

INTRODUCTION

The traditional product of seabed mapping on Canadian continental shelves was the surficial geology map depicting Quaternary sediment formations (e.g., Fader et al., 1982). With the advent of multibeam sonar, three principal types of 3000 m wide swaths were produced: 1) bathymetry; 2) backscatter strength; and 3) surficial geology. Geologists have long been aware of the links between surficial geology and bathymetry (Merritt, 1985) and the recent awareness by geologists (Todd et al., 2007) has led to multiple approaches to visualizing the links on maps. This map is merely one approach among many. It is based on synthetic aperture radar (SAR) mapping by the Geological Survey of Canada (GSC) and Canadian Hydrographic Service (CHS), supported by seismic data, sidescan sonograms, bottom photographs, video, sediment observations, and grab samples.

SEASCAPES

Our definition of a 'seascape' is based on the Australian Land-System approach, developed to manage agricultural land. To adequately understand the land and its use and management it was thought necessary to understand the relationships between soils and the soil parent materials, climate, and topography. Land-systems are areas or groups of areas, throughout which there is a recurring pattern of topography, soils, and vegetation (Christen and Stewart, 1963). Commonwealth Scientific and Industrial Research Organization, 1987). By analogy, our seascape definition is as follows: Seascapes are underwater landscapes characterized by unique combinations of geomorphology, texture, and biota.

MAKKOVIK BANK

Makkovik Bank consists of gently dipping Cenozoic rocks separated from largely igneous rocks of the Makkovik Province (Wardle et al., 1989) by the Labrador Trough (see Fig. 1). The Labrador Current travels south in this region, but a deep-water jet is evident in the mean circulation pattern at Makkovik Bank, at all depths down to the sea floor (Fig. 2). The bank top has low relief, is deepest in the west (130 m), and gradually deepens eastward. It is fringed by a steep escarpment to the west, gentle slopes extend towards Makkovik and Hamilton Saddles. To the north and south respectively, a prism of stacked tills on the western side of the bank attains a maximum thickness of 30 m. The tills have been compacted, for they form several minor escarpments and a major escarpment that is 30 m high. The report has been highly impacted by lobbing furrows and pits. The density of furrows and pits on the sea floor varies, as does the ratio between those two. Parts of the multiple tills have almost no lobbing furrows or pits, perhaps indicating a resistant substrate, or influence of some other factor. Figure 3 is a transect illustrating varying degrees of lobbing impacted on terrain.

MAKKOVIK BANK SEASCAPES

We classify the seascapes into four broad areas:

Makkovik Province Seascapes occur west of the Labrador Trough, adjacent to the mainland. Bedrock Terrain consists of crystalline basement rocks of the Makkovik Province (Wardle et al., 1997). It has regular relief that ranges up to 20 m. Bedrock outcrops are rare, and a veneer of Quaternary sediments is more typical at the seafloor. A small area mapped as bedrock terrain in the extreme southwest is marked by *Ca* Geer moraines (see below). The Grounding Line Wedge located on the flanks of the Labrador Trough consists of till heavily imprinted by lobbing furrows. It was formed by grounded ice advancing from the mainland. A mass of glaciomarine mud complexes gravely sandy clay deposited by meltwater plumes and imprinted by lobbing furrows. Due to winnowing, texture at the seafloor is muddy angular gravel. Postglacial mud up to 20 m thick occupies depressions in the seafloor, and is imprinted by recent lobbing scour.

Labrador Trough Seascapes occupy relatively deep water (>100 m) between Makkovik Bank and the Labrador Shelf. Bedrock terrain to the southwest, and also the saddle between Makkovik and Hamilton Banks. Bedrock Terrain is morphologically similar to that of the Makkovik Province seascapes but is marked by a thick drapage of glaciomarine mud (Fig. 1). Fluted Terrain was created by moulding of ice originating on the mainland and flowing towards the east. Glaciomarine Mud consists of gravely sandy mud deposited by meltwater plumes and draped over pre-existing terrain. Relatively sandy clay deposited by meltwater plumes and imprinted by lobbing furrows. Due to winnowing, texture at the seafloor is muddy angular gravel. Postglacial mud up to 20 m thick occupies depressions in the seafloor, and is imprinted by recent lobbing scour.

Makkovik Bank Seascapes occupy the gently sloping top and more steeply sloping flanks of Makkovik Bank. The western, northern and southern flanks of the bank are mapped as the Bank Margin Moraine-Lower Till seascapes unit. Multiple layers of compact till from deep recolonizations up to 30 m high with slopes up to 15 degrees. Lobbing furrowing is evident, although less so above -100 m, and indeed the moraine surface seems smooth above this depth. The Upper Till is located at the north and south flanks of the bank, has more subdued relief, is marked by lobbing furrows, and is younger (Jøssensen and Zvenkovskan, 1989).

Bank Top Terrain has low relief, low to intermediate backscatter, sandy and gravel textures at the sea floor, and is heavily imprinted by lobbing furrows. A veneer of Quaternary sediments is present to allow Cenozoic bedrock structure to be evident at the seafloor in excess of 10 m deep channels of unknown origin. Some gently arcuate what others are arranged in a parallel fashion. Perhaps they were created by water under grounded ice, and were excavated along weaknesses in the underlying bedrock. Possibly meandering channels on offshore banks others are arranged in parallel fashion by Fåhræus and Higgins (1982). Another anomaly in this seascape is the presence of 5 m or more moraine ridges up to 3 m high, and suggestive of a general retreat towards the east.

A strip of Sand Waves runs along the west side of the bank in depth of 100-115 m. Finally, in the extreme east, in relatively deep water (200-240 m), is the Outer Bank Sand Seascapes. The low-backscatter sand is organized into barochan dunes indicative of a current from the northeast.

Continental Slope Seascapes

The Upper Continental Slope seascapes consists of a network of deep dendritic channels in depth below 300 m.

BACKSCATTER

Backscatter mosaic was used to classify the region into seascape units. Figure 4 illustrates some of the variability in the regular sands and muds, have low backscatter while other terrain, bedrock, till, glaciomarine mud are characterized by higher, more reflective substrates.

BIOTA

Coarse substrates predominate in the map area, but at varying depths, with varying sizes and current influence, so the biota seems to vary also. In some areas boulder-cobble-pebble seafloor show sponges and soft corals distributed fairly evenly (Fig. 6), but elsewhere a greater diversity is apparent (Fig. 7, 9). Crinoids and basket stars occur in quite large numbers in some areas, but it is difficult to confirm on their own distribution based on the small number of photographs. Figure 8 is a proxy typical of areas of postglacial mud.

ACKNOWLEDGMENTS

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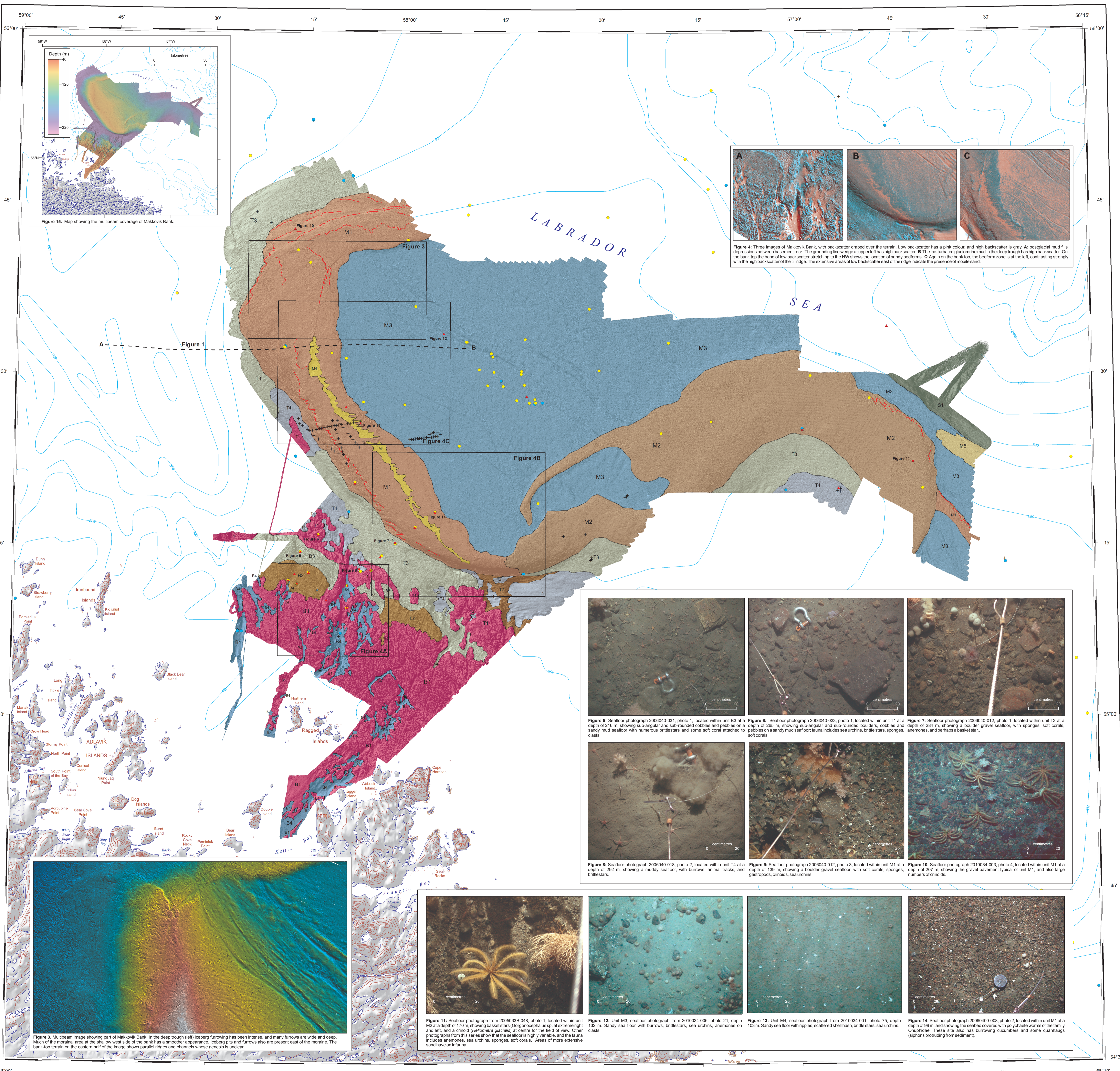


Figure 1: Airgun seismic reflection record showing the deep water of the Labrador Trough onto the bank. Figure 2: Mean annual current velocity at 100 m depth. Figure 3: Multibeam image showing part of Makkovik Bank. Figure 4: Bathymetry map. Figure 5: Seafloor photographs (B1, B2, B3, B4, T3, T4, M1, M2, M3, M4, M5) with detailed descriptions of each seascape unit's morphology, texture, and biota. Figure 6: Seafloor photograph showing sponges and soft corals. Figure 7: Seafloor photograph showing a boulder gravel seafloor. Figure 8: Seafloor photograph showing a muddy seafloor. Figure 9: Seafloor photograph showing a boulder gravel seafloor. Figure 10: Seafloor photograph showing a boulder gravel seafloor. Figure 11: Seafloor photograph showing a boulder gravel seafloor. Figure 12: Seafloor photograph showing a boulder gravel seafloor. Figure 13: Seafloor photograph showing a boulder gravel seafloor. Figure 14: Seafloor photograph showing a boulder gravel seafloor. 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