## INTRODUCTION

This Surficial Geology Map of NTS 94-O/10 (Canadian Geoscience Map 126) is the product of collaboration between the Geological Survey of Canada and the British Columbia Ministry of Energy, Mines and Natural Gas as part of the Geo-mapping for Energy and Minerals Program GEM-Energy Yukon Basins Project. The accompanying geodatabase includes field observation points and field photos, landform features as lines, and sufficial geology unit polygons. The map and geodatabase are essential baseline geoscience information for a range of potential end-users including resource explorationists, geotechnical engineers, land-use managers, terrestrial ecologists, archaeologists, geoscientists and communities in northern British Columbia. By providing new insight into the distribution and origins of surficial earth materials, CGM 126 will help to reduce the economic costs and risks associated with the sustainable development of energy and mineral resources in NTS 94-O/10. Environmental impact assessments for new access roads, work camps, well pads, pipeline and power transmission line corridors, water storage and waste management systems and other infrastructure will benefit from the geoscience information presented here. By identifying areas prone to geological hazards (e.g., landslides, permafrost, flooding), CGM 126 will also help to protect natural resources, infrastructure and communities vulnerable to climate change in Canada's north.

DESCRIPTIVE NOTES

APPROACHTO SURFICIAL GEOLOGY MAPPING Terrain mapping and field-based benchmarking studies have led to a better understanding of the regional distribution of surficial deposits. permafrost, landslides and other geomorphic processes in the NTS 94-O/10 map area (Huntley and Hickin, 2010; Huntley et al., 2011a-b). Surficial earth materials and landforms were interpreted using a combination of stereo-pair air photos (BCB97010, 15BCB97015, 15BCB97029, 15BCB97075 and 15BCB97088 series), LANDSAT 7 satellite imagery (http://glovis.usgs.gov/ [URL 2011]) and Shuttle Radar Topography Mission digital elevation models (http://dds.cr.usgs.gov/srtm/ [URL 2011]). The base map was generated from CANVEC shape files (http://geogratis.cgdi.gc.ca/geogratis/ [URL 2011]). Surficial geology polygons and landform line symbols were digitized using commercially available computer software packages (Global Mapper, ArcMap and ArcGIS) and compared to published maps, reports and archived digital data (e.g., Stott and Taylor, 1968; Bednarski, 2003a-d; Clement et al., 2004; Bednarski, 2005a-b). The geodatabase accompanying this map conforms to the Science Language for the Data Management component of the GEM Geological Map Flow process (cf. Huntley and Sidwell, 2010; Huntley et al., 2011a; Deblonde et al., 2012).

 $Field work was undertaken in 2009 \, and \, 2010 \, to \, ground \, truth \, surficial \, geology \, polygons \, interpreted \, from \, air \, photos \, and \, satellite \, imagery, \, and \, to \, property \, and \,$ gather characteristics that could not be determined through remote predictive mapping. Earth materials were defined on the basis of facies and landform associations, texture, sorting, colour, sedimentary structures, degree of consolidation, and stratigraphic contact relation field stations and remote observations from helicopters. The distribution of glacial and non-glacial landforms is depicted on the surficial geology map. Map units in the Legend are presented chronostratigraphically and include organic deposits, alluvial, colluvial, eolian, glaciolacustrine and glaciofluvial sediments, tills and areas of bedrock. INFERRED GEOLOGICAL HISTORY

The distinctive landscape of NTS 94-O/10 is largely a product of underlying bedrock and geological structures, with ornamentation by the Late Wisconsinan Laurentide Ice Sheet. Conglomerate, sandstone and carbonaceous shale of the Upper Cretaceous Dunvegan Formation underlie the southeast part of the map area, where the Etsho Plateau reaches its highest elevation. Elsewhere, exposures in borrow pits, and creek and river sections indicate that undifferentiated clastic bedrock (Lower Cretaceous Fort St. John Group) underlies the map area (Stott

Topography and drainage patterns were greatly modified during the phase of maximum ice cover (>18 <sup>4</sup>C ka BP or >21.4 calendar ka BP). olidated sediment thicknesses in excess of 2-5 m are observed in major valleys and it is suspected that similar drift thicknesses blanket bedrock across the map area. Laurentide tills have low clast contents (<20%) of proximally derived Cretaceous siliciclastic sedimentary rocks and distal exotic igneous and metamorphic clasts from the Canadian Shield exposed hundreds of kilometres to the northeast. Drumlins up to several kilometres in length, together with eskers and subglacial mellwater channels confined to the escarpment crest of the Etsho Plateau, imply tills were deposited beneath active, rapidly flowing warm-based glacial ice (Huntley and Hickin, 2010; Huntley et al., 2011b). Small lake basins were excavated by erosion and ice-thrusting as Laurentide ice and subglacial meltwater scoured and deformed older glacial deposits and weak bedrock, Multiple generations of ice flow are preserved in streamlined landforms across the map area (unit Ts). Southwest oriented drumlins are interpreted as a late glacial maximum ice flow direction across the map area. The distribution of younger drumlins suggests converging ice lobes flowed west-northwest from the Fort Nelson Lowland and west to southwest over the Etsho Plateau.

Deglaciation began sometime after 18 <sup>14</sup>C ka BP (or >21.4 calendar ka BP) and ended before 10 <sup>14</sup>C ka BP (ca. 12 calendar ka BP), with the retreating active Laurentide Ice Sheet, stagnant ice masses in lowlands, glaciofluvial outwash and landslide debris blocking and reordering the regional drainage system. The mapped distribution of major moraine ridges (unit Tm) implies that ice in the Fort Nelson Lowland retreated to the southeast. Ice over the Etsho Plateau retreated to the northeast into the Petitot River valley (Huntley and Hickin, 2010). Some large end moraines are deformed and streamlined suggesting that receding lobes remained active during retreat and occasionally rapidly advanced. Minor moraine ridges drape drumlins in cross-cutting patterns and are interpreted as crevasse fillings and squeeze moraines deposited shortly after drumlinization ended, or as ice retreated from the map area (Huntley et al., 2011b). Hummocky till (unit Th) is found in proximity to short segments of subareal-subglacial meltwater channels and eskers, indicating that bodies of stagnant glacier ice remained in lowland areas west of the Maxhamish Escarpment. Eskers (unit GFr) are composed of hummocky till and glaciofluvial gravelly sand, and likely exploited pre-existing crevasse patterns beneath the retreating ice sheet or stagnant ice bodies (Huntley et al., 2011a; b).

Post-placiation (10 <sup>14</sup>C ka BP, or ca. 12 calendar ka BP to present), changes in regional base-level led to episodes of channel incision and aggradation. In the early Holocene, pulses of fluvial terrace building followed initial valley incision by the Liard and other major rivers. Most streams and rivers have alluvial fans (unit Af) and terraces (unit At) < 5 m above active floodplains (unit Ap) consisting of gravel overlain by silt and sand. Poorly drained clay-rich till on the plateaux and glaciolacustrine sediments in lowland areas are covered by extensive postglacial peat deposits (unit Owb), fens (unit Owf) and undifferentiated wetlands (unit O). Discontinuous permafrost is sporadically encountered in glaciolacustrine and some peat deposits. Charcoal, observed in dug pits on alluvial terraces, suggest forest fires may have contributed to periods of landslide activity on slopes and local fluvial aggradation. Colluviated deposits (unit Cv) are restricted to slopes where underlain by shale and fine-grained glacial deposits. Stream networks and wetlands draining plateau watersheds are disrupted by beaver activity and, to a lesser extent, by roads and infrastructure where they cross streams, rivers and organic deposits (Huntley and Hickin, 2010; Huntley and

Canadian Geoscience Map 126 depicts the surficial geology over some 790 km<sup>2</sup> covered by the Tightfit Lake map sheet (NTS 94-O/10) in northeastern British Columbia. The map area lies at the western limit of the Etsho Plateau and northwest limit of the Fort Nelson Lowland. This boundary is a major drainage divide: Klenteh Creek and other tributaries to Kiwigana and Fort Nelson rivers drain the south-facing escarpment of the Etsho Plateau and Fort Nelson Lowland; to the north, the plateau and Tightfit Lake drain north into the Petitot River valley. Bedrock is mantled by unconsolidated earth materials that dates to the Late Pleistocene (Late Wisconsinan Glaciation, > 25 ka to ca. 10 ka) and non-glacial Holocene (ca. 10 ka to present). Deposits of till, green on the map, are generally suitable for placement of infrastructure. Glaciofluvial and eolian deposits with mineral, aggregate, and groundwater potential are coloured orange and buff. Slopes disturbed by landslides, debris

flows, and rock falls appear brown and pink.

Glaciolacustrine and organic deposits with sporadically

d'environ 790 km<sup>2</sup> couvert par le feuillet cartographique de Tightfit Lake (SNRC 94-O/10), dans le nord-est de la Colombie-Britannique. La région cartographique se situe à la bordure occidentale du plateau d'Etsho et à la bordure nord-ouest des basses terres de Fort Nelson. Cette limite correspond à une importante ligne de partage des eaux : le ruisseau Klenteh et d'autres affluents des rivières Kiwigana et Fort Nelson drainent l'escarpement à regard sud du plateau d'Etsho et les basses terres de Fort Nelson; au nord, le plateau et le lac Tightfit se drainent vers le nord dans la vallée de la rivière Petitot. Le socle rocheux est couvert de matériaux terrestres non consolidés remontant au Pléistocène supérieur (Glaciation du Wisconsinien supérieur, de > 25 ka à env. 10 ka) ainsi que de matériaux non glaciaires de l'Holocène (d'env. 10 ka jusqu'à nos jours). Les dépôts de till, de couleur verte sur la carte, sont généralement propices à l'établissement de l'infrastructure. Les dépôts discontinuous permafrost are coloured purple and grey. fluvioglaciaires et éoliens, qui recèlent un potentiel en Alluvial deposits prone to flooding, erosion, and versants dérangés par des glissements de terrain, des coulés de débris et des chutes de blocs sont représentés en brun et en rose. Les dépôts glaciolacustres et organiques, qui renferment sporadiquement du pergélisol discontinu, sont

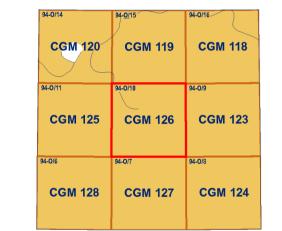
représentés en violet et en gris. Les dépôts alluviaux

sujets aux inondations, à l'érosion et à la sédimentation

apparaissent en jaune sur la carte.

La Carte géoscientifique du Canada 126 illustre la

géologie des matériaux superficiels d'un territoire



National Topographic System reference and index to adjoining published Geological Survey of Canada maps

Cover illustration Drumlin ridge on the Etsho Plateau in northeast British Columbia, view west in the direction of Photograph by D.H. Huntley. 2013-104

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Natural Resources Ressources naturelles Canada du Canada

**CANADIAN GEOSCIENCE MAP 126 SURFICIAL GEOLOGY** 

**TIGHTFIT LAKE** British Columbia 1:50 000



## **CANADIAN GEOSCIENCE MAP 126**

Holocene earth materials and landforms **Organic Deposits** Peat bogs: fibric to humic organic matter; massive to stratified accumulations; generally greater than 2 m thick; confined to topographic depressions or level areas: underlain by poorly drained till, glaciolacustrine and other unconsolidated sediments; formed by decomposition of plant material in wetland areas; bogs with sporadically discontinuous permafrost and thermokarst depressions potentially unstable if organic material is disturbed or Fens: fibric organic matter; massive to stratified; generally greater than 2 m thick; confined to topographic depressions, level areas and meltwater channels; underlain by poorly drained till, glaciolacustrine and other unconsolidated sediments; formed by decomposition of plant material in wetland areas; fens are prone to flooding following drainage damming by Undifferentiated peat bogs and fens: humic to fibric organic matter; massive to stratified accumulations; generally greater than 2 m thick; confined to topographic depressions, level areas or channels; underlain by poorly drained till, glaciolacustrine and other unconsolidated sediments; formed by decomposition of plant material in wetland areas; may contain sporadically discontinuous permafrost and thermokarst depressions; potentially unstable if disturbed or removed during development. Alluvial deposits Alluvial terraced sediments: boulders, gravel, sand and silt; generally massive to planar stratified; well to rapidly drained; greater than 2 m thick; may contain interbedded debris flows and buried organic material; underlain by outwash, till or bedrock; transported and deposited by modern rivers, streams and creeks; subject to rare flooding; potential source of aggregate; land use activities may adversely affect stream courses and conditions, and impact fish and wildlife resources. Alluvial floodplain sediments: gravel, sand and silt; massive, trough cross-bedded, rippled-bedded, planar stratified; well to rapidly drained; greater than 2 m thick; underlain by till or bedrock; transported and deposited by modern rivers, streams and creeks; subject to seasonal flooding; land use activities may adversely affect stream courses and conditions, and impact fish Colluvial deposits Colluvial veneer: clast-supported diamictons and rubble; massive to stratified, poorly-sorted; well to rapidly drained; deposits less than 2 m thick; landslide headscarps range from 300 m to 10.5 km; formed by the weathering and down-slope movement of earth materials by gravitational processes; bedrock and unconsolidated debris on slopes above 10-15° with greater than 5 m relief prone to mass-wasting; rock falls, topples, rock slides and debris flows occur where shale, sandstone and carbonate strata is exposed close to the surface; retrogressive rotational debris slides, debris flows and slumps occur in glaciolacustrine sediments and outwash containing sporadically discontinuous permafrost; where ground ice is found slope failure can occur on surfaces less than 5°; slope instability could present major problems for construction in Late Pleistocene earth materials and landforms Glaciofluvial deposits Kames and hummocky outwash: boulders, cobbles, pebble-gravel, sand, silt and diamicton; generally massive to stratified, some slump structures; moderately to well-drained; greater than 2 m thick; irregular hummocks and kettled topography; in contact with, and overlying till units, outwash and glaciolacustrine sediments; deposited by rivers and streams flowing from, or in contact with glacial ice; potential source of groundwater and granular aggregate when material is gravel rich. Esker ridges: boulders, cobbles, pebble-gravel, sand, silt and matrix-supported diamicton; generally massive to stratified, some slump structures; moderately to well-drained; greater than 2 m thick; range from 100 m to 8.8 km in length; in contact with, and overlying till units, outwash and glaciolacustrine sediments; deposited by subglacial meltwater in contact with glacial ice; potential source of groundwater and granular aggregate when material is gravel rich. Till blanket: sand, silt and clay-rich diamictons; massive, matrix-supported and compact; clast contents less than 20% and contain sub-rounded granitic erratic boulders with sources on the Canadian Shield; moderately to well-drained; greater than 2 m thick mantling bedrock and older glacial deposits; transported and deposited by the Laurentide Ice Sheet directly through lodgement, basal meltout, glacigenic deformation of sediment beneath active, warm-based ice and in situ melting from stagnant cold-based ice; stable terrain, generally suitable for infrastructure placement. Hummocky till: sand and silt-rich diamictons; massive to stratified, matrixand clast-supported; clast contents less than 20% and contain sub-rounded granitic erratic boulders with sources on the Canadian Shield: moderately to well-drained; greater than 2 m thick; drapes till and other glacial deposits; deposited by in situ melting from stagnant cold-based ice and modified by meltwater; evidence for ice collapse includes slump structures, kettle lakes and irregular topography; potential source of aggregate when material is gravel rich; generally suitable for infrastructure placement. **loraine ridges**: sand, silt and clay-rich diamictons; massiv matrix-supported; clast contents less than 20% and contain sub-rounded granitic erratic boulders with sources on the Canadian Shield; moderately to well-drained; greater than 2 m thick; minor moraines less than 1 km long and 5 m high; major moraines up to 12.5 km in length and 10 m high; ridges drape bedrock and older glacial deposits; minor moraines include crevasse-fill ridges and small recessional push moraines; major ridges features are large recessional end moraines and ice-thrust ridges; generally suitable for infrastructure placement. Streamlined till: silt and clay-rich diamictons; massive, matrix-supported and compact; clast contents less than 20% and contain sub-rounded granitic erratic boulders with sources on the Canadian Shield; moderately well-drained; greater than 2 m thick mantling bedrock and older glacial deposits; drumlins and fluted till ridges typically under 1 km long but can exceed 9 km in length; generally less than 50 m wide and 20 m high; formed beneath the Laurentide Ice Sheet directly through lodgement, basal meltout, glacigenic deformation of sediment beneath rapidly-flowing warm-based ice; generally suitable for infrastructure placement. ---- Geological boundary (Confidence: approximate) Bedrock scarp Major moraine ridge (end, interlobate, or unspecified) Other moraine ridge (DeGeer, minor lateral, recessional, rogen, washboard, other transverse or unspecified) <><><><> Esker ridge (sense: unknown or unspecified) Drumlin ridge Major meltwater channel scarp Hinor meltwater channel central axis (marginal, overflow, subglacial or unspecified; sense: known) Terrace scarp (environment: glaciofluvial) Station location (ground observation or stratigraphic section) 122°30'

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Canadian **Geoscience Maps** 

Authors: D.H. Huntley, A.S. Hickin, W. Chow, and M. Mirmohammadi Geology by D.H. Huntley and A.S. Hickin (2009–2010)

Geomatics by D.H. Huntley, W. Chow, and M. Mirmohammadi Cartography by W. Chow Initiative of the Geological Survey of Canada, Geological compilation by D.H. Huntley (2009–2011) conducted under the auspices of the Yukon Basir

**Preliminary** 

**SURFICIAL GEOLOGY TIGHTFIT LAKE British Columbia** mapping for Energy and Minerals (GEM) program

**CANADIAN GEOSCIENCE MAP 126** 

Map projection Universal Transverse Mercator, zone 10. North American Datum 1983 Base map at the scale of 1:50 000 from Natural Resources Canada, with modifications. Elevations in feet above mean sea level Magnetic declination 2013, 20°26'E

decreasing 22' annually.

**Preliminary** 

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this series have not been scientifically edited.