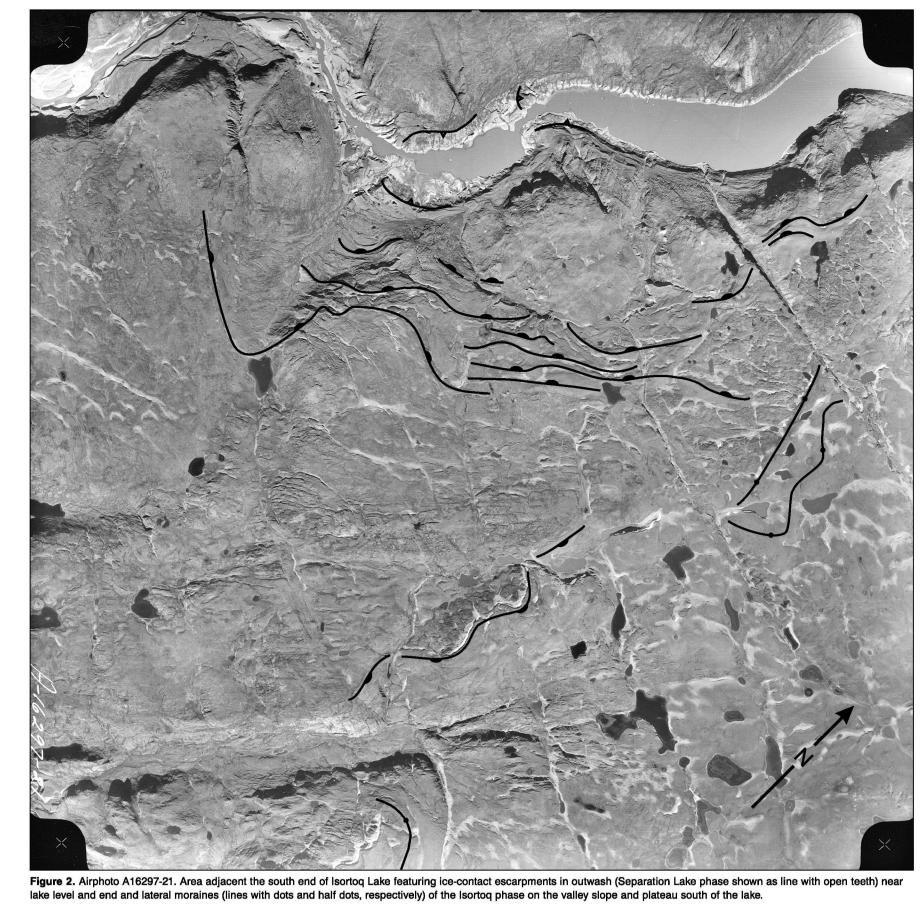
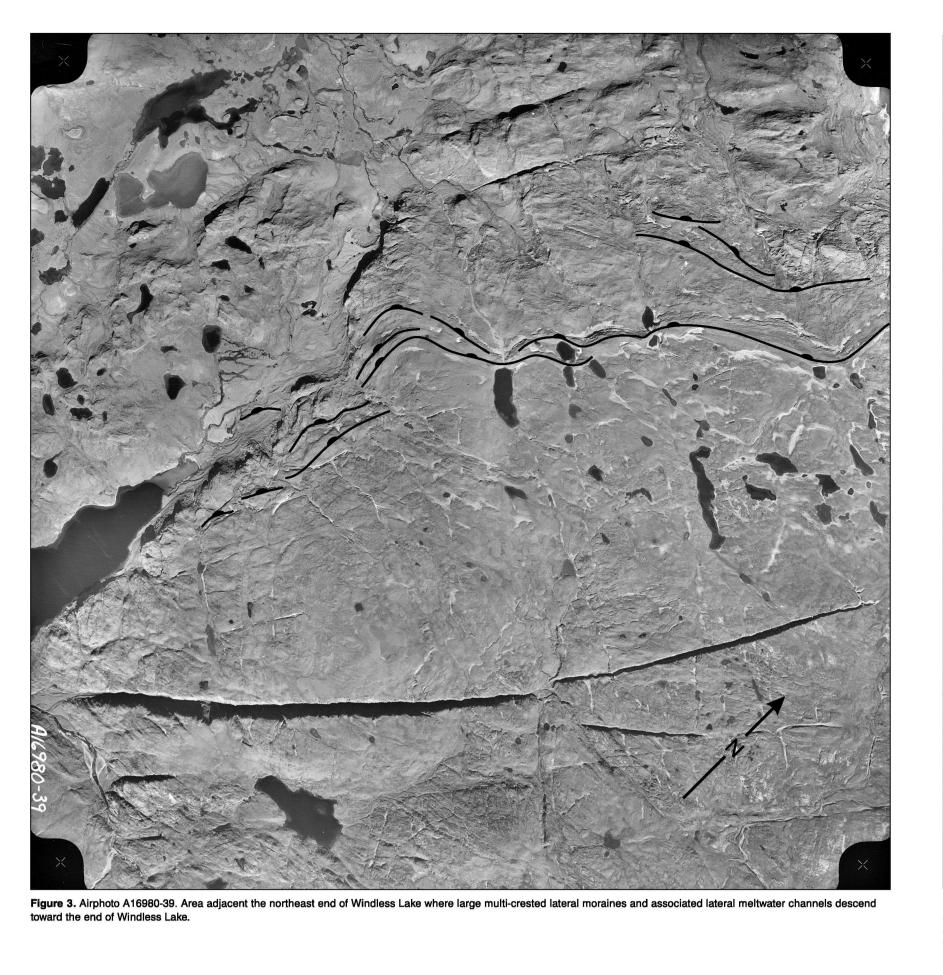
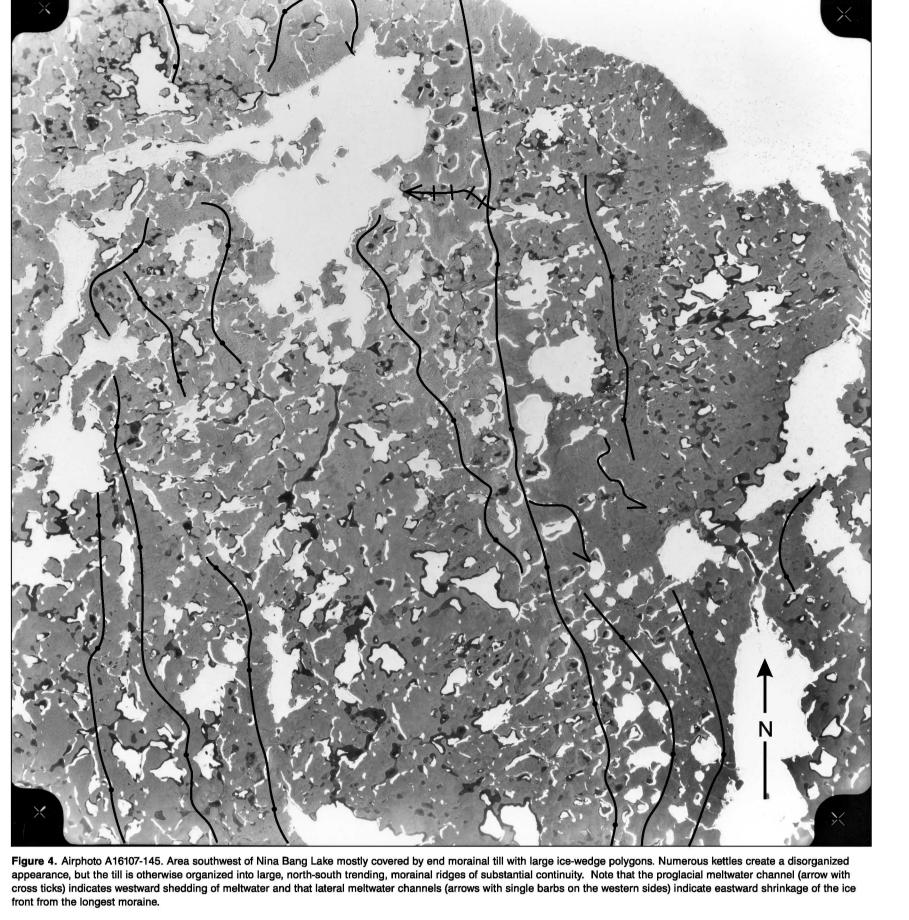


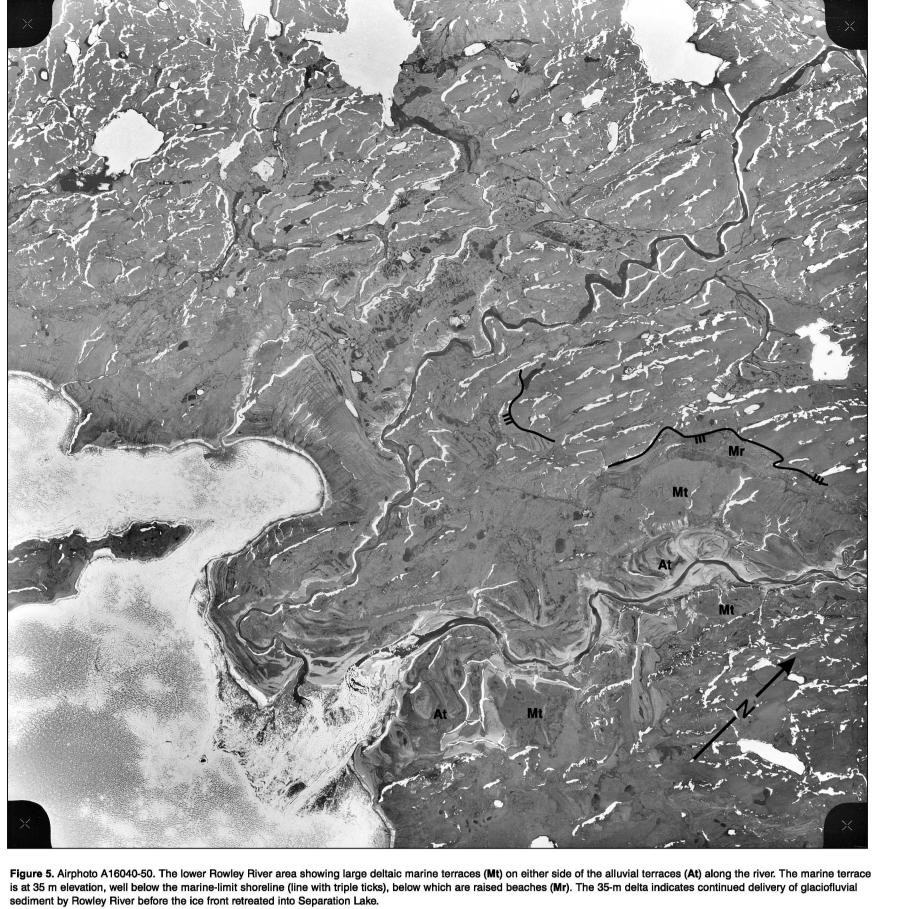
with open triangular teeth. Note the esker (> > symbols) feeding directly into the delta. Marine limit on either side of the delta is shown by the line with triple ticks and

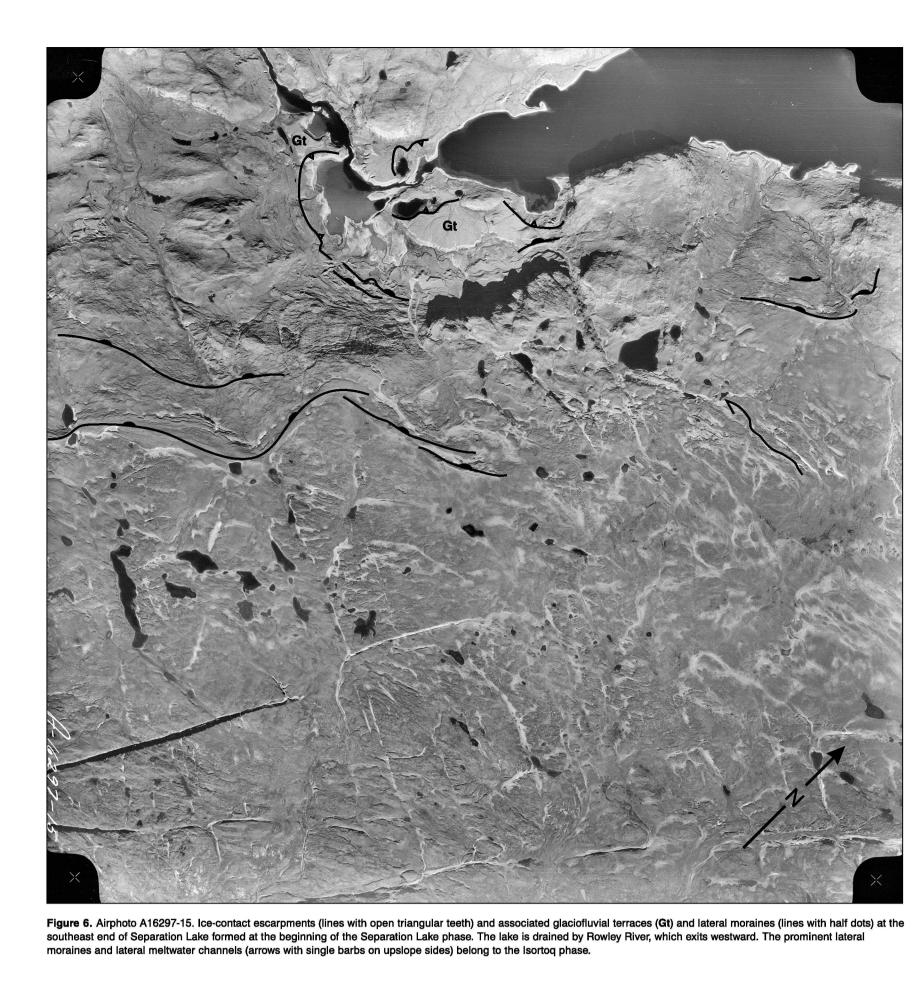
raised beaches (Mr) extend to this limit in places. Lines with two or more dots are end moraines. All map symbols and labels as on GSC Map 2095A (Dyke, 2005g).

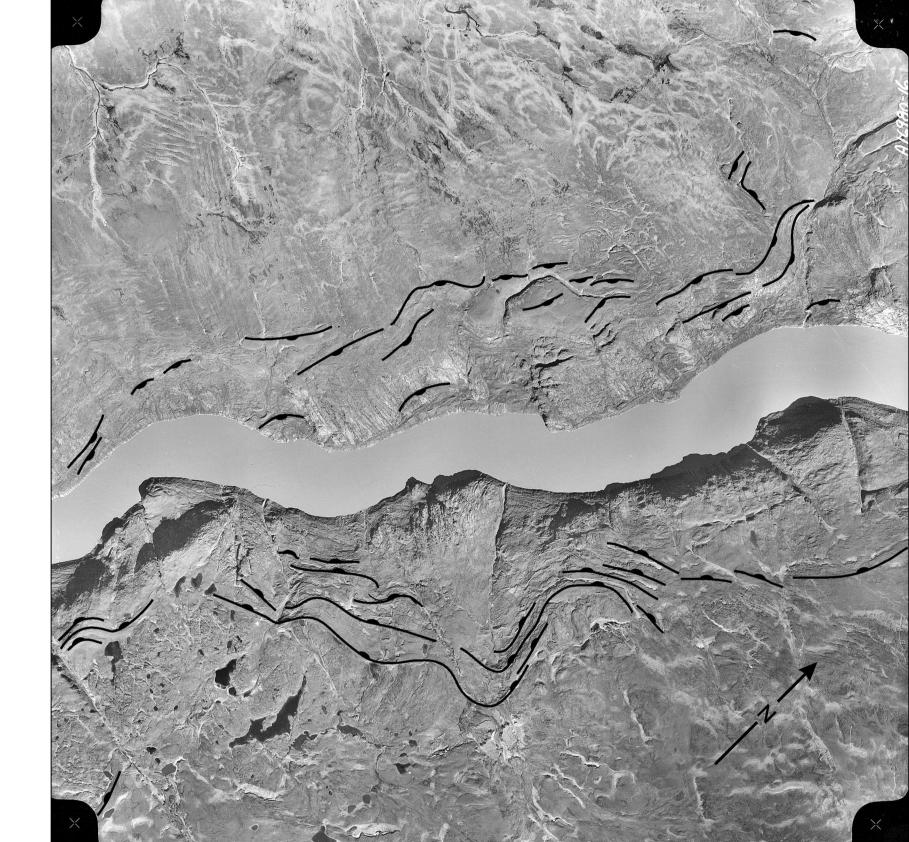












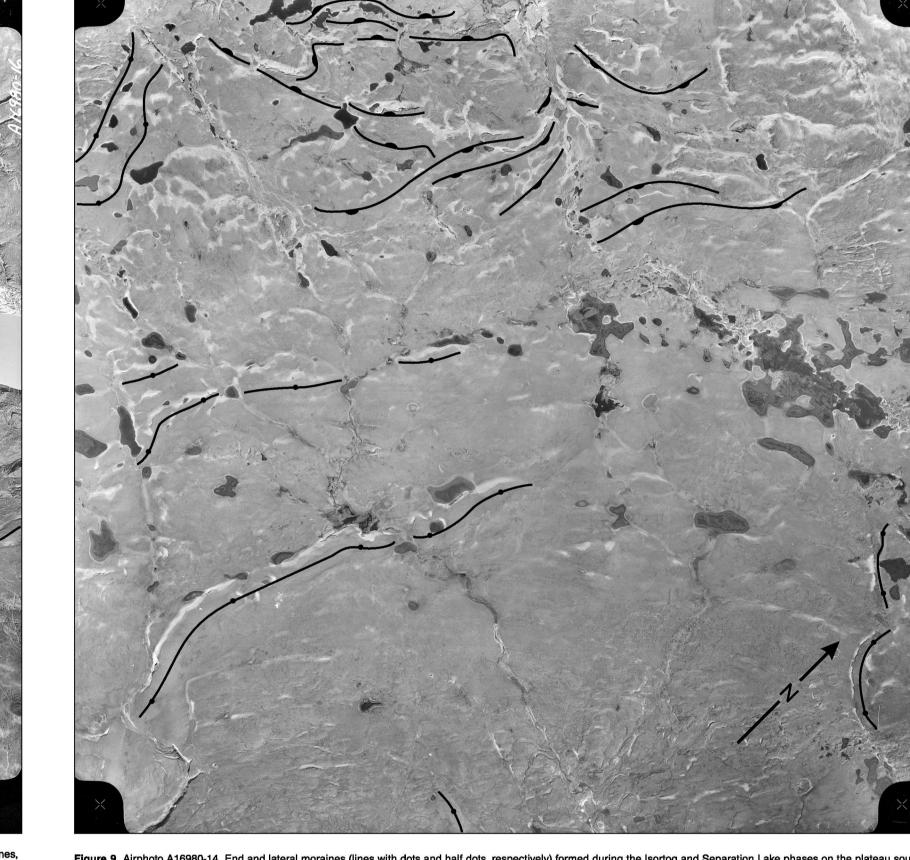
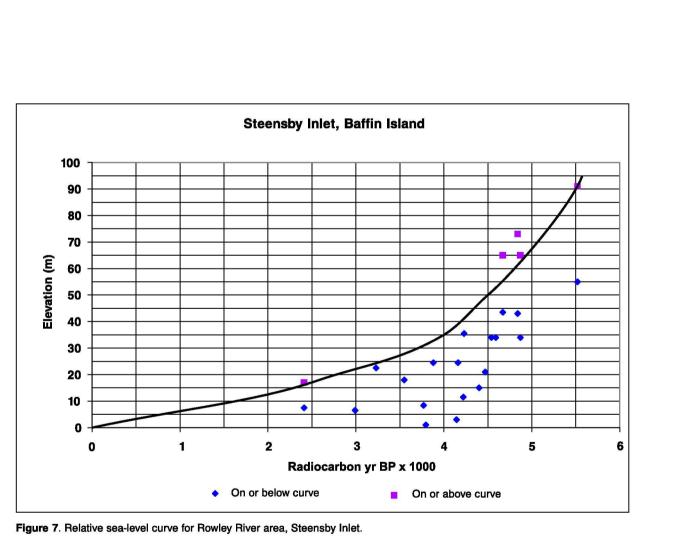
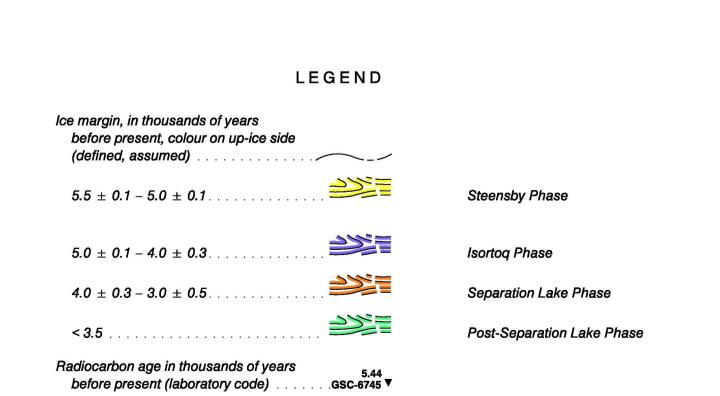


Figure 8. Airphoto A16980-16. Multiple lateral moraines (lines with half dots) formed high along the Isortoq Lake trough during the Separation Lake phase. The highest moraines,

Figure 9. Airphoto A16980-14. End and lateral moraines (lines with dots and half dots, respectively) formed during the Isortoq and Separation Lake phases on the plateau southeast





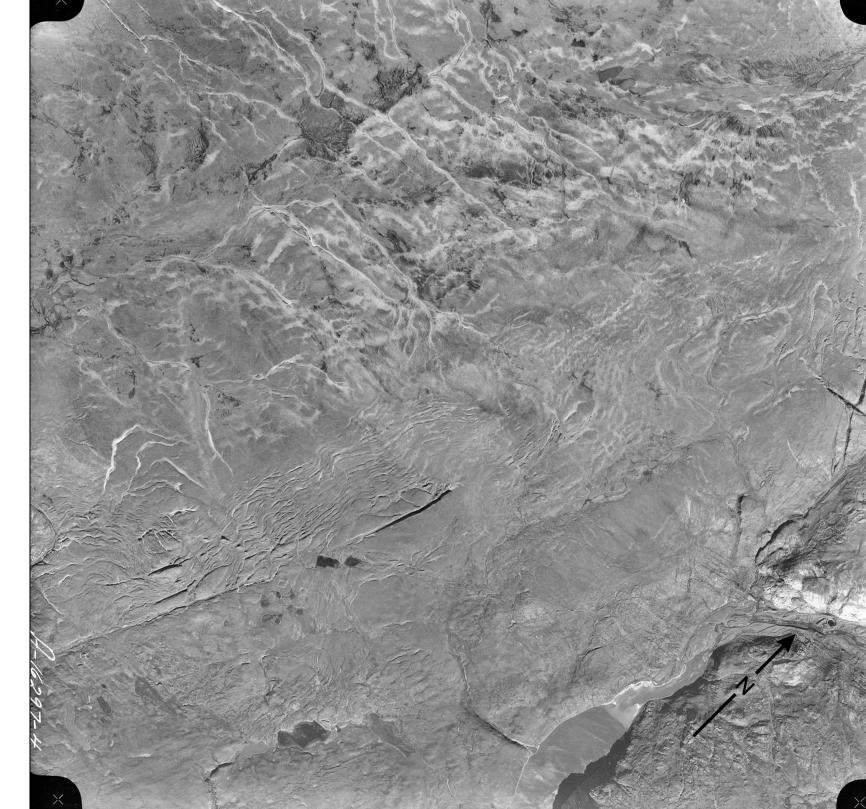


Figure 10. Airphoto A16297-4. Closely spaced lateral meltwater channels inscribed in till on gentle plateau slopes in the north-central part of the map area formed during the

Isortoq and Separation Lake phases. The ice margin is thought to have been cold based during formation of such features (Dyke, 1993).

## DESCRIPTIVE NOTES

This map at 1:250 000 scale displays the pattern of ice-marginal recession during the middle Holocene on part of northern Baffin Island. It is derived from 1:50 000 scale surficial geology maps of the same area by the author (Dyke, 2005a-p). A somewhat generalized compilation of these maps is also available at 1:250 000 scale (Dyke, 2005q). The pattern of deglaciation of all of Baffin Island west of 80°W is shown on a map at 1:500 000 scale (Dyke and Hooper, 2001). The notes below are best read as a companion of the 1:250 000 scale surficial geology map

Steensby Inlet is the northernmost arm of Foxe Basin, a shallow inland sea that occupies a Paleozoic carbonate basin. West and north of the inlet is a lowland, much of which was inundated by the postglacial sea, which generally attained elevations of 50–100 m. An escarpment along the which typically attains elevations of 400 to 600 m in the map area. The plateau is traversed by unnamed rivers. Over-deepened segments of these valley are occupied by fiord-like lakes. The plateau is largely drift mantied and ice-marginal landforms are common. During the last glaciation, which started about 120 000 years ago, this part of Baffin Island was covered entirely by the Laurentide Ice Sheet. The prevailing concept is that the plateau of Baffin Island was an area of ice-sheet nucleation at the beginning of the last glaciation (Ives and Andrews, 963) and that the inland sea of Foxe Basin was slowly filled by glacier ice flowing from Baffin Island and from Melville Peninsula in the early part of the last glaciation. It is further thought that one of the three major domes of the Laurentide Ice Sheet during the Last Glacial Maximum, about 18 000 adiocarbon (<sup>14</sup>C) years ago, was centred over Foxe Basin, south of the map area. At that time, ice is thought to have flowed across the map area from the Foxe Dome northward to Milne Inlet and northeastward to Buchan Gulf and adjacent fiords on the north and northeast coasts of Baffin Island, respectively (Ives and Andrews, 1963; Dyke and Prest, 1987; Dyke, 2004b). Carbonate erratics, which are spread northeastward from Steensby Inlet to the northeast corner of the map area, are evidence of this flow. During deglaciation, mainly between 7000 and 6000 <sup>14</sup>C years ago (Dyke et al., 2003; Dyke, 2004b), Foxe Basin became free of glacier ice as the sea produced a calving bay that propagated

northward across the basin from Hudson Strait. During this process, the ice surface over Foxe Basin was quickly lowered, which forced the ice dome and associated ice divides to shift northward (and landward) until eventually ice flow in the map area was directed southward into Foxe Basin. The mapped ice-flow patterns in the Steensby Inlet map area (Dyke, 2005q) and in the Erichser Lake map area to the west (Dyke, 2004a) date entirely from deglacial phases. Amongst the flow patterns is a strong convergence of flow toward and through the Steensby Inlet lowland, which is suggestive of a major ice stream filling and exiting the lowland just prior to its deglaciation. Later ice recession involved the shrinkage of an ice cap, the Early Barnes Ice Cap, which nearly covered Baffin Island, back to the present margin of the Barnes Ice Cap, which is located only 60 km east of the Steensby Inlet map area. The Quaternary geology of the area southeast of the Steensby Inlet area, including part of the Barnes Ice Cap, was mapped by Dredge (2004).

This map is one of a series of regional ice retreat maps that are based on similar methods and assumptions (Dyke and Hooper, 2001; Dyke et al., 2003b). The surficial geology, including all icemarginal landforms, was mapped on scale 1:60 000 airphotos, which were then scanned, georectified, and transferred to 1:50 000 scale topographic bases. The ice-marginal features were then traced and correlated at that scale on the basis of topographic position and extrapolation of ice-surface gradients along valley sides. Correlation here simply means the lateral connection of mapped ice-marginal features that are judged to be of the same age. This method allows reasonably secure, though by no means precise, correlation of features from valley to valley. orrelations across greater distances, for example across the width of Steensby Inlet, were made at :250 000 scale. These correlations are based on the assumption that ice lobes remained roughly symmetrical as they retreated, and on the assumption that the ice margins trended normally to the last ice-flow direction where that last flow was warm based as evidenced by mapped ice-flow features. The ages of ice margins are based on radiocarbon dates on marine molluscs, to which a local marine reservoir correction of -630 years is applied (Dyke et al., 2003a; McNeely et al., 2006), and on the estimated ages of shorelines based on local relative sea-level curves.

Ice recession is subdivided into four somewhat arbitrary phases. They are discussed below from

The Steensby phase (yellow) encompasses the latter part of the interval during which the Steensby Inlet Ice Stream was still operating. Early in this phase, a large ice-contact delta and associated er moraines (Dyke, 2005q) were deposited west of Steensby Inlet (Fig. 1). The delta top indicates that the marine-limit shoreline at this time was at 95 m above present sea level. The oldest dated shells from the area in front of the delta have a reservoir-corrected age of 5440  $\pm$  70  $^{14}$ C years BP (GSC-6745 on Mya truncata) and the relative sea-level curve for the Cape Thalbitzer-Koch Işland region (Dyke and Savelle, 2006) indicates that the 95-m shoreline should be about 5500 <sup>14</sup>C years old (6300 calendar years old). Lateral moraines on the east side of Steensby Inlet south of Fellside Lake are considered here to be correlative. A marine-limit delta deposited by outwash issuing southward from Fellside Lake has an uppermost terrace at about 88.5 m (Dyke, 2005q), which might indicate a slight eastward dip on the 5500 <sup>14</sup>C year BP shoreline. Andrews (1966) identified a marine-limit beach in this area at 91 m, which is probably a more accurate measurement because it was determined by levelling rather than by altimetry. The middle Steensby phase ice margin, which intersects the western map boundary at about 70°10'N, correlates westward, according to Dyke and Hooper (2001), with the largest end moraine in the Erichsen Lake (Dyke, 2004a) and Phillips Creek (Dyke, 2000) map areas (illustrated in Fig. 13 of Dyke and Hooper, 2001). This large end moraine, which is also figured on the Glacial Map of Canada (Prest et al., 1968) and which outlines the northwest nose of the Early Barnes Ice Cap, is centred on the Gifford River, which carried the greatest amount of outwash away from it. The name Gifford Moraine is here proposed for this and Hooper (2001) assigned it an age of less than 6000 <sup>14</sup>C years but placed it well behind their 6000 year BP ice margin. The new radiocarbon dates from Steensby Inlet indicate an age of slightly less than 5500 <sup>14</sup>C years. An earlier age determination on shells from a site west of Windless Lake (I-486; Ives, 1964;

Andrews, 1966; Andrews and Drapier, 1967) is somewhat problematic for the chronology proposed above. Surface shell fragments yielded an age of 5750 ± 250 years BP. If we assume that this age BP and the reservoir-corrected age would be 5520  $\pm$  250 years BP or 5520  $\pm$  500 years BP, if we use the two-sigma uncertainty that is applied to GSC dates. If we accept the reservoir-corrected mean date of I-486, all of the Steensby phase ice margins illustrated on the accompanying map are compressed into an unmeasurably short interval at about 5500 <sup>14</sup>C years BP. However, the poor quality of the shell sample and the large uncertainty on the age determination suggest that little use can be made of it at this time and an age closer to 5000 <sup>14</sup>C years BP is preferred for the end of the Steensby phase. By the end of the phase, the northwest nose of the Early Barnes Ice Cap had receded into the Steensby Inlet map area.

The Isortog phase (purple) is a name retained from Andrews (1966), who mapped conspicuous sets of end and lateral moraines in the vicinity of Isortoq River (Fig. 2) in the southeast part of the Steensby Inlet map area and correlated them with moraines farther to the southeast. Although the present mapping of moraines (Dyke, 2005q) differs in detail from that of Andrews (1966), both in umbers of moraines and in local correlations of moraine segments, these differences are probably due to the different scales of base maps used and to the availability now of georeferencing tools, which allow more accurate portrayal. All features of the Isortog phase of Andrews (1966) near Isortoq River (indicated by the blue line in his Fig. 3) are deliberately included in the Isortoq phase on the current map. The moraines of this phase are well developed in the Isortog River region, which Andrews (1966) designated as the type area, particularly south and east of the river (Fig. 2) and along moderately steep slopes between Separation Lake and Windless Lake (Fig. 3). Whether they are lateral moraines on valley sides or end moraines on intervening plateaux, they are typically multi-crested accumulations. Early in the Isortoq phase, ice extended to the mouth of the Rowley River, as indicated by the lateral moraines near the north end of Windless Lake. These moraines are at elevations well above that of the lowland to the west, which indicates that the lowland would have been ice covered when they were forming, contrary to the configuration illustrated in Andrews (1966, his Fig. 2). West of Nina Bang Lake, in the northwestern part of the map area, extensive errain is mantled by thick end morainal till, which is extensively kettled but otherwise organized into north-south trending ridges (Fig. 4). This belt of end moraines is here placed in the early part of the Isortoq phase. Proglacial meltwater channels shed westward from the ridges and lateral meltwater channels consistently indicate thinning of ice lobes into topographic lows and (or) general eastward These bulky, kettled, end moraines around Nina Bang Lake are conspicuous on airphotos and

Isortog phase

from the air, because they are ubiquitously patterned by large (100-m scale), rectilinear, ice-wedge polygons. This is the most common end moraine morphology in arctic Canada (Dyke and Evans, 2003). Where spectacularly developed, as along Bernier Bay and Berlinguet Inlet farther west on Baffin Island (Dyke and Hooper, 2001), they are consistently interpreted as end moraines (e.g., Craig, 1965; Falconer et al., 1965; Dyke and Hooper, 2000; Hooper and Dyke, 2000). However, where the morainal helts are broad and more extensively kettled, they are sometimes interpreted a ice-stagnation terrain, and this was the interpretation applied by Ives and Andrews (1963) to the terrain around Nina Bang Lake. Although extensive kettling is commonly taken as evidence of meltout of buried glacier ice following regional ice stagnation in areas south of permafrost, this process is not applicable in the continuous permafrost zone, because there has been no opportunity for deeply buried ice to melt during postglacial time unless it was submerged by water. Instead, the kettles are here interpreted as areas within generally ice-cored moraines where the ice was either left exposed or was buried by debris that was thinner than the active layer. The ice flow direction in the northwest part of the map area during the formation of the moraines west of Nina Bang Lake must have been generally westward. However, there are no westward flow bears south-southeastward flow traces dating from the previous flow of the Steensby Inlet Ice by a broad field of ribbed moraine, which also indicates southward ice flow. The westward ice flow to he moraines, therefore, must have been cold based in order for the earlier bedforms to have survived unaltered. Although the bulky end moraines may be extensively ice cored, the morainal lebris indicates that the cold-based ice still retained a substantial basal debris load, inherited from its earlier warm-based phase. Climbing flow is required in the terminal ice zone in order to form icecored moraines by elevating basal debris. Only in a small area between the head of Steensby Inlet and Neergaard Lake are there faint southwestward ice flow traces on till that formed under warmbased ice as the nose of the Early Barnes receded into the map area (Dyke, 2005q) Ice recession midway through the Isortog phase exposed the lower Rowley River valley. The narine-limit shoreline there is exceptionally clear (Fig. 5, Ives, 1964) and is at an elevation of 65 m. Shells from deepwater sediment exposed along Rowley River, which probably were deposited when the sea stood at the local marine limit, have reservoir-corrected ages of 4670  $\pm$  80  $^{14}$ C years BP (GSC-6713 on Hiatella arctica) and 4870  $\pm$  130  $^{14}$ C years BP (GSC-6761 on Portlandia arctica). Because Portlandia shells and shells of other deposit feeders in areas of calcareous drift commo date somewhat older that shells of suspension feeders such as Hiatella arctica (Dyke, 2004b; McNeely et al., 2006), the younger of the two dates is considered the more reliable age estimate for deglaciation. Hence, the beginning of the Isortoq phase probably dates to about 4800-4900 14C years BP, which is compatible with the preferred age stated above of 5000 14C years BP for the end of the Steensby phase but is substantially younger than the age of 5500–7000 <sup>14</sup>C years BP for the Isortog phase deduced by Andrews (1966). There is no direct age control for the end of the Isortog phase (an age of 4000 <sup>14</sup>C years BP is suggested below), and there are no marine-limit features associated with the youngest ice margins assigned to this phase. End and lateral moraines continued to form during intermittent recession throughout the phase.

Separation Lake phase The Separation Lake phase (brown) started with deposition of ice-contact gravels and associated lateral moraines at the south end of Separation Lake (Fig. 6) and the deposition of a large outwash train extending westward from there along the Rowley River. Outwash sediment continued to pour along this train until the ice margin retreated into the sediment trap of Separation Lake. A large marine terrace along the lower Rowley River at 34 m elevation, well below the 65-m marine lin may have formed during this period of sediment delivery (Fig. 5). If so, an age of about 4000 14

years BP is indicated for the early part of the Separation Lake phase, based on a relative sea-level curve for the Rowley River area (Fig. 7). The most prominent features assigned to the Separation ake phase are nested sets of lateral moraines high along the vallev sides along Isortog Lake (Fig . Correlative end moraines occur on the plateau to the southeast (Fig. 9). Here the ice seems to have been warm based during or just before moraine formation as indicated by fluted till, including oss-cutting till flutings, and by small eskers, and corridors of hummocky glaciofluvial gravel. Most ice-marginal features in the north-central part of the map area are placed in the Separation Lake phase. Here the northwest nose of the Early Barnes Ice Cap appears to have treated much faster than its western side, probably because the ice was thinner and was flowing divergently over this high ground. Ice-marginal features in that region are predominantly lateral meltwater channels (Fig. 10), which are common in areas where the ice-marginal zone was cold based during recession (Dyke, 1993). However, recessional moraines occur on the valley floor along the middle reaches of the Rowley River. The upper part of the Rowley River valley within the nap area and beyond was dammed by a southwestward retreating ice lobe, which retained a lake neir Fig. 1). The lake was about 100–120 m deep at its maximum and it probably drained in a series ce-marginal channels as the ice lobe in the Rowley River valley shrank and thinned late in the

The post-Separation Lake phase (green) ice margins, the youngest in the map area, illustrate the sequential retreat of the northwest nose of the Barnes Ice Cap eastward and out of the map area. A ubstantial moraine belt east of the upper end of Isortoq Lake is placed in this phase. Elsewhere, he pattern of recession is only sparsely defined by scattered lateral meltwater channels. Age he combined sequence probably started about 4000 <sup>14</sup>C years BP. Extrapolation of previous rates fice recession would suggest that the youngest ice-marginal features in the map area are about

The pattern of ice-marginal recession shown on this map is the expected one. It outlines the recession of the Early Barnes Ice Cap in greater detail than previously shown (Ives and Andrews, 1963; Andrews, 1966; Dyke, 1974) and adds age control to that event. Nevertheless, the chronology of recession after 4000 <sup>14</sup>C years BP remains unconstrained. As in areas farther northwest on affin Island (Dyke and Hooper, 2001), the pattern of ice-marginal recession can be reconstructed here with a resolution that is decadal to centennial. The large number of substantial end and lateral moraines that were deposited on Baffin Island during the Holocene makes it difficult to isolate single moraines as being particularly significant from a paleoclimatic point of view. For example, although the early Isortog phase moraines are well developed in their type area along the Isortog River, numerous equally well eveloped moraines were deposited along the same valley later during Isortog phase recession. during the two subsequent phases, and during recession across the map area to the east (Ives and Andrews, 1963). The moraines of the Steensby phase are equally well developed, or more so if we include the Gifford Moraine. No single moraine would seem to represent a great deal of time, for there are too many to be accommodated, and there are no cross-cutting relationships to indicate ignificant readvances, though smaller ones probably did occur. If the moraines are treated as vidence of stillstands or slight readvances, as seems reasonable, what emerges is a view of the Early Barnes Ice Cap as one whose mass balance was never more than slightly negative on a decadal to centennial average and was frequently positive on average for decades. Nevertheless, i moraine ages could be established precisely and moraines thus correlated more precisely than is now possible, a high-resolution mass balance history of the Holocene Barnes Ice Cap might be This view is rather different from that of Andrews (1966), who "...stressed that no major noraine systems occur on the immediate proximal or distal side of [the Isortoq] moraine system."

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Furthermore, all marine-limit features in the area appear to have formed contemporaneously with

ransgression brought on by Isortoq phase glacial readvance, as previous suggested by Andrews

ice recession, as is normally the case in North America, rather than during a subsequent marine

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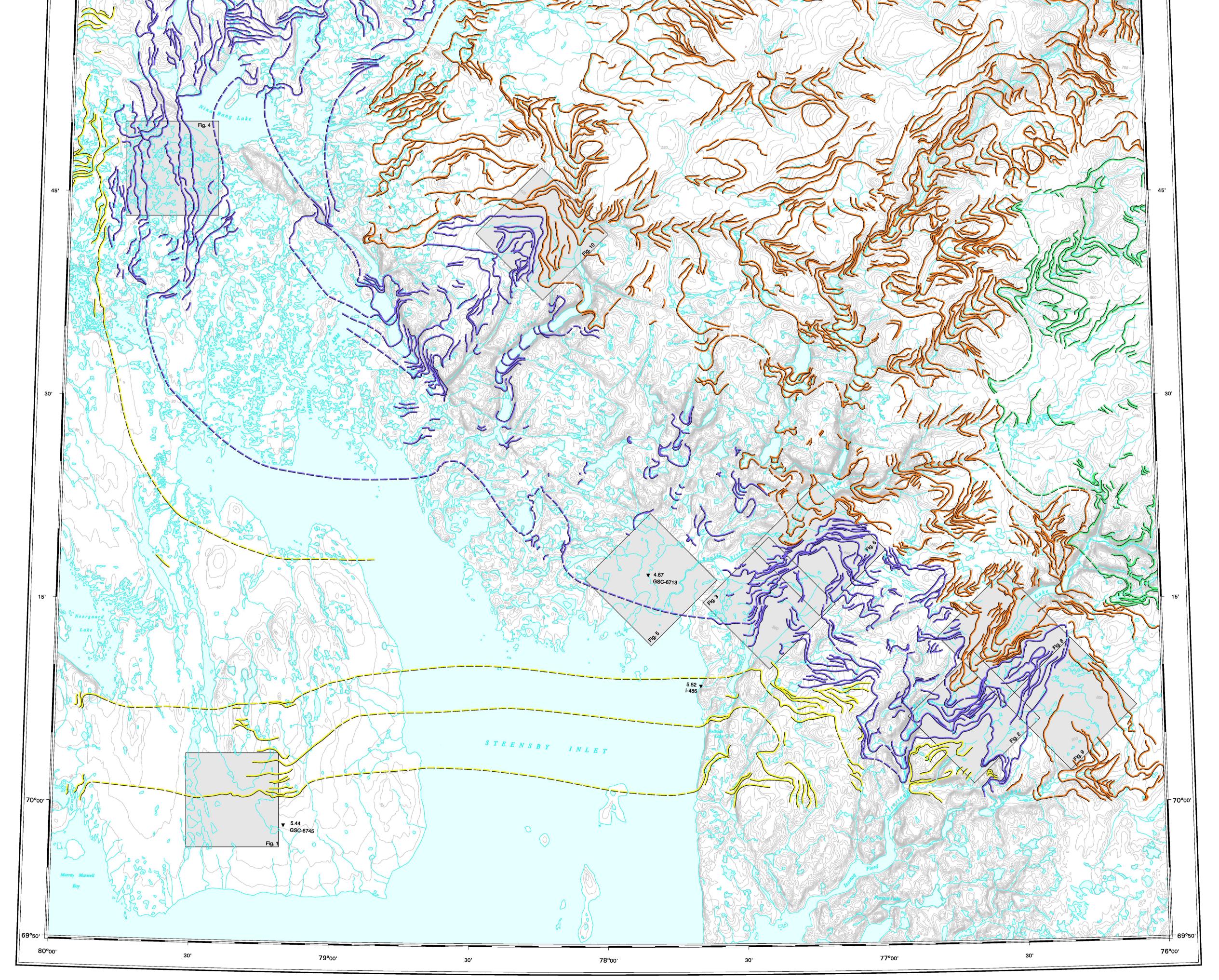
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Author: A.S. Dyke Geology by A.S. Dyke, 2002-03 Geological compilation by A.S. Dyke, 2006 Digital cartography by J.D. Narraway, Data Dissemination Division (DDD) This map was produced from processes that conform to the Scientific and 「echnical Publishing Services Subdivision (DDD) Quality Management System,

registered to the ISO 9001: 2000 standard

**DEGLACIATION OF STEENSBY INLET** BAFFIN ISLAND Scale 1:250 000/Échelle 1/250 000 Projection transverse universelle de Mercator North American Datum 1983 Système de référence géodésique nord-américain, 1983 © Her Majesty the Queen in Right of Canada 2007 © Sa Majesté la Reine du chef du Canada 2007

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Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada Digital base map from data compiled by Geomatics Canada, modified by DDD Proximity to the North Magnetic Pole causes the magnetic compass to be erratic in this area Mean magnetic declination 2007, 39°55' W, decreasing 47.5' annually. Readings vary from 36°42'W in the SW corner to 43°12'W in the NE corner of the map

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