

Figure 6. Hazards from a cinder cone eruption.

Conventional (height of

Nazko Earthquake Swarm October 2007 – May 2008 On Tuesday October 9, 2007, a swarm of small earthquakes commenced

of the Anahim Volcanic Belt, which stretches from central BC to the coast.

In order to more easily pinpoint the location and depth of future tremors, Natural Resources Canada personnel installed additional seismic stations in the Nazko region (Figure 1). An infrasound station was installed to record low frequency sound waves below the threshold of human hearing. Carbon dioxide levels were also measured in the epicentral area (Inset Map 1). Carbon dioxide is commonly released along faults and fractures in volcanically active regions and may accumulate in soil or beneath snow, so elevated levels provide information about subsurface magmatic activity. However, all gas measurements were within the normal range, indicating no volcanic release of carbon dioxide.

Analysis of the seismic data suggested the earthquake swarm was caused by an intrusion of magma deep within the earth's crust. Similar sequences of events have occurred in other volcanic areas without being followed by eruptions. For example, in 2003, a swarm of deep (29-33 km) earthquakes was recorded beneath Lake Tahoe and attributed to magma (Smith et al.,

Although the main Nazko earthquake swarm ended in late November 2007, the level of activity has not yet returned to normal. Intense seismic activity might again resume in the Nazko region, and be indicative of deep movement of magma, but it might also indicate the ascension of magma towards the surface. If the latter were to occur, it is anticipated that the size and number of earthquakes would increase significantly, providing a warning in the very unlikely event of an eruption. If an eruption were to occur, it would likely be a small cinder cone building event, similar to the most recent Nazko eruption about 7000 years ago. Given our present understanding of the seismicity, the volcanic vent would most likely be

within a few kilometres of the epicentre of the earthquake swarm.

Background

Nazko Cone (Figure 3) is the youngest known volcano in the Anahim Volcanic Belt, and was first mapped by Souther et al. (1987). The first volcanism at Nazko Cone started about 340,000 years ago, with the eruption of a basaltic lava flow. This was followed by eruptions of basalt and breccia into glacial meltwater during the Fraser Glaciation (25,000-10,000 years ago). A final stage of eruption took place about 7200 years ago, after the disappearance of the continental ice sheet. This eruption fountaining vents. Two separate streams of lava later breached this cone and were diverted around the north and south flanks of the central mound (Inset Map 1). Tephra associated with this final eruption was scattered to the north and east, in deposits up to 3 m thick near the cone and less than a few centimetres thick a few kilometres away (Inset Map 1; Figure 4). This limited distribution of volcanic ash suggests that the explosive fire fountaining phase was relatively brief. This last Nazko Cone eruption is probably a good analogue for what might happen if a new eruption were to occur in the vicinity of the earthquake swarm. Quarrying at Nazko Cone is ongoing, as the scoria is used as a lightweight aggregate, landscaping material, and agricultural soil additive (Figure 5). Outcrops of Chilcotin basalt assumed to be a few million years in age and a large post glacial lava flow of uncertain origin are also found in the area surrounding Nazko

Volcanic Hazards

Potential hazards posed by a volcanic eruption in the Nazko region (Figure 6) include volcanic ash (tephra), lava flows, lahars (volcanic mudflows), gas emissions, and minor steam explosions. Flooding may occur due to diversion of streams by lava flows. The most serious of these hazards is volcanic ash, because it can severely damage aircraft and other machinery and is commonly transported long distances. However, the style of eruption is expected to be "Hawaiian" (Figure 6), a low energy eruption that produces localized lava flows and ash plumes only a few hundreds of metres in height.

The hazard maps shown here are based on the premise that the eruption will occur close to the centre of the earthquake swarm. Given the spatial coincidence of spasmodic tremor (detected during the event) and volcanotectonic earthquakes, this is the most likely scenario. Any escalation of earthquake activity, and any accompanying ground deformation or gas release will be closely monitored. Using all available tools, it should be possible to identify the eruption site more closely if the magma moves closer to the surface. When the eruption site is better known, new hazard maps and wind trajectories will be released based on the revised location.

Atmospheric winds have a strong influence on the movement of airborne volcanic ash and where ashfall is likely to take place if an eruption occurs. The circular patterns shown in Inset Map 1 show the distance that could be mpacted by heavy ashfall and ballistic particles. The actual pattern on the ground will be influenced by the wind at the time of the eruption as is shown for Nazko Cone on Inset Map 2. During the eruption 7,200 years ago, wind carried the ash mostly to the north and east.

Ninds can quickly change in time and space, and predicting correctly how situations, rapid wind shifts take place near a volcano and dramatically alter the areas affected by ash over the course of a few hours. It is therefore important that the forecasts of winds and trajectories be updated twice daily. Inset Map 2 gives 4 examples of how the wind might behave near the cone. The lines depict how ash could be carried at elevations of 300, 900, and 1200 m above the cinder cone. The heights are typical of Hawaiian style eruptions and are well below the level of jet aircraft (Figure

Tables A and B provides one year of wind strength data taken from the closest two weather stations to Nazko. The wind roses (Figures 9 and 11) show how often and how forcefully the wind blows in a given direction. The bar graphs (Figures 8 and 10) show the proportion of time the air is calm to windy. Note that the air is calm more than 60% of the time in Quesnel and more than 80% of the time at Puntzi Lake. These stations are a long way from the Nazko swarm, but they give an indication of what

Lava flows, fire fountaining (the ejection of lava into the air), and noxious ases pose a threat only in the immediate vicinity of a volcanic vent, although they could ignite forest fires that could potentially spread great distances. Inset Map 1 shows the areas most likely to be affected by lava flows. The extent of the lava flow was determined using an estimation of the volume of the nearby postglacial flow, and local topographic variation in the vicinity of the likely eruption site. This probably represents the centre, the area indicated by lava could vary significantly from that shown on the map. In addition, the interaction of hot volcanic materials with snow and loose wet sediment, or the disruption of rivers by volcanic materials, could cause lahars (volcanic mudflows) and flooding, both of which would affect areas downstream of the volcano tens of kilometres away. Small steam explosions and volcanic gas emissions (most commonly water vapour, carbon dioxide, or sulphur dioxide) could potentially occur in the vicinity of the vent.

Smith, K.D., Blewitt, G., von Seggern, D.H., Preston, L.A., Anderson, J.G., and Kilb, D., 2005. Sierran uplift and lower crustal earthquake swarm; evidence for magma injection beneath north Lake Tahoe, Nevada-California in late 2003. Seismological Society of America 2005 annual meeting, Incline Village, Nevada, April 27-29. Seismological Research Letters, vol. 76, no. 2, p. 245

Souther, J.G., 1986. The western Anahim belt: root zone of a peralkaline magma system. Canadian Journal of Earth Sciences, 23, 895-908

Souther, J.G., Clague, J.J., and Mathewes, R.W., 1987. Nazko cone: a Quaternary volcano in the eastern Anahim Belt. Canadian Journal of Earth Sciences, 24, 2477-2485



Figure 1. Seismic station near the epicentre of the 2007

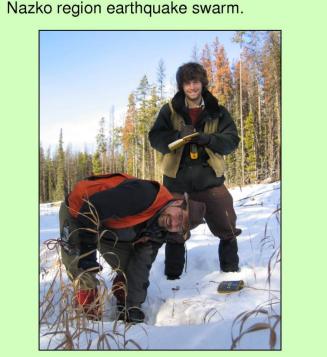


Figure 2. Geologists taking carbon dioxide measurements in the area of the Nazko region earthquake swarm, November 2007.

Recommendation citation:

2009: Nazko Region Volcanic Hazard Map

Universal Transerve Mercator Projection

North American Datum 1983

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Villiams-Jones, G.



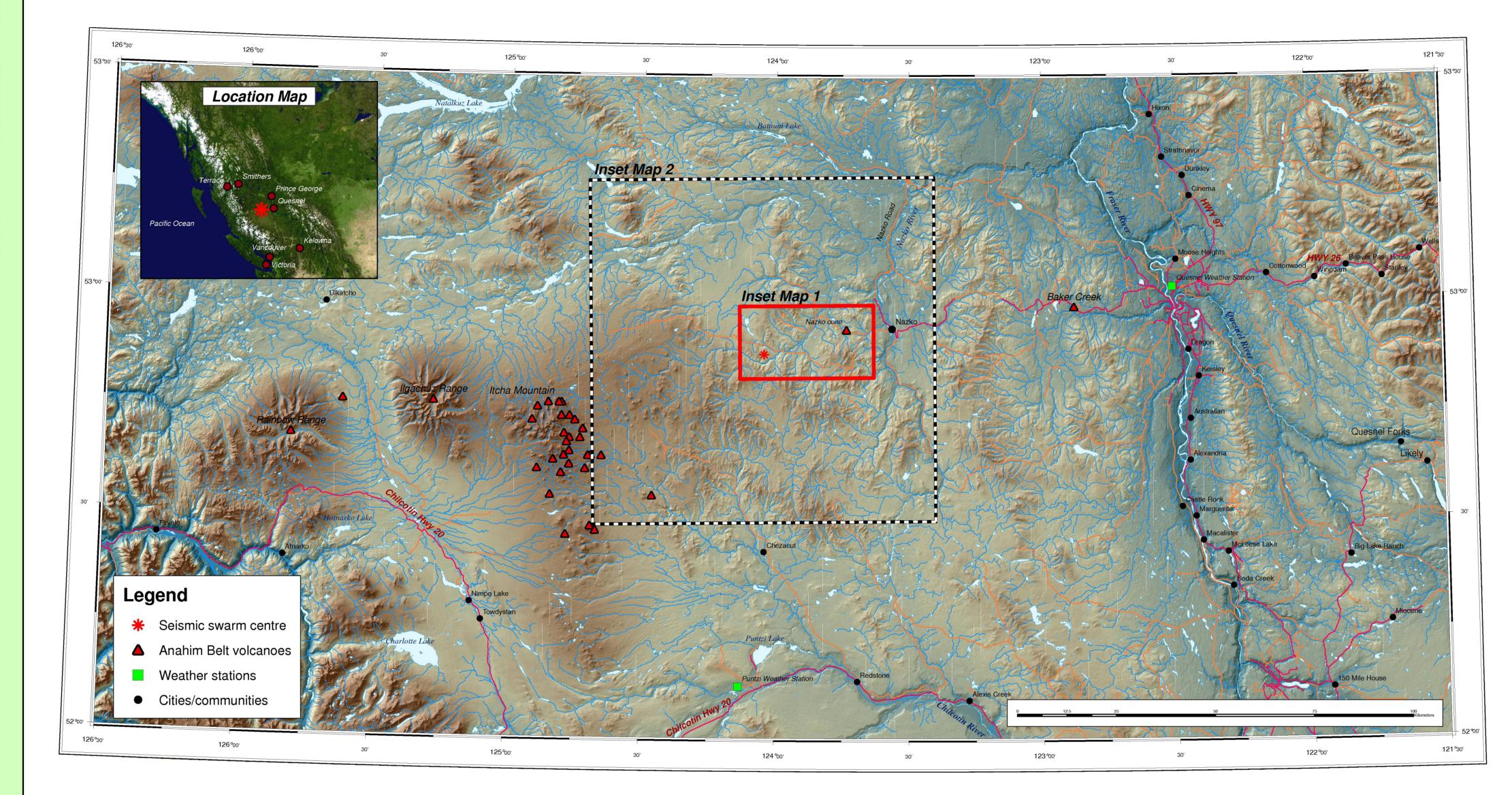
Figure 3. Nazko Cone, viewed from the north.

Figure 5. Mining of tephra at Nazko Cone.

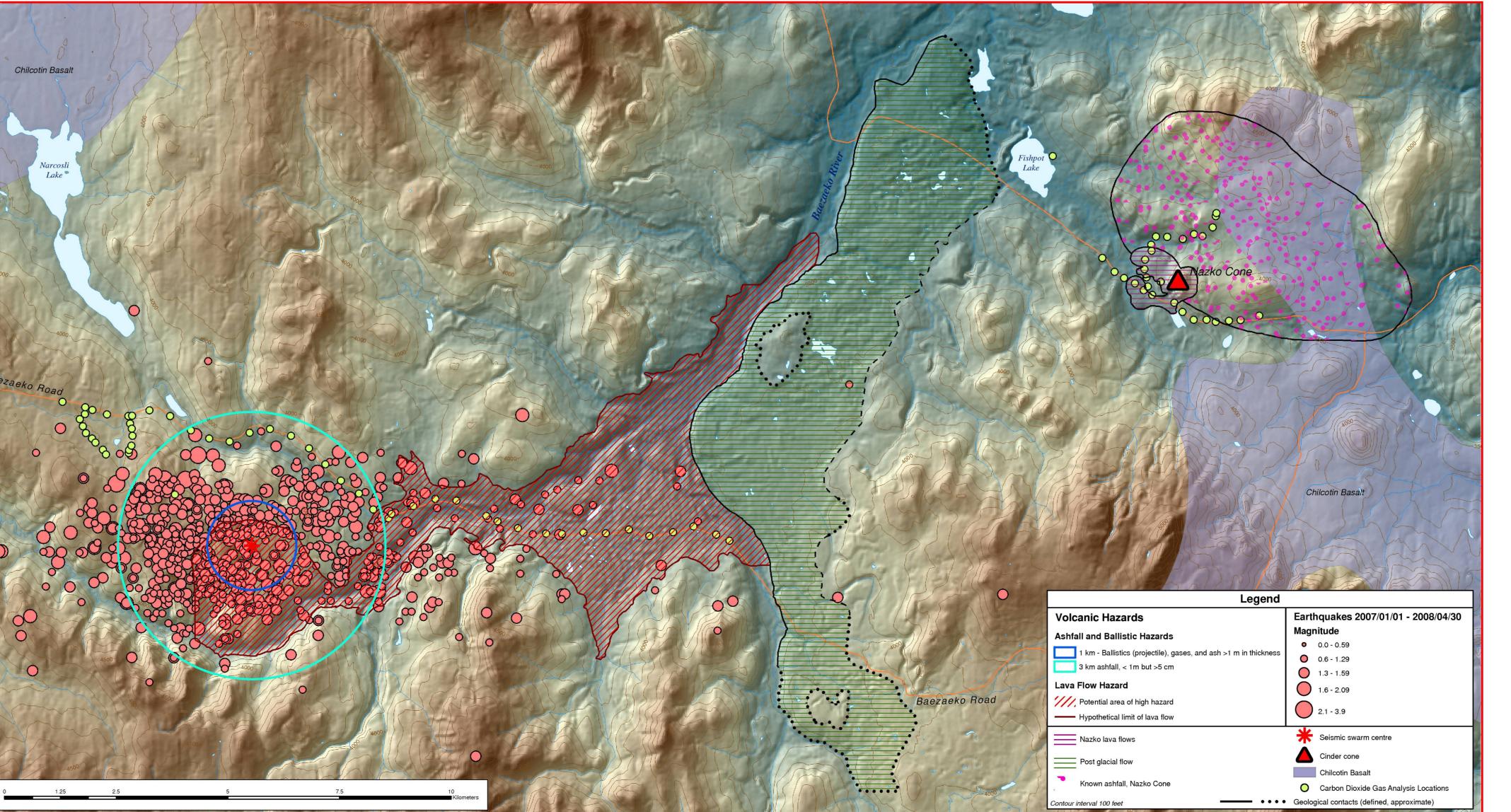
Nazko Region Volcanic Hazard Map

Hickson, C.J.¹, Kelman, M.C.¹, Chow, W.¹, Shimamura, K.¹, Servranckx, R.², Bensimon, D.², Cassidy, J.³, Trudel, S.², and Williams-Jones, G.⁴

- 1. Natural Resources Canada, GSC Vancouver, 625 Robson St, Vancouver, BC V6B 5J3
- 2. Canadian Meteorological Centre, Operations Division, 2121 Trans-Canada Hwy, Dorval, QC H9P 1J3
- 3. Natural Resources Canada, GSC Sidney, 9860 West Saanich Rd, North Saanich, BC V8L 4B2 4. Simon Fraser University, Department of Earth Sciences, 8888 University Dr, Burnaby, BC V5A 1S6



Inset Map 1: Close Up Region of the Recent Nazko Seismic Activity



Geological Survey of Canada, Open File 5978, scale 1: 650 000. Projection transverse universelle de Mercator © Sa Majesté la Reine du chef du Canada 2009

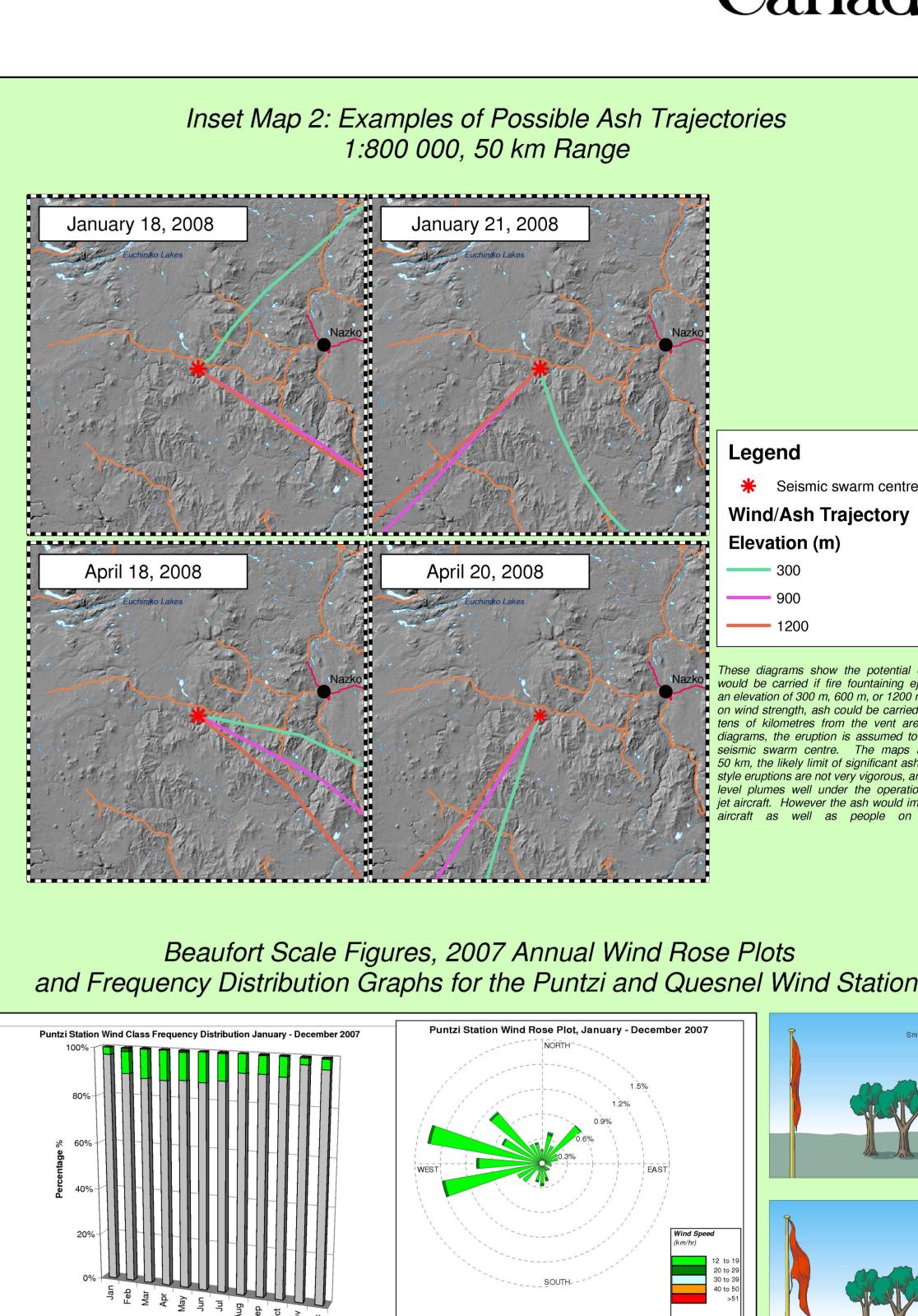
Digital base map from data compiled by Geomatics Canada Digital base map supplied by National Atlas Information Service (NAIS), Geomatics Canada, modified by ESS Info Mean Magnetic declination 2008, 19º 16' E, decreasing 15.2' annually Readings vary from 19º 24' E in the SW corner to 19º 60' E in the NE corner of the map Elevations in metres above mean sea level

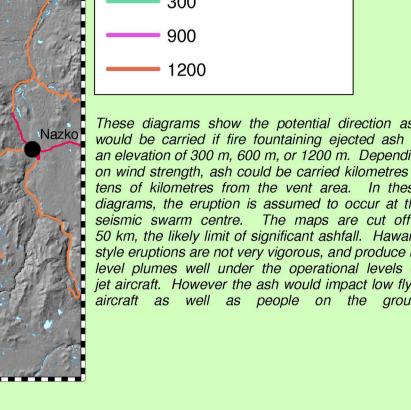
OPEN FILE Open files are products DOSSIER PUBLIC that have not gone through the GSC formal es dossiers publics son GEOLOGICAL SURVEY OF CANADA COMMISSION GÉOLOGIQUE DU CANADA des produits qui n'ont pas été soumis au 2009

Fresh Breeze 30-39

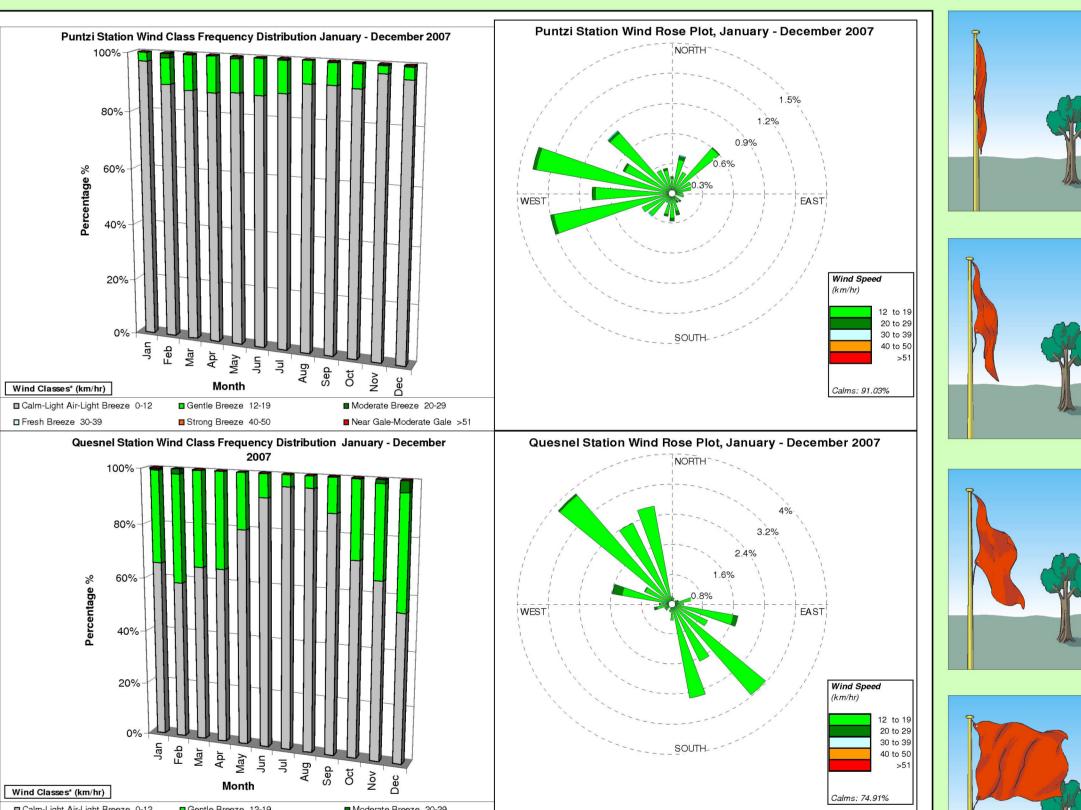
Strong Breeze 40-50

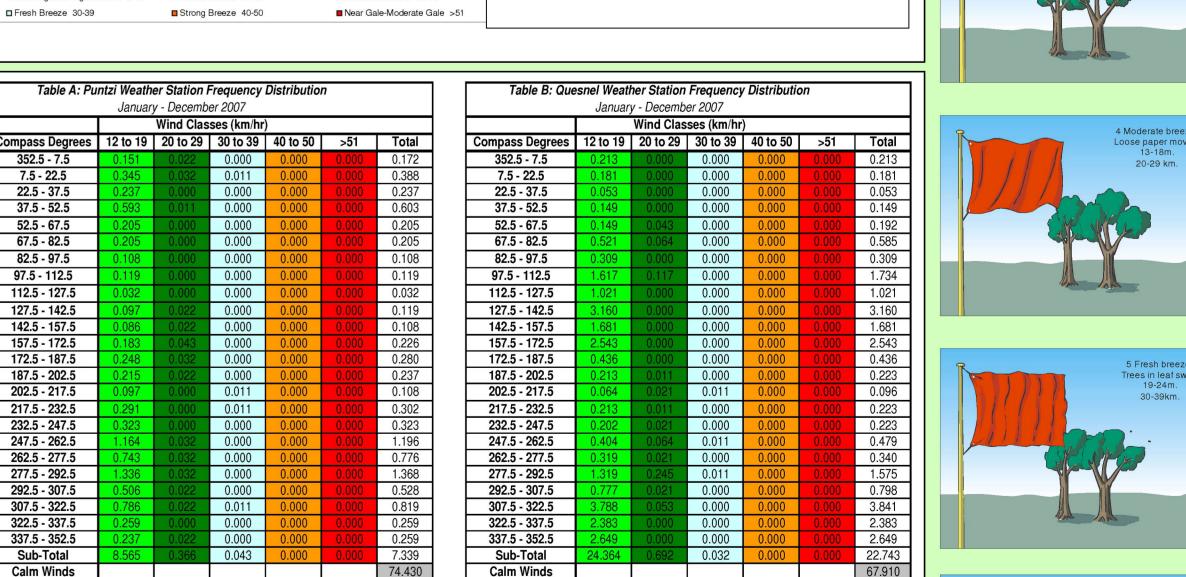
Near Gale-Moderate Gale >51

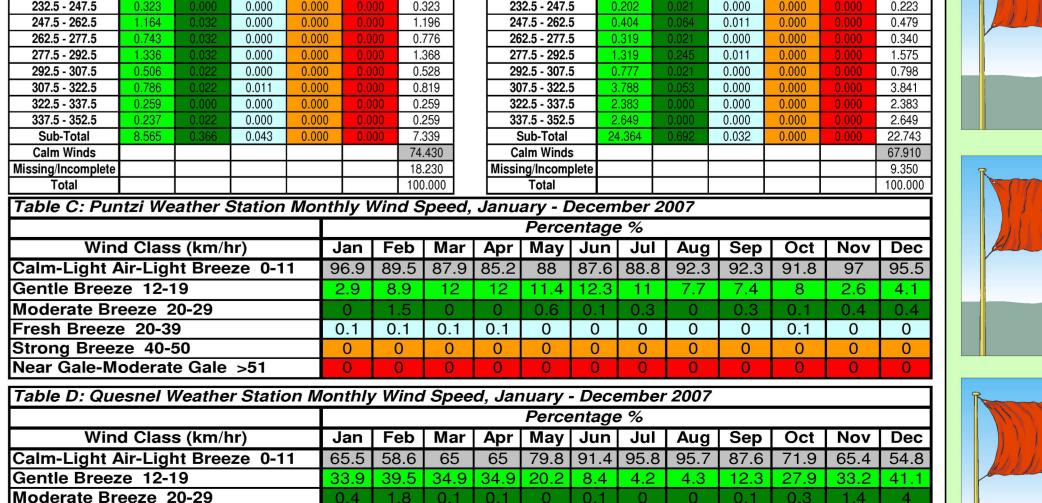




Beaufort Scale Figures, 2007 Annual Wind Rose Plots and Frequency Distribution Graphs for the Puntzi and Quesnel Wind Stations







* Wind classes and description based on the Beaufort Scale to give users a visual clue to wind strength. Calm winds (0-11 km/hr) are not shown on the wind rose plots. For more information on the Beaufort Scale please visit: http://www.metoffice.gov.uk/weather/marine/guide/beaufortscale.html