

Figure 8. View northwest towards inner-continental shelf in Portley Cove. Photograph by J. Shaw, 22 May 2006, 2010-062.

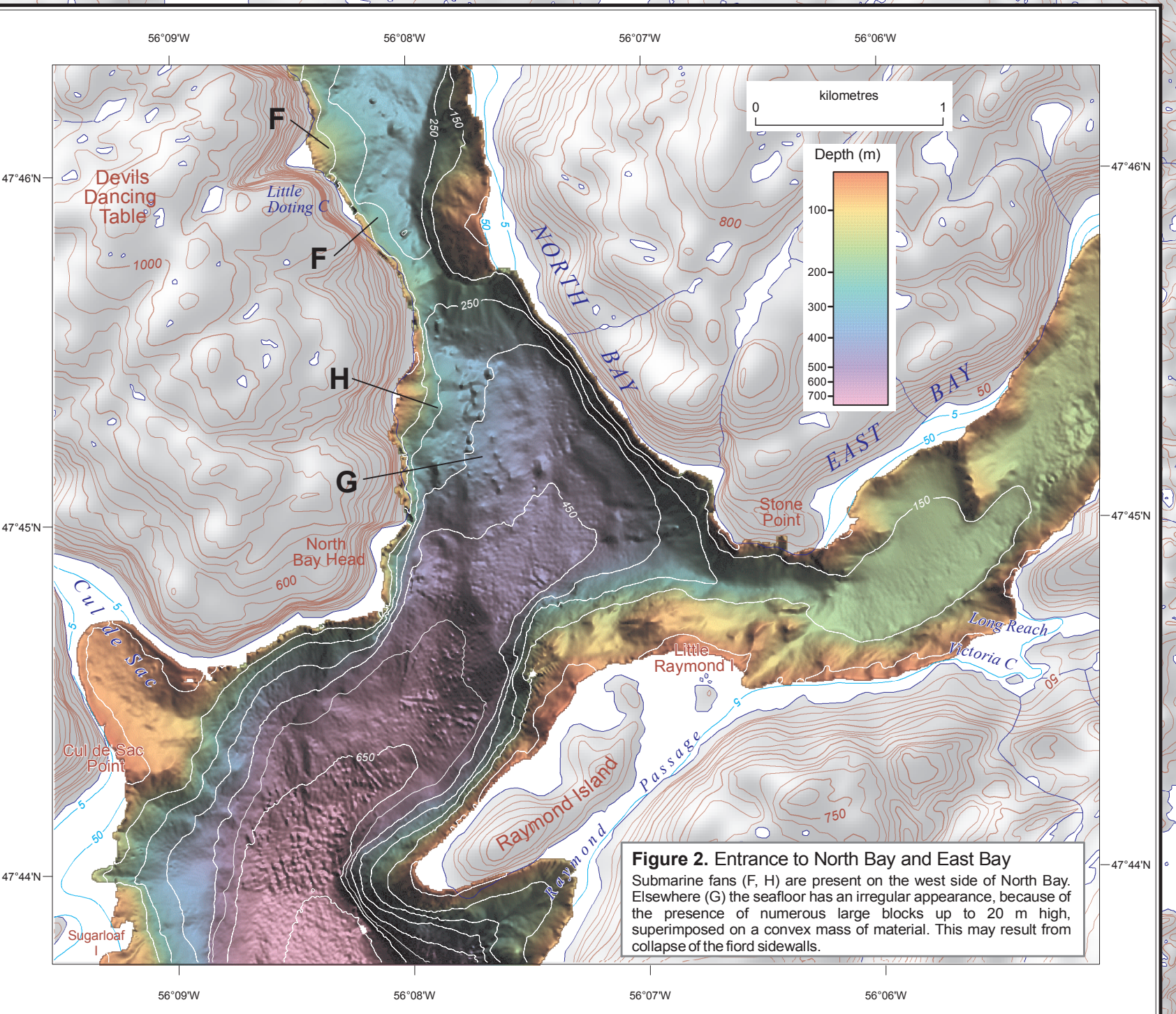


Figure 2. Entrance to North Bay and East Bay. Supratidal bars of 10 m are present on the west side of North Bay. Exposure of the bedrock is an unusual appearance because of the presence of numerous large stones 10 to 20 m high, superimposed on a coarse mass of material. This may result from collapse of the bed sediments.

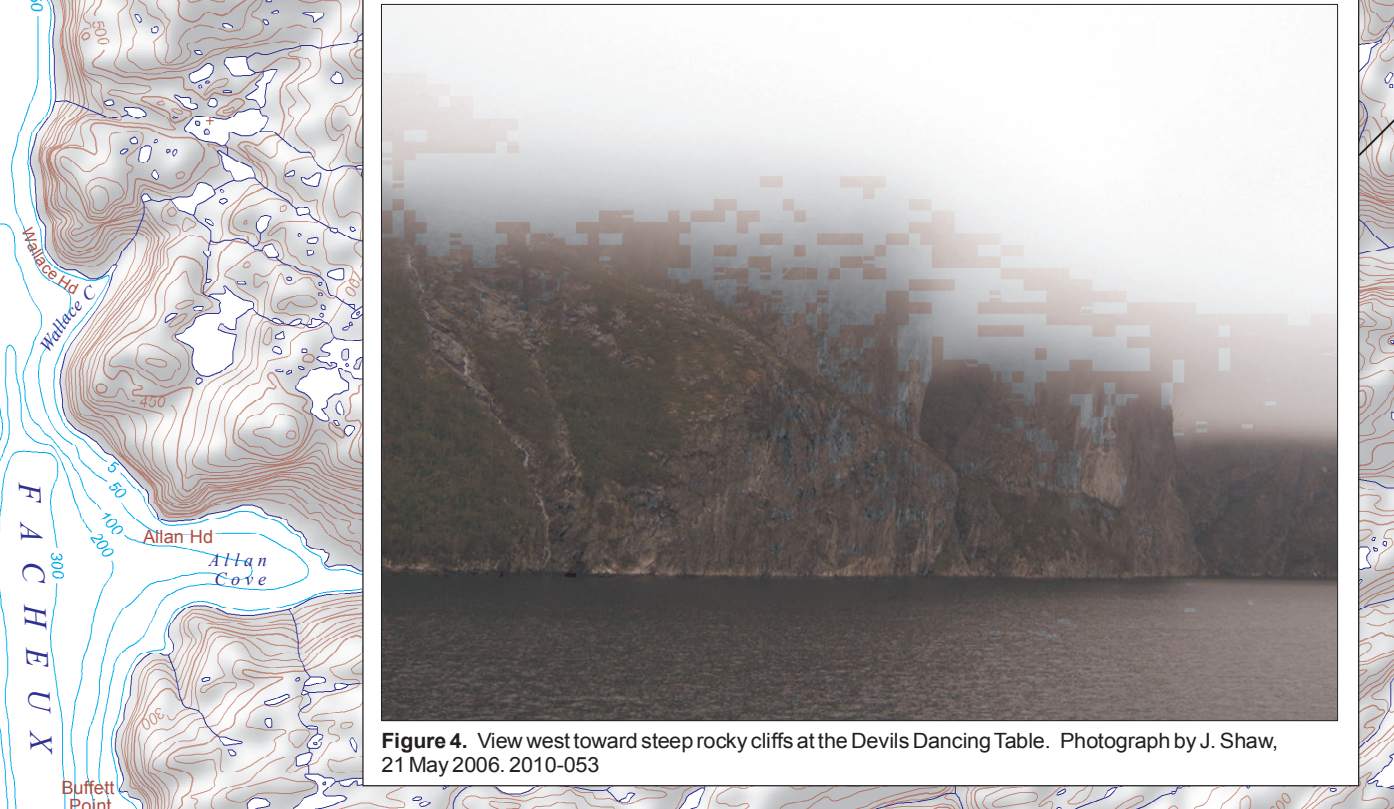


Figure 4. View west toward inner-shelf cliffs at the Devils Diving Table. Photograph by J. Shaw, 21 May 2006, 2010-063.



Figure 7. Aerial view of the bay entrance. Photograph by J. Shaw, 21 May 2006, 2010-054.

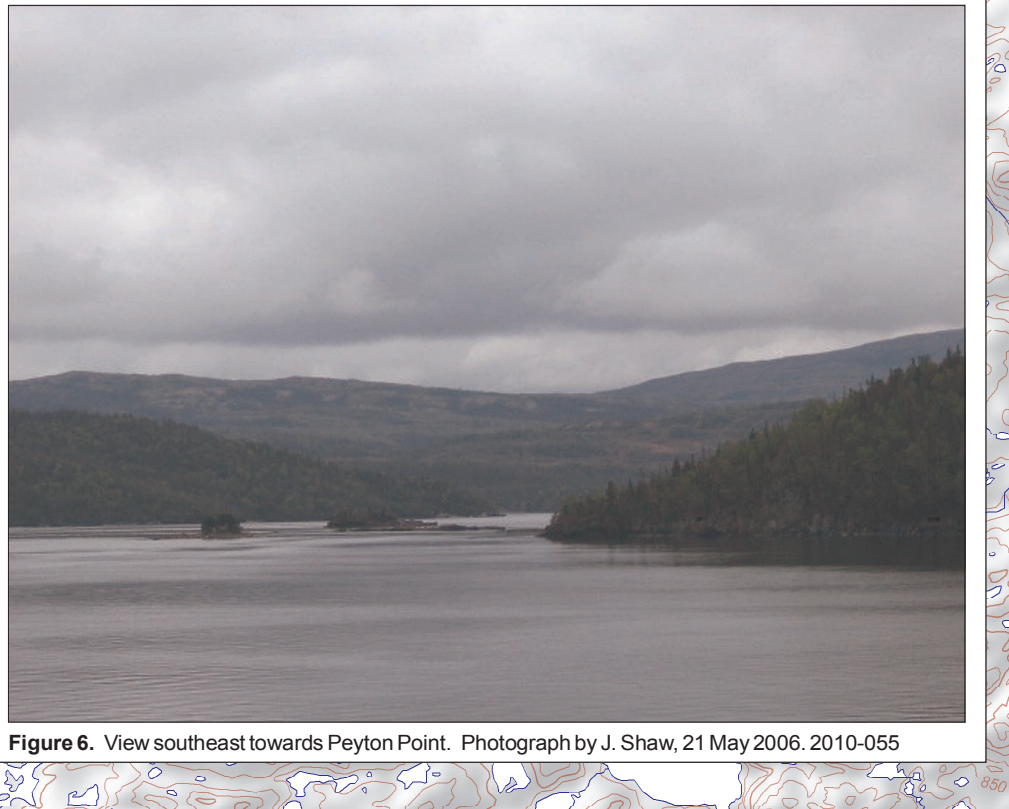


Figure 6. View southeast toward Peyton Point. Photograph by J. Shaw, 21 May 2006, 2010-056.

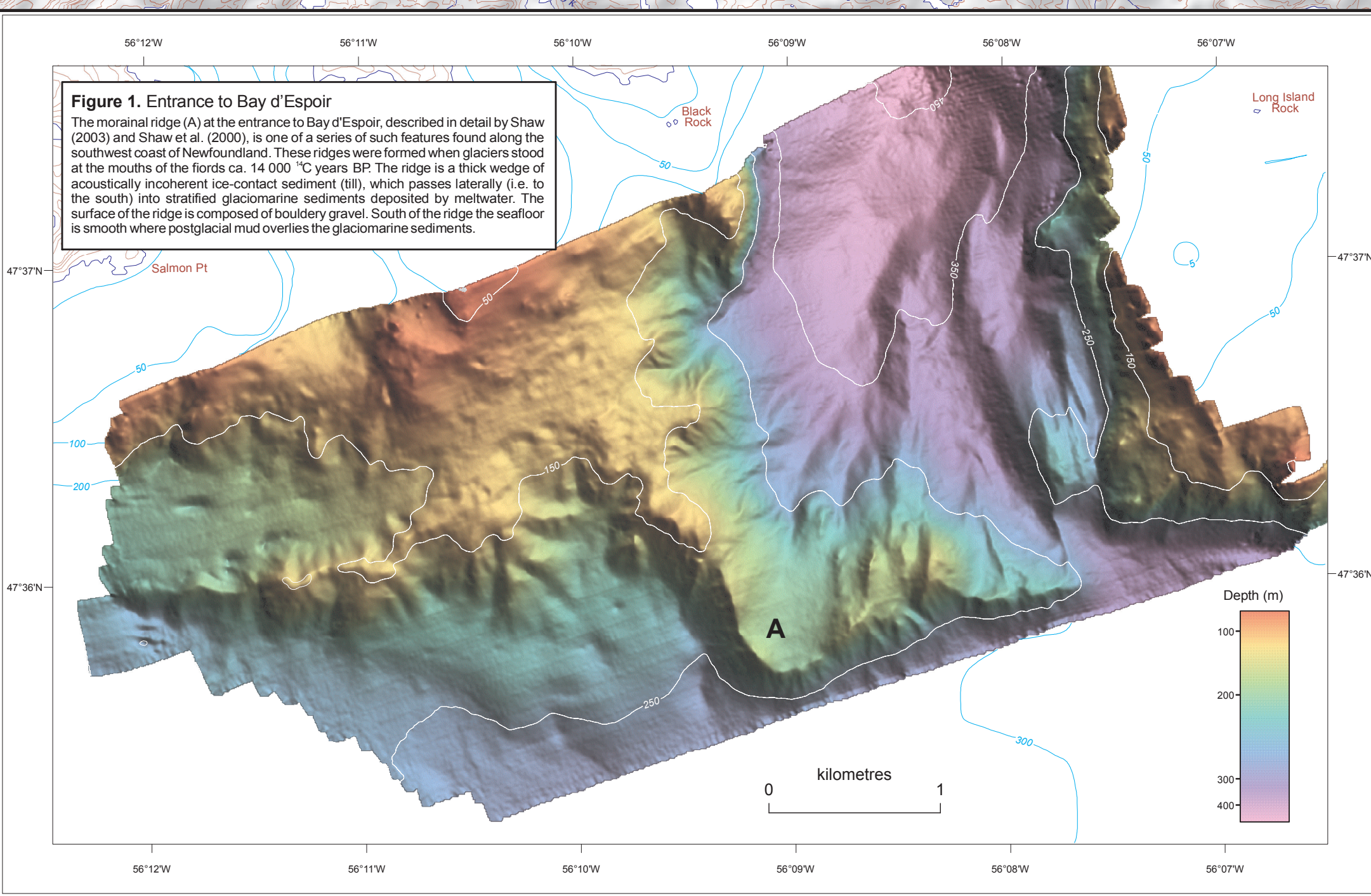


Figure 1. Entrance to Bay d'Espoir. The moraine ridge (A) at the entrance to Bay d'Espoir, described in detail by Shaw (2003) and Shaw et al. (2006), is a remnant of an ice advance from the southeast coast of Newfoundland. These ridges are formed when glacial ice at the margin of the Flandre Sea, 14 000 to 15 000 years BP, the ridge is a back wedge of acoustically laminated ice-contact sediments (B), which pass laterally (C) to the south into stratified glaciolacustrine deposits by meltwater. The surface of the ridge is composed of boulders (D). South of the ridge the seafloor is smooth where bedrock and covered by glaciolacustrine sediments.

DESCRIPTIVE NOTES

INTRODUCTION
This map shows the bathymetry of Bay d'Espoir, a major fjord on the south coast of the island of Newfoundland, located in the western part of the continental shelf of Atlantic Canada (170 m). Located on the Canadian Hydrographic Service Chart 466, the bay includes a shallow area at the entrance (Fig. 1) and a deep north-facing bay just inside. Immediately north of the entrance the bay narrows to East Bay and North Bay (Fig. 2). The narrow neck of the bay is a wide channel. A narrow neck of the bay is a wide channel. The inner bay terminates in an arm extending northeast to the St. Alban's area (Fig. 3). The narrow channel of the Passage connects the bay with Hermitage Bay, extending along inland from the mainland. The bay has incoming sediments, which are mostly in the form of a wide channel.

MULTIBEAM BATHYMETRIC DATA COLLECTION
The earliest data collection was in 1996, when CCGS Matthew surveyed the entrance to Bay d'Espoir using the full-mounted Simrad EK600 system. In 1999 surveys were conducted using the Furukawa O. Crowl, a SIMRAD (Small Inshore Area Towed) vessel equipped with a Simrad EK1000 multibeam echosounder with the transducer mounted in the catamaran position. The surveys were conducted by the Geological Survey of Canada (GSC) and covered most of the bay, although data quality in the deep trench was poor. In the fall of the same year the Canadian Hydrographic Service (CHS) conducted a survey of the bay using a Simrad EK600 system. The CHS survey was conducted using a Simrad EK600 system. The CHS survey was conducted using a Simrad EK600 system. The CHS survey was conducted using a Simrad EK600 system.

Table 1. Remarks on surveys carried out in study area. Columns include YEAR and REMARKS.

DATA DISPLAY
Artificial sun illumination from 000° azimuth and 30° inclination was applied in the GRASS GIS. Vertical exaggeration is 4:1. A colour palette was applied to the bathymetric data, with colours (e.g., red) representing shallow water and cool colours (e.g., blue) representing deep water. Histogram equalization was applied to the colour palette, i.e. bathymetric divisions between colours were weighted such that equal areas are covered by each colour in the palette.

MORPHOLOGY
The region is topographically complex, and shows evidence of strong bedrock structural control, particularly by glacially and the effects of lowered sea levels in the early Holocene. Cliffs and bay have steep rugged shorelines, with thin covers of unconsolidated sediments. Shallow bays at the mouth of the bay are a result of glacial erosion. The deep trench is a result of glacial erosion. The deep trench is a result of glacial erosion. The deep trench is a result of glacial erosion.

Mouth of Bay d'Espoir
The mouth (A) of the bay (see also Fig. 1) is an accurate submarine ridge 80 m high and composed of unconsolidated sediments. The ridge is 75 m thick wedge of acoustically laminated ice-contact sediments (B). The ridge is 75 m thick wedge of acoustically laminated ice-contact sediments (B). The ridge is 75 m thick wedge of acoustically laminated ice-contact sediments (B). The ridge is 75 m thick wedge of acoustically laminated ice-contact sediments (B).

North Bay and East Bay
North Bay (see also Fig. 2) is a shallow bay with a depth averaging 200 m, and has a maximum depth of 700 m. The bay is a shallow bay with a depth averaging 200 m, and has a maximum depth of 700 m. The bay is a shallow bay with a depth averaging 200 m, and has a maximum depth of 700 m. The bay is a shallow bay with a depth averaging 200 m, and has a maximum depth of 700 m.

Lampyris Passage
A narrow (K) is located at the entrance to Lampyris Passage. The ridge is 75 m high and composed of unconsolidated sediments. The ridge is 75 m high and composed of unconsolidated sediments. The ridge is 75 m high and composed of unconsolidated sediments. The ridge is 75 m high and composed of unconsolidated sediments.

Northwest Bay d'Espoir
This bay is a shallow area extensively covered by acoustics (Fig. 7). North of Little Cove Head the basin floor (C) is a shallow area extensively covered by acoustics (Fig. 7). North of Little Cove Head the basin floor (C) is a shallow area extensively covered by acoustics (Fig. 7). North of Little Cove Head the basin floor (C) is a shallow area extensively covered by acoustics (Fig. 7).

ACKNOWLEDGMENTS
The authors thank CHS for providing access to data collected in 1996. For surveys in 1999 the authors acknowledge the master and crew of CCGS Matthew O. Crowl and CCGS Matthew O. Crowl. The authors also thank a Geomatics Canada (GSC) for providing access to bathymetric data collected during the 1996 survey. The authors also thank the staff of the Geomatics Canada (GSC) for providing access to bathymetric data collected during the 1996 survey. The authors also thank the staff of the Geomatics Canada (GSC) for providing access to bathymetric data collected during the 1996 survey.

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