

DESCRIPTIVE NOTES

INTRODUCTION
 This colour-coded, multibeam sonar shaded-relief seafloor map is part of a three-map series of The Gully at a scale of 1:100 000. The series also includes a surficial geology map (Cameron et al., 2008) and a colour-coded, shaded seafloor slope analysis map (Cameron and King, 2009).

MULTIBEAM BATHYMETRIC DATA COLLECTION
 Multibeam surveys were co-ordinated by the Geological Survey of Canada (Atlantic) in 1996, 1997, and 1998 and by the Canadian Hydrographic Service in 2000 within The Gully Marine Protected Area. Surveys were conducted in collaboration with several industry partners and the Department of Fisheries and Oceans. The survey utilized a Simrad EK1000 multibeam sonar with operational depths from 10 m to 4500 m. The system provides a 350° bearing maximum arc of 150°. During the survey, the swath coverage was up to seven times the water depth, with survey speeds of 10–12 knots. A Differential Global Positioning System provided positioning accuracy of approximately 2 m.
 Figures 1–6 are bathymetric profiles derived using ArcMap™. They illustrate a variety of features found in The Gully.

DATA DISPLAY
 Shaded seafloor relief was produced by processing multibeam data at both 10 m and 20 m grid spacing intervals. The elevation model has been displayed using sun illumination from 315° and 45° above the horizon. Bathymetric values are colour-coded shades of blue and red for depths from deep to shallow, respectively. The vertical exaggeration is 10X.
 Linear acoustic artifacts oriented parallel to the survey direction are due to refraction errors. Horizontal resolution decreases with increasing water depth because of an increase in the acoustic footprint and beam spacing. This results in a smoothing effect with a local underestimation of steep slopes and an inability to resolve small features.

MORPHOLOGY
 Introduction
 Positioned between Sable Island Bank and Banquereau, The Gully is the largest submarine canyon on the outer Scotian Shelf. It appears morphologically similar to many other canyons on the Scotian Shelf; however, unlike the continental shelf much deeper than other canyons, connecting the middle shelf to the continental slope.

Mechanisms of formation
 It is generally recognized that The Gully formed by fluvial, glacial, and meltwater erosion (Fader et al., 1986) that cut deeply into Cenozoic mudstone and sandstone units. It was excavated over the past million years as successive continental glaciations lowered sea level, exposing the shelf and allowing rivers to erode and deposit sediments near the top of the continental slope. The Gully was partially eroded by turbidity currents that flowed down the canyon at times of glacial maxima.

Upper part of The Gully
 The upper part of The Gully was partially infilled by subglacial and proglacial sediments. A meandering main channel, large feeder channels, and retrogressive slumped canyon walls with terrace character morphology of the upper part of The Gully (shallow to 800 m). This morphology was shaped after ice retreat, allowing deltas, fans to build the glacial sediments. Easily eroded glacial marine sediment infilling the upper western flanks of The Gully has contributed to the growth of major side canyons. Till on the eastern flank has prevented large side canyons from forming.

Lower part of The Gully
 With slopes frequently greater than 50°, the lower part of The Gully (beyond 800 m) is characterized by ridge and valley morphology formed by submarine erosion. This ridge and valley morphology may have been inherited by the overlying proglacial and hemipelagic deposits and is older than the morphology of the upper part of The Gully.

Feeder canyons
 Nine major feeder canyons or channels and many smaller channels occur on the western flank of the upper part of The Gully. Large feeder canyons do not occur on the eastern flank of The Gully. All feeder canyons have sand-filled catchment basins at their heads that give way to bedrock-controlled sidewalls in deeper water. Ledges, terraces, steep slopes of exposed sandstone and mudstone, and minor gullies characterize the feeder canyons (Fader and Strang, 2002).

The halfway
 The Gully halfway (channel) has a uniform sand-covered surface, sloping at approximately 2° seaward from its head. Only at 2702 m water depth does the slope of the half-way begin to decrease slightly. Oversteepening of the present half-way channel floor does not occur. Aggrading initial glacial erosional effects in the deep channel and uniform processes of sediment deposition in recent time.

The canyon head
 The head of The Gully is sand formed and bifurcated into two channels, a minor one that extends to the north and a major channel extending to the northwest. The major channel floor consists of rippled sand with accumulated organic flocculation in the troughs of the bedforms. The sediment at the canyon head is similar in texture and lithology to fine-grained sand found on Sable Island Bank. A large megalite occurs on the seafloor to the northwest of The Gully, which suggests formation by strong and likely episodic bottom currents.

The canyon flanks
 The seabed is scoured and pitted by past iceberg impacts on both flanks of the continental shelf edge of The Gully, between 300 m and 500 m water depth. Sand and gravel blanket these features in water depths shallower than 300 m.

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