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Geological Survey of Canada

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New geochronological constraints from Mill, Salisbury, and Nottingham islands, Nunavut

N. Rayner, D.J. Scott, N. Wodicka, and A. Kassam

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Abstract: New U-Pb results refine our understanding of the area north of the exposed limit of the Superior Craton in eastern Nunavut. On Salisbury Island, gneissic tonalite was emplaced between 2983 and 2939 Ma. A crosscutting monzogranite dyke contains zircon grains dated at 1894 ± 7 and 1856 ± 3 Ma, and minor amounts of Archean inherited zircon. A late syenogranite dyke contains two age populations at 1843 ± 8 and 1804 ± 13 Ma.

On Mill Island, tonalite gneiss, crystallized at 2816 ± 31 Ma, contains 2.97 Ga inherited zircon grains and 1870 ± 5 Ma zircon grains. A crosscutting monzogranite contains 1.88 Ga zircon grains, a younger population at 1852 ± 5 Ma and evidence of a post-tectonic fluid event at 1.76 Ga.

Detrital zircon from Nottingham Island range in age from 2.3 to 1.85 Ga, with a prominent mode at ca. 1.90 Ga. The maximum age of deposition is constrained between 1.90 and 1.88 Ga. A similar provenance with the Lake Harbour Group and Tasiuyak paragneiss on Baffin Island is proposed.

Résumé : De nouveaux résultats de datation U-Pb viennent raffiner notre compréhension de la région située au nord de la limite d'affleurement du craton du lac Supérieur dans l'est du Nunavut. Dans l'île Salisbury, nous évaluons l'âge de mise en place de la tonalite gneissique entre 2983 et 2939 Ma. Un dyke de monzogranite recoupant l'unité précédente contient des zircons datés à 1894 ± 7 Ma et à 1856 ± 3 Ma, ainsi que de faibles quantités de zircons hérités de l'Archéen. Un dyke tardif de syénogranite contient deux populations âgées respectivement de 1843 ± 8 et de 1804 ± 13 Ma.

Dans l'île Mill, le gneiss tonalitique, qui a livré un âge de cristallisation de 2816 ± 31 Ma, contient des zircons hérités de 2,97 Ga ainsi que des zircons de 1870 ± 5 Ma. Un monzogranite qui recoupe le gneiss renferme des zircons datant de 1,88 Ga, une population plus récente datant de 1852 ± 5 Ma et des indices d'un épisode post-tectonique de circulation de fluides remontant à 1,76 Ga.

Les âges des zircons détritiques de l'île Nottingham varient de 2,3 à 1,85 Ga, avec une prédominance à environ 1,90 Ga. L'âge maximal de dépôt se situe entre 1,90 et 1,88 Ga. Nous émettons l'hypothèse d'une provenance similaire à celle déduite pour le Groupe de Lake Harbour et le paragneiss de Tasiuyak dans l'île de Baffin.

INTRODUCTION

Mill, Salisbury, and Nottingham islands are located in the Hudson Strait south of Cape Dorset (Kingnait) and were mapped by Blackadar in 1964 (Blackadar, 1970) (Fig. 1). In 1994 the islands were briefly visited by D.J. Scott for the purpose of sampling the dominant lithologies for U-Pb geochronology; the results of which have a bearing on our understanding of the pre- and syn-collisional history of the eastern Trans-Hudson Orogen (THO). Determination of basement ages may assist in assigning an affinity to one or the other of the bounding cratons (Superior or Meta Incognita; Fig. 1). Paleoproterozoic ages, such as magmatic crystallization or metamorphic overprint, help constrain the tectonic evolution of the major components of the THO and the location of the trace of major crustal boundaries, in particular the ca. 1845 Ma Soper River suture and the 1820 to 1795 Ma Bergeron suture (e.g. St-Onge et al., 2002, 2006b). The south-verging Bergeron suture separates the lower-plate Superior Craton from allochthonous crustal elements

(e.g. Narsajuaq terrane and Meta Incognita microcontinent) of the upper Churchill plate to the north. The Soper River Suture separates Narsajuaq arc from the Meta Incognita microcontinent (St-Onge et al., 1999, 2001, 2002). Recent compilation maps (St-Onge et al., 2006a) ascribe the islands to the Superior Craton in the immediate footwall of the Bergeron suture (Fig. 1). The association of Mill, Salisbury, and Nottingham islands with the Superior Craton is largely based on their position relative to Big Island, for which U-Pb geochronological results indicate a Superior affinity (Wodicka and Scott, 1997; Scott et al., 2002). New geochronological data presented here allows this hypothesis to be tested.

GEOLOGICAL BACKGROUND

The geology of Mill, Salisbury, and Nottingham islands has been subdivided into three main lithological units originally ascribed to the Paleoproterozoic (Blackadar, 1970),

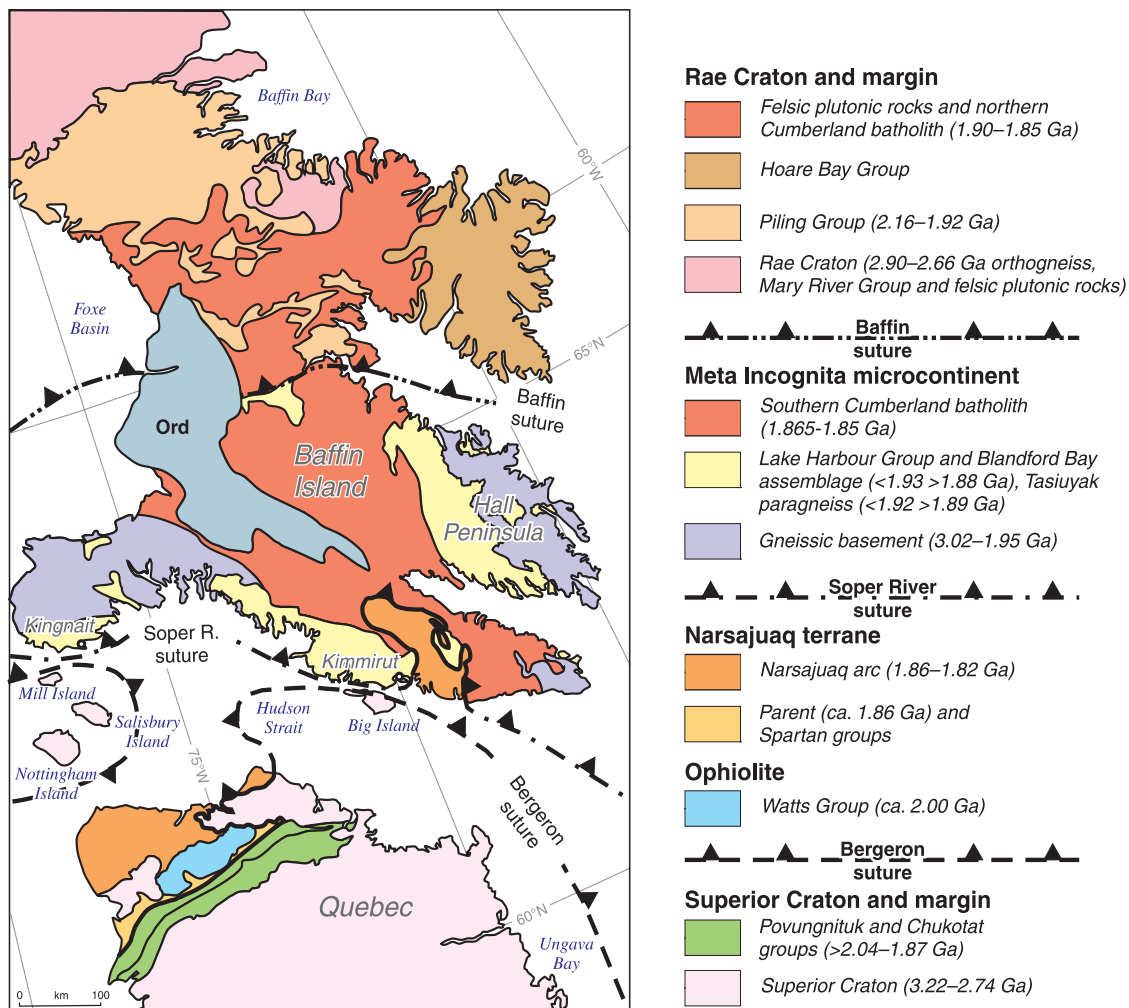


Figure 1. Regional geological map for central-southern Baffin Island and surroundings (adapted from St-Onge et al., 2006a). Major tectonostratigraphic assemblages and bounding crustal sutures are after St-Onge et al. (2006b). Ord, Ordovician cover.

but more recently they have been ascribed to the Archean (St-Onge et al., 2006a). Unit 1 is a quartz-feldspar gneiss composing the majority of the map area and is further divided in some areas into 5 subunits all containing variable amounts of garnet. Unit 2 is a hornblende-quartz-feldspar gneiss found mainly as localized domains on Mill Island; Blackadar (1970) suggested that it may be related to unit 1e (mafic-rich gneiss). Unit 3 consists of diabase dykes that occur as map-scale units on Mill Island and as several narrow bodies on Salisbury Island. None are reported on Nottingham Island.

Salisbury Island is largely mapped as granular gneiss or veined gneiss (Blackadar, 1970). Three samples were collected from one outcrop on this island on the basis of their relative ages. Sample SAL-A is a tonalite gneiss and represents the oldest recognized phase. SAL-B is a monzogranite cutting the fabric of the tonalite. SAL-C, a crosscutting syenogranite, is the youngest observed phase. Mill Island is mapped as veined gneiss (unit 1c) with two small enclaves of hornblende-quartz-feldspar gneiss (unit 2, Blackadar, 1970), one of which is within 2.5 km of the sampling site. A suite of three samples, similar to those from Salisbury Island, was collected from Mill Island. The oldest phase, MIL-A, is a tonalite gneiss. MIL-B is intermediate in relative age and monzogranitic in composition. A younger syenogranite phase was collected, but has not been dated. All five subunits of quartz-feldspar gneiss (Unit 1) have been observed on Nottingham Island (Blackadar, 1970) suggesting that it is more lithologically diverse than Mill or Salisbury Islands. Sample NOT-A was identified as a metasedimentary gneiss during the 1994 visit.

ANALYTICAL METHODS

All crushing and mineral separation was performed at the GEOTOP Laboratory at the Université du Québec à Montréal using standard crushing, heavy-liquid, and magnetic-separation techniques. Mineral selection and analytical work were performed at the Geological Survey of Canada Geochronology Laboratory in Ottawa, Ontario. Zircon grains were dated using the Sensitive High-Resolution Ion Micro Probe (SHRIMP II). SHRIMP analytical procedures followed those described by Stern (1997), with standards and U-Pb calibration methods following Stern and Amelin (2003). SHRIMP analytical and calibration details can be found in the footnotes to Table 1. The internal features of the zircon grains, such as zoning, structures, and alteration, were characterized with back-scattered electrons (BSE) using a Zeiss Evo 50 scanning electron microscope. Monazite and titanite were dated using the isotope dilution – thermal ionization mass spectrometry (ID-TIMS) technique. ID-TIMS analytical procedures for titanite followed those of Parrish et al. (1992) and Davis et al. (1997). Procedures for monazite ID-TIMS analyses followed the sample dissolution and chemical methods as described in Parrish et al. (1987) and Davis et al. (1998).

U-Pb RESULTS

SHRIMP data are presented in Table 1 with age errors reported at the 1σ uncertainty level. TIMS data are presented in Table 2 with age errors reported at 2σ uncertainty level. Isoplot v. 3.00 (Ludwig, 2003) was used to generate concordia plots (error ellipses 2σ). Errors quoted in the text are at the 95% confidence level unless otherwise noted.

SAL-A (GSC lab # 4111)

Twenty-eight analyses were carried out on 19 separate zircon grains from tonalite gneiss sample SAL-A and the data are presented in Figure 2a and Table 1. The data are arrayed in two clusters, as well as an indistinct smear of non-reproducible results. The six oldest analyses yield a weighted mean age of 2983 ± 8 Ma (unfilled ellipses, $n = 6$, MSWD = 1.6, probability of fit [POF] = 0.16). This is statistically distinct from a younger group at 2939 ± 8 Ma (light grey ellipses, $n = 9$, MSWD = 1.6, POF = 0.12 Ma). The remaining analyses not included in the calculation of the weighted mean (dark grey ellipses) include non-reproducible replicate analyses within individual zircon grains indicating that this sample has experienced ancient Pb-loss. These data can be interpreted as either crystallization at 2983 Ma with variable Pb-loss, or 2983 Ma inherited zircon with tonalite crystallization at 2939 Ma. The predominance of the 2939 Ma zircon population (*see also* SAL-B results) makes the latter the preferred interpretation. Two single-grain fractions of monazite give TIMS ages of 1862 and 1867 Ma (Fig. 2b, Table 2).

SAL-B (GSC lab # 4112)

Thirty-one analyses were carried out on individual zircon grains from monzogranite sample SAL-B (Fig. 2c). The four oldest analyses give a weighted mean age of 2947 ± 25 Ma (MSWD = 1.9, POF = 0.13). This age is identical within error to that of the younger population of zircon grains from SAL-A, and likely represents inheritance from that source. A subset of clear, colourless, low-U zircon gives a weighted mean age of 1894 ± 7 Ma (unfilled ellipses, $n = 12$, MSWD = 0.87, POF = 0.5), while pale brown, high-U zircon grains give a weighted mean age of 1856 ± 3 Ma (light grey ellipses, $n = 14$, MSWD = 0.94, POF = 0.5). These two ages are statistically distinct and suggest that the younger age cannot be the result of Pb-loss from the older, higher U population. One single grain monazite fraction analyzed by TIMS gives an age of 1875 ± 1 Ma (Fig. 2b).

SAL-C (GSC lab # 4113)

Thirty analyses were carried out on twenty-nine zircon grains from syenogranite sample SAL-C (Fig. 2d). The three oldest analyses yield ages between 2.65 and 1.98 Ga and are between 7 and 10% discordant. These analyses are

Table 1. U-Pb SHRIMP analytical data.

Spot name	U (ppm)	Th (ppm)	Th/U	Pb* (ppm)	204Pb/206Pb	±206Pb/206Pb	f(206)	208Pb/206Pb	±208Pb/206Pb	207Pb/206Pb	±207Pb/206Pb	206Pb/238U	±206Pb/238U	207Pb/206Pb	±207Pb/206Pb	Apparent ages (Ma)						
																206Pb/238U	207Pb/206Pb					
SAL-A (Z4111; 63.3831°N 76.8458°E)																						
4111-33.2	579	65	0.12	306	3	1.23E-05	5.58E-06	0.00021	0.02821	0.00067	13.334	0.182	0.5027	0.0061	0.9532	0.1924	0.0010	2625	26	2763	8	5.0
4111-35.1	505	157	0.32	314	4	1.52E-05	6.28E-06	0.00026	0.08381	0.00077	15.699	0.189	0.5605	0.0060	0.9314	0.2031	0.0009	2869	25	2852	7	-0.6
4111-15.1	700	126	0.19	412	3	1.00E-05	1.00E-05	0.00017	0.04896	0.00054	15.477	0.208	0.5454	0.0068	0.9604	0.2058	0.0008	2806	28	2873	6	2.3
4111-35.2	466	236	0.52	270	2	9.12E-06	4.27E-06	0.00016	0.13545	0.00096	14.341	0.180	0.5010	0.0058	0.9572	0.2076	0.0008	2618	25	2887	6	9.3
4111-20.2	347	134	0.38	227	2	1.00E-05	1.00E-05	0.00017	0.10456	0.00079	15.666	0.176	0.5447	0.0057	0.9566	0.2006	0.0007	2803	24	2895	5	3.2
4111-20.2	370	262	0.84	224	3	1.70E-05	1.07E-05	0.00029	0.22979	0.00158	15.079	0.243	0.5208	0.0077	0.9566	0.2100	0.0010	2702	33	2906	8	7.0
4111-33.1	562	76	0.14	327	4	1.47E-05	5.60E-06	0.00025	0.03890	0.00065	15.665	0.191	0.5409	0.0058	0.9290	0.2101	0.0010	2787	24	2906	7	4.1
4111-73.1	158	87	0.57	102	1	1.00E-05	1.00E-05	0.00017	0.15102	0.00176	16.007	0.231	0.5510	0.0064	0.8642	0.2107	0.0015	2829	27	2911	12	2.8
4111-1	241	168	0.72	162	0	4.06E-06	1.67E-05	0.00007	0.19356	0.00246	16.145	0.212	0.5540	0.0065	0.9402	0.2114	0.0010	2842	27	2916	7	2.5
4111-54.1	485	107	0.24	277	2	8.96E-06	8.71E-06	0.00016	0.06121	0.00058	15.920	0.191	0.5446	0.0057	0.9192	0.2120	0.0010	2803	24	2921	8	4.0
4111-10.1	116	68	0.60	77	1	1.78E-05	2.80E-05	0.00031	0.16189	0.00232	16.318	0.291	0.5570	0.0075	0.9104	0.2125	0.0021	2854	31	2924	16	2.4
4111-80.1	210	154	0.76	144	1	1.34E-05	8.92E-06	0.00023	0.20468	0.00228	16.472	0.256	0.5603	0.0075	0.9104	0.2132	0.0014	2868	31	2930	11	2.1
4111-15.2	383	220	0.59	248	0	2.27E-06	4.71E-06	0.00004	0.16305	0.00107	16.093	0.249	0.5469	0.0082	0.9853	0.2134	0.0006	2812	34	2932	4	4.1
4111-18.2	491	104	0.22	300	1	3.68E-06	4.22E-06	0.00006	0.05706	0.00053	16.516	0.236	0.5589	0.0071	0.9500	0.2143	0.0011	2862	29	2938	9	2.6
4111-44.1	609	189	0.32	392	2	5.51E-06	6.95E-06	0.00010	0.08415	0.00087	17.052	0.217	0.5760	0.0064	0.9222	0.2147	0.0011	2933	26	2941	8	0.3
4111-11.1	157	84	0.55	105	0	7.00E-08	1.60E-05	0.00000	0.15226	0.00210	16.743	0.220	0.5656	0.0065	0.9196	0.2147	0.0011	2890	27	2941	8	1.8
4111-63.1	149	70	0.48	96	1	1.97E-05	1.78E-05	0.00034	0.13334	0.00171	16.522	0.278	0.5570	0.0074	0.8530	0.2151	0.0019	2854	31	2945	14	3.1
4111-43.1	1733	28	0.02	558	1	1.17E-06	3.99E-06	0.00002	0.00477	0.00017	5.187	0.055	0.3351	0.0034	0.9790	0.1123	0.0003	1863	16	1836	4	-1.5
4112-58.1	725	249	0.35	246	4	1.91E-05	1.11E-05	0.00033	0.10248	0.00085	5.039	0.058	0.3246	0.0034	0.9543	0.1126	0.0004	1812	17	1842	6	1.6
4112-25.1	790	449	0.59	287	2	1.00E-05	1.00E-05	0.00017	0.17174	0.00096	5.120	0.057	0.3289	0.0034	0.9589	0.1129	0.0004	1833	17	1847	6	0.7
4112-23.1	664	377	0.59	234	1	7.64E-06	1.80E-05	0.00013	0.17135	0.00175	4.967	0.067	0.3189	0.0037	0.9174	0.1130	0.0006	1784	18	1848	10	3.4
4112-65.1	832	372	0.46	294	1	2.95E-06	8.76E-06	0.00005	0.13328	0.00084	5.150	0.062	0.3297	0.0035	0.9330	0.1133	0.0005	1837	17	1853	8	0.9
4112-62.1	1073	321	0.31	360	4	1.31E-05	7.53E-06	0.00023	0.08764	0.00068	5.072	0.064	0.3244	0.0034	0.8993	0.1134	0.0006	1811	17	1855	10	2.3
4112-34.1	613	257	0.43	217	3	1.57E-05	1.65E-05	0.00027	0.12228	0.00115	5.219	0.070	0.3335	0.0041	0.9553	0.1135	0.0005	1855	20	1856	7	0.1
4112-86.1	1387	990	0.74	524	1	2.40E-06	4.34E-06	0.00004	0.21578	0.00079	5.164	0.056	0.3299	0.0034	0.9771	0.1135	0.0003	1838	17	1857	4	1.0
4112-24.1	937	500	0.55	338	3	1.22E-05	1.12E-05	0.00021	0.16214	0.00087	5.152	0.058	0.3288	0.0034	0.9546	0.1136	0.0004	1833	17	1858	6	1.4
4112-33.1	488	934	1.98	231	1	8.51E-06	1.39E-05	0.00015	0.57555	0.00282	5.083	0.065	0.3243	0.0036	0.9186	0.1137	0.0006	1811	18	1859	9	2.6
4112-10.1	708	189	0.28	239	0	2.80E-07	7.69E-06	0.00000	0.07949	0.00080	5.166	0.060	0.3295	0.0034	0.9424	0.1137	0.0004	1836	17	1860	7	1.3
4112-32.1	804	313	0.40	281	1	4.67E-06	1.06E-05	0.00008	0.11706	0.00098	5.180	0.058	0.3303	0.0034	0.9590	0.1137	0.0004	1840	17	1860	6	1.1
4112-17.1	475	727	1.58	211	0	3.68E-06	2.33E-05	0.00006	0.46790	0.00241	5.120	0.065	0.3264	0.0036	0.9139	0.1138	0.0006	1821	18	1860	9	2.1
4112-61.1	1275	1032	0.84	491	0	2.70E-07	4.88E-06	0.00000	0.24480	0.00090	5.169	0.056	0.3295	0.0034	0.9740	0.1138	0.0003	1836	16	1861	4	1.3
4112-51.1	947	403	0.44	333	3	1.17E-05	6.21E-06	0.00020	0.12687	0.00070	5.179	0.074	0.3298	0.0045	0.9778	0.1139	0.0003	1837	22	1862	5	1.3
4112-4.1	97	114	1.21	40	2	5.88E-05	8.85E-05	0.00102	0.33779	0.00513	5.235	0.108	0.3326	0.0044	0.7214	0.1142	0.0017	1851	21	1867	26	0.8
4112-46.1	84	178	0.60	113	4	4.15E-05	1.91E-05	0.00072	0.17289	0.00162	5.288	0.082	0.3344	0.0038	0.8018	0.1147	0.0011	1847	18	1875	17	0.8
4112-79.1	88	73	0.85	34	3	1.34E-04	7.83E-05	0.00232	0.24487	0.00510	5.253	0.108	0.3319	0.0039	0.6651	0.1148	0.0018	1847	19	1877	28	1.6
4112-26.1	158	300	1.96	75	1	2.75E-05	4.11E-05	0.00048	0.55472	0.00393	5.232	0.072	0.3296	0.0036	0.8544	0.1151	0.0008	1836	17	1882	13	2.4
4112-11.1	469	253	0.56	173	1	7.62E-06	1.44E-05	0.00013	0.16070	0.00156	5.334	0.072	0.3358	0.0038	0.8923	0.1152	0.0007	1866	18	1883	11	0.9
4112-2.1	92	271	3.04	51	2	7.10E-05	8.55E-05	0.00123	0.66909	0.00734	5.190	0.105	0.3266	0.0042	0.7222	0.1153	0.0016	1822	20	1884	26	3.3
4112-63.1	84	322	3.97	54	0	1.00E-05	1.00E-05	0.00017	1.15512	0.01297	5.289	0.089	0.3321	0.0043	0.8351	0.1155	0.0011	1849	21	1888	17	2.1

Notes: (See Stem. 1997):
 Analytical details: ion probe mount IP394, spot size = 35µm, beam current = 10nA, 6 scans, no Pb fractionation correction
 Spot name follows the convention x-y-z; where x = sample number, y = grain number and z = spot number. Multiple analyses in an individual spot are labelled as x-y.z.z
 Uncertainties reported at 1σ (absolute) and are calculated by numerical propagation of all known sources of error
 f206* refers to mole fraction of total 206Pb that is due to common Pb, calculated using the 204Pb-method; common Pb composition used is the surface blank (4/6; 0.05770; 7/6; 0.89500; 8/6; 2.13840)
 * refers to radiogenic Pb (corrected for common Pb)
 Discordance relative to origin = 100 * (1 - (206Pb/206Pb) / (206Pb/206Pb))
 Calibration standard 6266; U = 910 ppm; Age = 559 Ma; 206Pb/238U = 0.09059
 Error in 206Pb/238U calibration 1.0%
 Th/U calibration: F = 0.03900 * UO + 0.85600

Table 1. (cont.)

Spot name	U (ppm)	Th (ppm)	Th (ppm)	Pb* (ppm)	204Pb (ppb)	206Pb/208Pb	\pm 204Pb/208Pb	t(206)	208Pb/206Pb	\pm 208Pb/206Pb	207Pb/238U	207Pb/235U	206Pb/238U	207Pb/206Pb	208Pb/206Pb	207Pb/206Pb	Apparent ages (Ma)		Disc. (%)		
																	206Pb/238U	207Pb/206Pb			
																	206Pb/238U	207Pb/206Pb			
4112-59.1	770	1036	1.39	339	5	2.09E-05	1.60E-05	0.00036	0.40768	0.00179	5.355	0.064	0.3352	0.0037	0.9478	0.1159	0.0005	18	1894	7	1.6
4112-11.1	206	408	2.04	100	2	3.69E-05	3.05E-05	0.00064	0.60001	0.00556	5.241	0.073	0.3279	0.0037	0.8670	0.1159	0.0008	18	1894	13	3.5
4112-85.1	102	316	3.19	59	1	2.31E-05	4.78E-05	0.00040	0.91368	0.01205	5.234	0.088	0.3269	0.0039	0.7898	0.1161	0.0012	18	1897	19	3.9
4112-64.1	158	315	2.06	78	2	4.84E-05	5.17E-05	0.00084	0.91962	0.00371	5.401	0.082	0.3363	0.0039	0.8272	0.1165	0.0010	18	1903	16	1.8
4112-50.1	172	95	0.57	62	0	1.00E-05	1.00E-05	0.00017	0.16580	0.00355	5.242	0.068	0.3247	0.0038	0.9323	0.1171	0.0006	18	1912	9	5.2
4112-28.1	244	186	0.79	164	3	2.38E-05	1.29E-05	0.00041	0.22005	0.00227	16.002	0.259	0.5455	0.0072	0.8708	0.2128	0.0017	30	2927	13	4.1
4112-8.1	73	59	0.84	48	6	1.75E-04	9.58E-05	0.00303	0.22966	0.00436	15.636	0.262	0.5322	0.0066	0.8165	0.2131	0.0021	28	2929	16	6.1
4112-19.1	122	135	1.15	92	0	8.90E-07	2.41E-05	0.00002	0.31848	0.00474	17.139	0.302	0.5746	0.0081	0.8602	0.2163	0.0020	29	2954	15	0.9
4112-20.1	166	171	1.07	120	1	8.06E-06	3.00E-05	0.00014	0.29940	0.00400	16.625	0.229	0.5561	0.0069	0.9392	0.2168	0.0010	29	2957	8	3.6
SAL-C (Z4113; 63.3831°N 76.8459°E)																					
4113-18.1	719	78	0.11	219	11	5.64E-05	1.35E-05	0.00098	0.02948	0.00063	4.594	0.070	0.3116	0.0043	0.9471	0.1069	0.0005	17	1748	9	-0.1
4113-7.1	258	93	0.37	89	46	6.22E-04	7.00E-05	0.01078	0.10837	0.00405	4.913	0.115	0.3320	0.0054	0.7723	0.1074	0.0016	18	1755	28	-5.3
4113-47.1	419	78	0.19	137	63	5.32E-04	9.64E-05	0.00922	0.05564	0.00367	4.875	0.097	0.3255	0.0038	0.6795	0.1086	0.0016	19	1776	27	-2.3
4113-14.1	991	93	0.10	312	23	8.52E-05	1.43E-05	0.00148	0.02672	0.00062	4.840	0.064	0.3222	0.0038	0.9352	0.1090	0.0005	18	1782	9	-1.0
4113-22.1	677	74	0.11	216	14	7.69E-05	1.82E-05	0.00133	0.03563	0.00086	4.908	0.060	0.3248	0.0036	0.9411	0.1096	0.0005	18	1793	8	-1.2
4113-35.1	1034	100	0.10	328	2	7.90E-06	7.28E-06	0.00014	0.02925	0.00073	4.919	0.060	0.3233	0.0036	0.9523	0.1104	0.0004	18	1805	7	0.0
4113-22.2	765	453	0.61	279	11	5.01E-05	1.03E-05	0.00087	0.17499	0.00096	5.026	0.060	0.3298	0.0035	0.9225	0.1106	0.0005	17	1808	9	-1.6
4113-33.1	467	322	0.72	172	60	4.66E-04	5.21E-05	0.00908	0.23011	0.00242	4.946	0.072	0.3241	0.0034	0.8009	0.1107	0.0010	17	1810	16	0.0
4113-17.1	450	727	1.64	193	18	1.43E-04	3.11E-05	0.00247	0.45082	0.00316	4.803	0.066	0.3146	0.0036	0.8852	0.1107	0.0007	17	1812	12	2.7
4113-19.1	179	63	0.36	60	2	3.75E-05	2.43E-05	0.00065	0.10389	0.00190	4.923	0.074	0.3216	0.0039	0.8590	0.1110	0.0009	17	1817	14	1.1
4113-1.1	784	1285	1.69	348	1	5.33E-06	8.00E-06	0.00009	0.49143	0.00277	4.923	0.057	0.3211	0.0034	0.9556	0.1112	0.0004	17	1819	6	1.3
4113-57.1	836	512	0.63	301	2	1.02E-05	7.96E-06	0.00018	0.18228	0.00172	4.961	0.058	0.3231	0.0035	0.9533	0.1114	0.0004	18	1822	7	0.9
4113-2.1	1018	89	0.09	319	6	2.02E-05	1.09E-05	0.00035	0.02529	0.00049	4.938	0.061	0.3208	0.0033	0.8926	0.1116	0.0006	16	1826	10	1.8
4113-52.1	418	461	1.14	166	73	6.45E-04	8.63E-05	0.01178	0.37751	0.00391	4.783	0.085	0.3107	0.0034	0.6940	0.1116	0.0015	17	1826	24	4.5
4113-46.1	252	195	0.80	92	3	4.48E-05	2.26E-05	0.00018	0.23078	0.00032	4.874	0.082	0.3166	0.0041	0.8438	0.1117	0.0010	20	1827	17	3.0
4113-6.1	44	46	1.07	17	1	1.01E-04	2.23E-04	0.00175	0.30847	0.01163	4.951	0.249	0.3215	0.0083	0.6101	0.1117	0.0045	17	1827	75	1.7
4113-16.1	254	109	0.50	79	2	2.76E-05	4.45E-05	0.00048	0.14671	0.00052	5.042	0.107	0.3262	0.0046	0.7461	0.1121	0.0016	18	1834	26	0.8
4113-34.1	55	71	1.35	23	1	5.11E-05	1.36E-04	0.00088	0.38677	0.00870	4.974	0.160	0.3216	0.0046	0.5473	0.1122	0.0031	22	1835	50	2.0
4113-4.1	258	263	1.05	99	1	1.00E-05	1.00E-05	0.00017	0.30350	0.00278	4.889	0.064	0.3159	0.0036	0.9141	0.1122	0.0006	17	1836	10	3.6
4113-30.1	852	151	0.18	278	0	2.70E-07	1.07E-05	0.00000	0.05329	0.00060	5.052	0.060	0.3261	0.0035	0.9440	0.1124	0.0004	18	1838	7	1.0
4113-64.1	569	257	0.47	193	4	2.68E-05	1.48E-05	0.00046	0.13203	0.00113	4.913	0.056	0.3168	0.0033	0.9380	0.1125	0.0005	17	1840	7	3.6
4113-8.1	162	88	0.56	57	1	2.67E-05	2.90E-05	0.00046	0.16049	0.00218	4.963	0.081	0.3197	0.0039	0.8191	0.1126	0.0011	17	1842	17	2.9
4113-59.1	37	44	1.22	15	0	3.47E-05	2.66E-04	0.00030	0.34452	0.01308	5.045	0.231	0.3242	0.0059	0.5072	0.1129	0.0045	19	1846	74	2.0
4113-42.1	60	60	1.03	24	0	1.72E-05	9.80E-05	0.00060	0.29708	0.00686	5.118	0.159	0.3266	0.0064	0.7172	0.1137	0.0025	18	1859	40	2.0
4113-20.1	82	116	1.46	35	0	1.66E-05	4.85E-05	0.00029	0.43151	0.00642	5.009	0.116	0.3191	0.0051	0.7680	0.1138	0.0017	25	1862	27	4.1
4113-11.1	160	104	0.67	58	0	1.00E-05	1.00E-05	0.00017	0.19818	0.00198	5.102	0.071	0.3223	0.0037	0.8769	0.1148	0.0008	18	1877	12	4.1
4113-37.1	93	55	0.61	33	0	1.55E-05	4.31E-05	0.00027	0.18298	0.00447	5.109	0.139	0.3211	0.0060	0.7661	0.1154	0.0020	29	1886	32	4.9
4113-31.1	779	121	0.16	262	82	3.67E-04	3.42E-05	0.00636	0.05565	0.00142	5.585	0.073	0.3324	0.0036	0.8842	0.1218	0.0008	17	1983	11	6.7
4113-28.1	171	110	0.66	67	2	4.36E-05	3.37E-05	0.00076	0.18948	0.00410	6.179	0.115	0.3429	0.0053	0.8925	0.1307	0.0011	26	2107	15	9.8
4113-26.1	486	102	0.22	240	31	1.60E-04	3.90E-05	0.00277	0.06116	0.00189	11.447	0.144	0.4623	0.0052	0.9319	0.1796	0.0008	23	2649	23	7.5
MIL-A (Z4114; 64.0047°N 77.8900°E)																					
4114-15.1	21	1	0.03	7	2	3.79E-04	1.77E-04	0.00647	-0.00147	0.00695	5.144	0.164	0.3290	0.0047	0.5487	0.1134	0.0031	23	1855	8	1.1
4114-11.1	2195	143	0.07	728	2	3.36E-06	5.59E-06	0.00006	0.01888	0.00029	5.351	0.062	0.3405	0.0037	0.9642	0.1140	0.0004	18	1864	6	-1.4
4114-8.1	578	32	0.06	180	6	3.95E-05	1.28E-05	0.00068	0.00843	0.00072	5.073	0.077	0.3228	0.0040	0.8857	0.1140	0.0008	20	1864	13	3.2
4114-28.1	4514	540	0.12	1581	1	8.70E-07	8.60E-07	0.00002	0.03518	0.00018	5.974	0.071	0.3545	0.0043	0.9985	0.1141	0.0002	21	1866	4	-4.8
4114-29.1	635	21	0.03	199	5	2.86E-05	1.23E-05	0.00050	0.00885	0.00053	5.094	0.081	0.3238	0.0043	0.8957	0.1141	0.0008	18	1866	13	3.1
4114-12.1	1609	242	0.16	540	11	2.29E-05	6.11E-06	0.00040	0.04536	0.00070	5.322	0.061	0.3367	0.0036	0.9711	0.1146	0.0003	17	1874	5	0.2
4114-90.1	32	1	0.02	9	2	1.88E-04	1.57E-04	0.00326	0.00540	0.00592	4.941	0.176	0.3120	0.0073	0.7449	0.1149	0.0028	36	1878	44	6.8
4114-67.1	706	62	0.09	223	1	2.75E-06	8.59E-06	0.00005	0.02675	0.00047	5.112	0.061	0.3223	0.0035	0.9614	0.1150	0.0004	17	1880	6	4.2
4114-83.1	368	103	0.29	192	0	1.60E-06	1.53E-05	0.00003	0.08239	0.00097	11.642	0.156	0.4818	0.0060	0.9620	0.1753	0.0007	26	2608	6	2.8
4114-47.1	423	105	0.26	220	41	2.34E-04	5.80E-05	0.00405	0.07446	0.00229	11.728	0.179	0.4843	0.0059	0.8556	0.1756	0.0014	25	2612	13	2.5
4114-45.1	477	70	0.15	237	3	1.55E-05	1.48E-05	0.00027	0.04041	0.00079	11.654	0.178	0.4730	0.0062	0.9088	0.1787	0.0012	24	2641	11	5.5
4114-21.1	293	124	0.44	160	4	3.23E-05	1.98E-05	0.00056	0.11959	0.00208	12.266	0.180	0.4886	0.0057	0.8581	0.1821	0.0014	25	2672	13	4.0
4114-40.1	278	109	0.41	150	3	2.80E-05	1.95E-05	0.00049	0.11300	0.00262	12.167	0.191	0.4840	0.0063	0.8913	0.1823	0.0013	28	2674	12	4.8
4114-48.1	521	82	0.16	266	4	2.08E-05	1.93E-05														

Table 1. (cont.)

Spot name	U (ppm)	Th (ppm)	Th U	Pb* (ppm)	204Pb (ppb)	206Pb/208Pb	207Pb/208Pb	f(206) ²⁰⁴	208Pb/206Pb	± 206Pb/208Pb	207Pb/208Pb	± 206Pb/208Pb	206Pb/238U	± 206Pb/238U	207Pb/206Pb	± 206Pb/208Pb	206Pb/238U	± 206Pb/238U	Apparent ages (Ma)		Disc. (%)	
																			207Pb/206Pb	± 207Pb/206Pb		
4114-47.2	367	99	0.28	181	1	1.00E-05	1.00E-05	0.00017	0.07604	0.00077	11.546	0.4543	0.0050	0.1843	0.0009	2414	22	2692	8	10.3		
4114-1.1	312	137	0.46	179	4	2.92E-05	1.82E-05	0.00051	0.12116	0.00129	12.946	0.180	0.5093	0.0061	0.9165	0.1844	0.0010	2653	26	2693	9	1.5
4114-4.1	245	101	0.43	136	13	1.22E-04	6.18E-05	0.00211	0.11859	0.00342	12.672	0.219	0.4935	0.0066	0.8392	0.1862	0.0018	2586	28	2709	16	4.6
4114-73.2	148	124	0.86	90	2	2.97E-05	3.32E-05	0.00051	0.22813	0.00259	12.810	0.215	0.4956	0.0075	0.9460	0.1875	0.0010	2595	33	2720	9	4.6
4114-13.1	545	68	0.13	289	5	2.09E-05	2.11E-05	0.00036	0.03422	0.00106	13.173	0.160	0.5035	0.0055	0.9414	0.1898	0.0008	2629	24	2740	7	4.1
4114-73.1	173	122	0.73	107	0	4.67E-06	1.79E-05	0.00008	0.20261	0.00191	13.569	0.194	0.5167	0.0058	0.8468	0.1905	0.0015	2685	25	2746	13	2.2
4114-23.1	785	67	0.09	432	4	1.00E-05	1.00E-05	0.00017	0.02396	0.00046	14.068	0.160	0.5252	0.0056	0.9710	0.1943	0.0005	2721	24	2779	4	2.1
4114-23.2	379	67	0.18	201	1	7.04E-06	9.47E-06	0.00012	0.04935	0.00096	13.753	0.198	0.4929	0.0067	0.9773	0.2024	0.0006	2583	29	2845	5	9.2
4114-13.2	361	46	0.13	201	1	8.68E-06	1.35E-05	0.00015	0.03526	0.00079	14.636	0.232	0.5243	0.0077	0.9622	0.2025	0.0009	2717	33	2846	7	4.5
4114-34.1	365	95	0.27	218	12	7.28E-05	2.18E-05	0.00126	0.06964	0.00108	15.534	0.208	0.5422	0.0062	0.9044	0.2078	0.0012	2792	26	2888	9	3.3
4114-25.1	248	88	0.37	148	3	2.65E-05	3.34E-05	0.00046	0.09612	0.00156	15.431	0.187	0.5281	0.0057	0.9386	0.2119	0.0009	2733	24	2920	7	6.4
4114-38.1	458	123	0.28	272	4	1.92E-05	1.36E-05	0.00033	0.07440	0.00087	15.718	0.209	0.5370	0.0065	0.9488	0.2123	0.0009	2771	27	2923	7	5.2
4114-77.1	276	113	0.42	175	1	1.00E-05	1.00E-05	0.00017	0.11248	0.00231	16.470	0.286	0.5552	0.0082	0.9021	0.2152	0.0016	2847	34	2945	12	3.3
4114-83.3	335	133	0.41	216	0	2.01E-06	9.02E-06	0.00003	0.11036	0.00105	16.914	0.242	0.5624	0.0073	0.9497	0.2181	0.0010	2876	30	2967	7	3.1
ML-B (Z4115; 64.0047°N, 77.8900°E)																						
4115-71.2	4144	216	0.05	1281	2	1.50E-06	2.16E-06	0.00003	0.01538	0.00014	4.746	0.050	0.3202	0.0033	0.9873	0.1075	0.0002	1790	16	1758	3	-1.9
4115-71.1	4590	192	0.04	1457	6	4.59E-06	2.00E-06	0.00008	0.01223	0.00016	4.916	0.063	0.3296	0.0041	0.9925	0.1082	0.0002	1836	20	1769	3	-3.8
4115-2.1.2	1084	131	0.13	332	6	2.03E-05	5.66E-06	0.00035	0.03294	0.00052	4.686	0.055	0.3125	0.0034	0.9564	0.1088	0.0004	1753	17	1779	6	1.4
4115-1.1	1294	153	0.12	394	4	1.22E-05	1.02E-05	0.00021	0.03446	0.00065	4.650	0.054	0.3098	0.0033	0.9613	0.1089	0.0004	1740	16	1781	6	2.3
4115-2.1	1040	121	0.12	316	2	8.21E-06	1.71E-05	0.00014	0.03446	0.00072	4.666	0.053	0.3095	0.0032	0.9438	0.1094	0.0004	1738	16	1789	7	2.8
4115-8.1	1167	131	0.12	361	1	4.69E-06	7.36E-06	0.00008	0.03328	0.00043	4.839	0.056	0.3146	0.0033	0.9476	0.1116	0.0004	1763	16	1825	5	3.4
4115-1.2	920	226	0.25	313	2	6.01E-06	4.78E-06	0.00010	0.06738	0.00091	5.160	0.078	0.3336	0.0042	0.9815	0.1122	0.0003	1856	20	1835	5	-1.1
4115-57.1	839	196	0.24	277	4	1.76E-05	7.04E-06	0.00031	0.07146	0.00091	5.031	0.070	0.3250	0.0037	0.9779	0.1123	0.0008	1814	18	1836	12	1.2
4115-13.1	696	197	0.29	231	3	1.38E-05	2.08E-05	0.00023	0.08432	0.00102	5.000	0.064	0.3226	0.0037	0.9306	0.1124	0.0005	1802	18	1839	9	2.0
4115-76.1	87	2	0.03	27	2	7.28E-05	1.27E-04	0.00126	0.00945	0.00488	5.019	0.140	0.3228	0.0052	0.6699	0.1128	0.0024	1803	25	1844	38	2.2
4115-9.1	1021	170	0.17	332	8	2.65E-05	1.32E-05	0.00046	0.04915	0.00082	5.058	0.061	0.3251	0.0034	0.9049	0.1129	0.0006	1814	16	1846	9	1.7
4115-3.1	1607	356	0.23	535	5	1.00E-05	1.00E-05	0.00017	0.06512	0.00060	5.127	0.060	0.3287	0.0035	0.9518	0.1131	0.0004	1832	17	1850	7	1.0
4115-18.1	1494	276	0.19	483	5	1.20E-05	1.20E-05	0.00021	0.05318	0.00059	5.023	0.065	0.3220	0.0035	0.8899	0.1131	0.0007	1800	17	1850	11	2.7
4115-78.1	671	232	0.36	224	2	1.09E-05	9.18E-06	0.00019	0.10423	0.00188	4.976	0.067	0.3189	0.0038	0.9192	0.1132	0.0006	1784	18	1851	10	3.6
4115-61.1	779	194	0.26	257	1	6.68E-06	5.23E-06	0.00012	0.07424	0.00106	5.053	0.057	0.3238	0.0034	0.9602	0.1132	0.0004	1808	17	1851	6	2.3
4115-21.1	627	231	0.38	211	1	5.15E-06	2.35E-05	0.00009	0.11145	0.00126	4.981	0.062	0.3189	0.0034	0.9017	0.1133	0.0006	1784	16	1852	10	3.7
4115-77.1	1573	361	0.24	543	5	1.00E-05	1.00E-05	0.00017	0.06870	0.00055	5.307	0.059	0.3395	0.0036	0.9713	0.1134	0.0003	1884	17	1854	5	-1.6
4115-65.1	1054	235	0.23	345	3	1.16E-05	1.02E-05	0.00020	0.06589	0.00123	5.039	0.066	0.3223	0.0037	0.9173	0.1134	0.0006	1801	18	1855	10	2.9
4115-55.1	1073	179	0.17	371	4	1.28E-05	5.01E-06	0.00022	0.04982	0.00045	5.399	0.076	0.3452	0.0047	0.9856	0.1134	0.0003	1912	23	1855	4	-3.0
4115-23.1	1179	237	0.21	386	0	5.60E-07	4.14E-06	0.00001	0.05910	0.00042	5.082	0.061	0.3248	0.0034	0.9220	0.1135	0.0005	1813	17	1856	8	2.3
4115-5.2	1115	160	0.15	371	3	9.95E-06	5.13E-06	0.00017	0.04237	0.00048	5.238	0.060	0.3346	0.0036	0.9594	0.1135	0.0004	1861	17	1856	6	-0.2
4115-67.1	933	236	0.26	308	3	1.00E-05	1.00E-05	0.00017	0.07398	0.00085	5.067	0.066	0.3229	0.0036	0.9099	0.1138	0.0006	1804	18	1861	10	3.1
4115-72.1	799	214	0.28	262	4	1.60E-05	1.19E-05	0.00028	0.07743	0.00086	5.028	0.066	0.3202	0.0038	0.9459	0.1139	0.0005	1791	19	1863	8	3.9
4115-48.1	95	4	0.04	30	1	2.98E-05	5.92E-05	0.00052	0.01338	0.00231	5.205	0.129	0.3312	0.0049	0.6917	0.1140	0.0021	1844	24	1864	33	1.1
4115-14.1	1102	218	0.20	363	2	6.65E-06	8.30E-06	0.00012	0.05904	0.00075	5.125	0.057	0.3260	0.0033	0.9539	0.1140	0.0004	1819	16	1864	6	2.4
4115-57.2	136	4	0.03	43	2	6.44E-05	2.86E-05	0.00112	0.00790	0.00144	5.213	0.083	0.3309	0.0046	0.9239	0.1143	0.0007	1843	23	1868	11	1.4
4115-5.1	130	2	0.01	40	0	1.00E-05	1.00E-05	0.00017	0.00560	0.00053	5.137	0.080	0.3191	0.0039	0.8517	0.1168	0.0010	1785	19	1907	15	6.4

Notes: (See Stern, 1997):

Analytical details: ion probe mount IP394, spot size = 35µm, beam current = 10nA, 6 scans, no Pb fractionation correction

Spot name follows the convention x-y-z, where x = sample number, y = grain number and z = spot number. Multiple analyses in an individual spot are labelled as x-y-z.z

Uncertainties reported at 1s (absolute) and are calculated by numerical propagation of all known sources of error

f(206)²⁰⁴ refers to mole fraction of total 206Pb that is due to common Pb, calculated using the 204Pb-method; common Pb composition used is the surface blank (4/6; 0.05770; 7/6; 0.89500; 8/6; 2.13840)

* refers to radiogenic Pb (corrected for common Pb)

Discordance relative to origin = 100 * (1 - (206Pb/208Pb) / (206Pb/208Pb) age)

Calibration standard 6266; U = 910 ppm; Age = 559 Ma; (207Pb/206Pb) / (207Pb/206Pb) = 0.09059

Error in 206Pb/208Pb calibration 1.0%

Th/U calibration: F = 0.03900 ± 0.85600

Table 1. (cont.)

Spot name	U (ppm)	Th (ppm)	Th/U	Pb* (ppm)	²⁰⁴ Pb (ppb)	²⁰⁴ Pb/ ²⁰⁶ Pb	$\pm \frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	f(206) ²⁰⁴	$\frac{^{208}\text{Pb}}{^{206}\text{Pb}}$	$\pm \frac{^{208}\text{Pb}}{^{206}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{238}\text{U}}$	$\pm \frac{^{207}\text{Pb}}{^{238}\text{U}}$	Corr Coeff	$\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	$\pm \frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	Apparent ages (Ma)						
																$\frac{^{206}\text{Pb}}{^{238}\text{U}}$	$\pm \frac{^{206}\text{Pb}}{^{238}\text{U}}$	$\frac{^{206}\text{Pb}}{^{238}\text{U}}$	$\pm \frac{^{206}\text{Pb}}{^{238}\text{U}}$			
NOT-A (Z4117: 63.1591°N 78.0119°E)																						
non-detrital zircon																						
4117-7-1	44	30	0.69	19	4	2.47E-04	1.40E-04	0.00429	0.18389	0.00843	5.937	0.273	0.3814	0.0088	0.6031	0.1129	0.0042	2083	41	1846	68	-12.8
4117-71.1	335	83	0.26	115	2	1.90E-05	2.20E-05	0.00033	0.07543	0.00152	5.239	0.083	0.3360	0.0046	0.9129	0.1131	0.0007	1868	22	1849	12	-1.0
4117-71.1.2	331	76	0.24	107	1	1.00E-05	0.00017	0.00976	0.00866	0.00086	4.994	0.090	0.3193	0.0053	0.9524	0.1135	0.0006	1786	26	1855	10	3.7
4117-71.1.3	278	64	0.24	86	5	6.51E-05	1.66E-05	0.00113	0.07010	0.00107	4.758	0.122	0.3035	0.0072	0.9612	0.1137	0.0008	1709	36	1859	13	8.1
4117-2.1	345	126	0.38	121	6	6.06E-05	1.97E-05	0.00105	0.10760	0.00120	5.202	0.080	0.3329	0.0042	0.8810	0.1133	0.0008	1853	20	1853	13	0.1
4117-2.1.2	335	123	0.38	109	3	3.03E-05	2.35E-05	0.00052	0.10907	0.00136	4.893	0.076	0.3107	0.0041	0.9009	0.1142	0.0008	1744	20	1867	12	6.6
4117-50.1	214	109	0.52	78	2	2.62E-05	2.14E-05	0.00045	0.15299	0.00202	5.270	0.080	0.3328	0.0044	0.9302	0.1149	0.0006	1852	22	1877	10	1.4
4117-50.2	225	114	0.52	83	2	2.93E-05	2.36E-05	0.00051	0.15085	0.00224	5.275	0.083	0.3372	0.0046	0.9166	0.1135	0.0007	1873	22	1856	11	-0.9
4117-45.2	967	13	0.01	310	4	1.30E-05	7.89E-06	0.00023	0.00360	0.00033	5.241	0.072	0.3334	0.0042	0.9487	0.1140	0.0005	1855	20	1865	8	0.5
detrital zircon																						
4117-4.1	1926	384	0.21	684	3	5.49E-06	3.74E-06	0.00010	0.06164	0.00080	5.565	0.077	0.3509	0.0047	0.9795	0.1150	0.0003	1939	22	1880	5	-3.1
4117-67.1	287	107	0.38	103	3	4.08E-05	1.43E-05	0.00071	0.11029	0.00210	5.437	0.078	0.3410	0.0044	0.9342	0.1156	0.0006	1892	21	1890	9	-0.1
4117-62.1	1702	260	0.16	593	3	5.67E-06	2.62E-06	0.00010	0.04544	0.00030	5.560	0.076	0.3487	0.0043	0.9354	0.1156	0.0006	1929	20	1890	9	-2.0
4117-98.1	620	24	0.04	203	5	2.96E-05	1.44E-05	0.00051	0.01134	0.00063	5.408	0.236	0.3388	0.0077	0.6201	0.1158	0.0040	1881	37	1892	63	0.6
4117-98.2	452	15	0.04	150	1	6.75E-06	1.32E-05	0.00012	0.01054	0.00055	5.476	0.083	0.3421	0.0048	0.9632	0.1161	0.0005	1897	23	1897	7	0.0
4117-28.1	470	178	0.39	169	4	2.71E-05	1.12E-05	0.00047	0.11422	0.00119	5.442	0.080	0.3402	0.0044	0.9267	0.1160	0.0006	1888	21	1896	10	0.4
4117-34.1	2168	1314	0.63	866	5	6.81E-06	4.36E-06	0.00012	0.18345	0.00492	5.715	0.181	0.3572	0.0083	0.8038	0.1161	0.0022	1969	39	1896	35	-3.8
4117-94.1	2718	754	0.29	993	3	4.10E-06	3.01E-06	0.00007	0.08242	0.00060	5.683	0.074	0.3544	0.0044	0.9830	0.1163	0.0003	1955	21	1900	4	-2.9
4117-24.1	2891	783	0.28	1060	1	1.64E-06	1.50E-06	0.00003	0.07970	0.00032	5.723	0.077	0.3564	0.0046	0.9850	0.1165	0.0003	1965	22	1903	4	-3.3
4117-48.1	3872	1199	0.32	1479	2	1.52E-06	1.63E-06	0.00003	0.09205	0.00105	5.904	0.128	0.3675	0.0063	0.8584	0.1165	0.0013	2018	30	1904	20	-6.0
4117-43.1	2670	401	0.16	940	5	5.59E-06	2.03E-06	0.00010	0.04427	0.00026	5.672	0.075	0.3525	0.0045	0.9792	0.1167	0.0003	1947	21	1906	5	-2.1
4117-52.1	474	186	0.40	172	1	1.06E-05	1.03E-05	0.00018	0.11739	0.00182	5.502	0.077	0.3417	0.0045	0.9659	0.1168	0.0004	1895	21	1907	7	0.7
4117-23.1	648	17	0.03	222	2	8.59E-06	2.00E-05	0.00015	0.00836	0.00077	5.853	0.081	0.3529	0.0044	0.9442	0.1203	0.0006	1948	21	1961	8	0.6
4117-23.3	647	20	0.03	231	5	2.25E-05	7.61E-06	0.00039	0.00928	0.00036	5.952	0.095	0.3681	0.0049	0.8922	0.1173	0.0009	2020	23	1915	13	-5.5
4117-23.2	796	29	0.04	270	2	1.00E-05	1.00E-05	0.00017	0.01087	0.00043	5.663	0.083	0.3496	0.0046	0.9362	0.1175	0.0006	1933	22	1918	9	-0.7
4117-23.1.3	526	13	0.03	150	6	4.75E-05	1.46E-05	0.00082	0.06638	0.00058	4.813	0.169	0.2956	0.0066	0.7227	0.1181	0.0029	1669	33	1928	44	13.4
4117-23.1.2	608	15	0.03	193	5	2.73E-05	1.35E-05	0.00047	0.00677	0.00054	5.351	0.108	0.3274	0.0056	0.9065	0.1185	0.0010	1826	27	1934	15	5.6
4117-3.1	109	3	0.02	37	7	2.03E-04	5.66E-05	0.00351	0.00260	0.00215	5.719	0.105	0.3498	0.0049	0.8350	0.1186	0.0012	1934	24	1935	18	0.1
4117-3.3	164	4	0.02	56	1	1.29E-05	2.10E-05	0.00022	0.00916	0.00102	5.825	0.113	0.3544	0.0054	0.8532	0.1192	0.0012	1956	26	1944	18	-0.6
4117-3.2	80	2	0.03	27	5	2.30E-04	6.42E-05	0.00398	0.00353	0.00247	5.461	0.129	0.3437	0.0059	0.7962	0.1152	0.0017	1905	28	1883	26	-1.1
4117-1.1	573	191	0.34	233	3	1.70E-05	6.03E-06	0.00029	0.09772	0.00114	6.816	0.107	0.3858	0.0055	0.9457	0.1281	0.0007	2103	25	2073	9	-1.5
4117-69.1	196	57	0.30	80	6	9.04E-05	2.80E-05	0.00157	0.08553	0.00356	7.004	0.425	0.3878	0.0139	0.6824	0.1310	0.0059	2112	65	2111	81	-0.1
4117-51.1	263	365	1.43	131	4	4.31E-05	2.27E-05	0.00075	0.39914	0.00521	6.812	0.105	0.3771	0.0050	0.9138	0.1310	0.0008	2063	23	2112	11	2.3
4117-49.1	191	102	0.55	84	5	7.90E-05	3.62E-05	0.00137	0.15676	0.00196	7.291	0.119	0.3965	0.0054	0.8878	0.1334	0.0010	2153	25	2143	13	-0.5
4117-93.1	216	101	0.48	95	4	5.47E-05	2.41E-05	0.00095	0.13827	0.00157	7.497	0.279	0.4019	0.0055	0.4803	0.1353	0.0045	2178	26	2168	59	-0.5
4117-68.1	242	194	0.83	119	1	1.16E-05	1.80E-05	0.00020	0.23439	0.00223	8.049	0.130	0.4135	0.0053	0.8558	0.1412	0.0012	2231	24	2242	15	0.5
4117-30.1	476	290	0.63	232	3	1.54E-05	9.19E-06	0.00027	0.18139	0.00173	8.595	0.122	0.4266	0.0055	0.9469	0.1461	0.0007	2290	25	2301	8	0.5
4117-45.1	42	34	0.84	22	3	2.12E-04	9.05E-05	0.00387	0.06855	0.00257	8.674	0.260	0.4237	0.0091	0.7948	0.1485	0.0027	2277	42	2329	32	2.2
4117-45.3	112	131	1.21	58	1	1.95E-05	4.82E-05	0.00034	0.35755	0.00409	7.811	0.161	0.4039	0.0071	0.9041	0.1403	0.0012	2187	33	2230	15	1.9

Notes (see Stern, 1997):

Analytical details: ion probe mount IP412, spot size = 25µm, beam current = 8nA, 6 scans, no Pb fractionation correction

Spot name follows the convention x-y-z; where x = sample number, y = grain number and z = spot number. Multiple analyses in an individual spot are labelled as x-y.z.z

Uncertainties reported at 1σ (absolute) and are calculated by numerical propagation of all known sources of error

¹²⁰⁶ refers to mole fraction of total ²⁰⁶Pb that is due to common Pb, calculated using the ²⁰⁴Pb-method; common Pb composition used is the surface blank (4/6: 0.05770; 7/6: 0.89500; 8/6: 2.13840)

* refers to radiogenic Pb (corrected for common Pb)

Discordance relative to origin = $100 \cdot (1 - \frac{^{206}\text{Pb}}{^{206}\text{Pb}}) / \frac{^{206}\text{Pb}}{^{206}\text{Pb}}$

Calibration standard 6266; U = 910 ppm; Age = 559 Ma; $\frac{^{206}\text{Pb}}{^{238}\text{U}}$ = 0.09059

Error in $\frac{^{206}\text{Pb}}{^{238}\text{U}}$ calibration 1.0%

Th/U calibration: F = 0.03900^{UO} + 0.85600

Table 2. U-Pb TIMS analytical data.

Fraction ¹	Wt. µg	U ppm	Pb ² ppm	Pb ³ / ₂₀₄ Pb	Pb ⁴ pg	Isotopic Ratios ⁵						Ages (Ma) ⁷																
						²⁰⁶ Pb/ ₂₀₆ Pb	²⁰⁷ Pb/ ₂₃₅ U	±1SE Abs	²⁰⁶ Pb/ ₂₃₈ U	±1SE Abs	²⁰⁷ Pb/ ₂₀₆ Pb	±1SE Abs	²⁰⁶ Pb/ ₂₃₈ U	²⁰⁷ Pb/ ₂₃₅ U	±2SE	²⁰⁷ Pb/ ₂₀₆ Pb	±2SE	% Disc										
SAL-A (Z4111; 63.3831°N 76.8458°E)																												
4111-1 (M)	8	562	3898	13277	7	22.82	5.26069	0.00518	0.33417	0.00028	0.938	0.11418	0.00004	1858.5	2.7	1862.5	1.7	1866.9	1.3	0.5								
4111-M2 (M)	9	311	4671	5555	11	50.70	5.24372	0.00523	0.33400	0.00029	0.926	0.11387	0.00004	1857.7	2.8	1859.8	1.7	1862.0	1.4	0.3								
SAL-B (Z4112; 63.3831°N 76.8459°E)																												
4112-1 (M)	5	674	6321	10564	6	30.93	5.34146	0.00532	0.33775	0.00029	0.930	0.11470	0.00004	1875.9	2.8	1875.5	1.7	1875.1	1.3	0.0								
SAL-C (Z4113; 63.3831°N 76.8459°E)																												
4113-T1 (T)	56	77	48	546	158	1.22	4.57094	0.01149	0.31116	0.00036	0.693	0.10654	0.00020	1746.4	3.6	1744.0	4.2	1741.0	7.0	-0.4								
4113-T2 (T)	46	70	47	325	201	1.42	4.48889	0.01775	0.30576	0.00039	0.674	0.10648	0.00034	1719.8	3.9	1728.9	6.6	1740.0	11.8	1.3								
MIL-A (Z4114; 64.0047°N 77.8900°E)																												
4114-T1 (T)	49	52	25	282	180	0.72	4.41920	0.02002	0.30326	0.00043	0.665	0.10569	0.00040	1707.4	4.3	1715.9	7.5	1726.3	13.7	1.3								
4114-T2 (T)	42	39	20	256	129	0.90	4.32654	0.02209	0.29927	0.00047	0.651	0.10485	0.00045	1687.7	4.7	1698.4	8.4	1711.7	15.6	1.6								
4114-T2 (T)	53	123	75	876	145	1.23	4.42152	0.00742	0.30440	0.00027	0.724	0.10535	0.00013	1713.1	2.7	1716.4	2.8	1720.4	4.4	0.5								
MIL-B (Z4115; 64.0047°N 77.8900°E)																												
4115-T1 (T)	70	73	37	637	155	0.86	4.35326	0.00997	0.30170	0.00037	0.719	0.10465	0.00017	1699.7	3.6	1703.5	3.8	1708.2	6.1	0.6								
4115-T3 (T)	24	101	54	620	77	0.93	4.43679	0.00995	0.30390	0.00031	0.668	0.10589	0.00018	1710.6	3.0	1719.2	3.7	1729.8	6.4	1.3								

Notes:
¹T=titanite fraction, M=monazite fraction
²Radiogenic Pb
³Measured ratio, corrected for spike and fractionation
⁴Total common Pb in analysis corrected for fractionation and spike
⁵Corrected for blank Pb and U and common Pb, errors quoted are 1 sigma absolute; procedural blank values for this study ranged from 0.3–0.5 pg U and 3–5 pg Pb for monazite analyses and 1–2 pg U and 10–14 pg Pb for titanite analyses; Pb blank isotopic composition is based on the analysis of procedural blanks; corrections for common Pb were made using Stacey-Kramers compositions.
⁶Correlation coefficient
⁷Corrected for blank and common Pb, errors quoted are 2 sigma in Ma
The error on the calibration of the GSC ²⁰⁶Pb-²³³U-²³⁵U spike utilized in this study is 0.22% (2σ).

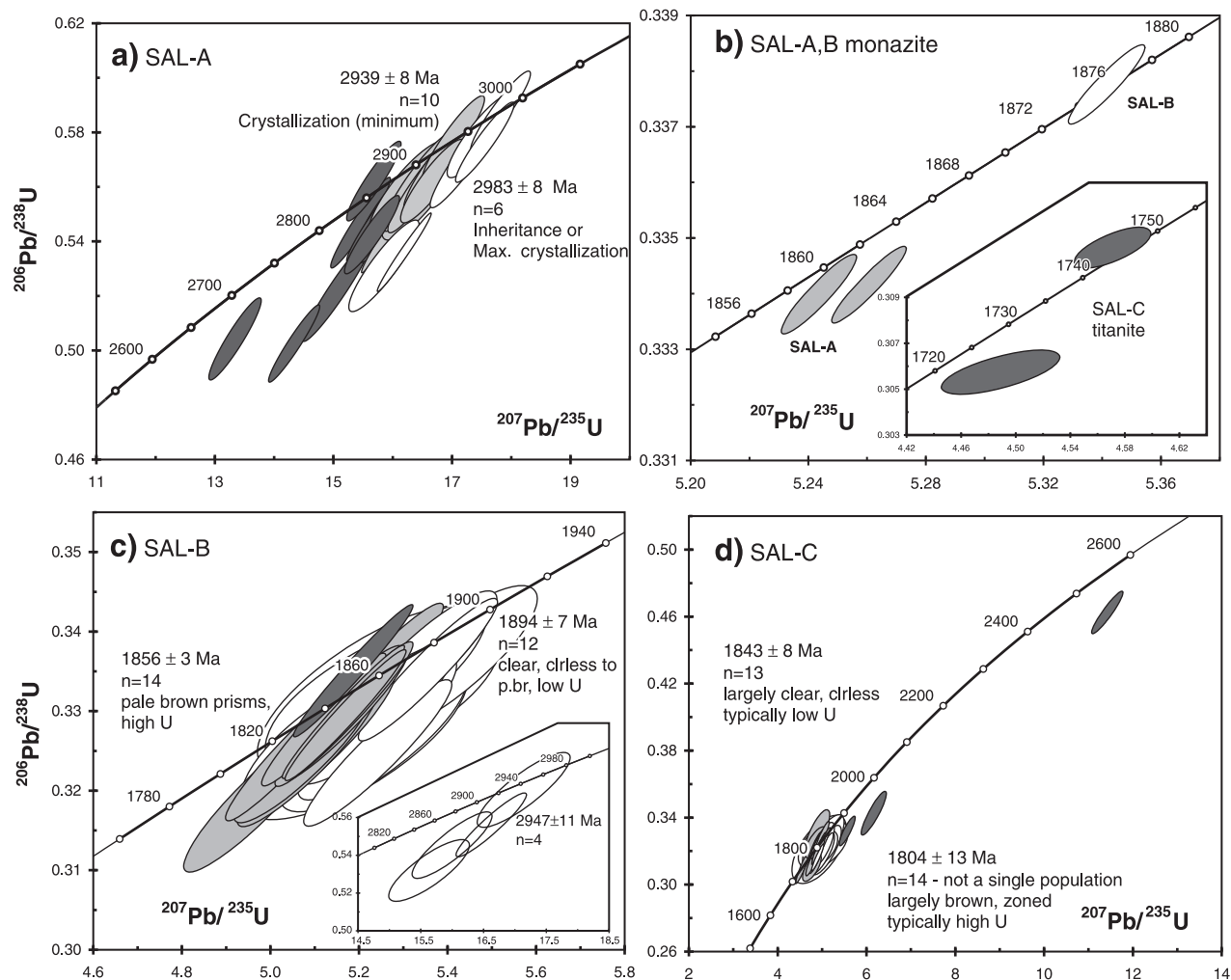


Figure 2. Concordia diagrams for U-Pb results from Salisbury Island. Error ellipses at 2σ . See text for interpretation. **a)** SHRIMP zircon results from tonalite SAL-A. White ellipses are plotted for oldest zircon population. Light grey ellipses illustrate younger zircon population. Dark grey ellipses represent analyses not included in the weighted mean and illustrate the effect of ancient Pb loss. **b)** TIMS monazite and titanite results from samples from Salisbury Island. Monazite data from SAL-A are shown as light grey ellipses, and those from SAL-B as a white ellipse. Titanite data from SAL-C are shown as dark grey ellipses. **c)** SHRIMP zircon results from monzogranite SAL-B. Older 2.9 Ga inherited zircon data are as inset. See data table for details. White ellipses indicate a clear, colourless, low-U zircon population. Light grey ellipses illustrate a younger, pale brown, high-U zircon population. One grain (dark grey ellipse) has exceptionally high U concentration and exceptionally low Th/U ratio. **d)** SHRIMP zircon results from syenogranite SAL-C. Oldest, discordant zircon grains are shown by dark grey ellipses. White ellipses represent a population of largely clear, colourless, low-U zircon grains. Light grey ellipses illustrate a group of dominantly brown, zoned, high-U zircon grains.

interpreted to represent inherited Archean zircon grains that have experienced Pb-loss during a Paleoproterozoic heating event. As observed in sample SAL-B, there are two morphologically and temporally distinct suites of zircon. Dominantly low-U, clear, colourless zircon grains yield a weighted mean age of 1843 ± 8 Ma (unfilled ellipses, $n = 13$, MSWD = 1.04, POF = 0.41), whereas typically higher U, pale brown zircon grains give a poorly constrained mean age of 1804 ± 13 Ma ($n = 14$, MSWD = 5.6, POF = 0.0, light grey ellipses). These younger zircon analyses exhibit excess scatter in their ages suggesting that they do not represent a single population. Two multigrain titanite TIMS analyses from sample SAL-C yield a weighted mean age of 1741 ± 6 Ma (MSWD = 0.021, POF = 0.88, Fig. 2b).

MIL-A (GSC lab # 4114)

Thirty analyses were carried out on twenty-five zircon grains from tonalite gneiss sample MIL-A (Fig. 3a). Three age groupings are observed. The oldest group extends up to 2.97 Ga and is interpreted as inherited. A regression through the slightly younger group of analyses and a forced lower intercept of 1.87 Ga (*see below*) gives an upper intercept age of 2816 ± 31 Ma ($n = 14$, MSWD = 2.0, POF = 0.01), which is interpreted as the best estimate for the crystallization age of the tonalite. The youngest group of 8 analyses gives a weighted mean age of 1870 ± 5 Ma (MSWD = 0.97, POF = 0.45). This group is composed of two low-U, imprecise analyses and six high-U, more precise analyses (light grey and dark grey ellipses, respectively; Fig. 3a). It is possible that these low-U analyses are distinct and slightly older in age (1.89 Ga), as observed in monzogranite SAL-B. However, this cannot be conclusively resolved due to the limited number and low precision of the analyses. Three multigrain titanite TIMS analyses from sample MIL-A yield a weighted mean $^{207}\text{Pb}/^{206}\text{Pb}$ age of 1720 ± 4 Ma by TIMS (MSWD = 0.99, POF = 0.37 Table 2, Fig. 3b).

MIL-B (GSC lab # 4115)

Twenty-seven analyses were carried out on 22 zircon grains from monzogranite sample MIL-B (Fig. 3c). A population of four clear, colourless, low-U zircon grains gives an imprecise weighted mean age of 1879 ± 37 Ma (unfilled ellipses, MSWD = 1.9, POF = 0.13). Pale brown, high-U zircons give a weighted mean age of 1852 ± 5 Ma (light grey

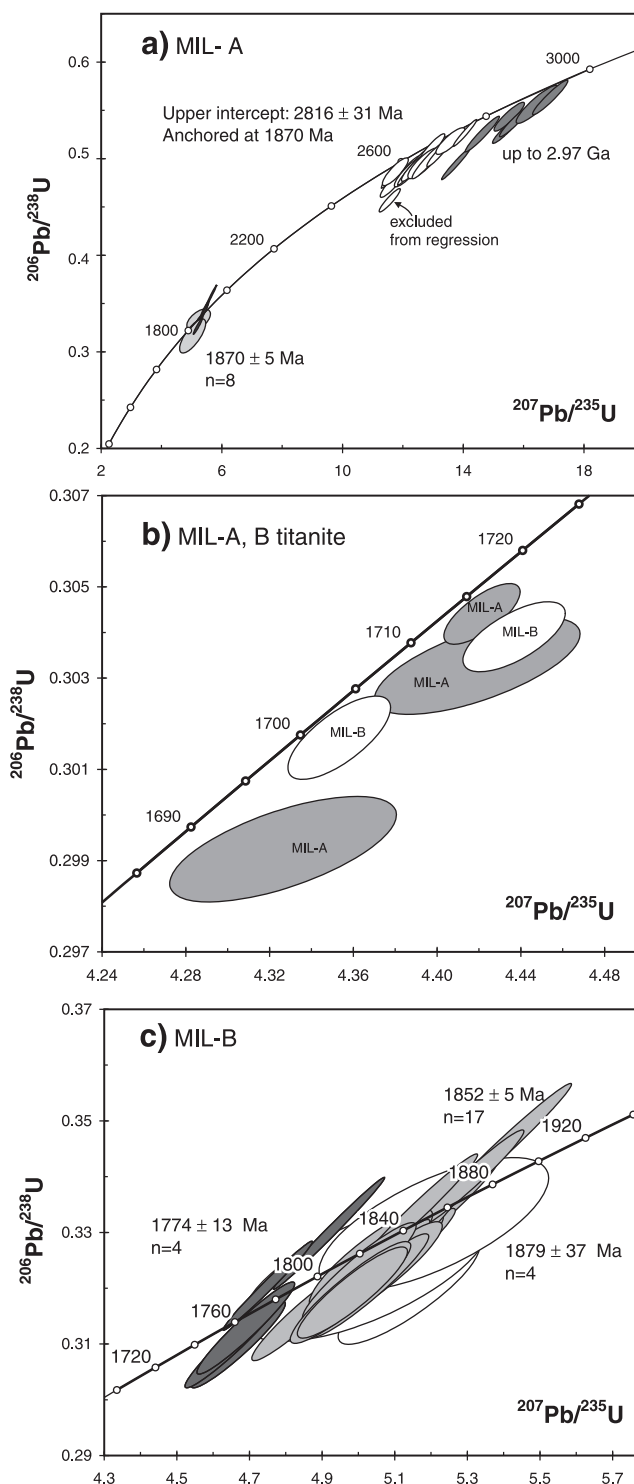


Figure 3. Concordia diagrams for U-Pb results from Mill Island. Error ellipses at 2σ . See text for interpretation. **a)** SHRIMP zircon results from tonalite MIL-A. Medium grey ellipses are plotted for an inherited zircon population. White ellipses illustrate the population regressed to define the crystallization age of MIL-A. Light grey ellipses indicate two clear, colourless, low-U zircon grains which may be distinctly older than a population of brown, high-U zircon shown as dark grey ellipses. **b)** TIMS titanite results from samples from Mill Island. Data from MIL-A are shown as light grey ellipses and from MIL-B as white ellipses. **c)** SHRIMP zircon results from monzogranite MIL-B. White ellipses indicate a clear, colourless, low-U zircon population. Light grey ellipses illustrate a younger, pale brown, high-U zircon population. Dark grey ellipses are also high U but give significantly younger $^{207}\text{Pb}/^{206}\text{Pb}$ ages.

ellipses, $n = 17$, $MSWD = 1.7$, $POF = 0.05$). As observed in SAL-B and MIL-A, a slightly older (ca. 1.89 Ga), chemically distinct component is observed, but low U has hampered our ability to precisely determine this age. Within the group of high-U zircon grains there is a further subcomponent that gives even younger $^{207}\text{Pb}/^{206}\text{Pb}$ ages ranging from 1.83 to 1.76 Ga (Fig. 3c dark grey ellipses). This subcomponent typically has lower Th/U than the 1852 Ma population and is interpreted to represent high-U zircon from that population that has lost Pb by partial resetting during a Paleoproterozoic metamorphic fluid event. The timing of this late heating/fluid event is constrained between 1.76 Ga (the youngest zircon date) and 1708 Ma, a cooling age determined by TIMS titanite analysis (Table 2, Fig. 3b).

NOT-A (GSC lab # 4117)

The dominant population of zircon grains from metasedimentary gneiss sample NOT-A consists of large, dark brown, elongate to stubby prisms, most having large quartz inclusions and fractures. Less common are dark to pale brown, slightly faceted to rounded, equant crystals commonly with cores or concentric zoning. Thirty-eight analyses were carried out on twenty-four zircon grains from sample NOT-A (Table 1, Fig. 4a, b). The probability distribution curve for detrital zircon grains (Fig. 4b) displays a dominant mode at 1.90 Ga and

older detritus ranging from 1.93 to 2.35 Ga. The youngest five zircon grains are all morphologically similar (subhedral, equant prisms), dominantly pale brown and are not shown on the probability density diagram. They may represent either the youngest detrital population or some younger addition of zircon, such as metamorphic growth or geological contamination from the thin leucosomes observed at the sampling site. These zircon grains form a single statistical population with a weighted mean age of 1861 ± 8 Ma ($n = 9$ includes replicates, $MSWD = 0.62$, $POF = 0.76$, light grey ellipses in Fig. 4a). This age is similar to the emplacement age of the Cumberland Batholith, an extensive 1865 to 1848 Ma felsic plutonic suite documented across southern and central Baffin Island, and a known time of tectonometamorphism (Jackson et al., 1990; Scott and Wodicka, 1998; Wodicka and Scott, 1997; Rayner et al., 2008). We interpret the 1.86 Ga zircon as non-detrital and likely leucosome-related on the basis of the presence of partial melt in the sample and well known 1.86 Ga plutonism and metamorphism in the region. The youngest detrital zircon (grain 4117-4.1) is 1880 ± 5 Ma in age (Table 1). This analysis falls within a group of zircon grains whose age falls between 1.90 and 1.88 Ga and currently represents our best estimate of the maximum age of deposition of the metasedimentary gneiss on Nottingham Island.

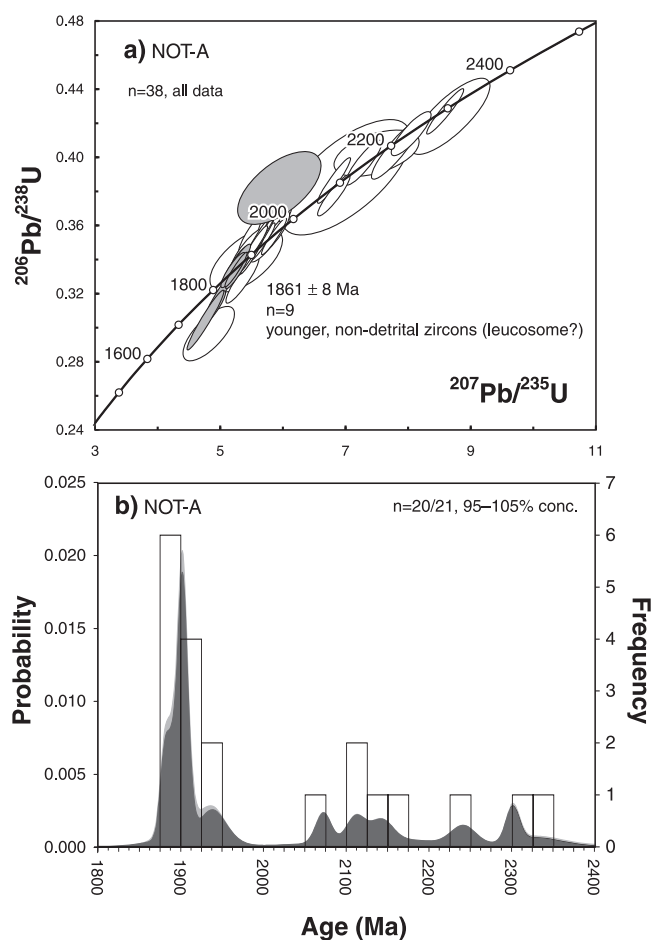
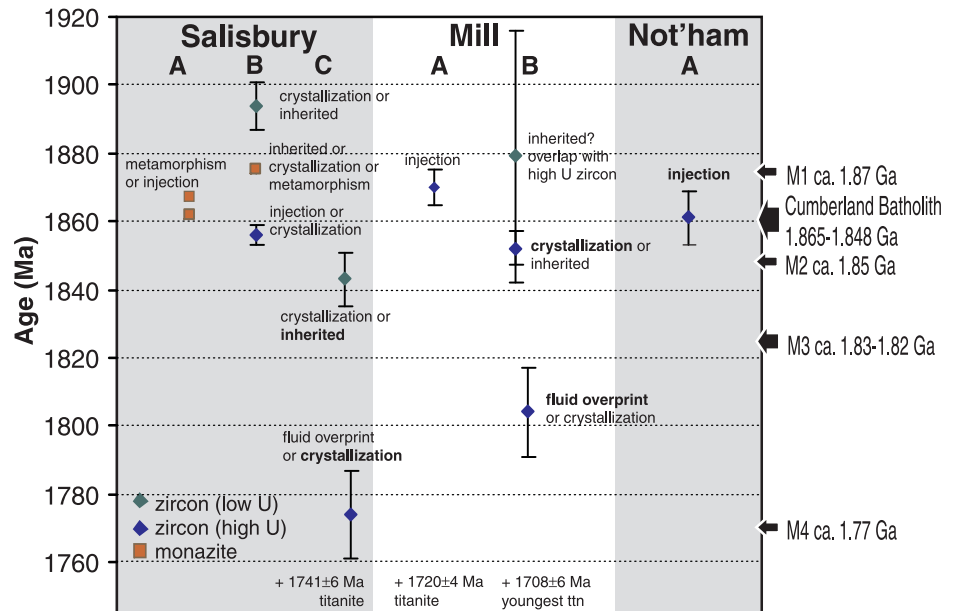


Figure 4. Concordia diagram and cumulative probability curve of U-Pb results from Nottingham Island, sample NOT-A. Error ellipses at 2σ . See text for interpretation. **a)** Detrital zircon data shown as white ellipses and leucosome data as grey ellipses. All data illustrated including replicates and discordant results. **b)** Cumulative probability curve and overlain histogram of detrital zircon results. Ca. 1861 Ma, non-detrital analyses are excluded in the probability distribution curve – binned frequency histogram plot. All data are illustrated in the light grey curve. Only data within $\pm 5\%$ of concordia are presented in the dark grey curve and histogram. In the instance of multiple analyses of a single detrital crystal, only the oldest analysis or the average was included in the diagram.

Figure 5. Summary of Paleoproterozoic ages from Mill, Salisbury and Nottingham islands. Text in bold indicates preferred interpretation; ttn = titanite. See text for explanation. Reference ages of Cumberland batholith across Baffin Island (Jackson et al., 1990; Scott and Wodicka, 1998; Wodicka and Scott, 1997) and metamorphism in southwest Baffin Island (Rayner et al., 2008) are provided for comparison.



DISCUSSION AND INTERPRETATION

The analytical data presented above for samples from Mill and Salisbury islands are extremely complex, with alternate interpretations possible for each phase. The data for the more complex Paleoproterozoic components are summarized in Figure 5. Tonalite crystallization on Salisbury Island occurred at 2939 Ma, or possibly as early as 2983 Ma (SAL-A). The SAL-B monzogranite may have intruded the tonalite at 1894 Ma, with addition of 1856 Ma zircon through the injection of a subsequent melt phase, the generation of in situ melt, or the effects of hydrothermal fluids. Alternatively, the 1894 Ma component is inherited within a 1856 Ma monzogranite. Both scenarios are equally plausible. The oldest coherent age group from the undeformed SAL-C syenogranite is 1843 Ma and may represent the crystallization age of the unit. However, on southwest Baffin Island, regional deformation is ongoing at this time (Rayner et al., 2008). Assuming a correlation between these units, the lack of strain in SAL-C is inconsistent with this observation and would suggest that these zircon grains are inherited. A poorly defined group of ages at ca. 1804 Ma is considered to be a better estimate for the time of syenogranite emplacement, with a lower limit defined by a titanite cooling age of 1741 Ma.

On Mill Island we interpret the tonalite MIL-A to have crystallized at 2816 Ma with 2.97 Ga inheritance. The presence of 1870 Ma zircon (\pm minor, low U, 1.89 Ga zircon) in MIL-A is attributed to the injection of a later igneous phase, the generation of in situ melt or the effects of hydrothermal fluids. In MIL-B the dominant zircon age of 1852 Ma is the best estimate for the crystallization age of the monzogranite. An imprecise grouping of ca. 1.88 Ga, low U zircon grains cannot be confidently separated from the 1852 Ma population but, if taken to be distinct based on unique chemistry

would be interpreted as inherited. A subset of the high-U zircon grains from MIL-B show ages as young as 1.76 Ga. This younger population is interpreted to record a fluid overprint or the age of the late syenogranite intrusion (not dated). The minimum age of this late event is constrained by titanite cooling ages as young as 1708 Ma.

Despite the difficulty in assigning specific ages to particular intrusive phases from these complex outcrops, the Paleoproterozoic zircon ages (1894, 1852–1856, and 1843 Ma) recorded in the meta-igneous rocks provide some insight into the cratonic affinity of Mill and Salisbury islands. These ages contrast markedly with an absence of precollisional (i.e. \geq ca. 1820 Ma) deformation, magmatic, and metamorphic events of Paleoproterozoic age in the Archean Superior Craton of northern Quebec (e.g. St-Onge et al., 2006b; their Fig. 8), thereby precluding a lower plate position for Mill and Salisbury islands with respect to the Bergeron suture. However, the possibility that these islands represent a rifted fragment of the Superior Craton cannot be excluded. Alternatively, similarities in age and lithologies between Mill and Salisbury islands and Hall Peninsula (Scott, 1999) provide evidence that the islands may form part of the upper plate Meta Incognita microcontinent as recently defined by St-Onge et al. (2006b). Current data are insufficient to distinguish between these two hypotheses.

The detrital profile from the Nottingham Island metasedimentary gneiss, dominated by Paleoproterozoic ages at ca. 1.90 Ga with sparse older ages, is similar to that for a Tasiuyak paragneiss from Hall Peninsula and to those for Lake Harbour Group siliciclastic rocks, which are characterized by dominantly 1.9 to 2.0 Ga detritus and a sparse early Paleoproterozoic to late Archean component (Scott et al., 2002; their Fig. 9). These similarities suggest that similar source region(s) were involved in the generation of the sedimentary detritus from these samples, but the

linkage in space between these metasedimentary packages at the time of deposition, if any, is not yet well understood. Due to their proximity, it may be tempting to suggest that the Archean rocks on Mill and Salisbury islands form the depositional basement to the metasedimentary rocks on Nottingham Island. However, this is not consistent with the detrital profile for NOT-A, which shows a distinct absence of basement-aged detritus (2.8–2.98 Ga). This apparent conflict suggests that the Nottingham Island metasedimentary rocks either are detached from their basement or developed in an off-craton position.

SUMMARY AND CONCLUSIONS

Archean inherited and basement ages of 2.97 Ga, 2983 to 2935 Ma, and 2819 Ma are not uniquely associated to one of the bounding cratons (Superior or Meta Incognita; Fig. 1). Multiple Paleoproterozoic events between 1.89 and 1.71 Ga have been identified. Evidence for an 1894 Ma event is recorded in low-U zircon grains from monzogranite SAL-B. Weak indications of this same event on Mill Island is present as few, low-U, imprecise analyses. Statistically indistinguishable ages of 1852 ± 5 Ma and 1856 ± 3 Ma representing igneous crystallization from in situ melt or mobilizate are present on both Mill and Salisbury islands. There is evidence for still younger zircon growth at 1843 ± 8 Ma from syenogranite SAL-C, and zircon growth and/or resetting at ca. 1.80 Ga (SAL-C) to as young as 1.76 Ga (MIL-B). Monazite ages range from 1875 to 1862 Ma and could represent inheritance, peak metamorphic, or cooling ages. Titanite ages, which likely represent cooling ages, range from 1740 Ma on Salisbury Island to 1730 to 1708 Ma on Mill Island. Intrusive and metamorphic ages older than 1.82 Ga exclude a direct linkage of the Archean rocks on Mill and Salisbury islands with the contiguous Superior Craton. The detrital profile from the Nottingham Island metasedimentary gneiss compares well with Tasiuyak paragneiss from Hall Peninsula and Lake Harbour Group siliciclastic rocks on southern Baffin Island.

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