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GEOLOGICAL SURVEY OF CANADA | COMMISSION GÉOLOGIQUE DU CANADA

DESCRIPTIVE NOTES

INTRODUCTION

This 1:100 000-scale map and the accompanying Open File 1794 (Hanmer 1 (NTS 56 M (North)) present results of bedrock mapping undertaken by the Canada-Nunavut Geoscience Office and Geological Survey of Canada under the Canada-Nunavut Geoscience Agreement (2000–2004). The purpose of the project was to provide a new level of geoscience knowledge of this prospective Committee Bay area (Fig. 1) and provide modern, 1:100 000 scale geological maps for the region. The Ellice Hills map area (NTS 56 P/ is underlain by northeast-trending rocks of the Archean Rae Province. The Rae Province in this area consists of a variety of granitoid rocks that record the association of supracrustal and ironstone. Previous investigations of the Committee Bay belt (Figs. 1,2), include reconnaissance mapping (1:506 880) of Heywood (1962) and the subsequent 1:250 000 scale mapping in parts of the Committee Bay belt by Eade (1985, 1986, 1987, 1988, 1989, 1990, 1991, 1992). This mapping was undertaken as a new national park in Nunavut (Fig. 1), including the headwaters of the Ellice River, but is a mineral resource assessment of the Committee Bay belt. Mapping in the Ellice Hills area was undertaken as part of the Committee Bay Targeted Geoscience Initiative Project, bedrock maps and related petrological and structural studies (Fig. 1). Other recent studies in the Committee Bay belt include Sanborn-Barrie et al., 2000; Skulski et al., 2000; Machado, 2002; Chapter Lake and Arrowsmith River (56 J/5 and 56 O/8), respectively (Sanborn-Barrie et al., 2000; Skulski et al., 2002; Skulski et al., 2003a), and S-Bibba Hills (56 P, this map; Sanborn-Barrie et al., 2003; Skulski et al., 2003b). Geological mapping in the Committee Bay belt is part of the Slave–Rae–Thelon Tectonic Project (Berman et al., 2003; Tella and Eamon, 2003), and thematic studies, as are results of surficial mapping and drift prospecting (Utsi, 2001; Utsi et al., 2002; Utsi and Kerr, 2003). Airborne geophysical data for the region consist of 800 m line-spaced aeromagnetic surveys that crossed 80,260nN/m, was flown along north-northeast-trending airline-perpendicular flight lines (Geological Survey of Canada, Committee Bay Project, Aeromagnetic Survey, Open File 4243, 1:50 000; Machado, 2002). Chapter Lake and Arrowsmith River (56 J/5 and 56 O/8), respectively (Sanborn-Barrie et al., 2000; Skulski et al., 2002; Skulski et al., 2003a), and S-Bibba Hills (56 P, this map; Sanborn-Barrie et al., 2003; Skulski et al., 2003b). Geological mapping in the Committee Bay belt is part of the Slave–Rae–Thelon Tectonic Project (Berman et al., 2003; Tella and Eamon, 2003), and thematic studies, as are results of surficial mapping and drift prospecting (Utsi, 2001; Utsi et al., 2002; Utsi and Kerr, 2003). Airborne geophysical data for the region consist of 800 m line-spaced aeromagnetic surveys that crossed 80,260nN/m, was flown along north-northeast-trending airline-perpendicular flight lines (Geological Survey of Canada, Committee Bay Project, Aeromagnetic Survey, Open File 4243, 1:50 000; Machado, 2002). These datasets have been integral in the production of the map, particularly in regions of extensive Quaternary drift.

REGIONAL GEOLOGICAL SETTING

The southern mid-Rae domain of the western Churchill Province, long-west of Hudson Bay, contains a discontinuous, ca. 1800 km long, northeast-trending zone of polydeformed and metamorphosed Archean supracrustal belts and associated granitoid rocks (Fig. 1). Supracrustal belts in this zone, including the Committee Bay belt, are dominated by supracrustal rocks are inferred to have formed in a continental setting, reflected by local evidence for deposition on continental crust (Eade, 1985, 1986, 1987, 1988, 1989, 1990, 1991, 1992; Sanborn-Barrie et al., 2000). Supracrustal rocks occur in widespread quartzite (eg., Ashton, 1984; Davis and Zaleski, 1998; Skulski et al., 2003b), volcanic rocks within this northeast-trending corridor include komatiite, basalt, andesite, dacite, and rhyolite. The latter two (utilizing Th/Hf–Ta/Hf ratios) are interpreted to have formed in a continental setting (Eade, 1985, 1986, 1987, 1988, 1989, 1990, 1991, 1992; Sanborn-Barrie et al., 2000; and references therein; Skulski et al., unpublished). Archean plutonic rocks in the region appear to be dominated by calc-alkaline granitoid plutons emplaced between ca. 2.64 and 2.58 Ga (Ashton, 1986; LeCheminant et al., 1991; Sanborn-Barrie et al., 2000; Skulski et al., unpublished; Machado, 2002).

Sutton, in possibly figure of platform to basinal facies. Paleoproterozoic metasedimentary rocks are found in the Committee Bay belt, and in the Rankin Inlet belt, to the west (Ashton, 1984; Davis and Zaleski, 1998; Skulski et al., 2003b). Volcanic rocks are dominated by calc-alkaline granitoid plutons emplaced between ca. 2.64 and 2.58 Ga (Ashton, 1986; LeCheminant et al., 1991; Sanborn-Barrie et al., 2000; Skulski et al., unpublished; Machado, 2002).

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In the Committee Bay area, Paleoproterozoic intrusive rocks include late to post-tectonic ca. 1.835 Ga calc-alkaline granitoid plutons (Ashton, 1986; Sanborn-Barrie et al., 2000; Skulski et al., 2003b), and ca. 1.83 Ga magnetite–ilmenite monzogranites of the Kivalliq suite (this map area (Skulski et al., 2000b); Flotorin–Uvlik equivalent to the ca. 1790 Ma Nueltin Suite (Peterson et al., 2002) have not been documented in this part of the Rae domain, however, a small (ca. 1 km diameter) pluton of granite, dated at 1.835 Ga, was encountered in drill core in the Ellice Hills area, and is inferred to be correlative with the Nueltin Suite. The northernmost mapping, ca. 1,200 Ga gabbroic Murdoch dykes cut all Archean and Proterozoic rocks.

GEOLOGY OF THE COMMITTEE BAY AREA

The Committee Bay area is generally underlain by three, lithologically distinct, northeast-trending crustal domains (Sanborn-Barrie et al., 2000; Fig. 2). The Rae domain, in the northwest, consists of a variety of granitoid rocks and a metasedimentary belt of supracrustal rocks belonging to the Prince Albert group (Pkg), herein called the Committee Bay supracrustal belt, that extends from the Laughland Lake area in the southwest to the Curtis River area in the northeast. Supracrustal rocks in the Committee Bay belt are dominated by quartzite (Eade, 1985, 1986, 1987). The sheet are mainly exposed in the southern part of the map sheet, although thin, discontinuous sills, lenses and pendants of quartzite are found throughout the area. Quartzite is commonly interlayered with pelitic to semipelitic psammite and quartz arenite, less common iron-formation and ultramafic rocks and uncommon pelagites, including graphitic pelite and chemical sediments. Quartzite is fine- to medium-grained, commonly massive to northwest and southeast to medium-grained, variably foliated (ion- to post-tectonic) and metamorphosed to the supracrustal rocks. Collectively, these rocks are intruded by weakly extensive, equigranular and microlite-bearing granite and pegmatite (Sanborn-Barrie et al., 2000).

The southern domain is underlain by the regionally extensive, pinkish, Walker Lake intrusive complex, which is dominated by potassium-feldspar porphyritic to augen granodiorite and monzogranite. The areal extent of the Walker Lake intrusive complex is delineated by Skulski et al. (2000b, 2002, 2003a, b) and Machado (2002). The Walker Bay supracrustal belt is underlain by voluminous, homogeneous, moderately foliated, and intruded potassium-feldspar porphyritic granodiorite and monzogranite of the Walker Lake intrusive complex. The Walker Bay supracrustal belt is underlain by voluminous, homogeneous, moderately foliated, and intruded potassium-feldspar porphyritic granodiorite and monzogranite of the Walker Lake intrusive complex. The Walker Bay supracrustal belt is underlain by voluminous, homogeneous, moderately foliated, and intruded potassium-feldspar porphyritic granodiorite and monzogranite of the Walker Lake intrusive complex. The Walker Bay supracrustal unit, is widely intruded by a range of younger, variably foliated granitoid through pyroxenegranite, Barrie et al., 2000; Skulski et al., 2000b, 2002, 2003a, b; Machado, 2002). Rocks of the Walker Lake intrusive complex are dominated by potassium-feldspar porphyritic granodiorite and monzogranite (Skulski et al., 2000b).

The metamorphic grade in the main supracrustal belt ranges from upper greenschist in the southeast to amphibolite facies in the northwest. The Walker Lake intrusive complex is dominated by rocks that are inferred to be a lack of rocks that offer diagnostic mineral assemblages. However, to the northwest (NAD), the grade increases to upper amphibolite facies with consequent development of metasedimentary diatexites and paragneisses. Metamorphic grade in the