Abstract

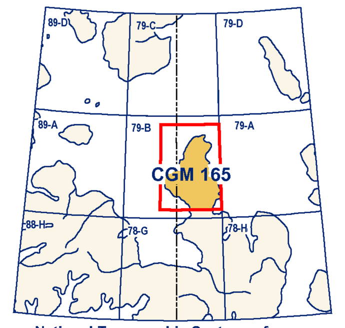
Sabine Peninsula of Melville Island was the subject of an oil and gas exploration boom from 1961 to 1985, during which time seismic-reflection data were collected and wells were drilled. As a result, the two largest conventional natural gas fields in Canada were discovered.

Seismic-reflection methods use sound waves to image the internal structure of the Earth. Waves are emitted at the surface before being reflected back to the surface by geological interfaces and recorded. Modern analysis methods were used to reinvestigate existing seismic data. In doing so, eight seismic unit boundaries identified on seismic profiles in two-way traveltime were correlated to the regional geological framework and gridded to provide subsurface maps. Each map approximates the structures preserved at that particular time or depth allowing the enhancement of the geological knowledge of Sabine Peninsula and better delimitation of elements of the petroleum systems therein.

Résumé

La péninsule de Sabine de l'île de Melville a connu un boum d'exploration gazière et pétrolière entre 1961-1985 pendant lequel des données de sismique-réflexion furent acquises et des puits forés. Il en résultat la découverte des deux plus grands champs de gaz naturel conventionnels du Canada.

La sismique-réflexion utilise des ondes sonores pour imager la structure interne de la Terre. Les ondes sont émises en surface avant d'être réfléchies de nouveau vers la surface par des interfaces géologiques où elles sont enregistrées. Des méthodes d'analyse modernes furent utilisées pour ré-investiguer des données sismiques existantes. Ainsi, huit limites d'unités sismiques identifiées sur les profils sismiques en temps de parcours aller-retour furent corrélées au cadre géologique régional et maillées afin de produire des cartes de la sous-surface. Chaque carte est une approximation des structures préservées à un certain temps ou une certaine profondeur nous permettant d'améliorer les connaissances géologiques de la péninsule de Sabine et de mieux délimiter les éléments des systèmes pétroliers s'y trouvant.



National Topographic System reference

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Permian sandstone hoodoos, Sabine Peninsula, Melville Island, Nunavut. Photograph by T.A. Brent.
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CANADIAN GEOSCIENCE MAP 165

TIME- AND DEPTH-STRUCTURE MAP

GROSVENOR ISLAND FORMATION

Sabine Peninsula, Melville Island
Nunavut–Northwest Territories

1:200 000

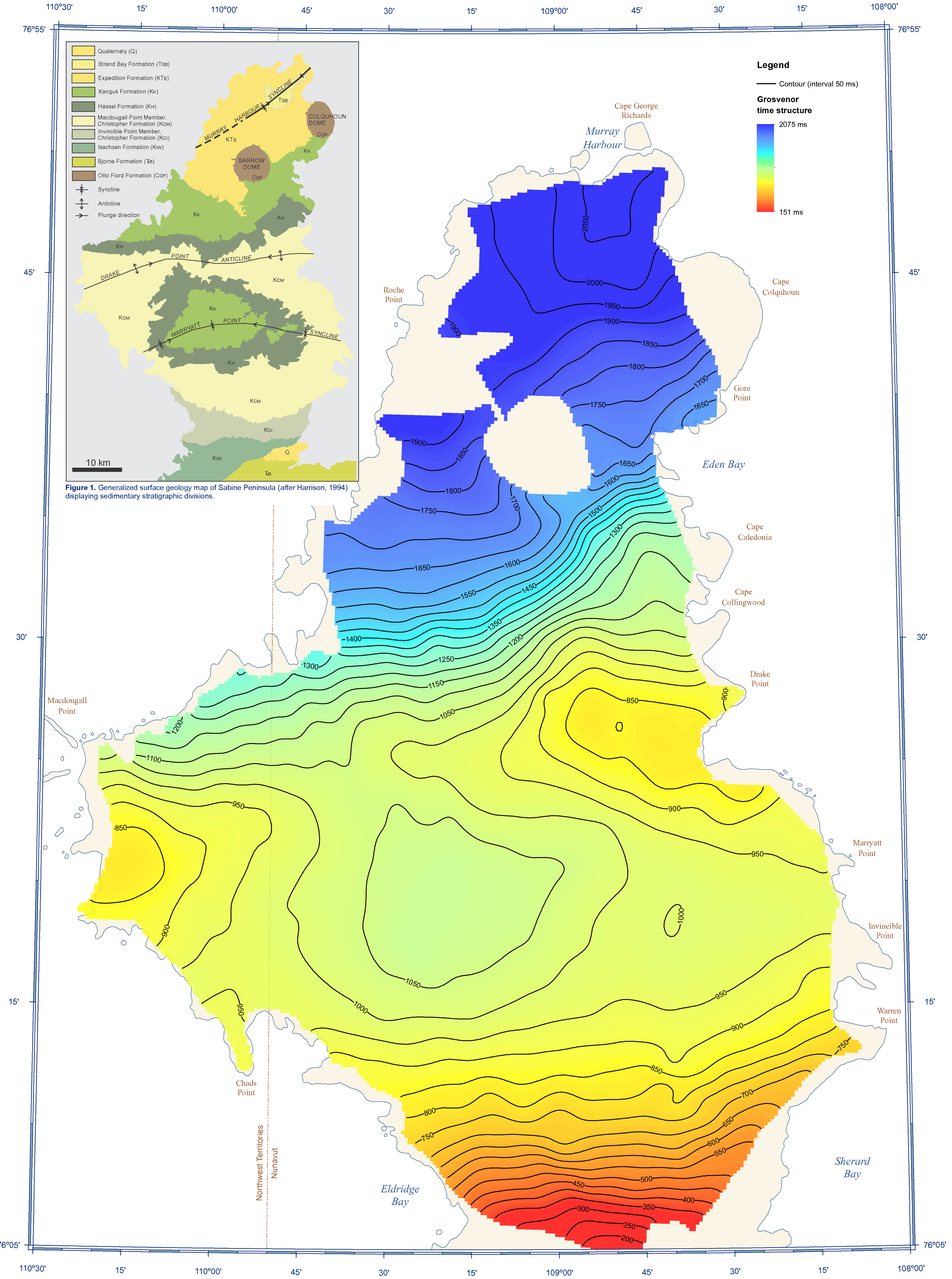

Figure 1. Generalized surface geology map of Sabine Peninsula (*after* Harrison, 1994) displaying sedimentary stratigraphic divisions.

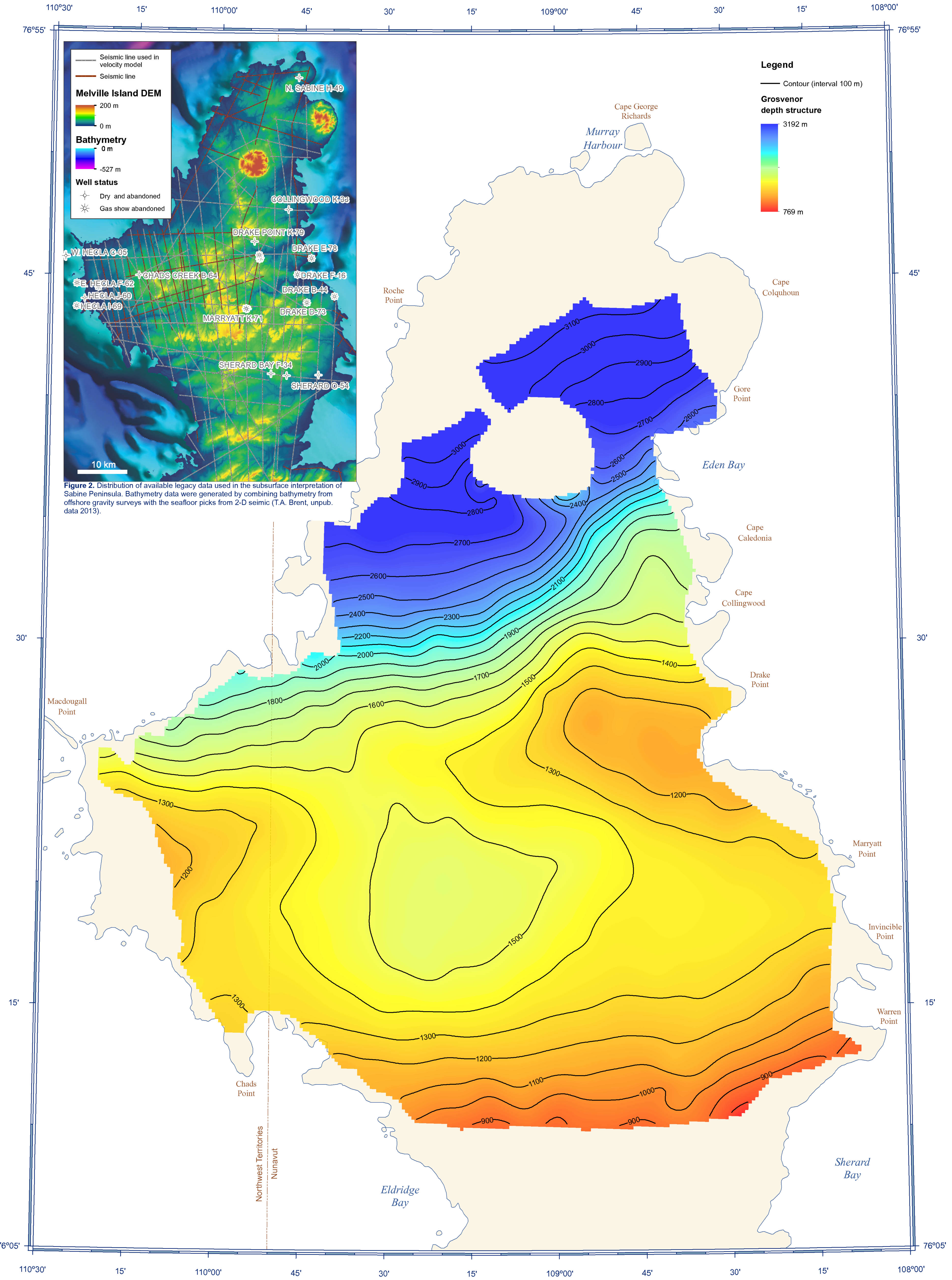

Figure 2. Distribution of available legacy data used in the subsurface interpretation of Sabine Peninsula. Bathymetry data were generated by combining bathymetry from offshore gravity surveys with the seafloor picks from 2-D seimic (T.A. Brent, unpub. data 2013).

DESCRIPTIVE NOTES

INTRODUCTION

The time- and depth-structure maps presented herein are part of an eight-map series of the subsurface of Sabine Peninsula spanning the Early Permian through Early Cretaceous interval.

These maps are the product of the application of modern geoscientific methods of processing and interpretation to a suite of legacy seismic-reflection data from onshore Sabine Peninsula (Melville Island, Western Arctic Islands). The resultant processed seismic lines were interpreted using the existing regional geological framework (*see* Harrison, 1995) by integrating existing regional well data, geophysical logs, age control, and lithological information through synthetic seismograms.

REGIONAL SETTING

The Sabine Peninsula of Melville Island is located within the Sverdrup Basin in the Queen Elizabeth Islands of the western Arctic. The Sverdrup Basin extends for about 1300 km in a northeast-southwest direction and is up to 350 km wide. The basin contains up to 13 km of sedimentary strata (Embry and Beauchamp, 2008). The Sverdrup Basin is separated from the underlying Franklinian Basin by an unconformity at the base of the Carbonifreous strata. The Franklinian Basin was superseded by widespread rifting following the Late Devonian–earliest Carboniferous Ellesmerian Orogeny. The resulting rift-related structural depression acted as a major depocentre from the Carboniferous through the Paleogene (Embry and Beauchamp, 2008). The Sverdrup Basin succession was uplifted and deformed during the early Cenozoic Eurekan Orogeny.

The surface geology of Melville Island is dominated by Lower Paleozoic strata of the Franklinian Basin. The Sabine Peninsula is an exception to this, as surface strata are part of the Sverdrup Basin. The geology of the Sabine Peninsula consists of deformed Late Carboniferous to Paleocene sandstone, siltstone, shale, and minor amounts of carbonate. Additionally, evaporitic rocks are exposed in two diapirs on northern Sabine Peninsula — the Barrow and Colquhoun domes, which consist of deformed anhydrite and gypsum. The strata of the Sverdrup Basin succession on Melville Island were deformed into a series of folds, including the Murray Harbour syncline in the northern part of the peninsula and the Drake Point anticline and the Marryatt Point syncline to the south (Harrison, 1994) (Fig. 1).

During a 1961 to 1985 phase of petroleum exploration, companies drilled 52 wells on Melville Island and surrounding waters (22 of which were on Sabine Peninsula) and acquired about 3,400 line-kilometres of onshore seismic-reflection data (Fig. 2).

Three separate gas fields were discovered in the Sabine Peninsula area: Drake Point, Hecla, and Roche Point. Feasibility studies for the development of the gas fields were conducted in the early 1980s; however, due to low gas prices and the lack of gas markets, the gas fields on Melville Island (and elsewhere in the Canadian Arctic) were not developed (Harrison, 1995).

SEISMIC DATA SET AND PROCESSING

Data access was obtained through a Memorandum of Understanding signed in 1997 by the Geological Survey of Canada (GSC), Panarctic Oils, the Arctic Islands Exploration Group, and the Offshore Arctic Exploration Group joint-venture parties. The data sets consist of original land seismic-reflection field tapes transcribed from 21-, 7-, and 9-track media. Data were collected using a dynamite charge of 20–30 kg per shot at about 20 m below the surface. Shot-point spacing ranged from 67 m to 300 m, the shorter spacing being used for most surveys. The majority of the seismic-reflection data were recorded using 48- or 96-channel systems. Channel stations were generally deployed using nine receivers spaced at about 8 m and station intervals varying from 50 m to 70 m. The common-midpoint multiplicity of the data sets range from single to 12-fold coverage. The most common recording length was 6 s.

The processing consisted of three main steps: 1) principal component decomposition was used to remove both coherent and random noise, 2) data were migrated utilizing poststack Kirchhoff migration, and 3) seismic bandwidths were extended to increase vertical resolution (Claprood et al., 2011; Duchesne et al., 2012).

Velocity model

A 3-D velocity model was built using about 1300 km of linear seismic data (78 lines) and 13 wells spread over an area of about 2800 km2 (Fig. 2). The velocity model was then used for poststack migration processing and to convert seismic horizon surfaces from time to depth. The primary assumption behind the velocity model is that the coherent high-amplitude reflections that were picked to build the model correspond to important acoustic impedance contrasts caused by significant and abrupt velocity changes. This assumption was confirmed by tying seismic picks to well sonic logs (Duchesne et al., 2012). The geostatistical approach of kriging with an external drift (KED) was applied to both the reflection time of the picked seismic horizons and time-depth pairs derived from check shot data to compute the 3-D velocity field. Kriging interpolates values between the known positions based on weighted spatial correlations. The KED technique was specifically developed for the integration of seismic data into the kriging process where the number of wells is insufficient for the computation of adequate depth statistics (Hass and Dubrule, 1994). Hence, it uses the information provided by the time horizon picks to improve estimates where depth control is sparse. For seismic migration, root mean squared (RMS) velocity values are first estimated by KED from time-to-depth conversion of seismic horizon surfaces mapped as important velocity boundaries (Duchesne et al., 2012). Then, once the approximate depths of the surfaces are known, the interval velocities (V_{int}) for all time intervals delimited by two consecutive horizons is computed from:

$$V_{int} = \frac{\Delta z_i}{\Delta t_i}$$

where z and t are the depth and time intervals between two successive horizons i. Once V_{int} is obtained the RMS velocity (V_{rms}) is calculated using:

$$V_{rms} = \sqrt{\frac{1}{t_0}\sum_{i=1}^{N} V_{int}^2 \Delta t_i}$$

in which N is the total number of horizons and t_0 is the sum of all time intervals.

SEISMIC INTERPRETATION AND VISUALIZATION METHODS

Processed seismic lines were loaded into IHS-Kingdom® seismic and geological interpretation software. Prominent seismic-reflection horizons, tied to well formation-top information, were manually correlated. Seed points were generated at seismic line intersections, thereby permitting the interpretation of adjacent lines.

The map would benefit from a detailed structural interpretation; however, confidence of this interpretation is minimized due to minor vertical offsets (about 0.1 s) attributed to faulting and the large line spacing. Thus reflections are readily identified across faults despite offset.

Time-structure maps of the key seismic horizons were computed using universal kriging. Universal kriging permits the interpolation of a nonstationary, random field by adding a term in the kriging equation that accommodates any linear trends present in a scattered point set (Chilès and Delfiner, 1999). Given that all picked horizons showed a strong linear trend for time versus depth over distance, universal kriging provided the best fit to the picked horizons.

TIME TO DEPTH CONVERSION

All time surfaces are converted to depth using the following procedure. First V_{int} of the 3-D velocity model are calculated using Dix equation:

$$V_{int} = \left[\frac{V_{n_{rms}}^2 t_n - V_{n-1_{rms}}^2 t_{n-1_{rms}}}{t_n - t_{n-1}}\right]^{1/2}$$

where t is the zero-offset arrival time of the nth reflection. Interval limits corresponded to seismic horizons that are picked and tied to geological interfaces. Then V_{int} are extracted from the velocity model along picked horizons. Velocity maps are then computed using Universal kriging at a cell size of 250 m. Finally, the time-structure surfaces of the various seismic horizons are converted to depth (Z) using:

$$Z = \frac{V_{int}}{2} t$$

Because the depth-conversion process is a function of the velocity model, the lateral extent of depth maps is confined to the lateral extent of the model. The final depth-structure maps were imported into ArcGIS for visualization using the Arc extension Team-GIS KBridge.

UNCERTAINTY

Quantifying the uncertainty of seismic subsurface maps is difficult since several sources of data, each with their unique level of uncertainty, are used in the map generation. Sources of error may arise from limitations in acquisition, processing, and interpretation. Moreover, seismic data are collected remotely and the images they provide are derived from generalized mathematical and physical concepts. Constraints in acquisition that increase the uncertainty include gaps in coverage because of obstacles to source and receiver deployment, and effect of direction of shooting on data quality (Sheriff and Geldart, 1995). Processing errors may result from inadequate static corrections, inaccurate velocity analysis, and inappropriate parameter determination.

More specifically to this data set, errors may have also been introduced by the velocity model and the ability to tie formation tops to seismic horizons. The velocity model represents an estimation of the velocity fluctuations for which the accuracy depends on the number of wells and the good fit between time picks and corresponding depths at the well locations. A regression analysis shows that time picks and their corresponding depths at the wells have a strong linearity (r^2 = 0.98), meaning that the use of time picks as the external drift in the kriging strategy is justified and trustworthy. Nevertheless, the uncertainty of the velocity model increases when the distance between the well and any points where velocity is predicted exceed the range of the variogram expressing the spatial dependence between depth and time. In the present case, the range of the different horizons is between 9.5 km and 34 km. The ability to tie formation tops to seismic horizons relies on the successful use of well sonic and density logs, since it is the contrast between the product of these properties for two successive geological layers that generates reflections recorded in seismic exploration. Formation tops used in this study are from Dewing and Embry (2007), for which they mainly utilized gamma-ray logs to position the upper limit of the formations in depth. Thus errors may have been introduced by projecting the formation tops on seismic sections recorded in time.

TIME- AND DEPTH-STRUCTURE DATA DISPLAY

The time- and depth-structure data shown on this map were gridded at a cell size of 250 m using Universal kriging. Each map presents a grid with a stretched colour ramp at 20% transparency. Time contours generated from the time-structure grids are shown in black at a 50 ms interval, whereas depth contours derived from the depth-structure grid are presented at 100 m intervals.

GROSVENOR ISLAND MAP DESCRIPTIONS

The Early Jurassic Grosvenor Island Formation (possibly of the Heiberg Group) consists of shale, chert, carbonate, and olistostromes (Dewing and Embry, 2007, *see also* Fig. 3). According to Embry (1983a), the Heiberg Group is widespread, but variable across the Sverdrup Basin. In the eastern and central areas of the basin the group is divisible into five units: Skybattle, Grosvenor Island, Maclean Strait, Lougheed, and King Christian (Embry, 1983a). In contrast, the Heiberg Group is elsewhere divided into three members: Romulus, Fosheim, and Remus (Embry, 1983b). Available formation-top data indicate that the Grosvenor Island Formation overlies the shale units of the Barrow Formation or occasionally those of the Hoyle Bay Formation. The upper contact of the Grosvenor Island Formation is typically with the shale and siltstone of the Lougheed Island Formation or less often, the sandstone of the Maclean Strait Formation (Dewing and Embry, 2007).

The mapped Grosvenor Island Formation reflection extends from the narrowest point of the peninsula near Eldridge and Sherard bays to the northern limit of the peninsula. The data gap west of Eden Bay marks the location of Barrow Dome. Two-way traveltimes of the Grosvenor Island reflection increase northward from 151 ms to 2075 ms, or from 769 m to 3192 m. The average slope of the Grosvenor Island Formation horizon is generally between 0.5° and 2.5° although steeper slopes up to 10° are observed north of the Drake Point anticline and in the southern part of the peninsula. A large depression is observed at the centre of Sabine Peninsula between the Hecla and Drake gas fields. This feature is bounded by the Drake Point anticline to the north and the Marryatt Point syncline to the south (Harrison, 1994).

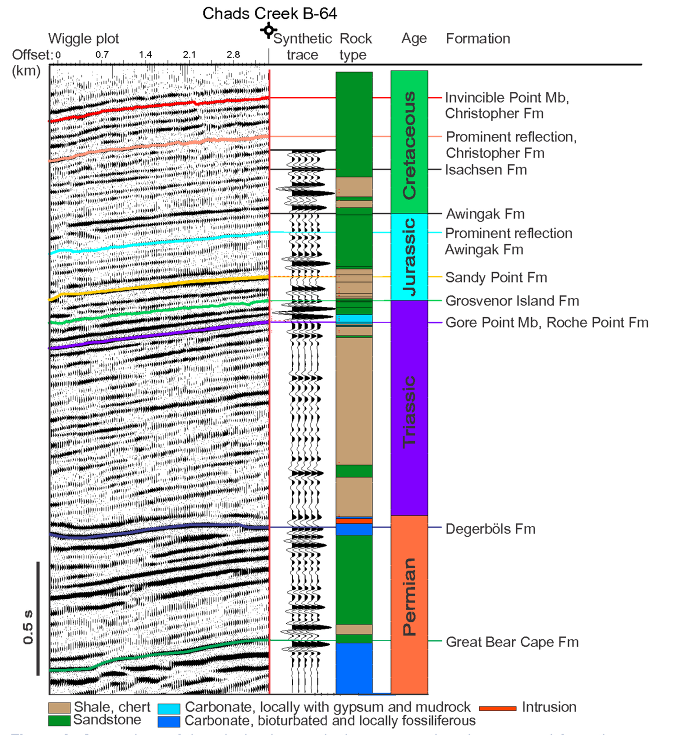

Figure 3. Comparison of the wiggle plot, synthetic trace, stratigraphy, age, and formation-top data for the Chads Creek B-64 well.

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CANADIAN GEOSCIENCE MAP 165

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Canadian Geoscience Maps

Authors: M.J. Duchesne, V.I. Brake, K. Dewing, M. Claprood, E. Gloaguen, and T.A. Brent

Time-structure map by V.I. Brake and M.J. Duchesne, Geological Survey of Canada, 2013

Depth-structure map by M.J. Duchesne and V.I. Brake, Geological Survey of Canada, 2013

Seismic interpretation by V.I. Brake and M.J. Duchesne, Geological Survey of Canada, 2010–2013

Geomatics by V.I. Brake, Geological Survey of Canada and G. Huot-Vézina, Institut national de la recherche scientifique

Cartography by R. Boivin

Scientific editing by E. Inglis

Initiative of the Geological Survey of Canada, conducted under the auspices of the Western Arctic Islands' project as part of Natural Resources Canada's Geo-mapping for Energy and Minerals (GEM) program.

TIME- AND DEPTH-STRUCTURE MAP
GROSVENOR ISLAND FORMATION
Sabine Peninsula, Melville Island
Nunavut–Northwest Territories
1:200 000

5 0 5 10 15 20 km

Map projection Universal Transverse Mercator, zone 12
North American Datum 1983

Base map at the scale of 1:250 000 from Natural Resource Canada, with modifications.

Proximity to the North Magnetic Pole causes the magnetic compass to be useless in this area.

The Geological Survey of Canada welcomes corrections or additional information from users.

The data may include additional observations not portrayed on this map. *See* documentation accompanying the data.

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This map is not to be used for navigational purposes.

Canada

CANADIAN GEOSCIENCE MAP 165
TIME- AND DEPTH-STRUCTURE MAP
GROSVENOR ISLAND FORMATION
Sabine Peninsula, Melville Island
Nunavut–Northwest Territories