



**GEOLOGICAL SURVEY OF CANADA
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Sue-Dianne IOCG deposit, Northwest Territories and the Pipe
Ni-Cu deposit, Manitoba**

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ABSTRACT

This open file reports the chemical compositions of Fe oxide mineral grains (mainly magnetite and minor hematite) from the ferromagnetic fraction of bedrock and till from the iron oxide copper-gold (IOCG) Sue-Dianne deposit in the Great Bear magmatic zone, NWT and from till around the Pipe magmatic Ni-Cu deposit in the Thompson Nickel Belt, Manitoba. These data were collected in support of research by the DIVEX research network and the Geological Survey of Canada into the use of magnetite as an indicator mineral for mineral exploration in glaciated terrain.

INTRODUCTION

The composition of Fe-oxides, especially magnetite and hematite, has been shown to be useful to distinguish between different ore deposit types (Carew 2004; Gosselin *et al.* 2006; Singoyi *et al.* 2006; Nadoll *et al.* 2009; Rusk *et al.* 2009, 2010; Dupuis & Beaudoin, 2011). The method can be applied to ferromagnetic fractions of till samples to assist in the detection of mineral deposits up-ice in glaciated terrains. Beaudoin *et al.* (2009) showed that the composition of magnetite (and hematite) in a small aliquot of the total ferromagnetic fraction of till could be representative of the magnetite (and hematite) compositional range of the entire sample.

In this report, we present the optimized electron microprobe (EMP) method developed to analyze iron oxide grains, and new geochemical data for Fe oxides in till and bedrock samples collected up-ice, overland down-ice from the Cu-Ag-(Au) IOCG Sue-Dianne deposit in the Great Bear magmatic zone (GBmz), NWT and from the magmatic Ni-Cu deposits in the Thompson Nickel Belt (TNB), Manitoba. Both test sites are part of joint government-industry-academia research projects. The ongoing objective is to determine the optimum size fraction and number of grains for best representing the range in composition of magnetite and hematite contained in the till ferromagnetic fractions in order to document their compositional variations along ice-flow paths crossing mineral deposits.

METHODS

Eight till samples (7-19 kg of <2 mm material) containing ferromagnetic heavy minerals and five bedrock samples (1 to 3 kg) disaggregated to <2 mm from the Sue-Dianne IOCG deposit area along a 10-km long transect parallel to the predominant ice-flow direction (Fig. 1) were used in this study. Magnetite and hematite in four bedrock thin sections from the Sue-Dianne deposit and nearby Brooke Zone were also studied for comparison. Fourteen till samples (15-20 kg) from the TNB were used in this study, including samples proximal to the Pipe Ni-Cu deposit (Figs. 2 and 3), which collectively refers to the Pipe 1 and 2 underground mines and the Pipe open pit (Fig. 3). The samples are from a 180 km-long transect across the TNB along the general westward direction of the last major ice flow, as well as southwest of the Pipe deposit (Fig. 2) along the older SW direction of ice flow.

Till and bedrock samples were processed using a combination of tabling and heavy liquids (specific gravity 3.2) to produce ferro- and non-ferromagnetic heavy mineral

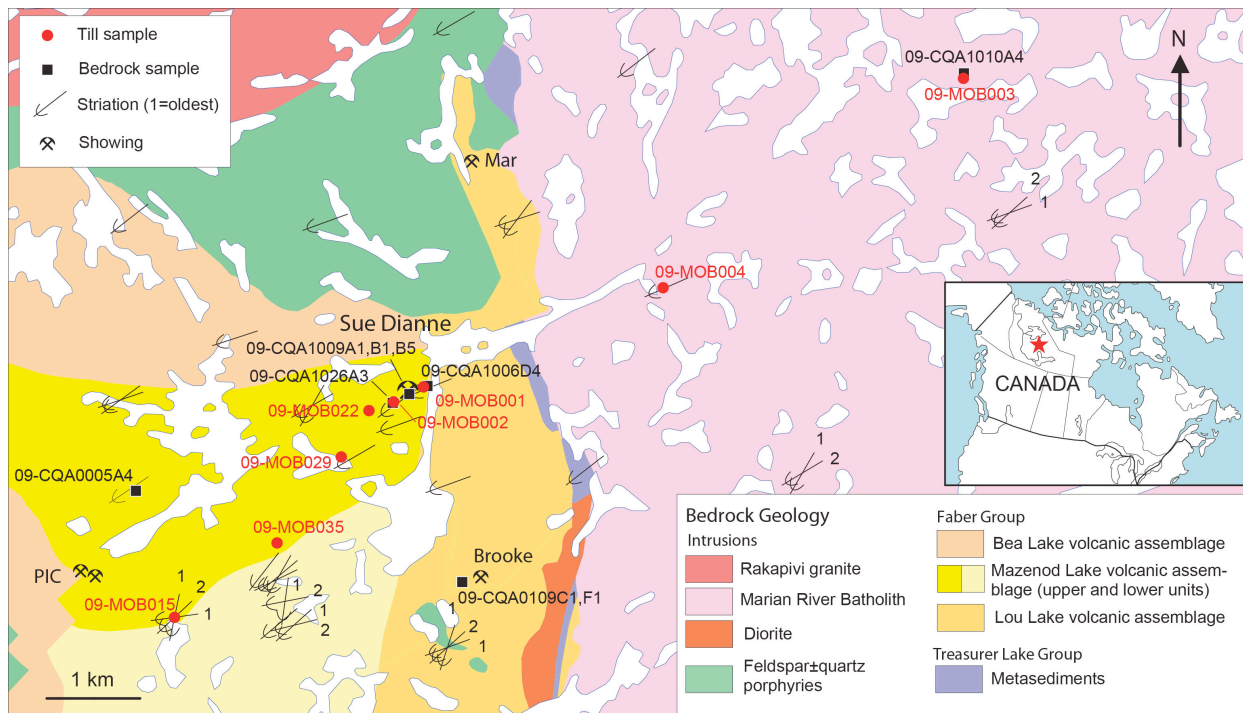


Figure 1. Mineral occurrences and regional bedrock geology of the Sue-Dianne deposit area (from NORMIN database: www.nwtgeoscience.ca/normin) and till/bedrock sample locations. Trend and relative age of striations are from McMartin *et al.* (in prep).

concentrates for picking of indicator minerals. Sample processing procedures and sample weights of the various fractions produced are reported for the Sue-Dianne samples by McMartin *et al.* (in prep.) and for the TNB samples by McClenaghan *et al.* (in press).

Individual grains from the ferromagnetic fraction of the samples were mounted in 25 mm epoxy grain mounts and the minor and trace element contents were determined for magnetite and hematite grains using a CAMECA SX-100 five-spectrometer electron microprobe (EMP) at the Université Laval, Québec. Analytical conditions are modified from those reported by Dupuis and Beaudoin (2011). The optimized EMP analytical conditions for magnetite and hematite are summarized in Table 1. EMP data for grains from the Sue-Dianne and Thompson are presented in Appendix A and B, respectively.

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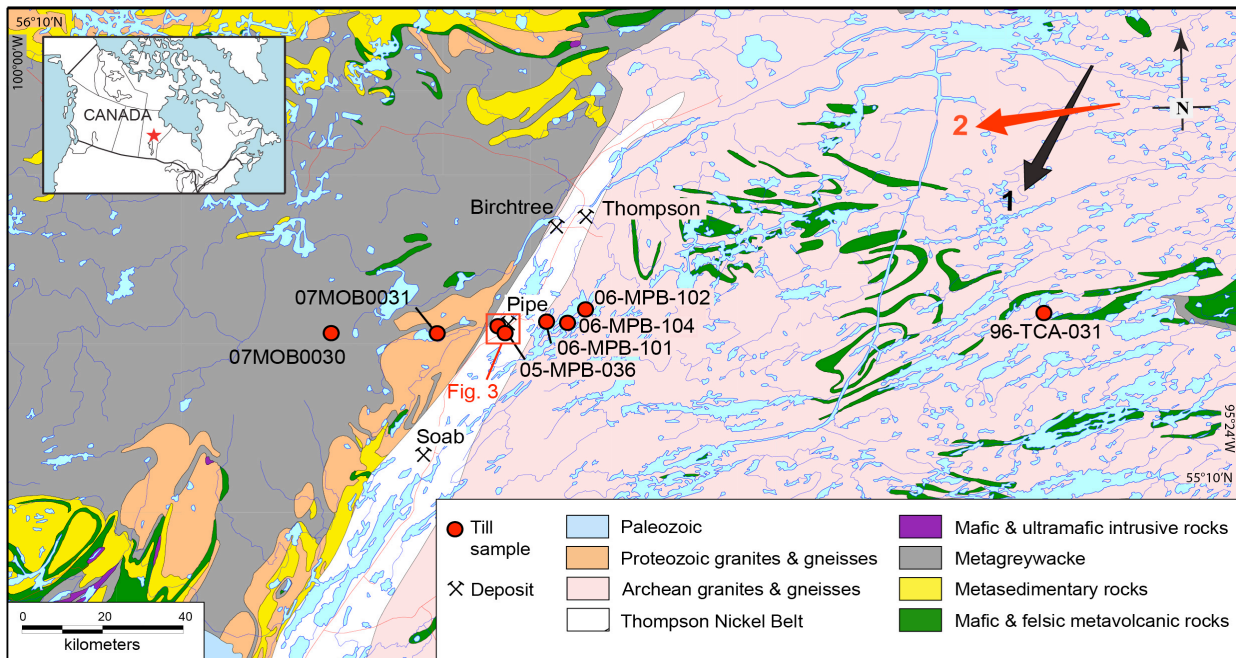


Figure 2. Simplified regional bedrock geology of the northern Thompson Nickel Belt region (Manitoba, Canada), showing the location of till samples (red dots) and the two ice-flow directions (1= older flow). Box outlines area shown in Figure 3 (modified from McClenaghan et al. 2011).

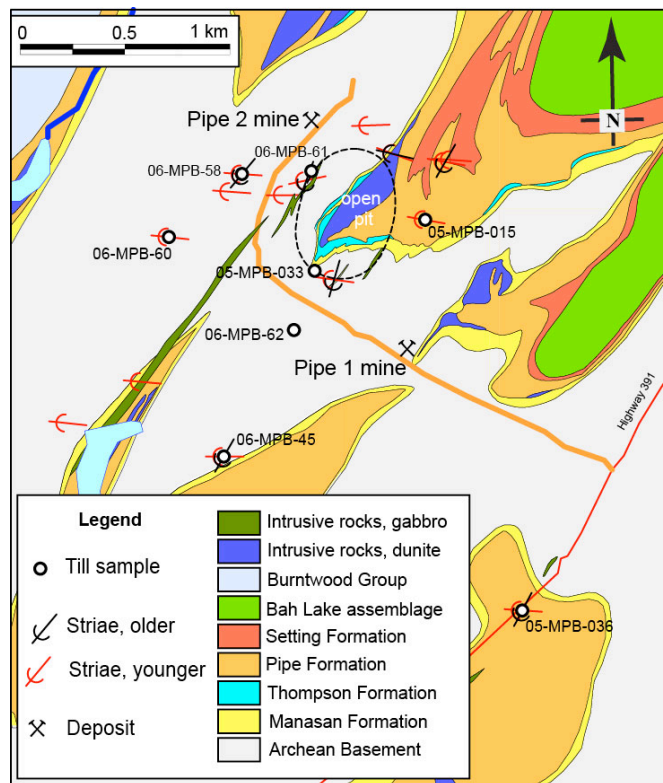


Figure 3. Bedrock geology of the Pipe deposit outlined in Figure 2, showing till sample locations (white dots) and the two ice flow directions (1=older flow) (from McClenaghan et al. 2011).

Table 1. Optimized analytical conditions for electron microprobe analysis of magnetite and hematite grains from till.

Element	Crystal	Line	Sinq*		Counting time (s)		Range of Detection Limits (ppm)
			Peak	Background	Peak	Background	
Zn	LIF	Ka	0.35635	0.35135	40	15	44-144
Cu	LIF	Ka	0.38256	0.37756	40	20	37-90
Ni	LIF	Ka	0.41173	0.40673	40	20	27-67
Mn	LLIF	Ka	0.52207	0.51607	40	20	17-45
Cr	LLIF	Ka	0.52207	0.51607	80	20	47-62
V	LLIF	Ka	0.62197	0.62797	80	20	45-65
K	LPET	Ka	0.42742	0.42142	40	20	14-16
Ca	LPET	Ka	0.38387	0.37787	40	20	15-17
Ti	LPET	Ka	0.31423	0.30823	40	20	19-25
Al	TAP	Ka	0.32462	0.33062	40	20	18-20
Si	TAP	Ka	0.27737	0.28337	60	20	14-16
Mg	TAP	Ka	0.38502	0.39102	40	20	20-23

* $\lambda = 2d \sin \theta$, where λ is the wavelength and d is the interplanar distance of the analysing crystal.

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