



INTRODUCTION

The Bay of Fundy, located on the east coast of Canada between the provinces of New Scotia and New Brunswick (Fig. 1), is a macrotidal estuarine embayment (Amos et al., 1980) with the highest recorded tides in the world (17 m (O'Reilly et al., 2009; Bishop, 2008)). This map is one of a series of maps that show seafloor relief of the Bay of Fundy and topography of the surrounding area in shaded-relief view (color by contour) at a scale of 1:50 000. The maps are based on multibeam-sensor surveys completed between 1993 and 2008 to map 13 000 km² of the seafloor. Water-depth contours generated from the multibeam-sensor data are shown in white on the color-coded water-depth image at a depth interval of 20 m. Bathymetric contours in blue outline the multibeam survey area, presented as a depth interval of 50 m, are from the Natural Resource Map series (Canadian Hydrographic Service, 1967, 1974a, b, c). The broad intertidal zone in the Bay of Fundy presented a particular surveying challenge: the collection of water-depth data. Historically, the intertidal zone was not surveyed due to the danger involved in operating vessels in coastal areas that dry between tides. As part of the multibeam-sensor mapping, the intertidal zone was surveyed at high tide using shallow-draft survey vessels, thus overcoming operational challenges associated with deeper draft survey vessels.

DESCRIPTIVE NOTES

MULTIBEAM BATHYMETRY DATA COLLECTION

- The Canadian Coast Guard Ship (CCGS) Frederick G. Creed, a 5940 t St. John's Waterkeeper Area Tug, hull vessel equipped with a Kongsberg EM1000 (prior to 2003) and a Kongsberg EM1002 (post-2003) multibeam-sensor bathymetric survey system with 111 beams operating at 45 kHz with the transducer mounted in the starboard position.
- The CCGS Mathews equipped with a Kongsberg EM100 multibeam-sensor bathymetric survey system with 200 or 400 beams operating at 70–80 kHz with the transducer mounted near the centre of the vessel.
- Hydrographic survey launches *Plouffe*, *Pilot*, and *Heron* equipped with Kongsberg EM3000 (prior to 2003) and Kongsberg EM1002 (post-2003) multibeam-sensor bathymetric survey systems with 254 beams operating at 300 kHz.

BATHYMETRIC DATA DISPLAY

The multibeam-sensor bathymetric data are presented at 5 m per pixel horizontal resolution. The shaded-relief image is presented with a vertical exaggeration of the bathymetry of 30 times and an artificial illumination of the seafloor by a virtual light source positioned 45° above the horizon at an azimuth of 315°. In the resulting image, bathymetric features are enhanced by strong illumination on the northward-facing slopes and by shadows cast on the southward-facing slopes. Superimposed on the shaded-relief image are colour-coded water depths ranging from red (shallow) to violet (deep). In order to apply the widest colour range to the most frequently occurring water depths, hypsometric analysis was used to calculate the cumulative frequency of water depth. The resulting colour ramp highlights subtle variations in water depth that would otherwise be obscured.

BAY OF FUNDY GEOMORPHOLOGY

The Bay of Fundy is a southwest-trending funnel-shaped bay 155 km long and 70 km wide at its entrance and tapers to 48 km at its northeastern end where it bifurcates into Chignecto Bay and Minas Channel (Fig. 1). The floor of the bay, although hummocky in detail, presents a gently dipping profile along its axis from northeast to southwest. Grand Marais Island is adjacent south of the bay and Grand Marais Island is several isolated depressions that together form Grand Marais Basin. The maximum water depth within this depression is 233 m and the depth to the sill between Grand Marais Basin and the adjoining deeper parts of the Gulf of Maine is 160 m.

The large tidal oscillations within this depression are due to the near resonance between the principal lunar semidiurnal (M₂) component of the tide (impressing 90% of the tidal energy) and the natural period (about 13 hours) of the Bay of Fundy-Gulf of Maine system. Tidal current speeds are about 0.7–1 m/s over much of the outer and central portions of the bay, but are considerably higher within constricted channels and passages to the northeast (Greenberg, 1990).

Geological History

Geomorphological features revealed through mapping of the Bay of Fundy seafloor reflect the geological history of the region. The Bay of Fundy is situated within the Cretaceous-Tertiary boundary (Goldsworthy, 1992; Crosby, 1962; Williams et al., 1972) and is underlain by Tertiary and Early Jurassic sandstone, shale, and basalt (Wells et al., 1968). Exposed bedrock has been modified by glacial erosion and exhibits a rugged surface.

During the late Wisconsinan glacial maximum, culminating in the Gulf of Maine region at approximately 20 ka (20 000 BP), the Bay of Fundy was covered by a regional ice sheet that terminated to the south on the Scotian Shelf (Schroeder et al., 2001; Shaw et al., 2007). The glacial maximum was followed by a multiphased retreat of the ice front. In the Gulf of Maine, ice-front retreat and glaciomarine deposition began as early as 15 ka (Schroeder et al., 2001; Shaw et al., 2007). The Bay of Fundy exhibits geomorphological features formed during the Quaternary glacial and deglacial periods: the area contains moraines, drumlins, and megaripples are topographically prominent. After glacial retreat from the area, icebergs scoured the seafloor in the waters east of Grand Marais Island.

After deglaciation, relative sea level fell rapidly to a lowstand of about 30 m at ca. 7 ka (Amos and Zaitch, 1980; Shaw et al., 2002) and then rose (Green, 1970). These high tides have resulted in large zones of erosion in areas with high current velocities (e.g., Cape D'Or and Cap de Sable, Fig. 1). Tidal eddies produced by headlands have created barrier banks (Dyer and Huntley, 1999) on both sides of coastal promontories. Coastal erosion is up to 8 m in some areas (Amos et al., 1981). Sediment derived from this coastal erosion, coupled with sediment from seafloor erosion and sediment delivered by rivers, has contributed to the development of intertidal mudflats and salt marshes in the Bay of Fundy. The coasts of the bay also host salt marshes and dykelands (Glenn, 1993; Gordon et al., 1985). Seaward of the mud flats in the subtidal zone is a thin layer of sand consisting of eroded bedrock, gravel, sand, and mud. In places, strong tidal currents create sand waves several metres in height and hundreds of metres in length (Greenberg et al., 1997).

Geomorphology of this map

A series of detailed maps at a scale of 1:50 000 (Fig. 2–7) highlights geomorphological features of Digby, Nova Scotia. For each of these detailed maps, the colour-range values are hypsometrically optimized and differ from the 1:50 000 map sheet colour-range values. This map shows the bathymetry of the Bay of Fundy offshore Digby, Nova Scotia (Fig. 1). Bathymetric data were also collected in Digby Gut and the Annapolis Basin.

In the northwest portion of the map, irregular pits in the seafloor are evident on ridges in water depths of approximately 60 m and length (Fig. 2). A pit is formed by a single, discrete impact of an isobath land into the seabed sediment (Fader and King, 1981). Beams greater than 60 m water depth appear little impacted, suggesting that the ridges provided a degree of bathymetric shielding of the seabed from isobath levels. The number of pits in the seabed suggests this area witnessed a substantial flux of icebergs during the retreat of the ice front.

Within 7–10 km offshore Digby Neck and North Mountain, the seafloor morphology is dominated by glacial megaripples and channelled ridges oriented northeast-southwest (Fig. 3, 4). Similar glacial geomorphology has been mapped elsewhere in the Gulf of Maine and is interpreted as indicating ice-flow direction and speed (Todd et al., 2007).

Offshore Digby Gut is a morphologically complex ridge striking approximately northwest for 8 km (Fig. 5). A sub-seafloor ridge, prominent in the bathymetry of Digby Gut, the ridge exhibits beamed crests with the northern ridge reaching 30 m in height and 300 m in width at the base. To the northwest, the ridge broadens to a small mound 10 m in height. At the distal end of the ridge is a ridge a few metres above the seabed and 100 m wide. This ridge is interpreted to be a segmented landfill (e.g., Warren and Ashley, 1999) formed during post-glacial retreat.

North of the mouth of Digby Gut is a set of four topographic sediment bedforms (Fig. 6). These bedforms have pronounced linear crests to 18 m high, 150 m wide, and 1400 m long, oriented northeast-southwest, normal to the southwest-northeast flow direction of water into the bay and the northeast-southwest flow direction of water out of the bay (Canadian Hydrographic Service, 1967; Greenberg, 1980). The ends of the crests terminate in arcuate moatings giving the visual impression that the bedforms are inclined to the seabed.

Within Digby Gut are sediment bedforms on the floor of the channel (Fig. 7). In contrast with bedforms in the open Bay of Fundy, the Digby Gut bedforms are conical and are only tens of metres in length with typical bedform heights of 4 m and widths of 50 m. The morphology of these channel bedforms is a function of the hydrodynamics of Digby Gut.



Figure 1. Location map showing seventeen 1:50 000 map sheets covering the Bay of Fundy. Sheet 6 (outlined by red box) is in central Bay of Fundy, offshore Digby, Nova Scotia.

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