

Canada

LOCATION MAP

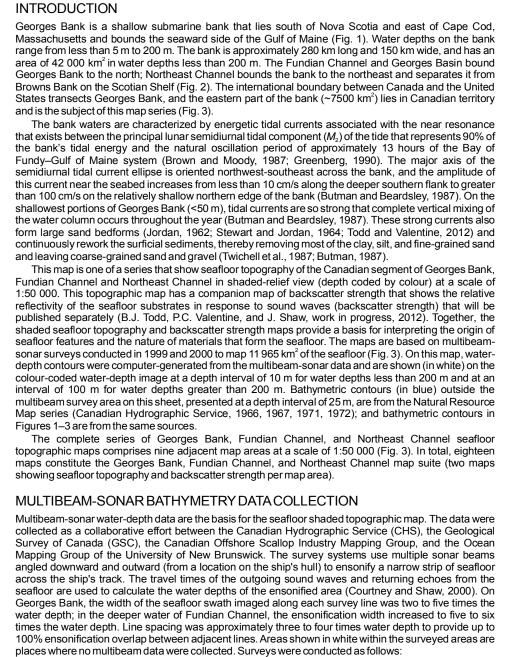
Multibeam bathymetric data compiled by Canadian Hydrographic Service and Geological Survey of Canada, 1999–2007

Digital cartography by P. O'Regan, Data Dissemination Division (DDD) and S. Hayward, GSC (Atlantic)

Universal Transverse Mercator Projection Projection transverse universelle de Mercator North American Datum 1983 Système de référence géodésigue nord-américain, 1983 © Her Majesty the Queen in Right of Canada 2013 © Sa Majesté la Reine du chef du Canada 2013 This map is not to be used for navigational purposes Cette carte ne doit pas être utilisée aux fins de navigation

Depth in metres below mean sea level

DESCRIPTIVE NOTES



- by the CHS in Northeast Channel in 1999 (area A, Fig. 2), using the Canadian Coast Guard Ship (CCGS) Frederick G. Creed, a SWATH (Small Waterplane Area Twin Hull) vessel equipped with a Kongsberg EM1000 multibeam-sonar bathymetric survey system with 60 beams operating at 95 kHz arrayed over an arc of 150° with the transducer mounted in the starboard pontoon, • by the CHS and the Canadian Offshore Scallop Industry Mapping Group on Georges Bank in 1999 and 2000 (area B, Fig. 2), using the MV Anne S. Pierce equipped with a Kongsberg EM1002 multibeam-sonar bathymetric survey system with 111 beams operating at 95 kHz configurable over an arc of up to 150° with the transducer mounted on a ram extended beneath the hull, and • by Clearwater Fine Foods, Inc. in Fundian Channel in 2000 (area C, Fig. 2), using the MV Anne S. Pierce equipped as described above. The Global Positioning System (with differential corrections) was used for navigation and provided a positional accuracy of ±3 m. Survey speeds averaged 14 knots (26 km/h) on the CCGS Creed resulting in an average data collection rate of about 48 km<sup>2</sup>/h in water depths of 250 m. Survey speeds averaged
- in water depths of 80 m. Sound velocity profiles were collected at regular intervals during multibeamsonar bathymetric data collection and were used to correct the data for refraction of the sonar beams due to density stratification in the water column. The bathymetric data were adjusted for tidal variation using an innovative tidal model that was developed for the seafloor mapping project by the Ocean Sciences Division of Fisheries and Oceans Canada at the Bedford Institute of Oceanography. The bathymetric data were processed using software developed by the Ocean Mapping Group at the University of New Brunswick. BATHYMETRIC DATA DISPLAY The multibeam-sonar bathymetric data are presented at 5 m per pixel horizontal resolution on Georges Bank and 10 m per pixel horizontal resolution in Fundian Channel and Northeast Channel. The shaded-

6 knots (11 km/h) on the MV Anne S. Pierce resulting in an average data collection rate of about 6.6 km²/h

- relief topographic image is presented with a vertical exaggeration of the bathymetry of 10 times and an artificial illumination of the relief by a virtual light source positioned 45° above the horizon at an azimuth of 315°. In the resulting image, topographic features are enhanced by strong illumination on the northwestfacing slopes and by shadows cast on the southeast-facing slopes. Superimposed on the topographic image are colours assigned to water depth, ranging from red (shallow) to violet (deep). In order to apply the widest colour range to the most frequently occurring water depths, hypsometric analysis was used calculate the cumulative frequency of water depth. Each figure has a unique colour ramp, and the colour ramp for an inset map showing topographic detail may differ from the colour ramp used for the main map. The resulting colour ramp highlights subtle variations in water depth that would otherwise be obscured. Some features in the multibeam data are artifacts of data collection and environmental conditions during the survey periods. The orientation of the survey track lines can, in some instances, be identified by faint parallel stripes in the image. Because these artifacts are usually regular and geometric in appearance on the map, the human eye can disregard them and distinguish real topographic features. GEOMORPHOLOGY OF GEORGES BANK, FUNDIAN CHANNEL, AND
- Physiographic setting Georges Bank flanks the seaward side of the Gulf of Maine and rises more than 300 m above the Gulf of Maine seafloor (Fig. 1, 2). Most of the Canadian portion of Georges Bank has water depths between 60 m and 90 m (Canadian Hydrographic Service, 1990, 1997; Valentine et al., 1992), but depths shallow to 42 m near 42°N, 67°W. The bank surface gradually deepens seaward and has an average slope of less than 0.05° (0.9 m/km). The 200 m isobath along the southeastern margin of Georges Bank approximates the continental shelf break (Fig. 1). Seaward, water depths increase down the continental slope, which has an average slope of 7° (123 m/km). The seaward margin of the bank in Canada is incised by a number of submarine canyons, the largest of which are Corsair and Georges canyons (Fig. 2). Georges Bank is bounded to the north by the Fundian Channel which comprises Northeast Channel to the east and Georges Basin to the west (Fig. 2). Georges Basin is the deepest portion of the Gulf of Maine with a maximum depth of 377 m. The Fundian Channel north of Georges Bank is approximately 45 km wide and is bounded to the north by Browns Bank on the Scotian Shelf. The Fundian Channel narrows southeastward into Northeast Channel, which is approximately 28 km wide. The seaward sill

depth in Northeast Channel is 232 m, and the channel mouth is incised by submarine canyons (Fig. 2).

NORTHEAST CHANNEL

Recent geological history Seismic-reflection profiles show that beneath the surface of Georges Bank there is a prominent unconformity formed on late Cretaceous and Tertiary sedimentary rocks (King and MacLean, 1976 Lewis et al., 1980). The surficial sediment overlying the unconformity is glacial debris transported to Georges Bank and other Gulf of Maine banks during the late Pleistocene epoch from continental areas to the north (Shepard et al., 1934, Knott and Hoskins, 1968; Oldale and Uchupi, 1970; Schlee, 1973; Schlee and Pratt, 1970; Fader, 1984; Fader et al., 1988; Todd et al., 2007). During the postglacial Holocene epoch (~12 000 years before present), sea level rose from a low stand 120 m below the present sea level (Emery and Garrison, 1967; King and Fader, 1986), and the bank was submerged about 6000 BP (radiocarbon years) (Shaw et al., 2002). Georges Bank surficial sediments were reworked by marine processes during sea-level transgression and continue to be reworked under the modern oceanic regime (Twichell et al., 1987; Valentine et al., 1993). The present morphology of the Gulf of Maine, Georges Bank, and the Fundian and Northeast channels mapped here displays the imprint of multiple glaciations during the Pleistocene epoch (~12 000 years ago to 2.5 million years ago). During the last glaciation of the Pleistocene, the Wisconsinan, the Laurentide Ice Sheet extended southeastward from central Canada across Maritime Canada and New England to the present northern margin of Georges Bank and the continental shelf edge off Nova Scotia and the Fundian and Northeast channels were a major outlet for glacial ice to the Atlantic Ocean (Shaw

et al., 2006). Based on the mapping of glacial gravel collected from the seabed in the Gulf of Maine region, Schlee and Pratt (1970) reported that glacial ice lapped onto the northern margin of Georges Bank.

- Geomorphology of Sheet 8 The east-central region of the Canadian portion of Georges Bank is shown on this map at a scale of 1:50 000. Inset maps at a scale of 1:25 000 (Fig. 4–9) highlight geomorphological features typical of the region. For each of these detailed maps, the colour-range values are hypsometrically optimized and differ from the colour-range values of the 1:50 000 scale map. A prominent feature on this map is the head of an unnamed submarine canyon that incises the continental shelf and slope of Georges Bank. The canyon head extends to the 130 m isobath on the shelf, and water depths in the canyon reach more than 700 m in the mapped region. The walls of the canyon are incised by gullies with interposed ridges (Fig. 4). Along the bank margin northeast of the canyon to 41°41'N, subparallel rounded ridges up to 6 m high are present below the shelf break in the 200-300 m depth interval (Fig. 5). These features are interpreted to represent the early stages of sediment slumping. Similar features are present in the southwestern part of the canyon head. The Georges Bank seafloor both south and north of the canyon has been incised with linear to curvilinear depressions and circular to oval pits (Fig. 6, 7). The depressions, or scours, and pits are interpreted as evidence of the dragging of the keels of icebergs on the seabed. Most of the scour marks have a muted appearance, whereas a few of the longest scours (3-8 km) show a well defined morphology (Fig. 7). The scours are oriented generally northeast-southwest. As iceberg trajectory is dictated mainly by ocean current direction (Todd et al., 1988; Bigg et al., 1996), a northeast to southwest current direction is inferred for the continental margin of Georges Bank during the last deglaciation. These icebergs may have calved from the front of the floating ice shelf in the Gulf of Maine and drifted out Northeast Channel
- North of 41°43′N, terminations of scour marks are associated with an irregular scarp-like linear feature of low relief at 120 m water depth that may delineate the position of a Wisconsinan glacial or postglacial shoreline. This feature extends northward into Sheet 5 of this map series (Todd et al., 2013a). South of the canyon, iceberg scour marks are present in the 130 m to more than 300 m depth interval. A large area of this map, in the 90–110 m depth interval, dips very slightly seaward and is covered with sediment bedforms. In the northwest portion of the map at depths from 90 m to 100 m, sediment bedforms with well defined crestlines form a simple anastomosing pattern in plan view with an overall northeast-southwest orientation, normal to the direction of dominant residual current flow, which is generally to the southeast. These bedforms range in height from a few decimetres to over 10 m (Fig. 8). At somewhat deeper depths to 110 m, bedforms are smaller and form 'horsetail' patterns of subparallel curvilinear crestlines that converge to the northeast and have heights of less than 1 m (Fig. 9). A small isolated field of low-relief, short-wavelength bedforms is present southwest of the canyon head on the outer part of the bank (Fig. 10). A small portion of the head of Corsair Canyon is present in the southwestern corner of the map. See Sheet 9 of this map series for a map of the canyon (Todd et al., 2013b).

North of the submarine canyon, iceberg keel marks are present in the 110 m to 160 m depth interval.

and thence southwestward along Georges Bank.

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