small flow. There had been no overflow of the dam at Moose Jaw, however, for two years previous to August, 1930. The dam is below the junction of Moosejaw and Thunder creeks. It is obvious, therefore, that water impounded by a dam higher up in the valley of Moosejaw creek would be stagnant and unfit for use as a city supply. There are other objections: the cost of the undertaking would be large and the Canadian Pacific railway have water rights on the creek.

Storage of surface water in Snowdy Springs valley is already developed to the capacity of the valley. The quality of the water, however, could be greatly improved by cleaning of the reservoir. Thunder Creek valley is wide and flat-bottomed; there is practically no surface water in the valley in dry seasons except in the marsh known as Goose lake in the part of the valley 4 miles northwest of Caron, and this water is saline and unfit for use. Flood water impounded in the valley by a dam or subsurface water held back by a pile dam across the valley would become saline in character, as the soil contains considerable amounts of salts.

The only other localities where important quantities of surface water can be impounded within reasonable distances of the pipe-line are on Forsythe creek and on Sandy creek above the present pumping station. The Forsythe Creek water supply is being developed from a well and it is doubtful whether the expense of damming the valley and impounding flood water, although the valley is narrow and deep in the lower part, would be justified owing to the small drainage area of the creek. Sandy creek has a large drainage area and flood water in spring passes down the valley, but there is a difficulty in holding the water by means of a dam much

above the level of the existing pile dam, because of the presence of the railway that extends through the valley and is only slightly above the level of the valley flat. The flood water, at least in part, could be stored by excavating a basin in the surface sandy deposits and by putting down perforated casings screened at the top to carry the surface water down into the lower beds. The presence of thin clay layers in the upper part of the sandy beds prevents the flood water in spring from seeping downwards, so that it lies on the surface and is lost partly by run off and partly by evaporation.

Bedrock Waters

The deep well at the powerhouse, Moose Jaw, extended to a depth of 3,302 feet. No important flows of water were found until a depth of about 900 feet was reached. A flow of water amounting to about 4,000 gallons per hour was obtained at 980 feet, and somewhat larger flows at greater depths. The water flowing from the well in August, 1930, had a temperature of 60 degrees F. and probably comes from a depth of about 1,000 feet. The water from a depth of 2,950 feet is said to have had a temperature of 81 degrees F. According to an analysis of the well water made in 1928 by the Provincial Laboratories, Regina, for the city of Moose Jaw, the water contains 6,280 parts per million of soluble salts consisting chiefly

of sodium chloride or common salt.

A well on the farm of Mr. Hans, 2 miles north of Caron, is reported to be 735 feet deep and to be mostly in the bedrock. Water was struck at 316 feet, but the supply proved insufficient. Deepening of the well provided a good supply, but the water is salty as was also that at 316 feet. Another well in Thunder Creek valley about 4 miles east of Caron is said to have been sunk to a depth of 500 feet without finding potable water. A well in Sandy Creek valley on the north bank of the creek about threefifths of a mile above the pump house is said to be 400 feet deep. casing of this well appears to be in good condition but is open to a depth of only 20 feet, a plug apparently having been inserted. The results of the drilling are not known. In the areas north and south of Thunder Creek valley a number of wells have been drilled in the bedrock to shallower depths, and in places have obtained small supplies of potable water, which, however, is fairly high in soluble salts. A number of test wells in Thunder Creek valley in the vicinity of Moose Jaw were put down by the Layne Canadian Water Supply Company without finding potable water. One of the wells was 200 feet deep and was in the bedrock except for the upper Some of the wells in the lower part of Moose Jaw valley above the city appear to be partly in the bedrock and the water is of poor quality.

It is improbable, therefore, that potable water can be obtained by drilling in the bedrock in Thunder Creek valley or in adjacent areas. Large supplies of water, however, can be obtained in wells drilled into the bedrock in the Darmody area northwest of Moose Jaw. This possible source of

water is described in a later part of this report.

Waters from the Glacial Drift

The present water supplies are derived from sandy areas forming part of the glacial drift or unconsolidated deposits overlying the bedrock. There are sandy and gravelly areas and numerous springs in the hilly Coteau country to the southwest of Moose Jaw, but these areas are small and so far as known there is no large supply of water at any one locality. There are no large streams coming from the Coteau; the bedrock forms a considerable part of the upland and its surface slopes towards the southwest, so that there is very little underground drainage towards the northeast. There are very few springs along the northeast face of the Coteau. The water from many of the springs on the upland is potable, but nearly all of the numerous ponds and lakes are saline, as are also Johnstone and Chaplin lakes. It is improbable, therefore, that any large supplies of potable water can be derived from the Coteau region, although it is high above Moose Jaw so that water derived from that region would flow into the city under a good head.

Water can be obtained from the alluvial sand and gravel in the valley flat of Moosejaw creek below the city, but there would be danger of pollution of the water from sewage. The water is mainly the creek water draining down the valley underground and is filtered to some extent by passing through the sand and gravel deposits at the surface, which, however, are only a few feet thick, and can hardly be regarded as a safe water for use.

The only large, sandy areas favourable for the storage of ground water and within a reasonable distance of the city are those drained by Snowdy Springs, Sandy, and Forsythe creeks. Any new developments to augment the present water supply must be undertaken in these areas unless a distant source is sought. Moreover, the present sources are widely scattered, and, therefore, the cost of supplying water is high, so that any new source sought should be a permanent one and sufficient for many years. It does not appear that any very large additional supplies can be obtained from the present sources, though sufficient water probably can be obtained to meet the requirements for a few years, so as to give time for careful consideration of the most economical and satisfactory future source of supply.

Snowdy springs, as already stated, were developed to the limit of their capacity a number of years ago and independently of the storage water produced only about 150,000 gallons a day. In the winter of 1929-30 the production was about 100,000 gallons a day, the water having been drawn directly from the spring pond through a 6-inch pipe leading to the pumping station below the lower dam. It is probable that the flow of the springs might be increased somewhat by excavating part of the face of the bank along which the springs occur and thereby removing the clay that has slidden down over the sand and gravel from which the spring Lowering of the water in the spring pond by pumping might also increase the flow by increasing the head. A larger yield probably could be obtained by pumping from a gravel-walled well at the spring. As the springs are fed by water coming from the sandy area to the southeast it is probable that pumping from wells in the sandy area would rob the springs and not greatly increase the total yield. Test borings in the valley bottom below the lower dam were made in 1930 by the Layne Canadian Water Supply Company and showed that there is only a foot or so of gravel at the surface and no water in the clay below.

Sandy creek is the largest spring creek in the region and is the most important source of water supply. More water than is made available by the collecting gallery and by the new well can be obtained from the surface sand either by deepening the gallery or by a gravel-walled well near

the lower end of the gallery. Probably the latter would be preferable as it would be difficult to deepen the gallery on account of the wet, caving nature of the sand. It would be advisable to put down a new well in the surface sand rather than to replace with screen the blank casing in the upper part of No. 1 well, as the depth of the surface sand in this well is only about 25 feet, whereas the depth in the central part of the valley near the lower end of the gallery is reported to be about 40 feet. Test borings should be made at the springs in the upper part of the valley, on the northwest quarter of sec. 19, tp. 17, range 29, near the highway where it turns north, and possibly on the adjoining section to the west. and pumping tests made in order to determine whether any large flow of water can be obtained. It is possible that the spring water comes from the sand and gravel above the clay and that there is no very large supply, but the water is of good quality and the area, because of its sandy character and the large extent of the drainage basin, is the most favourable one found anywhere in the district. The water obtained from one or more gravelwalled wells at the springs could be pumped into the creek supplying the collecting gallery or piped down stream to the upper end of the gallery. It is impossible to determine how much additional water can be obtained from the suggested shallow wells near the lower end of the gallery and at the springs but, judging by the fact that the springs continued to flow and spring ponds existed in the valley bottom after two exceptionally dry summers, the supply may be fairly large.

Darmody-Mawer Artesian Area

There are a number of flowing wells in the area along the Canadian National railways from Darmody nearly to Central Butte and north to within about 3 miles of Bridgeford (See Figure 5). The wells are 400 to 450 feet deep and derive their water from sandy beds overlain by shale, which forms an impervious cover. The thickness of the drift overlying the shale is about 60 feet in places and somewhat more in other places. Wells sunk to the same horizon are also obtained in the area extending west to South Saskatchewan river and even beyond the river, but in most of these the water does not rise quite to the surface because of the higher The water in the wells near Mawer has a temperature of 48 degrees F. and rises to about 1,987 feet above sea-level, but the height varies somewhat in different parts of the area. Several of the wells have been flowing for 10 to 15 years without any great diminution in flow. Most of them were put down with a jetting machine and have about 3-inch casings. As the material passed through is soft a well was put down ordinarily in 4 or 5 days. The Canadian National Railways well at Mawer, put down in 1913, is a 6-inch well and has a large flow which, however, is throttled down to a very small flow. The smaller wells have flows of 10 to 30 gallons a minute from a 1-inch pipe and there is no question that a very large supply exists. There is a "hardpan" layer a few inches to several feet thick at the top of sandy, water-bearing beds. The layer consists of concretionary iron pyrites and probably other salts and gives some trouble in drilling because of its hardness. The sand is fine; about 42 per cent of one sample passed a 40-mesh to the inch screen and is held on a 60-mesh screen, and the balance is finer. It is so loose that large quantities are brought out of the wells by the flow of the water. There is comparatively little trouble, however, from the sand choking the wells so long as the water is allowed to flow fairly freely.

An analysis of the water from the Canadian National Railways well at Mawer was made by Andrews and Cruickshank, Regina, and is as follows:

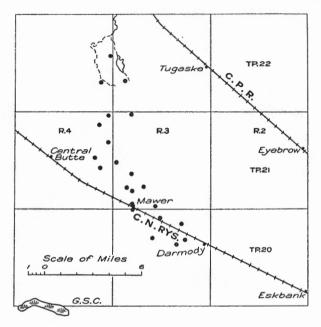


Figure 5. Location of flowing water wells in Darmody-Mawer area (west of 3rd meridian), Saskatchewan.

Analysis of Water from Mawer Well

(Results in parts per million)

Calcium bicarbonate (as carbonate)	
Calcium chloride	
Magnesium bicarbonate (as carbonate)	10.00
Magnesium sulphate	
Magnesium chloride	
Sodium carbonate	498.00
Sodium sulphate	789-00
Sodium chloride	163.00
Silica	2.00
Iron and alumina	2.00
Total solids by calculation	1,499.00
Total solids by evaporation	
Temporary hardness	5.00
Permanent hardness	Nil
Total hardness	5.00

The almost entire absence of hardening salts in the water is a remarkable feature, especially as most of the ground waters found in southern Saskatchewan are objectionable because of their hardness. The water comes from a sand that contains appreciable quantities of green sand, a water softening mineral, and, therefore, is naturally softened. tunately the water contains a fairly large proportion of non-hardening salts, is corrosive to some extent, although it does not appear to greatly affect the casings in the old wells, many of which are in good condition, and is unsuitable for irrigation because of the presence of large amounts of sodium carbonate in the water. It has, however, been used for many years at Mawer as the town supply and at other places with no apparent bad effects. In some of the wells that have a smaller flow than the one at Mawer the amount of salts present appears to be less, but only one analysis is available. It is possible that the water could be used satisfactorily when mixed with a proportion of hard water, which would reduce the corrosiveness and remove the flat taste. If, for example, the water was mixed in equal proportions with the water from the Sandy Creek gallery the mineral character of the mixture would be practically the same as that of the water from the new well at Sandy creek, and this water is considered to be

very good.

In order to determine whether the water is suitable as a city supply it would be advisable to have analyses made of the water from other wells in the vicinity of Darmody. The point nearest to the city at which the artesian water can be obtained is just west of Darmody, about 27 miles from the reservoir at the end of the pipe-line at Sandy creek in a direction measured along the railway and south to the reservoir. The most easterly well in which the artesian water has been obtained is about 1 mile west of Darmody, but the basin may extend a short distance east of this village. Numerous deep borings near Eskbank, the next town to the east on the railway, showed that the basin does not extend that far to the east, as only salt water was found. If the water is to be used, one or more test wells would have to be put down and a large well packed with gravel and screened in the lower part and a pump installed. A gravel-walled well and screen are necessary in order to avoid difficulties with sand coming into the well. The wells that are several feet down in the sand—which is said to be about 60 feet thick—and have a small flow of water with little or no sand, appear to produce better water than those that have a large flow of water from the upper part of the sand and from which large quantities Pumping will be necessary, for, although the of sand have issued. water in a large well probably will rise to the surface, if the well is placed at an elevation of about 1,970 feet above sea-level or lower, the fall to the reservoir at Sandy Creek from this elevation is only about 75 feet; if the pipe-line is placed along the railway there would be a slight rise at Lake Valley amounting to about 30 feet. There is no doubt that the supply would be permanent. The amount probably would be limited only by the capacity of the wells put down.

INTERGLACIAL DEPOSITS IN SOUTHERN SASKAT-CHEWAN

By R. T. D. Wickenden

CONTENTS

Introduction. Descriptions of occurrences. Age and extent of deposits.	65
Illustrations	
Figure 6. Locations of interglacial deposits in southern Saskatchewan	

INTRODUCTION

The Pleistocene deposits in southern Saskatchewan in many places are exceptionally thick—the region was one of accumulation of drift rather than erosion—and, therefore, the occurrence of interglacial deposits, known in many other regions, is to be expected. The ice-sheet extended beyond the area and precipitation throughout Pleistocene time may have been small, as it is at present; erosion during interglacial times probably would be not very great and records of these times may have been preserved. Interglacial deposits have not previously been reported from the region, although parts of it have been examined by several geologists. Good sections of the drift are not numerous and are found chiefly along newly constructed railways.

Investigation of the region southwest and west of Regina has shown that evidences of interglacial periods, such as sands and gravels containing organic matter and other fossil remains or exhibiting weathered zones, occur at seven widely separated localities (See Figure 6). At a number of other places two tills are shown in sections, but these are not separated by other deposits and are distinguished by merely a break in structure. In one locality two sets of interglacial beds are shown and numerous well sections also appear to show this relation. The interglacial sands and gravels are of importance as a source of ground water supply at many places, for example, at Regina where the city supply is derived from deposits of this character.

DESCRIPTIONS OF OCCURRENCES

Most of the interglacial deposits are in, or just in front of, the moraine covering the Missouri coteau, which extends from the International Boundary, at a point about 40 miles west of Portal, to South Saskatchewan river near Riverhurst. The only exception is a deposit in a small moraine about 55 miles west of the coteau near Lac Pelletier. All the exposures except those in the lower part of the valley of Swift Current creek are in railway cuts, and, therefore, not very large.

Beaubier

The most southerly occurrence is in Willowbunch area about a mile east of Beaubier station on the Canadian Pacific railway. In a cut on the south side of the track, 75 feet west of the road crossing, there is a lens of sand and silt surrounded by till. Near the top of the lens, there is a thin band of white, weathered silt and clay containing fossils. The fauna is made up chiefly of gasteropods with a few species of ostracods. A number of plant impressions were also observed and examination under

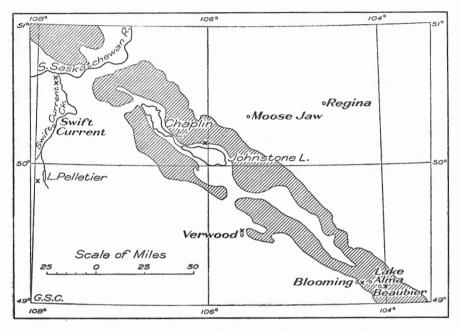


Figure 6. Locations of interglacial deposits in southern Saskatchewan; exposures of interglacial deposits are shown by crosses, the Missouri Coteau moraine by inclined ruling.

the microscope showed chara seeds. The deposit is somewhat distorted, probably owing to overriding by the ice-sheet. The lens is only about 20 feet long and 4 feet thick. Where the two tills are not separated by stratified deposits they show slight differences in structure.

Blooming

Several cuts along the same railway line between Lake Alma and Blooming show interglacial beds. The best section is on the north side of the track about a mile west of the main road south from Radville (NW. ½ sec. 25, tp. 2, range 18, W. 2nd mer.), and is as follows:

Till	12-15 feet
White, calcareous clay with shells	8-12 inches
Sand and gravel	2- 3 feet
Dark grey to brown till	4- 5 feet

The shells include small pelecypods, several species of gasteropods, nd ostracods. Some chara seeds and other traces of plant remains were also observed. A few limestone and granite pebbles in the sand and gravel are partly disintegrated, but this cannot be taken as evidence of interglacial weathering as most of the pebbles are unweathered. In places the beds are somewhat distorted like those near Beaubier.

In the vicinity of these sections there are cuts showing shell beds apparently in place on top of the upper till. These beds do not appear to be related to a higher stage of the present sloughs and ponds. The shells are much larger and many more species are present than in the proven interglacial deposits. Whether they represent deposits, which were not covered, of another interglacial period, or whether they were laid down during a much higher stage of water in ponds during post-Glacial time is not known.

Verwood

A section showing two tills, the lower of which is weathered at the top, occurs on the east side of the railway track about 2 miles northwest of the village of Verwood (SE. ½ sec. 6, tp. 7, range 7, W. 2nd mer.). The cut is through a ridge, probably part of the same moraine as at Beaubier, and exposes the following section:

Brown till.	3-50 feet
Weathered zone, buff till with plant roots	8-12 inches
Dark grey till	

The pebbles in the buff till are not weathered any more than are the pebbles elsewhere in the section.

Lake Johnstone

North of lake Johnstone and west of Oldwives station on the new Canadian Pacific branch line there is a cut (SW. ½ sec. 25, tp. 14, range 1, W. 3rd mer.), in which good sections showing interglacial deposits are exposed. The section (Figure 7) on the south side of the track is as follows:

Brown and dark grey till	5-10 feet
Sand and silt containing shells and plant remains	0-15 inches
Gravel and sand	0-10 feet
Dark grey till	0- 3 feet

The lower till in places appears to have been ploughed up by the icesheet and it is difficult to see the contact of the upper and lower tills where they are not separated by other deposits. However, on the north side of the track such a contact could be made out. The lower till is more compact and shows some banding. In places a slight amount of buff or creamcoloured soil occurs at the contact. This soil resembles a present day leached forest soil. Pieces of tree trunks and conifer branches occur in the upper till and in the reworked material. Several cones and branches with impressions of needles which resemble those of spruce were found. A piece of a tree about 6 inches in diameter was dug out and it is reported that some similar pieces up to 12 inches were seen when the cut was made. The wood which the writer found was identified at the Forest Products laboratory of the Department of Interior at Ottawa as belonging to the genus Picea. It resembles the species *Picea canadensis* very closely but the material is not well enough preserved to be certain of this.

The small band of sand and silt which lies on top of the lower till contains a few small gasteropods and pelecypod shells and numerous specimens of at least six species of ostracods. The impression of a moss-

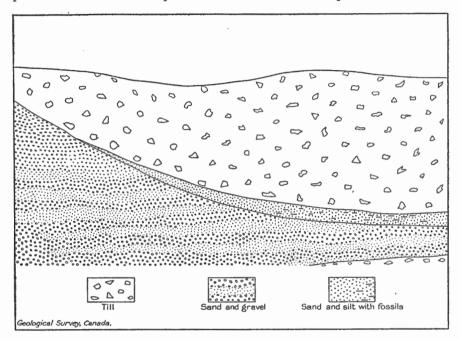


Figure 7. Section showing interglacial deposits near lake Johnstone, Saskatchewan.

like plant is also very common, as well as numerous chara seeds. The sand and silt layer lies on the surface of the gravel lens in an irregular manner, as shown in the diagram (Figure 7), but the bedding of the two does not merge.

Swift Current Valley

The best exposures are in the lower part of the valley of Swift Current creek and a small tributary gulch. They are just southwest of Missouri coteau in an area the surface of which, except for the stream valleys, is nearly level and is underlain by boulder clay. The surface may be in part an ancient terrace of South Saskatchewan river. The sections in places show three tills and two series of interglacial deposits.

A section on the west side of the valley in the east half of sec. 21, tp. 19, range 13, W. 3rd mer., shows, for nearly half a mile, three layers of till.

Measurements were made of this section at two places near the middle of the exposures, A being about 400 feet north of B.

	A	В
	Feet	Feet
Brown till Dark grey till. Stratified sand, silt, and gravel with plant remains. Brown to yellow till.	39 3 72 25	30 3 57 96
Total	139	186

The two upper tills are separated by only a break in the structure. Elsewhere in a different series of exposures in the valley these two tills are separated by water-laid deposits which contain plant remains.

The stratified deposits in sections A and B show discoloration at the bottom, probably caused by weathering or leaching by ground water. Some poorly preserved plant remains were observed in silt layers near the middle of the water-laid material.

A section, referred to as section C, is located on the east side of a tributary valley in SW. $\frac{1}{4}$ sec. 15, tp. 19, range 13, W. 3rd mer., about 2 miles south of these exposures, and is as follows:

Section C	Feet
Brown till. White sand and gravel; traces of plants. Dark grey till. Stratified sands and gravel; traces of lignite.	70 23 7 225
Total	355

The upper series of stratified deposits contains a few indefinite plant remains. In the lower series the lignite occurs in thin bands. In both of the stratified deposits much of the sand is pure white: examination of the heavy minerals in the sand shows the same abundance of hornblende and other easily weathered minerals as in other Pleistocene deposits in the region, which indicates that the sand was not derived by concentration from the drift, but its mode of origin is not apparent.

A comparison of the sections shows that the break in structure between the two upper tills in sections A and B is represented in section C by 23 feet of sand and gravel. The lower series of sands and gravels in section C represents the sands and gravels in sections A and B. The lower till of the latter sections probably was completely eroded at this point, for the stratified material appears to rest directly on Cretaceous shales.

About 1,000 feet farther up the same tributary valley on the northwest side an exposure, referred to as section D, was observed.

Section D	Feet
Yellowish brown till	30 35
Dark grey till. Light grey to white weathered till.	45
Dark brown till	75

193

The upper two tills and the sands and gravel seem to represent the two tills in section C. The thick sands and gravels of section C are apparently represented by 8 feet of weathered material and the lowest till of section D is the same as that seen in sections A and B. The weathered zone is full of traces of plants and roots. The thick zone of weathering indicates a long period, although the pebbles in this zone do not show any more effect of weathering than do those in the till above or below.

Lac Pelletier

About 4 miles due west of Lac Pelletier, in SE. ½ sec. 7, tp. 12, range 15, W. 3rd mer., a fresh railroad cut shows about 30 feet of drift. The following section was observed:

	Feet
Brown till	20
White sand	3-5
Don't mark Ail	7 10
Dark grey till	7-10

Some traces of plants were found in the sand beds, but no shells. This section is interesting in that it lies nearly 40 miles beyond the outermost moraine of Missouri coteau.

Moose Jaw

Along the railroad track at the northern edge of the city of Moose Jaw, in NE. ½ sec. 4, tp. 17, range 26, W. 2nd mer., several exposures show two tills distinguished by differences in structure. The lower till is somewhat darker and more compact than the upper and shows columnar jointing in the upper part, but is not greatly weathered. The sections are northeast of Missouri coteau.

Several wells in the vicinity of Regina and Moose Jaw pass through two or three tills separated in places by sands and gravel. Lignite is reported from the sand and gravel at the base in some of the wells, but in most cases the sands appear to be free from organic matter. It is reasonable, however, to consider the sands and gravel as interglacial. The evidences in Swift Current valley indicate how such a condition is possible. Moreover, glacial outwash associated with the moraines of the last glaciation is not abundant in the region and this suggests that the buried sands and gravels are interglacial deposits and not entirely glacial outwash.

AGE AND EXTENT OF DEPOSITS

The main question in connexion with the age of the deposits is, were they formed during interglacial warm climate periods and can they be correlated with other deposits of similar character, the age of which has been approximately determined. It is fairly certain that the Missouri Coteau moraine, which is a continuation of the Altamont moraine of North Dakota, is Winsconsin in age. It is probable, therefore, that the uppermost till at all the localities, except possibly near Lac Pelletier, is Wisconsin. There are two lower tills the ages of which are unknown.

The faura and flora of the upper series of interglacial deposits do not appear to indicate a climate as warm as the present. The shells are small and thin. The fauna found in the Assiniboine delta in glacial Lake Agassiz

Leonard, A. G.: "Geology and Natural Resources of North Dakota"; Jour. Geol., vol. 24 (1916).

basin, Manitoba, may be taken as an example of the development of life in a region during or shortly after the retreat of the ice-sheet. The fauna shows much larger forms than those found in the Saskatchewan deposits and very similar types of ostracods to those found in the deposits near lake Johnstone. The occurrence of well-developed spruce trees indicates a more moist climate than the present, as the area is nearly treeless owing to

deficiency in moisture.

The thickness of the weathered zone at the top of the lowest till in Swift Current valley and the stratified sands and gravels containing lignite appear to indicate a long period during which the region was free of ice. The smaller, weathered zone observed in the exposure near Verwood may represent a period as great as post-Glacial time in this region; similar zones in cuts at the surface in this area show weathering down to only about one foot. As the interglacial climate, however, probably was more moist than the present, weathering may have been more rapid. Verwood exposure probably represents the same stage as the deposits near Beaubier, Blooming, and lake Johnstone, as all these are nearly similar in The fact that the later deposits do not have a very thick covering of till is not of much significance because the thickness of till is variable. The younger interglacial beds evidently correspond to a period of deglaciation of no great length and of cool climate between the Wisconsin and Iowan or Illinoian or between Early and Late Wisconsin; the older beds represented in the exposures near Swift Current probably belong to an earlier period of considerable length. A fairly warm and moist climate may be indicated by the lignite, which, however, is too much altered to determine the character of the organic material. It is not clear to which stage the Lac Pelletier deposits belong as it is not certain whether the uppermost till in this area is Wisconsin in age or belongs to an earlier stage.

The deposits occur at widely separated localities and may be of considerable extent, but appear to be confined to buried stream valleys and to former basins of small ponds. They are most extensive in the morainic areas and on the southwest side of the moraines. As the deposits largely consist of well-washed sands and gravels they are of importance as a source of ground water. Prospecting for them is necessarily done by drilling and by tracing the courses of the ancient valleys by means of records of borings and by the regional slopes which indicate the direction of drainage.

VARIATIONS IN THICKNESS OF CRETACEOUS FORMATIONS IN SOUTHERN MANITOBA

By. R. T. D. Wickenden

A study of well sections in southern Manitoba reveals marked differences in thickness of some of the Cretaceous formations in different areas. studies were based on an examination of the micro-fossils in the well samples and the lithology of the samples from the Deloraine well drilled in 1892, and the Pembina Valley well drilled in 1930 by the Commonwealth Petroleum Company, Limited. The subdivisions of the Cretaceous of the Manitoba escarpment as described by S. R. Kirk¹ have been followed, and comparison is made with the thicknesses of the various formations as determined by him. Examination of foraminifera from outcrop and well samples indicates that identification of horizons for correlation purposes and structural determinations can be based on the micro-faunas of the Boyne and Assiniboine beds. Over fifty species of foraminifera have been observed in material from the Boyne beds. Many of these species are easily recognized, are abundant, and have a short stratigraphic range. Close correlations in limited areas can, therefore, be made by means of the micro-fauna of the Boyne beds. The Assiniboine beds, on the other hand, have a small micro-fauna of only four species, which occur throughout the beds. In some parts of the beds, particularly near the top of the series, the foraminifera are so abundant that the rock can be called a fossil globi-The fauna and its occurrence are consistent and characteristic of these beds over a wide area, thus making the Assiniboine beds a dependable and easily recognized horizon.

The samples of the Pembina Valley well down to 740 feet were received from the Calgary office of the Supervisory Mining Engineer, Department of the Interior, and those from 750 feet to 1,070 feet from Hon. S. Michener,

President of the Commonwealth Oil Company, Limited.

Log of Pembina Valley Well No. 1, L.S.D. 4, Sec. 23, Tp. 2, Range 9, W. Prin. Mer.

Formation	Depth in feet	Description
Pembina beds	20	Dark grey to brown, non-calcareous shales
Boyne and Morden beds	40 to 425	Dark grey, calcareous shales Foraminifera: large fauna of calcareous forms in the upper beds, arenaceous forms in the lower beds
Assiniboine and Keld beds	445-545 565 585-620 630-640	Dark grey, speckled and unspeckled, calcareous shales, with pelagic micro-fauna Limestone band, inoceramus prisms and foraminifera Dark grey, calcareous shales with foraminifera Dark grey, calcareous shales with bentonite
Ashville beds	650-740	Dark grey, calcareous and non-calcareous shales; some green sand and arenaceous foraminifera in the lower samples
760-780 790		Fine, white, calcareous sandstone with much pyrite Medium to coarse, white, quartz sand Dark grey, non-calcareous shale with some coarse quartz sand Very coarse sand and quartz pebbles up to 8 mm.

¹Kirk, S. R.: Geol. Surv., Canada, Sum. Rept. 1929, pt. B, p. 114.

Log of Pembina Valley Well No. 1-Continued

Formations	Depth in feet	Description
Palæozoic (probably Devonian)	812 820 830 840 850-60 912 990—1,020 1,030 1,040	Red and pistachio-coloured shale with some coarse sand White and pink shale with coarse sand Bright brick-red shale Light grey shale Pink shale Terra-cotta and green shale Fine, white, calcareous sand Medium grey shale Dark buff to terra-cotta shale Very light grey and dark buff to terra-cotta shale Light grey and brownish red shale

The Boyne beds grade into the Morden beds so that they cannot be separated in well samples, but the former are characterized by a fauna of calcareous foraminifera. The Assiniboine beds also grade into the Keld beds. The base of the Keld is arbitrarily placed at the base of the bentonite layers in accordance with Kirk's classification, though the calcareous shale in the upper part of the Ashville is more typical of the Keld and the characteristic foraminifera occur in the lower part of the Ashville only.

The Deloraine well is reported to be 1,943 feet deep, the lower 121 feet having been in the Basal beds.¹ Only samples from 1,070 feet to 1,800 feet are available for study, but these represent the greater part of the formations passed through in Pembina Valley well No. 1.

Partial Log of Deloraine Well

Formations Depth in feet		Description
Boyne and Morden beds	1,075 to 1,460	Calcareous shales; non-calcareous near the base. Foraminifera in the upper samples only
Assiniboine and Keld beds	1,470 to 1,580	Speckled calcareous shales. Pelagic foraminifera
≯shville beds 1,585 to 1,800+		Dark grey, non-calcareous shales. Arenaceous foraminifera

The thicknesses of the formations at different places are shown in the following table. The sections at Pembina mountain and at Riding mountain (Vermilion River beds) are as given by Kirk.²

	TI	Total				
Localities	Boyne and Morden	nd boine Ashville bade			thick- ness	
Pembina valley, well No. 1 Deloraine well Pembina mountain Riding mountain	385 340	190+ 110 160 160-165	90 215+ 100-150 170	50 121 90 19–90	715+ 816+ 670-740 599-725	

The table shows that the formations comprising the lower part of the marine Cretaceous rocks in southern Manitoba vary considerably in thickness over long distances. In limited areas the variation is small and if series like the Morden and Boyne beds are taken together the stratigraphic thickness is fairly uniform.

Tyrrell, J. B.: Geol. Surv., Canada, Ann. Rept., N.S., vol. VI, pt. A, p. 2 (1895).

Kirk, S. R.: Geol. Surv., Canada, Sum. Rept. 1929, pt. B, p. 114.

DEEP BORINGS IN THE PRAIRIE PROVINCES

By W. A. Johnston

(Chief, Division of Pleistocene Geology, Water Supply, and Borings)

The importance of obtaining dependable supplies of water, and the increasing necessity for providing information as to the probability of obtaining water from underground sources as the population of Canada increases both in numbers and density have led to the formation of the Division of Pleistocene Geology, Water Supply, and Borings. In the Prairie Provinces during 1930 the necessity for obtaining additional water supplies was emphasized by the very dry seasons that some parts of that area have recently experienced. The work of the Borings Division being so closely related to that of the new division it was considered advisable to include it in the newly formed division. The work which the Borings Division has been carrying on for over 20 years, involving the collection of samples and records from wells drilled for oil, gas, and water, will continue unaltered.

The necessity for the collecting of records of wells drilled for water is readily seen when the conditions prevailing in the Prairie Provinces are considered. It is an area almost entirely drift covered, supporting a large farming population, and data as to sources of water supply are generally only obtainable as a result of well-drilling operations. The rather irregular distribution in the drift of the sand and gravel deposits from which water is obtained also makes well records of particular value. The division is specially organized and operated to deal with the work of the collection of samples and records and it is hoped that drillers and owners of water wells will co-operate in the collection of data of this nature. The assistance and advice of the field officers of the Geological Survey are always available in the interpretation of the stratigraphic records.

A report on water conditions at Moose Jaw is given elsewhere. With the exception of Regina district, nothing of special interest seems to have transpired elsewhere in connexion with drilling for water. Thanks are due to the Layne Canadian Water Supply Company for information as to water wells drilled by them, to Messrs. McPherson and Thom of Regina for similar information in Balcarres district, and to the Clark Drilling Company of Wolseley, Saskatchewan, for information as to wells sunk. The Canadian National railways forwarded records and samples from a large number of wells drilled for them in connexion with branch lines built in the west. Mr. S. Toye forwarded the results of some drilling in the gypsum deposits in Manitoba, on lot 74 one mile from Charleswood, Manitoba. Mr. J. W. D. Farrell, of the Regina Waterworks Department, forwarded logs of wells drilled in Regina district. Mr. H. J. McLean of the Calgary Power Company and Dr. J. A. Allan of the University of Alberta forwarded information on drilling done in connexion with the Ghost River dam of the Calgary Power Company. Mr. F. J. Malkinson, secretary of the town of Fairview, Alberta, forwarded information as to water conditions in that area. The names of individual drillers who cooperated will be found in the tabulated list of shallow wells which follows. That the interest in the oil and gas possibilities of the Prairie Provinces was maintained during 1930 is shown by the number of samples received— 28,859, only 207 less than in 1929. The arrangement by which samples were forwarded from the Calgary office of the Supervisory Engineer of the Department of the Interior was continued during 1929, and much information relative to well logs, oil, gas, and water analyses, etc., was also received from that office.

In Manitoba work on the well at Dauphin was continued during 1930 and a deep test is being put down near Pembina by the Pembina Valley Gas and Oil Corporation. Neither well has obtained commercial production to date.

In Saskatchewan further drilling was done in Eden Valley, Pike Lake, Simpson, and Riverhurst districts. Tests are also going down near Biggar

and at Pilot Butte near Regina.

In Alberta, Turner valley continued to be the centre of interest, many new wells coming in and a large production of high-grade naphtha and crude oil resulting. This remarkable field made by far the largest contribution to the oil and gas production of Canada during 1930. Considerable interest was shown and some development took place in the Red Coulée, Irma-Wainwright, and Ribstone fields. In Pincher Creek area, south-western Alberta, the Alberta Gas and Fuel Company and the Weymarn interests continued drilling. Many other wells were put down at various points. Of the new areas tested those at Spring Coulée, Paintearth creek, and Keho lake have not been completed, and the well at Duvernay has been abandoned. Some drilling was done for gas in Medicine Hat area. The second deep test in Cypress Hills area was put down to 3,900 feet. Many other test holes, both shallow and deep, were put down in other areas.

The system of bottling samples was continued during 1930 and several thousands were so treated. The laboratory work by Mr. F. J. Fraser on samples brought in from the field by Mr. F. H. McLearn was continued, about 200 being examined and many permanent slides of the heavy minerals made. Samples from the diamond drill cores obtained from the well drilled near Ralph, Saskatchewan, many years back, were re-examined. The first of the three well logs presented on following pages, provides

information as to conditions at depth in an area in which no deep drilling seems to have been done. The success of this Experimental Station well extends the previously known area of the artesian water basin of southern Alberta about 20 miles to the east. A notable feature is that the water flows at an elevation of 3,050 feet, which is about 50 feet higher than that of any other flowing wells in the basin as reported by Dowling,1 and this height is only about 150 feet below the elevation of the intake beds on Milk river 50 miles to the west. Probably the gas coming with the water causes the water to rise higher than it otherwise would. Unfortunately, the water has a high mineral content, as it has in other outlying parts of the artesian basin. A partial analysis by the Division of Chemistry of the Dominion Experimental Farms, Ottawa, showed a mineral content of 1,494 parts per million, consisting chiefly of sodium bicarbonate, 986 parts per million, and sodium chloride, 283 parts per million. The presence of large amounts of sodium bicarbonate in the water renders it unsuitable for irrigation, but it is a usable water for stock. Flow is 6 G.P.M.; temperature 54.5 degrees F.; a little gas.

The second well log provides information as to the nature of the basal Ordovician limestones and the underlying Winnipeg sandstone in the district west of lake Winnipeg.

¹Dowling, D. B.: "Investigation of Artesian Water, Coal, Petroleum, and Natural Gas in Alberta"; Geol. Surv., Canada, Sum. Rept. 1922, pt. B, pp. 110-114.

The third well log provides detailed information on a section of limestone, probably Palæozoic in age, nearly 1,000 feet thick in the Ribstone field in Alberta in an area in which these limestones are buried beneath over 2,000 feet of sediments and where the nearest outcrop of these limestones is about 175 miles distant.

Log of Well, Dominion Range Experimental Station, Sec. 15, Tp. 2, Range 4, W. 4th Mer.

Depth (in feet)		Rate of efferves- cence with acid Cold	· Notes
10	Sand, light brownish grey	0	Medium grained
20	", Ight browning grey	ž	Medium to coarse grained
30- 60	Clay, " "	ĩ	Considerable sand
70	medium grey	Î	Commission banks
80	Shale, "Shale,	î	
85-110	Sandstone, " "	Ī	Fine grained
120	Shale, " "	Î	Considerable sand
130-140	" " "	l ī	C CALLAGO DIANA
150-160	Sandstone, " "	î	Fine to medium grained
170-190	Shale. "	ī	Considerable sandstone
200-220	" " "	ī	Company Sandstone
230	" " "	ĺ	66 66
240	Sandstone, " "	Ī	Fine to medium grained
250	Sandy shale. "	1	
260-290	Sandstone, " "	ī	66 66
300-310	Shalv sand-	_	
	stone " "	1	
320	Sandy shale. " "	1	
330-370	Sandstone, " "	1	66 66
380	Shale, very dark grey	1	Plant remains
390	" " "	1	Coal fragments
400-4-10	" dark grey	1	"
450-470		1	Very few coal fragments
480-630	" brown-grey	1	
640-670	" dark brown-grey	1	66 66
680-710	" " "	1	
720	66 66 66	1	Coal fragments
730-750	******	1	
760-780	" medium grey	1	White shell fragments
790-830	" "	1	1
840		2	A few coarse sandstone fragments. White shell fragments
850	" "	1	
860	66 66	1	A few coal fragments
870-920	" " "	1	
930	" " " ········	1	Iridescent shell fragments
940	" "	1	4.6
950-960	" "	1	A few coal fragments
970	" "	1	C1-11 6
980	" "	1	Shell fragment
990		1	Iridescent shell fragments
1,000	Shaly sandstone, medium grey	1	Fine grained
1,010-1,020	Sandy shale,	1	Fine to madium emined
1,030	Sandstone, "	1 1	Fine to medium grained a little shale
1,040-1,130	"		Medium to coarse grained
1,136	1	1	bureathur on coarse Stattlen

Percentage of coal in unwashed sample from depths noted: 380 feet, 5 per cent; 390 feet, 45 per cent; 400 feet, 5 per cent; 410 feet, 3 per cent; 420 feet, 4 per cent; 430 feet, 15 per cent; 440 feet, 2.5 per cent; 660 feet, less than 1 per cent.

2.5 per cent; 400 feet, 18 per cent; 420 feet, 2 per cent; 420 feet, 18 per cent; 420 feet, 220 feet, 18 per cent; 420 feet, 18 per cent; 630 feet, 18 per cent; 640 feet, 65 per cent; 650 feet, 650 feet and 660 feet and 660 feet and 650 feet and

Log of R. J. McGuckin No. 1 Test Well, Southwest Part Tp. 33, Range 1, W. 1st Mer.

Depth (in feet)		Rat effer cence	with	Notes
		Cold	Hot	
20- 30 30- 60 60- 70 70-142 142-165	Limestone, brown-grey	4 6 4 2 4 0	6 6 6 6 0	Medium to coarse grained, fairly well rounded, some etched grains
180-232	" light grey	0	0	rounded, some etched grains

This log does not show the base of the Winnipeg sandstone, but hole No. 3 records a thickness of 112 feet which would make the lower contact in well No. 1 about 250 feet.

Log of Ribstone Oils No. 2 Well, Ls. 5, Sec. 25, Tp. 46, Range 1, W. 4th Mer.

Depth (in feet)		Rate of efferves- cence with acid	Notes
		Cold Hot	
70- 90 100 111 120- 180 190- 200 210- 250 260 270 280 290	Sandstone, grey "brown-grey	3	With pebbles "" "" Medium grained Fine grained
310	Sandstone, grey Sandy shale, brown-grey	3	66
510- 600 610- 620	Shale, grey		Somewhat sandy A few sandy partings Trace of bentonite Pearly shell fragments at 680 feet
630- 820 830- 850 860- 960			Some ironstone At 890 feet trace of very fine sandstone and fragments of Baculites (?)
970- 980 990-1,060 1,070-1,100	" dark grey		Ironstone at 970 feet Trace of bentonite at 990 feet Shell fragments
1,110-1,150 1,160-1,290	grey		Shell fragments. Bentonite 1,260-1,270 feet
1,300-1,450 1,460-1,540 1,550 1,560 1,570-1,650 1,660	« « « « « « « « « « « « « « « « « « «		Fissile, shell fragments, traces bentonite Thin, sandy partings Ironstone Fish tooth Trace fine-grained sandstone at 1,650 feet A little medium-grained sandstone

Log of Ribstone Oils No. 2 Well-Continued

Depth (in feet)		Rate of efferves- cence with acid		Notes
		Cold	Hot	
1,710 1,720-1,739 1,740-1,743	" " " brown-grey	1		Bentonite
1,750 1,753	Sandy shale, dark grey Sandstone, grey Shale, grey	1 1 1		Medium grained, much shale, fibrous calcite Pebbles, coal
	Sandstone, brown-grey	0		Fine grained, pebbles "pyrite at 1,765 and 1,775 feet
1,779 1,785-1,798 1,802-1,835		0 0 0		Fine grained, pebbles, coal fragments Fine to medium grained Medium grained, little coal at 1,802, 1,835 feet
1,840 1,845–1,852 1,859–1,900	" brown-grey	0 0 1		Medium grained, pebbles, pyrite "Fine to medium grained. A little coal
1,908-1,932	Coal, blackSandstone, medium grey	0		at 1,889 feet Fine to medium grained. Salt water at 1,903 feet
1,946 1,956 1,970	Sandy shale, brown-grey Shale, dark grey brown-grey	0 0 0 1 0		Fine to medium grained Medium grained. Salt water at 1,953 feet Coarse sand and pebbles Sandy. A little coal
2,033	Sandstone, brown-grey	0 0 0 1 1		Medium grained. Salt water and oil "" much shale Pyrite Medium grained. Considerable shale Fine to medium grained. Considerable
2,065 2,069	Sandy shale "Shale, dark grey	0 1 4		shaly towards top and at 2,051 feet Considerable sandstone
2,090-2,100 2,105	Sandstone, medium grey " " Pyrite, brown-grey Sandstone, medium grey	3 2 0 1		Fine to medium grained. Considerable shale " " Some shale and sandstone Fine to medium grained, considerable shale
2, 119 2, 127-2, 133 2, 137	Shale, dark grey	0 1 1 1		CC CC CC
2,178 2,186	Sandstone, medium grey " " Conglomerate, grey and dark grey	0 0 0		Medium to coarse grained Coarse grained
	8-73	,	l	

79B

Log of Ribstone Oils No. 2 Well—Continued

Cold Hot Cold Hot Cold Hot Cold Hot Cold Cold Hot Cold							
Cold Hot Cold Hot Cold Hot Considerable Cold Hot Considerable Cold Coarse grained. Coa			effer	rves- with		residue after	tate with barium
Sandstone medium gray			Cold	Hot		2.01	CILICIAGO
2,229	2, 197-2, 216		0				
2,236	2,229	Sandstone, medium	0		Coarse grained		
2,280 Conglomerate, medium grey 3	2,237	1 " "			Medium grained		
2, 270-2, 271 2, 275 2, 276 2, 277 2, 276 2, 286 2, 279 2, 286 2, 292-2, 298 2, 306-2, 312 2, 320-2, 330 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 371 2, 380 2, 380 2, 380 2,	2,250				Some coarse grains		
2,275							
2,279	2,270-2,271 2,275			4			0
2, 286 Limestone, browngrey 2, 306-2, 312 Sandstone, browngrey 2, 320-2, 330 Limestone, browngrey 2, 340-2, 351 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 360 2, 371 2, 370 2,				ł	rounded sand grains		•
2, 292-2, 298	2,286	Limestone, brown-			Considerable shale		
2,306-2,312 Sandstone, browngrey 2,320-2,330 2,340-2,351 2,360 2,371 Dolomite, browngrey 2,376-2,387 2,388 2,411-2,429 2,444 3,42 2,462-2,548 2,660-2,574 2,630 2,630-2,610 2,630 2,630-2,610 2,630-2,630 2,630-2,830 2,701-2,732 2,101-2,		grey	4				0
2,320-2,330		Sandstone, brown-			Grains well rounded.		
2, 320-2, 330 Limestone, browngrey		grey			Considerable lime-		
2,340-2,351 2,360 2,371 Dolomite, brown- grey Limestone, brown- grey 2,600-2,610 2,620 2,633-2,683 2,692 Limestone, brown- grey 2,701-2,732 Limestone, brown- grey 2,775 Limestone, brown- grey 2,775 Limestone, brown- grey 2,775 Limestone, brown- grey 2,776 Limestone, brown- grey 2,777 Limestone, brown- grey 2,778 Limestone, brown- grey Limestone, brown- grey 2,778 Limestone, brown- grey Limestone, brown- grey 2,778 Limestone, brown- grey Limestone, brown- Grey Coolomite, residue	2,320-2,330		3	4	Well-rounded sand		0
2,360 2,371 Dolomite, brown- grey 2,376-2,387 Limestone, brown- grey 2,444 2,445-2,548 2,465-2,548 2,560-2,574 Dolomite, brown- grey Limestone, brown- grey 2,1444 2,465-2,548 2,560-2,574 Dolomite, brown- grey Limestone, brown- grey Limestone, brown- grey 2,600-2,610 2,620 2,633-2,683 2,635-2,830 2,775 Limestone, brown- grey 4	2 340-2 351	grey "	5	6	grains Shaly residue		0
2,376-2,387 2,398 2,411-2,429 3 4 Dolomitic, residue very low 2,411-2,429 4 5 Dolomitic, residue very low 0 Dolomitic Very little 0 Dolomitic 0 Orgrey	2,360		5	6	inary residue		0
2,376-2,387 2,398 2,411-2,429 2,444 2,465-2,548 2,560-2,574 2,000-2,610 2,630-2,630 2,632-2,683 2,701-2,732 2,100estone, brown- grey 2,744-2,763 2,775 2,784 2,795 2,800-2,800 2,805-2,830 2,775 2,805-2,830 2,355-2,830 2,	2,371		3	5			0
2,411-2,429	2,376-2,387	Limestone, brown-	4	5			0
2,411-2,429	2,398	grey "	3	4			0
2,444	2,411-2,429	1	4	6	very low		0
2,586-2,593	2,444	1			Dolomitic	Vone little	
2,586-2,593 Limestone, browngrey	2,560-2,574	Dolomite, brown-		5		very little	
2,620 Dolomite, medium grey Limestone, medium grey Limestone, brown-grey are grey Limestone, brown-grey are grey	2,586-2,593	Limestone, brown-	4	6		"	0
2,620 Dolomite, medium grey Considerable Co	2.600-2.610	grey "	3	5	Dolomitic	*	0
2,633-2,683 Limestone, medium grey Considerable Considerable	2,620					Bulky clay	ŏ
2,701-2,732 Limestone, browngrey 2,744-2,763 Limestone, browngrey 2,744-2,763 Limestone, browngrey 2,775 Limestone, browngrey 2,775 Limestone, brown 3 4 Bituminous. Dolomitic 2,784 Limestone, dark brown Limestone, brown 5 6 Medium 0 2,790 Limestone, brown 5 6 Medium 0 2,790 Limestone, brown 5 6 Medium 0 2,790 Limestone, medium 5 6 Medium 0 2,795 Limestone, brown 5 6 Medium 0 2,795 Limestone, brown 5 6 Medium 0 2,795 Limestone, brown 5 6 Medium 0 2,805-2,805 Limestone, brown 5 6 Medium 0 2,805-2,805 Medium	2,633-2,683	Limestone, medium	5	6		"	0
2,701-2,732 Limestone, medium grey 5 6 Small clay 0 2,744-2,763 Limestone, browngrey Limestone, brown 3 4 Bituminous Dolomitic 2,784 Limestone, dark brown 5 6 Medium 0 2,790 Limestone, brown 5 6 Medium 0 2,790 Limestone, medium grey 5 6 Bulky 0 2,795 2,800-2,805 " " 5 6 Small 0 2,805-2,830 2,855-2,850 " " 5 6 Trace 2,875 Limestone, medium grey 5 6 Trace 2,875 Limestone, medium grey 5 6 Trace 2,875 Limestone, medium grey 7 5 6 2,875 Limestone, medium grey 7 5 6 2,875 Limestone, medium grey 7 5 6 2,875 Limestone, medium grey 7 7 3 Limestone, medium grey 7 7 3 Limestone, medium grey 7 7 4 Bituminous Dolomitic 0 4 Bituminous Dolomitic 0 5 Medium 0 6 0 7 Medium 0 7 Considerable 0	2,692	Limestone, brown-	5	6		"	0
2,744-2,763 Limestone, browngrey 5 6 Small clay 0 2,775 Limestone, brown 3 4 Bituminous Dolomitic Free oil, quartz grains Considerable 0 2,784 Limestone, dark 4 4 Free oil, quartz grains Considerable 0 2,795 Limestone, brown 5 6 Medium 0 2,795 Considerable 0 Medium	2,701-2,732	Limestone, medium	5	6		"	0
2,784 Limestone, brown. 2,784 Limestone, brown. 3 4 Bituminous. Dolomitic Free oil, quartz grains Considerable. 0 Considerable. 1 0 Considerable. 2,790 Limestone, brown. 5 6 Medium. 0 grey 6 Small. 0 0 grey 7 6 Small. 0 0 Small. 0	2,744-2,763	Limestone, brown-	5	6		Small clay	0
2,784 Limestone, dark brown 2,790 Limestone, brown. 5 6	2,775		3	4		Medium	0
2,790 Limestone, brown. 5 6	2,784		4	4		Considerable.	0
2,795 Limestone, medium 5 6	2,790	Limestone, brown				Medium	
2,805—2,805 Limestone, brown-grey	2,793	Limestone, medium	5	6		"	0
2,805-2,830	2,795 2,800-2,805						
2,835-2,850				_			
2,855 " " 5 6 " " Present Trace 2,875 Limestone, medium 5 6 Wery little 0	2,805-2,830 2,835-2,850	44 44	5				
2,875 Limestone, medium 5 6 6	2.855		5	6		"	Present
grey	2,860-2,870						
2,000-2,000 0 7		grev					
	2,000-2,090	1	1 0	' '	1	Luiny Clay	, 0

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Log of Ribstone Oils No. 2 Well-Continued

Depth		effer	with	· .	Insoluble residue after HCl	Precipi- tate with barium
		Cold	Hot		HOI	chloride
2,895	Limestone, medium	6	7		Very little	0
2,900-2,905	Limestone, brown-	5	6		«	Present
2,910-2,915	Dolomite, brown-	2	7		"	**
2, 920-2, 930	Limestone, medium	5	7		"	**
2,935	Dolomite, medium grey	3	7		"	**
2,940	Limestone, medium	4	7		«	ce
2,945-2,950	Dolomite, medium grey	3	6			- 46
2,955-2,960	Limestone, brown-	5	7		***	ш
2,965	Dolomite, brown-	3	7		"	' «
2,970	Limestone, brown- grey	5	6			66
2,975	Dolomite, brown-	3	. 6		ee	
2,980	Limestone, brown-	4	6	.,	eš	: "
2,985 2,990–2 ,995	grey	6 4	7 6		"	66 ',
3,000 3,005-3,010	Missing Limestone, brown-	5	6		"	. 66
3,005-3,015	grey Dolomite, brown-	2	7		Considerable.	"
3,020-3,025	grey	3	6 7		Very little	. "
3,030 - 3,035	Limestone, brown- grey Dolomite, brown-	3	6	,,		88
	grey	5	6	Salt propert	u	
	Limestone, brown-	2	6	Salt present	66	"
3,050 3,055	Dolomite, brown-	2	6			· a
3,060	66 66	2 3	6			"
3,065	Limestone, brown- grey	5	6		66	"
3,070-3,075	Dolomite, brown- grey	3	6		"	66
3,080-3,085	Limestone, brown-	6	7			"
3,090-3,095	grey	5	6	Trace of salt at 3,095	"	Absent
3,100		- 5	6		Considerable.	"
3,105-3,130	" "	5 6	67		Very little	Trace
3,135-3,165	44 44	1			· · · ·	3,145
3,170		6	7	_	Bulky argilla- ceous	Absent
3, 175-3, 185	Limestone, medium	6	7		Very little	
3, 190-3, 195	" " " " " " " " " " " " " " " " " " "	6	7	Coarse sand grains		"
3,200-3,220 3,230		6 6	7 7		Very little Bulky argilla-	. "
,			· ·	l	ceous	

Water at 120 feet, 140 feet, 301 feet; salt water at 1,743 feet, 1,953 feet, 2,023 feet; sulphur water

at 2,166-2,170 feet.

Gas, trace, at 275 feet; gas shows, at 1,372 to 1,392 and 1,462 feet; gas, at 1,600 feet and 1,740 to 1,743 feet.

Oil show at 2,023 feet.

Dept.

Summary:	
Ribstone Creek-Lea Park contact (?). Sandstone in Lea Park. Base of Upper Cretaceous. Top of limestone.	Feet 140 250–320 1,739 2,275
An analysis made by E. A. Thompson, of the Mines Bra of Mines, of the brine found at 3,010 feet, follows:	anch,
Specific gravity at 60 degrees F	1.194 26.98 22.59
Composition of Solid Matter Sodium chloride Potassium chloride. Calcium chloride. Magnesium chloride Calcium sulphate. Calcium carbonate.	Per cent 43.62 2.48 40.36 12.97 0.21 0.30 99.94
Composition of Brine per Barrel (35 Imperial gallons) 26.98 grs. of solid matter in 100 cc. 2,698 lbs. of solid matter in 1,000 Imperial gallons 2.698 lbs. of solid matter in 1 Imperial gallon 94.43 lbs. of solid matter in 1 barrel of 35 Imperial gallons Pounds of Salt Constituent Recoverable from One Barrel of 35 Imperial Gallons	Pounds
Sodium chloride. Potassium chloride. Calcium chloride. Magnesium chloride. Calcium sulphate. Calcium carbonate.	41·19 2·34 36·11 12·25 0·20 0·28
	94.37

Deep Wells ((Samples and Logs Received During the Year)

	Remarks	LL	245 C.N.W. L. H. and P. Company, Burdett No. 40	неек	233 Moose Oils Limited No. 1	NO	64 Carbon Exploration Company, Ltd., No. 1	ION	61 Champion No. 1, Hudson Bay Marland Oil Company	HILLS	405 Eagle Butte Co. No. 2	NII	34 Dauphin Oil Co. No. 1
	Number of samples received	BURDETT	245	CANYON CREEK	233	CARBON	75	CHAMPION	. 61	CYPRESS HILLS	405	DAUPHIN	34
Description	Depth Number in feet covered samples by records		2,670	CA	2,820		099		2,740	O	4,070		1,256
	Year		1930		1930		1930		1930		1930		1930
	Mer.		4th		5th		4th		4th		4th		lst
	Range		12		9		23		23		69		80
Location	Tp.		6		22		88		16		4		*
	Sec.		14		28		18		6		30		14
	Ls.		15		16		7		6		4		

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Deep Wells (Samples and Logs Received During the Year)—Continued	:	Remarks	VER	High River Oilfields Ltd. No. 1 Ranchmen's Oil and Gas Co. Ltd. No. 1	ООО	Northwest Calgary and Edmonton No. 2 Hudson Bay Marland Oil Co., Highwood No. 1 Western Alberta No. 1 Warner Oils Ltd. No. 2 No. 3	AKE	209 Hudson Bay Marland Oil Co., Keho Lake No. 1	A.A.	1 J. P. McDonald No. 1	DGE	Lethbridge Petroleum No. 3 " " No. 1
		Number of samples received	HIGH RIVER	10	нісниоор	241 172 418 558 47 278	KEHO LAKE	208	KINSELLA	н	LETHBRIDGE	280
s and Lo	Description	Depth Number in feet of covered samples by records		1930 143 1980 3,990		2,260 3,310 4,010 3,872 2,770	-	2,130	;	2,080		358 358 350
Samples a		m Year drilled				1930 1930 1930 1930 1930 1930		1930		1930		1930 1930 1930
Wells		Mer.	*	4th		5th 5th 5th 5th		4th		4th		4th 4th
Deep		Range		29		ක ක ක ක ක		22		10		222
	Location	Tp.		18		1888		11		48		00 00 00
		Sec.		34		35.7 200.7 2		11		29		23 4 23
3		Ls.		91		∞∞∞±4∞		23		80		401

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MEDICINE HAT	96 Medicine, Hat Brick and Tile Company No. 1	MILK RIVER	141 Commonwealth Petroleum Ltd., Milk River No. i	MOOSE MOUNTAIN	48 Signal Hill No. 2	MORLEY RESERVE	1,588 7 Norcon No. 1 NEW BLACK DIAMOND		32 Outwest Petroleum No. 1 33 New Black Diamond, Weymarn No. 1	NEW VALLEY	2 New Valley Oil Ltd. No. 1	PAINTEARTH CREEK	173 Hudson Bay Marland Oil Company, Paintearth Creek No. 1	PEACE RIVER	19 Four X Oils Co. Led. No. 1							
M	1,025		4,346 MOO	MOC	4,090	MOI	1,588	NEW	3,690 3,910	-	2,210	PAIN	1,840	Д	230	To the second						
						,																
,	1930		1930		1930			٠					1930		1930 1930	. :	1930		1930		1930	-
	4th		4th		5th		erve		5th		5th		14 4th		5th							
	70 70		15		ъĊ.		Morley reserve		8 8 954 554		67		14		12							
	12		60		23				19 20		231		88		88							
	33 33		6	-	34				33.5		9		27		36							
	44		6		6				1011		4		r¢		2							

Deep Wells (Samples and Logs Received During the Year)—Continued

Description	Remarks	AKE	83 Pike Lake Oil and Gas Development and Exploration Co. Ltd. No. 1	CREEK	230 Weymarn Petroleums, Pincher Creek No. 2 265 Alberta Gas and Fuel Co., Yarrow No. 1	JFF	71 Redoliff Premier Brick Co. No. 1	ULÉE	Dalco Oil and Gag Co, No. 2 Capital Oil and Natural Gas Co. Ltd. No. 1A Capital Oils Ltd. No. 1 Celtic Oils Ltd. No. 1 Commonwealth Petroleum, Ltd., Red Coulée No. 1 Dixie Oil Co. No. 1A Alberta Pacific Consolidated, Red Coulée No. 1 Devonabire Oil Co. Ltd. No. 1 Ko-Top Oils Ltd. No. 1 Southern Alberta Exploration Co. No. 1 Taylor Oil and Gas Syndicate No. 1 Vanalta, Milk River No. 2
	Number of samples received	PIKE LAKE	88	PINCHER CREEK	230 265	REDCLIFF	71	RED COULÉE	193 275 275 275 280 280 280 211 113 218 225 255 255 255
Dosomintion	Depth Number of in feet covered samples by received		1930 3,170	PI	2,250		200		2, 970 2, 970 2, 970 2, 975 2, 975 2, 975 2, 977 2, 977 2, 977 2, 977 2, 977
	Year				1930 1930		1930		1930 1930 1930 1930 1930 1930 1930 1930
	Mer.		3rd		5th		4th		4444444444444
	Range		2		30		8		100
Lonetion	Tp.		3 ₹		3 8		13		8-8
	Sec.		55		34		6		811 412 100 100 100 100 100 100 100 100 100 1
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	Meridian Olls Ltd. No. 3. Oxville Oil and Gas Co. No. 1 Imperial Ribstone Oils Ltd. No. 2		Paramount Oils Ltd. No. 1		Simpson Oil Co., Roycroft No. 1		206 Dauntless Oils Ltd. No. 1 81 Devenish Petroleums Ltd. No. 1 172 No. 4	2	376 Alberta Pacific Consolidated Ltd., Spring Coulée No. 1		Ontario-Alberta Oil and Development Co. Ltd. No. 1	
T.	feridian (xville Oi mperial I	L1	aramoun	Z) aosdani		Jauntless Jevenish	ULÉE	lberta P	ď,	ntario-A	
KIDSIONE	193 N 187 C 242 C 182 I	RICKERT	30 I	SIMPSON	8 88	SKIFF	206 81 172	SPRING COULÉE	376	SUFFIELD	44	
	1,885 1,803 2,260 2,545		2,239		2,410		3,090 3,040 2,000	8P	3,978		3,520	
	1930 1930 1930 1930		1930		1930		1930 1930 1930		1930		1930	
	4th 4th 4th		5th		2nd		4th 4th		4th		4th	
	8		4		25		15 41 44		23		80	
	45 45 46 45		19		29		ಹಿಸುಗು		*		17	
	16 10 9		88		6		888		15		4	
	44 4		61		61		9 13 13		11		15	

Deep Wells (Samples and Logs Received During the Year)—Continued

	Remarks	SR	Cole Hunter, Taber Test No. 1 " No. 2 " No. 3 Majestic Mines, Test Hole No. 1 " No. 2 " No. 2 " No. 2 " No. 2	TURNER VALLEY	Advance No. 5A Alberta, Pacific Consolidated No. 1 Anaconda No. 1, Northwest Co. Ltd. Baltac No. 1, Northwest Co. Ltd. British Dominion Oil and Development, No. 2 Calmont, Oils Ltd. No. 1 No. 2 No. 4 No. 8 No. 10 No. 11 Dalfin Oils Ltd. No. 1 Dahlousie Oils Ltd. No. 1 Dannouse Oils Ltd. No. 1 Dannouse Oils Ltd. No. 1 Dannouse Oils Ltd. No. 2 East Crest No. 1 Rast Crest No. 1 Rast Crest No. 1 Rast Crest No. 1 Rast Crest No. 2
	Number of samples received	TABER	335 55		373 190 190 201 171 171 105 1128 113 113 114 40 40 40 40 40 40 40 40 40 40 40 40 40
Description	Depth Number in feet ocovered samples		73 91 97 130 95 96 96 3,317		6 020 5 795 5 795 5 795 5 705 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7
	Year		1930 1930 1930 1930 1930 1930 1930		1930 1930 1930 1930 1930 1930 1930 1930
	Range Mer.		16 4th 16 4th 17 4th 17 4th 17 4th 17 4th 17 4th 17 4th		
Location	Tp.		00 10 9 9 9 9 10		200 200 200 200 200 200 200 200 200 200
	Sec.		25 34 9 24 24 11		12222222 12222222222222222222222222222
	Ls.		co : 60 : co	4	500050000HHHHHH9005044

Foothills No. 2 189 Foothills No. 3 189 Freehold No. 1 189 Freehold No. 2 184 Freehold No. 1 180 18
4,0,4,4,4,4,4,4,4,4,4,4,4,4,4,4,4,4,4,4
1930 1930 1930 1930 1930 1930 1930 1930
म्बर्धाः स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स्वतंत्रम् स् स्वतंत्रम् स्वतंत्रम्
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85488000 r 8 1 r 4 3 1 1 1 0 8 8 5 7 7 8 1 8 8 8 4 4 4 1 8 4 4 4 6 9 8 8 5 5 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0

Deep Wells (Samples and Logs Received During the Year)—Continued

	Remarks	TURNER VALLEY—Continued	Royalite No. 20 No. 24 No. 24 No. 25 Sentinel Oils Ld. No. 1 Southern Lowery No. 1 Spooner Oils Ltd. No. 2 Spooner Oils Ltd. No. 1 Spray Oils Ltd. No. 1 Sterling Pacific No. 1 Sterling Pacific No. 1 Structure Oils Ltd. No. 1 Vanalta, Turner Valley No. 1 Wellington No. 1 Widney Oils Ltd. No. 1	OME	Twin Dome Oils Ltd. No. 1		Unity Valley Oils Ltd. No. 3
	Number of samples received	R VALLE	196 101 101 364 127 128 355 114 114 114 114 114 114 117 117 118 119 119 119 119 119 119 119 119 119	TWIN DOME	29	VERA	142
Description	Depth Number in feet covered samples by records	TURNE	7, 180 9, 180		1,250		1,850
	Year		1930 1930 1930 1930 1930 1930 1930 1930		1930		1930
	Mer.		5th 55th 55th 55th 55th 55th 55th 55th		4th		3rd
	Range		ର ପ ସ ସ ସ ପ ପ ପ ସ ସ ସ ସ ପ ପ ପ ସ ସ ସ ପ ପ ପ		88		**
Location	Tp.		812881188811888118881188811888118881188811888118881188811888118888		21		41
	Sec.		255850-10008451-85548351		16		14
	L8.		22222333333327		10		13

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Shallow Wells (Logs Received During the Year)

Remarks		≥ 22	E	All coarse sand be-	Water salty, a little gas at 273 feet)		Dry		Gravel bands in clay					
Driller		Friesen, C. K McCauley, J Rahrick and Mainville	Price, A Buckoly, E	. 33	Rahrick and Mainville	Buckoly, E Buckoly, E.	Buckoly, E	Buckoly, E Buckoly, E Rahrick, and	Mainville Price, A Buckoly, E	Byass, D	Buckoly, E	McPherson and Thom	McPherson and Thom	McPherson and	McPherson and Thom
Depth from surface to water	Ft. Ins.	147 17 6	320 14	17	234	260 23 27		: :		30 15				150	50
Depth to bed- rock	Feet														
Depth cover- ed by records	Ft. Ins.	154 42 136	360 26	27	374	360 35	35	52 27 340	208 31	83.0	34	130	200	200	100
At or near		Marchand Ninga	HirschSturgis	Hinchcliffe jct	Kakwa	Hirsch Lady Lake Hincheliffe	Endeavour	Usherville Tallpines	Bienfait	Copeau Wolseley	Chelan	Balcarres	Balcarres	Balcarres	Okanese reserveBalcarres
Mer.		E.P.M. W.P.M. 2nd	2 2	¥	3	222	3 3	3 3 3	3 3 3	: : :	2 2	:		Little Black Bear reserve	:
Range		18	44	4	4	יט יט יט	; ;	10 10 10	96	9 2 2	221		Little Black Bear reserve	erve	
Tp.		3 45	34.2	34	42	33.03	300	8884	42.2	17 6	333	e	ear res	ear res	
Sec.		19	30	30	9	23.3	0.4	15	36	9 11	16	reserv	lack B	ack B	reserv
Ls.		S S S S E E E E E E E E E E E E E E E E E E E	NW.	NW.	:	NW.	E Z Z	E Z	NE.	SW.	NW.	Okanese reserve	Little B	Little B	Okanese
Owner		Davidson, W. P Can. Pac. ry Can. Nat. rys	Gienn, Ervin Can. Nat. rys	33	33	Schopp, M Can. Nat. rys	"	3 3 3	Mandt, C. P	Wolseley town	Can. Nat. rys	Indian Affairs	Indian Affairs	Department Indian Affairs	Department Indian Affairs Department

McPherson and Clay throughout	Good supply		BB.	<u>M</u>	Water hard	1991	Water in sand	gravel 240-202 leeb					Water sand 26-31 feet			Water in coarse	Entirley in sand and	Very soft water
McPherson and	McPherson and	Clark, W. E	Clark, W. E Can. Well Sup-	ply Co. Buckoly, E Buckoly, E Buckoly, E Kech, H. J	Dempsey, S Dempsey, S Can. Well Sup-	Dempsey, S	Hay, T McPherson and	шопт			Rahrick and	Rahrick and	Buckoly, E	Buckoly, E.	Buckoly, E.	Buckoly, E	Buckoly, E	Buckoly, E Buckoly, E Fortier, C
:		96	69 6 87 6	35 0 17 9 6 99 6	20	116	165 205	117		26	80	22	10 6	35	. 9	35	28 4	16 58 6 24 405
	:			108							:							
526	291	186	159 6 121 6	45 0 26 25 6 121 6	200 105 301 5	146	275 265	176	06	163	259	232	31	45	34	60	42, 4	68 178 6 42 500
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Balcarres	Balcarres	Sintaluta	Sintaluta Zénon Park	Bjorkdale Bjorkdale Crooked River Zénon Park	Odessa Odessa New Osgoode	Lestock	Tehner Regina	Regina	Regina	Regina	Smuts	Aberdeen	Shellbrook	Ordale.	Mount Nebo	Shell Lake	Shell Lake	Amiens Amiens Mildred
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Okanese reserve.	Peepeekisis reserve	:	SW. ‡	NW.	SW.	NW.	SW. 1	SW.	SW.	NW.	NE.	SW.	SE. 1	W. W.	NW.	SW.	SE.	SW. NW.
Indian Affairs	Department Indian Affairs	Scott, J. G.	osley, naminond and Nanton Can. Nat. rys	***	Schroeder, J. P Can. Nat. rys	Musconequan	Rattray, G. T Crystal Brewing	Co., Ltd. Regina water-	Regina water-	Regina water-	works Smuts town	Aberdeen town	Can. Nat. rys	23	***	: 3	, ,	" " " " Greenhorn, A

Shallow Wells (Logs Received During the Year)—Continued

Remarks		Used for stock	Water larative		Water hard Water in gravel Water hard and	88	332	Water medium hard	Water in coarse sand	15	Water soft with
Driller		Buckoly, E Dom. Well	Dom. Well	Dom. Well	Robins, J. M Buckoly, E Elliott, R	Robins, J. M Fortier, C	Fortier, C. Elliott, R. Dom. Drilling	Robins, J. M Fortier, C. Buckoly, E Buckoly, E Layne Can. Water Supply	Buckoly, E Buckoly, E Buckoly, E	Hannah, P	Bengston, S
Depth from surface to water	Ft. Ins.	51 6 10 6 240	99	361	220 35 9 70	270	288	150 240 41 6 40 6 87	17 47	99	200
Depth to bed- rock	Feet			:							
Depth cover- ed by records	Ft. Ins.	61 45 300	266	441	300 44 9 294	272 360	490 198 426	240 320 50 6 50 6 167	55 61	100	808
At or near		Spiritwood Spiritwood	Sovereign	Sovereign	Sovereign Bapaume Sovereign	Sovereign	Sovereign Sovereign	Sovereign Sovereign Belbutte Belbutte	Medstead Medstead Medstead	Bracken	Bracken
Mer.		3rd "	2	*	* * *	2 2	222	* * * * *	* * *	2	*
Range		===	12	13	222	133	222	252224	15	16	16
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Sec.		888	31	26	9118	33	221 10	១ស្តូតតេច	222	*	4
Ls.		SE. P	NW.	8W. 1	SW.	SE. A	NZK EKK EKK	SE. F.	SE.	SE. 1	SE. 4
Owner		Can. Nat. 178 Head, Mrs. A	Turner, L. B	Sprott, W. J	Greer, J. V Can. Nat. rys Bailey, H	Piercy, B. F.	Miller, J. S Thornton, G. W McCauley, W. C	Millsop, J Lampman, B Can. Nat. rys Biggar town	Can. Nat. rys	Bracken village	,,

ot				of	15 76			축	ė ė	, 1 9	pu	ĵ¢		пе
Buckoly, E Hole dry until shot		See log on page 76	107	large supply of water in quicksand below 65 feet of hardpan	Alkaline water at 15 feet; sand at 56 feet; clayey gravel above sand	Lower water good	rift	Con- Entire-	ly in drift Dry. Black shale 50- 300 feet	ntirely in drift. Blue clay below water sand, water	hard Water soft, In sand	Dry sand 14-85 feet	Entirely in drift.	water soft. Nine barrels per hour
ry un		on pa	oft	supplith ith indu	e war sand slaye;	water	y in d	hard.	drift lack s	y in clay	soft.	nd 14	y in	soft.
ole dry until si		se log	Water soft	large supply Soft with a wateringuic below 65 fe	lkaline wate feet; sand feet; clayey above sand	ower	Entirely in drift	Water hard.	ly in drift Dry. Black sh 300 feet	Entirely in drift, Blue clay belov water sand, wa	hard ater soft.	ry sa	ntirely in	Water soft. barrels per
Η		S	: :	<u> </u>	- Y	ŭ			:	<u> </u>	<u>*</u>	<u> </u>	闰	1
Ξ.	Oszust, J	Colbert, J		Mr. Edwards			Ronnes, H	Oszust, J Barker, F	<u> </u>	- E4	ΕΉ	Ē	T	Schieck, T
skoly	rust,	bert,		Edw.			nnes,	rust,	Barker,	Barker, F	Barker, F.	Barker, F	Schieck, T	nieck,
Bu	Osz		:	Mr			: : <u>8</u>	Osz Baj	Baı	Baj	Ba	. Ba	Sel	Sch
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Can. Nat. rys	nders	perimental Sta- tion. Carlson, A.	Lindsay, A	Chisholm, F.F.O.	МсВаіп, Ј	**	Cross, S. Experimental	White, W. W	*	×	¥	×	Heath school	Reid, R. B
314	Anderson, E. W.	OB	H	O	X		OH	SC)					H	щ

Shallow Wells (Logs Received During the Year)—Continued

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Veaver and Co. Water at 387 feet oster, C	od	/ hard	good	Water soft in gravel; quicksand at	feet soft good	oft	
ft ard ow so	ard ft, go	ightly ift	hard,	oft in	100 fe ery sc oft,	oft irly 8	
ater at 387 feet ater soft ater hard ater now soft, we hard for a while	Water hard Water soft,	Water slightly Water soft	ter	Vater soft in	about 100 feet Water very soft Water soft, go	ater sc	
Wa Wa Wa	. Wa	× × × × × × × × × × × × × × × × × × ×	Wa	· -	W W W	 W8 	
Weaver and Co. Water at 38 Foster, C. Water soft Foster, C. Water hard Oszust, J. Water now	Wr. Mefstead Water hard Foster, C Water soft, good	Foster, C	Foster, C Water hard,	szust, J	Foster, C. Water very soft Oszust, J. Water soft, go	Foster, C Water fairly soft	
aver a ter, C ter, C ter, J	Mr. Mefstead Foster, C	ster, C	ter, C	ust, J	ster, C	ster, C	
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Weaver and Co. Water at 387 feet Foster, C. Water soft 112 Foster, C. Water hard 10 Oszust, J. Water now soft, y.	20	141	45	125	130	133	
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Rairview Rio Grande Halcourt	Beaverlodge	Beaverlodge	Beaverlodge	Beaverlodge.	Rio Grande	Rio Grande Beaverlodge	
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Eairview village Liberty, W Swanson, E Cleland, B.O	Lossing, C. F Stone, G	Hodges, V. C	r. A.	Lay, H. R	Cook, J	O'Connell, D Pack, G. W	
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#### OTHER FIELD WORK

### Geological

- G. S. Hume. Mr. Hume geologically mapped part of Jumpingpound 1-mile quadrangle (latitudes 51° 00′ to 51° 15′, longitudes 114° 00′ to 114° 30′) and part of Calgary Southwest 1-mile quadrangle (latitudes 50° 45′ to 51° 00′, longitudes 114° 00′ to 114° 30′), Alberta.
- W. A. Johnston. Mr. Johnston continued the investigation of the surface geology, including the soils, of Regina 8-mile map-area (latitudes 49° to 52°, longitudes 102° to 109°), Saskatchewan. He also investigated and mapped a block of seventy townships in southeastern Alberta and southwestern Saskatchewan; this work was done in co-operation with the Manyberries station of the Department of Agriculture.
- P. S. Warren. Mr. Warren completed the study and mapping of the geology of the northern part of Regina 8-mile map-area (latitudes 49° to 52°, longitudes 102° to 109°), Saskatchewan.
- S. R. Kirk. Mr. Kirk continued the study and mapping of the geology of the western part of Winnipeg 8-mile map-area (latitudes 49° to 52°, longitudes 95° to 102°).

### Topographical

- W. H. MILLER. Mr. Miller continued the topographical survey of Nordegg 1-mile quadrangle (latitudes 52° 15′ to 52° 30′, longitudes 116° 00′ to 116° 30′), Alberta.
- J. W. Spence. Mr. Spence completed the topographical survey of Wildcat Hills 1-mile quadrangle (latitudes 51° 15′ to 51° 30′, longitudes 114° 30′ to 115° 00′), Alberta.
- J. A. MACDONALD. Mr. Macdonald made a detailed topographical survey (field scale 800 feet to 1 inch) of the coal-mining area at Hillcrest, Alberta.

## INDEX

PA	GΕ	P.	\GE
Aberdeen, Sask	93	Boyne beds72,	73
Acknowledgments	74	Bracken, Sask	94
Admiral Oils, Limited	91	Brenner coal m	34
Agassiz Lake (Glacial)	70	Brine, Ribstone area, Alta., anal	81
Alberta		Britannia park, Sask	57
See also Milk River area		British Dominion Oil and Development	
Red coulée oil field		Company	88
Boring for oil and gas 75-	-07	Brock Petroleum Company	91
Alberta shales	-14	Buckley coulée, Mont	
Alberta Gas and Fuel Company	86	Buffalo gap, Sask	46
Alberta Pacific Consolidated Oils, Limi-		Bullopora-Clavulina horizon21,	. 22
ted	-88	Burdett, Alta	82
See also Figure 1	00	Buttress, Sask	51
	0	Duttiess, bask	91
Alkali lakes	_2	O.1 T.73	_
Allan, J. A	74	Caley, J. F.	2
Altamont moraine, N.D	70	Calgary Development and Producers	
Altitudes		Company	91
Milk River area, Alta	1	Calmont Oils, Limited	88
		Calvert, W. R. Cameo, Sask Campbell, W. P.	11
Amiens, Sask	93	Comes Carl	
Analysis		Cameo, Sask	93
Brine, Ribstone area, Alta	81	Campbell, W. P	21
Water, Colorado sand	21	Canadian National rys.	
Moose Jaw area58,	63	Wells 92	2-96
Red Coulée field	21	Mawer, Sask., anal	63
		Canadian Pacific ry., wells92,	
Anchor coal seam	35	Canadian Lacine Ly., wells	
Anderson, E. W	95	Canyon Creek, Alta	82
Angus Oils, Limited	91	Capital oil well13,	
Anxiety Butte coal seam, horizon33,	35	Carbon, Alta	82
Ashville beds	73	Carbon co., Mont	11
Ai_iL oing Lodg	73	Carlson, A.	95
Assiniboine beds72,		Caron Coals	80
Assiniboine delta, Lake Agassiz	70	Caron, Sask	60
Athabaska dist., Alta., wells	97	Celtic Oils, Limited2, 6, 13, 29,	86
Atlantic Keystone well	83	See also Figure 1	
•		Central Butte, Sask	62
Backman, O. L	31	Chambers, R	32
Bailey, H	94	Champion, Alta	82
Balcarres, Sask	0.2	Charlin 1 Carla	04
Daicarres, Sask92,	93	Chaplin I., Sask51,	61
Bapaume, Sask	94	Chelan, Sask	92
Basal beds72,	73	Chisholm, F. F. O	95
Bearpaw formation4, 19,	33	Chugwater formation	11
Bearpaw mt., Mont	12	Clark Drilling Company	74
Beaubier, Sask	$\overline{71}$	Clay	1 X
Deardner, bask			
Beaument Oils well	91	Moose Jaw area, Sask. See Moose Jaw	
Beaverlodge, Alta	97	Willowbunch area, Sask., rept. by F.	
Belbutte, Sask	94	H. McLearn 31	1-49
Belly River formation 4,	18	Claybank, Sask	31
Bengough, Sask		Cleland, B. O.	97
Bentonite	78	Climate	01
Dentonite14, 10, 12, 11,			*0
Berentin, Sask	95	Moose Jaw area	53
Biggar, Sask	94	Coal, Willowbunch area, horizons of	
Big Muddy, Sask31,	34	seams 32	2–35
Big Muddy coal seam33,	34	Cole Hunter well	88
Big Muddy I., Sask	43	Colgate formation 32	
Dig Muddy 11, Dask	42	Coloredo wester and	. 00
Big Muddy valley, Sask31, 35, 36,		Colorado water sand	1-20
Birch r., Mont	11	Commonwealth Petroleum, Limited 2, 4,	5, 9,
Bituminous shales	13	10, 16, 20, 24, 28, 72, 85	, 86
Bjorkdale, Sask	93	See also Figure 1	
Blooming, Sask	71	Contact coal seam33	. 34
Bonneau coal seam	34	Cook, J.	97
	36	Copeau, Sask	92
Bonneau I., Sask	90	Coronach Carls	
See also Twelvemile Lake valley	00	Coronach, Sask	44
Border oil field, Mont	22	Cosmos formation	22
Bore-hole records. See Deep borings		Coteau, The. See Missouri coteau	
Walls lors of		Cretaceous	
Borings Division	74	Man., southern, thicknesses72,	72
Boulder clay		Milk River area, and Red coulée oil	,
	71		3 10
Moose Jaw area, Sask55, 66	-11	field	2-19
31470—8			

PA	GE	P	AGE
Crooked River, Sask	93	Geology, general Milk River area and the Red Coulée	
Cross, S	95	Milk River area and the Red Coulée	
Crystal Brewing Company	93		219
Cunningham, R. A	97	Moose Jaw area54,	55
Cypress hills, Alta75,	84	Geology, structural Milk River area and Red coulée oil	
D-1 Oil and Car Commons 0 7 10 12	10		9-24
Dalco Oil and Gas Company2, 7, 10, 13, 20, 22, 30	26	field	83
See also Figure 1	, 00	Gibraltar well	91
Dalfin Oils, Limited	88	Gienn, Ervin	92
See also Figure 1 Dalfin Oils, Limited Dalhousie Oils, Limited	88	Gilbert, C	96
139TH1001V. 088K	-64	Giroux Bros	96
Dauntless Oils, Limited	87	Glauconite	3~13
Dauphin, Man	82	Glendon, Alta	96
Davis, N. B	32	Goldman, M	12 59
Dead Horse coulée, Alta	83 -97	Gravel. See Sand and gravel	99
Deep borings in Prairie Provinces	-97 73	Great Falls coal field, Mont	11
Deloraine well, Man	87	Greenhorn, A	93
Devonshire Oil Company	86	Greer, J. V	94
Dinosaur remains	32	Groundwater. See Water	
Dirt hills, Sask	32	Gye, Sask., coal seam33	, 34
Diver, Daniel	32	TT-1 A14-	97
Dixie Oil Company	86	Halcourt, AltaHans, Mr	60
See also Figure 2	0.5	Hargal Oils, Limited	89
Dobbin, C. E32,	35	Harptree, Sask	
Dome Oils, Limited	88 75	Head, Mrs. A	94
Dowling, D. B		Heath, Alta	95
Duvernay, Alta	18	Heinsburg, Alta	95
Dyc1, W. S	10	Herron oil well, Alta	83
Eagle formation	16	Herronton, Alta	96
Eagle Butte Company	82	High River, Alta	84
East Crest oil wells	88	Hinchcliffe jct	84 92
East End, Sask	35	Hirsch, Sask	92
Edalta Oils, Limited	91	Hodges, V. C	97
Eden Valley, Sask75,	83	Hoffmann, G. C.	31
Elbow oil well, Alta	$\frac{83}{22}$	Holgar oil well, Alta. See Figure 1	-
Ellis formation	92	Home Oils, Limited	89
Erdmann, C. E.	22	Homestead Oils, Limited	89
Erickson coulée, Alta 4,		Horizons of coal seams, Willowbunch area, Sask	
Eskbank, Sask	64	area, Sask	2-35
Evans, C. S.		Howe, M. A	12 83
Rept. by, on Milk River area and the		Hume, G. S	
	-30	atume, or beautiful and a second	, 00
Experimental farm, Alta	76	Illinois-Alberta well	89
Eyremore, Alta	83	Imperial Ribstone Oils, Limited	87
Toinview Alto	97	Indian Affairs Dept., wells 9	2-96
Fairview, Alta	74	Innerfold Oils, Limited	91
Featherston, W. E.	96	Interglacial deposits S. Sask., rept. by Wickenden 6	5 71
Ferris coal seam	35	Invaders oil well	89
Field work	98	Invaders oil wellIrma-Wainwright oil field, Alta	75
Fisher, C. A.	11		, ,
Foothills oil wells	89	Johnston, W. A	98
Foremost beds	18	Reports by, on:	
Forsythe ck., Sask	61		4-97
Fossils Alberta shales	21	Groundwater resources of Moose	n e4
Belly River formation	18	Jaw area, Sask	U-04
Montana oil and gas fields5, 11-13		Jurassic	, 11
Moose Jaw area, Sask67	68	Milk River area and Red coulée oil	
Pakowki formation	. 20		3-12
Southern Manitoba72	73	TZ I G I	
Willowbunch area, Sask 32	3-35	Kakwa, Sask	92
Four X Oils Company	85 75	Keele, Joseph	32
Fraser, F. J	31	Keld beds	
Frechette, HFreehold oil wells	89	Keogh coal seam, horizon33	34
Freeman Lundy oil well	89	Kevin-Sunburst oil field, Mont5, 11	1, 22
Frontier Oils, Limited	91	Kinsella, Alta	84

# 101B

PA		P	AGE
Kirk, S. R	98	Milk River, Alta 2	, 19
Knappen, R. S	11	Milk River area and the Red Coulée oil	,
Kootenay formation Milk River area		field, Alta.	
Alto 2 19	13	Pont by C S France	1 20
Alta	86	Rept. by C. S. Evans	T~90
Car Jac Elman 1	80	Milk River formation2, 5, 15, 19	, 20
See also Figure 1	0.1	Milk River ridge, Alta., physiographic	
Krieg coal seam33,	34	notes	1
		Mill City oil well	89
Lac Pelletier, Sask	71	Miller, J. S.	94
Lady Lake, Sask	92	Miller, W. H	98
Lake Alma, Sask	66	Millsop, J	94
Lake of the Rivers valley, Sask35,	36	Miracle oil well.	89
Area included in	31	Mississippian (Madison) limestone	5
Lake Valley, Sask		Missouri coteau, Sask53, 54, 56, 61, 65	
Tamera D	64	Model Offer I td	, 89
Lampman, B	94	Model Oils, Ltd	
Lance formation	32	Moose mt., Alta., oil well	85
Lay, H. R	97	Moose Jaw, Sask	31
Lay, H. R.  Layne-Canadian Water Supply Company			0-64
pany	94	Till	70
Leith, Edward	31	Moosejaw ck., Sask	, 61
Leonard, A. G	70	Moraine, Missouri coteau54, 65	, 66
Lestock, Sask	93	North Dakota	70
Lestock, Sask	84	Morden beds72	
Lindbergh, Alta	96	Morley reserve, Alta	85
Lindson A	95	Morrison formation	11
Lindsay, A. Lloyd's Petroleum, Limited Logs of wells. See Wells, logs of	91	Mortlach, Sask	
Lorg of wells - Vac Wells lorg of	91	Moulton, G. F.	10
Logs of wells. See Wells, logs of	07	Mount Nebo, Sask	93
Lossing, C. F	97	Mount Nebo, Bask	90
Lowery Petroleum, Limited	89	37 1 1	-
		Naphtha	75
McBain, J	95	Natural gas	
McCauley, W. C	94	Boring for. See Deep borings	
McConnell, R. G	31	New Black Diamond wells	85
McBain, J.  McCauley, W. C.  McConnell, R. G.  Macdonald, J. A.  McLean, H. J.  McLearn, F. H.	98	New McDougall-Segur wells	89
McLean, H. J.	74	New Osgoode, Sask	93
McLearn, F. H	75	New Valley Oil, Ltd	85
McLennan, Alta	96	Ninga, Man	92
McLeod oil well	89	Norcon oil well	85
McManamna, L.	51	Northwest Associated well	89
McNeely, D. M	51	Northwest Company	
McPherson, Messrs	74	, , ,	,
Madison limestone	11	O'Connell, D	97
Majestic Mines oil well	88	Odessa, Sask	93
Malkinson, F. J.	74	Oil	00
Mandt, C. P.	92	Milk River area and the Red Coulée	
Manitoba	34		1-30
Cretaceous formations, thickness of		field	
cretateous formations, timexness of	73		89
southern		Okalta Oils, Ltd	
Manyberries, Alta	95	Oldwives, Sask	67
Map, contour	10	Olemen, Chris	51
Milk River area, Alta	19	Onalta Oils, Ltd	91
Red Coulée oil field, Alta	23	Onefour, Alta	95
Map, sketch		Ontario-Alberta Oil and Development	
Milk River area and Red coulée oil		Company	87
wells	2–3	Ordale, Sask	93
wells	52	Ordale, Sask Orton, Edward, jun	32
Maple Leaf oil well	91	Outwest Petroleum well	85
Marchand, Man	92	Oxville Oil and Gas Company	87
Mawer, Sask	-64	Oyen, Alta	95
Mayland oil well	89		
Medicine Hat, Alta75,	85	Pack, G. W	97
Medstead. Sask	94	Paintearth ck., Alta75	, 85
Mercury Oils, Ltd	89	Pakowki formation	5-20
Meridian Oils, Ltd	87	Pale Beds	, 18
Merland oil well	89	Paramount Oils, Ltd	87
Midfield oil well	89	Parkbeg, Sask	54
Mid-Royal oil well	89	Peace River dist., Alta., wells85, 96	. 97
Mildred, Sask	93	Pembina mt., Man	73
Milk r., Alta., rocks		Pembina valley, Man	. 73
Springs	2	Pembina valley, Man	75
Wells17,	85	Perry, E. S	. 11
logs	4	Petroleum. See Oil	,
	-	a outocountries ovo out	

PA	GE	P	AGE
	31	*** *** ***	
Phillips, J. G Piercy, B. F		Skiff, Alta	87
Piercy, B. F	94	Slipper, S. E	, 17
Pierre shale	55	Snowdy Springs, Sask	55
Pike Lake, Sask	86	Snowdy Springs ck., Sask., 51, 54, 56, 59	. 61
Pincher Creek oilfield, Alta75,	86	Snowfall. See Climate	,
Distance of Carl	00	C. 1 D T	
Pleistocene, S. Sask	65	Souch, B. E	2
Police coulée, Alta	19	Souch, John	2
Population, Moose Jaw, Sask	58	Southern Alberta Exploration Company.2	13
Potlatch oil well, U.S	5	00 0	4 07
	Ð	20, 2	±, 21
Prairie Provinces		See also Figure 1	
Deep borings in 74	-97	Southern Lowery wells	90
See also Alta.; Man.; Sask.		Cough Coalestohomon m 50 51 65	
Dee also Alta., Mall., Dask.		South Saskatchewan r50, 51, 65	
Precipitation. See Climate		Southwest Petroleums, Ltd	90
Pre-Ellis strata	3-5	Sovereign Sask	94
Purple Springs, Alta	96	Sovereign, Sask	98
I urple oprings, Arta	90	Spence, J. W	
		Spiritwood, Man	94
Quadrant formation	11	Spooner Oils, Ltd	90
O-2 A11 C1-		Canasa Oila T td	90
Qu'Appelle r., Sask	51	Spray Oils, Ltd	
		Springs	2
Radville, Sask	66	See also Water	
	00	Spring Coulée, Alta75	07
Rain. See Climate			
Ralph, Sask	75	Sprott, W. J	94
Ranchmen's Oil and Gas Company	84	Stebinger, E	. 16
Transmitted a On and Gas Company	0.0	Steen, Sask	92
Ranger oil wells83,		C. P. C. 11	
Rattray, G. T	93	Sterling Pacific wells	90
Ravenscrag formation32-36,		Stone, G	97
D116- C1-	20	Stratigraphy	٠.
Red coulée, Sask			0 02
Redcliff, Alta	86		2–35
Red Coulée oil field, Alta, See Milk		Structure Oils, Ltd	90
Pirron area and the Ped Coulée field		Sturgia Sask	92
Triver area and the fred Codies held		Sturgis, Sask	87
Reeside, John B	35	Sumen, Ana	
Reeves, F	11	Sundance formation	11
Roging Sock 50 52 65 70	93	Surface deposits	
Reeves, F	90	Surface deposits	
Reid, R. B	95	moose Jaw area, Dask., map and notes	
Ribstone, Alta	87	52, 54	, 55
Ribstone Oils, Ltd., well No. 2	-80	Swanson, E	97
D:-1-G-1-J11	00	Sweet Green hills Mont	
Richfield oil well	89	Swanson, E	11
Richfield oil well	89 87	Sweet Grass hills, Mont	
Richfield oil well	87	Sweet Grass hills, Mont	11
Richfield oil well	87 73	Swiit Current r., Sask	11 8–71
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich.	87 73 32	Taber, Alta., oil wells	8–71 88–88
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta.	87 73 32	Taber, Alta., oil wells	8-71 8-71 88 92
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta.	87 73 32 97	Taber, Alta., oil wells	8-71 8-71 88 92
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta.	87 73 32	Taber, Alta., oil wells Tallpines, Sask Taylor Oil and Gas Company	8-71 8-71 88 92
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C Rivard l., Sask. See Twelvemile Lake	87 73 32 97	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1	88 92 , 86
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley	87 73 32 97 51	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company.  See also Figure 1 Tehner oil well.	88 92 8, 86 93
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley	87 73 32 97 51	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well.	88 92 8, 86 93
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65,	87 73 32 97 51	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well.	88 92 , 86 , 93
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard l., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads.	87 73 32 97 51	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well.	88 92 , 86 , 74 , 74 81
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard I., Sask. See Twelvemile Lake valley Riverhurst, Sask. Roads. Rockglen, Sask. 31, 36, 47,	87 73 32 97 51 75 2 48	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W.	88 92 , 86 , 74 81 94
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard I., Sask. See Twelvemile Lake valley Riverhurst, Sask. Roads. Rockglen, Sask. 31, 36, 47,	87 73 32 97 51 75 2 48	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59	88 92 , 86 , 74 81 94
Richfield oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce.	87 73 32 97 51 75 2 48 32	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59	11 8-71 88 92 , 86 , 74 81 94 9, 60
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well.	87 73 32 97 51 75 2 48 32 91	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J.	88 92 , 86 , 74 81 94
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard l., Sask. See Twelvemile Lake valley Riverhurst, Sask. Soek Twelvemile Lake valley Riverhurst, Sask. Roekglen, Sask. Roekglen, Sask. Rosel, Bruce. Roselite oil well. Ross, C. C.	75 2 48 32 91 2	Talbines, Sask	88 92 86 93 74 81 94 60 94
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. S.	87 73 32 97 51 75 2 48 32 91	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A Thornton, G. W. Thunder ck., Sask. Thurston, J. J. Till. See Boulder clay Toye, S.	88 92 86 93 74 81 94 60 94
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. S.	75 2 48 32 91 2 35	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company	88 92 , 86 , 74 , 81 , 60 , 94 , 74 32
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Soads. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. S. Royalite oil wells. Royalite oil wells.	75 2 48 32 97 51 75 2 48 32 91 2 35 90	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company	88 92 , 86 , 74 , 81 , 60 , 94 , 74 32
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. S.	75 2 48 32 97 51 75 2 48 32 91 2 35 90	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company	88 92 , 86 , 74 , 81 , 60 , 94 , 74 32
Richfield oil well Rickert oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17,	75 2 48 32 97 51 75 2 48 32 91 2 35 90	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company	88 92 , 86 , 74 , 81 , 60 , 94 , 74 32
Richfield oil well Rickert oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17,	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner walley, Alta. 75, 8 Turner Basin Oils, Ltd.	88 92 , 86 , 93 , 74 , 81 , 60 , 94 , 74 , 32 , 89 , 90
Richfield oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Sockglen, Sask. Rockglen, Sask. Rose, Bruce. Rose, C. C. Ross, C. C. Ross, C. C. Ross, C. S. Royalite oil well. Ross, C. S. Royalite oil wells. Royalite oil wells. Rosell, L. S. Royalite oil wells.	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner walley, Alta. 75, 8 Turner Basin Oils, Ltd.	88 92 , 86 , 93 , 74 , 81 , 60 , 94 , 74 , 32 , 89 , 90
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Roselite oil well.  Ross, C. C.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner walley, Alta. 75, 8 Turner Basin Oils, Ltd.	88 92 , 86 , 93 , 74 , 81 , 60 , 94 , 74 , 32 , 89 , 90
Richfield oil well Rickert oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard I., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17, St. Paul, Alta. St. Victor coal seam. 33, Sand and gravel	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta 75, 8 Turner Basin Oils, Ltd. Twelvemile L, Sask 34 Twelvemile Lake valley Sask 34	88 922 86 93 94 88 990 94 88 990 95 95 95 95 95 95 95 95 95 95 95 95 95
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Rose, Ituce.  Ross, C. C.  Ross, C. S.  Royalite oil well.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,  Sand and gravel  See also Interplacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta 75, 8 Turner Basin Oils, Ltd. Twelvemile L, Sask 34 Twelvemile Lake valley Sask 34	111 8-71 88 922 928 94 81 94 74 32 94 8-90 95 35 44 90
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Rose, Ituce.  Ross, C. C.  Ross, C. S.  Royalite oil well.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,  Sand and gravel  See also Interplacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta 75, 8 Turner Basin Oils, Ltd. Twelvemile L, Sask 34 Twelvemile Lake valley Sask 34	111 8-71 88 922 928 94 81 94 74 32 94 8-90 95 35 44 90
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Rose, Ituce.  Ross, C. C.  Ross, C. S.  Royalite oil well.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,  Sand and gravel  See also Interplacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner walley, Alta. 75, 8 Turner Basin Oils, Ltd.	88 922 86 93 94 88 990 94 88 990 95 95 95 95 95 95 95 95 95 95 95 95 95
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Rose, Ituce.  Ross, C. C.  Ross, C. S.  Royalite oil well.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,  Sand and gravel  See also Interplacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Talbines, Sask	111 88-71 88 92 92 94 86 93 94 94 8-90 90 90 90 97 73
Richfield oil well.  Rickert oil well.  Riding mt., Man.  Ries, Heinrich.  Rio Grande, Alta.  Ritchie, H. C.  Rivard 1., Sask. See Twelvemile Lake valley  Riverhurst, Sask. 51, 65,  Roads.  Rockglen, Sask. 31, 36, 47,  Rose, Bruce.  Rose, Ituce.  Ross, C. C.  Ross, C. S.  Royalite oil well.  Ross, C. S.  Royalite oil wells. 89,  Russell, L. S. 17,  St. Paul, Alta.  St. Victor coal seam. 33,  Sand and gravel  See also Interplacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Basin Oils, Ltd. Twelvemile Lake valley, Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd.	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well.   Rickert oil well.   Rickert oil well.   Riding mt., Man.   Ries, Heinrich.   Rio Grande, Alta.   Ritchie, H. C.   Rivard l., Sask.   See Twelvemile Lake valley   Riverhurst, Sask.   S1, 65, Roads.   Rockglen, Sask.   31, 36, 47, Rose, Bruce.   Rose, C. C.   Ross, C. C.   Ross, C. C.   Ross, C. S.   Royalite oil well.   Ross, C. S.   Royalite oil wells.   89, Russell, L. S.   17,   St. Paul, Alta.   St. Victor coal seam.   33, Sand and gravel   See also Interglacial deposits   Moose Jaw area, Sask.   See Moose Jaw Sandy ck., Sask.   50-57, 60, 61, Sandy Creek, Sask.   55-57,   Sandy Creek, Sask.   55-57,   Sandy Creek, Sask.   S55-57,   Sandy Creek, S	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Basin Oils, Ltd. Twelvemile Lake valley, Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd.	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well.  Rickert oil well.  Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask.  Rockglen, Sask.  Rose, Bruce.  Rose, Bruce.  Ross, C. C.  Ross, C. S.  Royalite oil well. Ross, C. S.  Royalite oil wells.  St. Victor coal seam.  St. Victor coal seam.  See Moose Jaw area, Sask.  See Moose Jaw Sandy Creek, Sask.  See Interglacial deposits  Moose Jaw area, Sask.  See Moose Jaw Sandy Creek, Sask.  See Interglacial deposits  Sandy Creek, Sask.  See Interglacial deposits  Sandy Creek, Sask.  See Interglacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Basin Oils, Ltd. Twelvemile Lake valley, Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd.	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well Rickert oil well Rickert oil well Riding mt., Man. Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17, St. Paul, Alta. St. Victor coal seam 33, Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. 55–57, Saskatchewan. See Interglacial deposits Moose Jaw See Interglacial deposits	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Talbines, Sask	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well	87 73 32 97 51 75 2 48 32 91 2 35 90 31	Taber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Basin Oils, Ltd. Twelvemile Lake valley, Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd.	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34	Talbines, Sask	1118-7188-7188-7188-7188-7188-7188-7188
Richfield oil well   Rickert oil well   Rickert oil well   Riding mt., Man   Ries, Heinrich   Rio Grande, Alta   Ritchie, H. C   Rivard l., Sask. See Twelvemile Lake valley   Riverhurst, Sask   51, 65, Roads   Rockglen, Sask   31, 36, 47, Rose, Bruce   Rose, C   C   Ross, C   C   Ross, C   C   Ross, C   C   Ross, C   Rose,	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34	Talber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Pasin Oils, Ltd. Twelvemile l., Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd. Urban Oils, Ltd. 4, 8, 10 See also Figure 1 Usherville, Sask.	1118-71888
Richfield oil well   Rickert oil well   Rickert oil well   Riding mt., Man   Ries, Heinrich   Rio Grande, Alta   Ritchie, H. C   Rivard   L, Sask   See Twelvemile Lake valley   Riverhurst, Sask   S1, 65, Roads   Roskglen, Sask   31, 36, 47, Rose, Bruce   Roselite oil well   Ross, C. C   Ross, C. C   Ross, C. S   Royalite oil wells   89, Russell, L. S   17, St. Paul, Alta   33, Sand and gravel   See also Interglacial deposits   Moose Jaw area, Sask   See Moose Jaw Sandy ck., Sask   55-57, Saskatchewan   See Interglacial deposits   Moose Jaw   Willowbunch area   Schroeder, J. P.   Scott, J. G.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34	Talber, Alta., oil wells. Tallpines, Sask. Taylor Oil and Gas Company. 24 See also Figure 1 Tehner oil well. Thom, W. T. 32 Thompson, E. A. Thornton, G. W. Thunder ck., Sask. 51, 54, 55, 59 Thurston, J. J. Till. See Boulder clay Toye, S. Triceratop zone. Turner, L. B. Turner valley, Alta. 75, 8 Turner Basin Oils, Ltd. Twelvemile Lake valley, Sask. 34 Twelvemile Lake valley, Sask. 31 Twin Dome Oils, Ltd. Tyrrell, J. B. Unity Valley Oils, Ltd. Unity Valley Oils, Ltd. 4, 8, 10 See also Figure 1 Usherville, Sask.	1118-7188-7188-7188-7188-71894488-90953-738-74999-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-738-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-909-708-
Richfield oil well	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34	Talbines, Sask	111 8-71 88 92 92 92 92 92 92 92 92 92 92 92 92 92
Richfield oil well Rickert oil well Riding mt., Man. Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17, St. Paul, Alta. St. Victor coal seam. 33, Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. 50-57, 60, 61, Sandy Creek, Sask. 50-57, 60, 61, Sandy Creek, Sask. See Interglacial deposits Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Scentinel Oils, Ltd Shell Lake, Sask.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34	Talbines, Sask	111 8-71 88 92 92 92 92 92 92 92 92 92 92 92 92 92
Richfield oil well Rickert oil well Riding mt., Man. Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17, St. Paul, Alta. St. Victor coal seam. 33, Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. 50-57, 60, 61, Sandy Creek, Sask. 50-57, 60, 61, Sandy Creek, Sask. See Interglacial deposits Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Scentinel Oils, Ltd Shell Lake, Sask.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34 93 99 99 99 99 99 99 99 99 99 99 99 99	Taber, Alta., oil wells.  Tallpines, Sask.  Taylor Oil and Gas Company	111 88-71 88 92 92 92 94 94 94 74 92 90 90 94 8-90 90 73 90 73 90 91 91 90 92 92 90 90 92 90 90 90
Richfield oil well Rickert oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads. Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17,  St. Paul, Alta. St. Victor coal seam 33, Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. 56–57, 60, 61, Sandy Creek, Sask. 55–57, Saskatchewan. See Interglacial deposits Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Sentinel Oils, Ltd. Shell Lake, Sask. Sibbald, Alta.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 4 64 64 93 93 93 99 95	Talbries, Sask	111 88-71 88 92 92 86 93 74 81 94 81 90 94 8-90 90 73 90 73 90 91 91 92 92 92 92 92 92 92 92 92 92 92 92 92 92 9
Richfield oil well. Rickert oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. C. Royalite oil wells. Ross, C. S. Royalite oil wells. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. So-57, 60, 61, Sandy Creek, Sask. See Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Sentinel Oils, Ltd. Shell Lake, Sask. Sibbald, Alta. Simpson, Sask. 75.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 4 64 64 93 93 99 99 93 98 95 87	Talbries, Sask	111 88-71 88 92 92 86 93 74 81 94 81 90 94 8-90 90 73 90 73 90 91 91 92 92 92 92 92 92 92 92 92 92 92 92 92 92 9
Richfield oil well. Rickert oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. C. Royalite oil wells. Ross, C. S. Royalite oil wells. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. So-57, 60, 61, Sandy Creek, Sask. See Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Sentinel Oils, Ltd. Shell Lake, Sask. Sibbald, Alta. Simpson, Sask. 75.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 4 64 64 93 93 93 99 95	Talbries, Sask	111 88-71 88 92 92 86 93 74 81 94 81 90 94 8-90 90 73 90 73 90 91 91 92 92 92 92 92 92 92 92 92 92 92 92 92 92 9
Richfield oil well. Rickert oil well. Rickert oil well. Riding mt., Man. Ries, Heinrich. Rio Grande, Alta. Ritchie, H. C. Rivard 1., Sask. See Twelvemile Lake valley Riverhurst, Sask. Rockglen, Sask. Rockglen, Sask. Rose, Bruce. Roselite oil well. Ross, C. C. Ross, C. C. Royalite oil wells. Ross, C. S. Royalite oil wells. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. St. Victor coal seam. Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy ck., Sask. So-57, 60, 61, Sandy Creek, Sask. See Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Sentinel Oils, Ltd. Shell Lake, Sask. Sibbald, Alta. Simpson, Sask. 75.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 34 93 93 93 95 87 93	Taber, Alta., oil wells.  Tallpines, Sask.  Taylor Oil and Gas Company. 24  See also Figure 1  Tehner oil well.  Thom, W. T.  Thompson, E. A.  Thornton, G. W.  Thunder ck., Sask. 51, 54, 55, 59  Thurston, J. J.  Till. See Boulder clay  Toye, S.  Triceratop zone.  Turner, L. B.  Turner valley, Alta. 75, 8  Turner Basin Oils, Ltd.  Twelvemile Lake valley, Sask. 31  Twin Dome Oils, Ltd.  Tyrrell, J. B.  Unity Valley Oils, Ltd.  Urban Oils, Ltd. 4, 8, 10  See also Figure 1  Usherville, Sask. 86  Vanalta oil wells. 86  Vanalta Oils, Ltd. 9  Vanalta Oils, Ltd. 10, 12, 13, 21  Vanalta Oils, Ltd. 2  See also Figure 1  Vanalta Oils, Ltd. 2  See also Figure 1  Vanalta Oils, Ltd. 32  Vanalta Oile, Alta. 34  Vanalta Oile, Alta. 36  Vandleyerigis coulée, Alta. 36	111 88-71 88 92 92 86 93 74 81 94 94 74 32 94 88-90 90 73 90 73 90 73 91 92 92 94 94 90 73 90 73 91 92 94 91 94 94
Richfield oil well Rickert oil well Rickert oil well Rickert oil well Riding mt., Man Ries, Heinrich Rio Grande, Alta Ritchie, H. C. Rivard I., Sask. See Twelvemile Lake valley Riverhurst, Sask. 51, 65, Roads Rockglen, Sask. 31, 36, 47, Rose, Bruce. Roselite oil well Ross, C. C. Ross, C. C. Ross, C. S. Royalite oil wells. 89, Russell, L. S. 17,  St. Paul, Alta. St. Victor coal seam 33, Sand and gravel See also Interglacial deposits Moose Jaw area, Sask. See Moose Jaw Sandy Creek, Sask. 50–57, 60, 61, Sandy Creek, Sask. 55–57, Saskatchewan. See Interglacial deposits Moose Jaw Willowbunch area Schroeder, J. P. Scott, J. G. Sentinel Oils, Ltd. Shell Lake, Sask. Sibbald, Alta.	87 73 32 97 51 75 2 48 32 91 2 35 90 31 96 4 64 64 93 93 99 99 93 98 95 87	Talbries, Sask	88 92 94 84 90 73 92 92 92 92 92 92 92 92 92 92 92 92 92

## 103в

PAGE	PAGE
Verwood, Sask	
Viking, Alta 91	
Virgelle formation	
Vogelberg coal seam	5 Wickenden, R. T. D
	Repts. by, on:
Wainwright, Alta91, 95	
Wainwright Petroleum, Ltd 91	
Waite valley, Alta 91	
Warner Oils, Ltd 84	
Warren, P. S 98	
Water	Wildcat hills, Alta 91
See also Alkali lakes	Williams, M. Y
Springs	Willowbunch, Sask34, 43
Moose Jaw dist., Sask., rept. on ground 50-64	
Regina, Sask	
Waterhole, Alta 96	
Wellington oil well 90	
Wells, logs of	Willowbunch I., Sask31, 35, 36, 40, 41
Milk River area and Red Coulée oil	See also Lake of the Rivers valley
field	
Ribstone field, Alta 77–80	
Willowbunch area, Sask 27-49	Wood Mountain, Sask
Winnipeg Lake area, Man 77	
Western Alberta oil well 84	
Wetherelt, J	
Weymarn Petroleum, Ltd 86	
White, W. W 98	5 Zénon Park, Sask 95



The Annual Summary Report of the Geological Survey is issued in parts, referring to particular subjects or districts. This year there are four parts, A, B, C, and D. A review of the work of the Geological Survey for the year forms part of the Annual Report of the Department of Mines.