Natural Resources Canada | Ressources naturelles Canada

GEOLOGICAL SURVEY OF CANADA | COMMISSION GÉOLOGIQUE DU CANADA

DESCRIPTIVE NOTES

INTRODUCTION

The Committee Bay belt is a 650 km long belt of northeast-striking, Archean supracrustal rocks and granitoid plutons that form part of the Rae Province in central Nunavut (Fig. 1). The Archean supracrustal rocks of the belt belong to the Woodburn Lake Group (Fig. 2) and host the Meadowbank gold deposit and the Amer gold deposit (Fig. 3), both a product of the 2000-2002 Targeted Geoscience Initiative that was instigated by the Geological Survey of Canada, in collaboration with the Canada-Nunavut Geoscience Office, to encourage mineral exploration in the Committee Bay area. The project involved an aeromagnetic survey, helicopter-supported, 1:100,000 scale bedrock mapping, and a study of the Amer gold deposit. This Open File is a digital publication of the bedrock mapping studies and a drift prospecting survey. The aeromagnetic survey covered 69,300 line km and was flown along 400 m spaced, northeast - southeast flight lines (perpendicular to strike) on a pre-calculated drape surface with a mean terrain clearance of 300 m. The survey was flown by Fugro Airborne Surveys in 2001. Results of the bedrock and petrological/structural studies are presented in the following: 1) Laughland Lake area (NW) - Johnstone (2002), Johnstone et al. (2003), MacHattie (2003) and Sanderson et al. (2001a-c); 2) Walker Lake and Arrowsmith River (SE) - Sanderson et al. (2001d-f), Sanderson and Panagapko (2002), Sanderson, Thurston et al. (2003); 3) Wilson Hills (SW) - Sanborn-Barrie et al. (2003) and Skulski et al., 2003); 4) Regional petrochemical prospecting results for the study area are presented in Dayell and Sherlock (2003), Hyde et al. (2003) and Sherlock and Dayell (2002, 2003), with the regional geochronological framework established by Sanborn-Barrie et al. (2003), Lobo-Dihn Lake et al. (2003), McMartin et al. (2002, 2003), Dyer and Hynes (2003) and Ulrich and Wood (2003).

REGIONAL GEOLOGICAL SETTING

The eastern part of the Rae Province contains a discontinuous, 2690 km long, northeast - trending zone of Archean supracrustal rocks and granitoid plutons (Fig. 1; Sanborn-Barrie et al., 2003). The zone is characterized by reflective N-trending fabrics that are discordant on an older suite (oldest?) of granitoid plutons (Ashton et al., 2001; Zaleski et al., 2001) and the Paleoproterozoic Hoare Bay Group (Archean - Paleoproterozoic) unconformably overlies the zone (Jackson and Berman, 1999; Hoffman et al., 2001). The latter group, which includes the Ivakdlit, Penrhyn, Lake and Hills (Jackson and Berman, 1999; Hoffman et al., 2001). The latter group, which includes the Ivakdlit, Penrhyn, Lake and Hills (SAP) - Sanborn-Barrie et al. (2003) and Skulski et al., 2003), Chiperwhyan metasedimentary rocks include <2.62 Ga quartzite in the Laughland Lake area and ~2.58 Ga arkose and salt-arenite in the Ellice Hills (SHRIMP U/Pb detrital zircon, T. Skulski, unpublished Fig. 2). Archean supracrustal rocks are metamorphosed to amphibolite facies, have a subsolidus deformation and are locally metamorphosed to greenschist facies.

Archean and Paleoproterozoic strata in the eastern Rae Province have a prominent northeasterly striking tectonic grain that is truncated by the Paleoproterozoic Thelon Orogen (Hoffman et al., 2001; Sanborn-Barrie et al., 2003; Zaleski et al., 2001; Sanderson et al., 2001c; Sanborn-Barrie et al., 2002, 2003). The prominent northeasterly fabric is interpreted to be the result of D2 deformation in the hinterland of the Trans-Hudson orogen, and is locally overprinted by northwest - striking, Paleoproterozoic D3 folds (Fig. 1; Hoffman et al., 2001; Zaleski et al., 2001). D2 deformation was accompanied by regional amphibolite grade metamorphism. Prominent dextral, and locally sinistral-oblique, east-striking shear zones (e.g., the Amer, Wager Bay, and Walker Lake shear zones) overprint D2 fabrics in the eastern and central parts of the belt.

Paleoproterozoic iron-oxide rocks in the Rae Province include lake- to post-tectonic (~1825 Ga old) alkaline granites and granodiorites of the Pond Lake batholith (LaChemiant et al., 1987) and similarly aged biotite – magnetite (-) fluorite veins (Sanborn-Barrie et al., 2003; Sanborn-Barrie et al., 2002, 2003). Northwest-trending, Mesoproterozoic Mackenzie diabase dykes cross-cut all Archean and Hoffman and Hoffman rocks.

GEOLOGY OF THE COMMITTEE BAY AREA

The Committee Bay area contains three northeast-trending supracrustal subdomains (Boulter et al., 2002; 2003; Fig. 2). The Prince Albert group subdomain (central subdomain of Boulter et al, 2002) contains abundant supracrustal rocks (Sanborn-Barrie et al., 2002, 2003; Fig. 2). The Hoare Bay Group subdomain (SE subdomain of Boulter et al., 2002; Sanborn-Barrie et al., 2003), whereas the Hoffman Bay Complex subdomain (NW subdomain of Boulter et al., 2002) contains mainly granitoid rocks. The Prince Albert group subdomain contains the Northern Migmatite and Prince Albert group subdomains south of the Arrowsmith River (Sanborn-Barrie et al., 2002). The Walker Lake intrusive complex intrudes the Prince Albert group subdomain (Fig. 2). The Hoare Bay Group subdomain contains the Wilson Hills metasedimentary rock monogranite, which is cut by late to post-tectonic, ~1821 Ma biotite - magnetite ± fluorite monzogranite (Sanborn-Barrie et al., 2001b; Skulski et al., 2003; Fig. 2). Plutonic rocks of the Walker Lake intrusive complex intrude the Prince Albert group subdomain or are in fault contact with it along the central Walker Lake shear zone (Fig. 2).

WALKER LAKE – ARROWSMITH RIVER REGION

The Walker Lake and Arrowsmith River map area has an excellent bedrock exposure (generally in the order of 90%), with the exception of the Laughland Lake area, which has extensive cover east of Walker Lake (Dyer and Hynes, 2002; Utting and Wood, 2003). Aeromagnetic data, in particular the second vertical derivative, was used to extrapolate geological contacts into areas of overburden cover. The Walker Lake area contains a diverse suite of Archean to Paleoproterozoic supracrustal belts, migmatitic northwest-striking folds and the distribution of non foliation, komatiitic, dioritic metasedimentary and mafic volcanic rocks. The Pkg in the Walker Lake – Arrowsmith River area includes the Penrhyn Group (Jackson and Berman, 1999; Ellis Bay Formation, Fig. 2) to the south and the Northern Hoare River belt (NHRB) to the northwest (Fig. 2).

LOWER AND MIDDLE PRINCE ALBERT GROUP

The lower Pkg is dominated by basalt and komatiite-bearing volcanic successions that show local northwest younging directions (Sanborn-Barrie et al., 2002). Komatiite-bearing volcanic successions in the Prince Albert group comprise: amphibolite biotite, massive to pillowed, non-migmatic basalt (unit Avb) that is overlain by komatiite, sulphidic pelitic sedimentary rocks and a thin upper basalt unit. The komatiite sequence (unit Akv) includes 50 cm to 1 m thick, olivine ± pyroxene spinifex-textured flows, olivine ± pyroxene cumulate flows and local, olivine ± pyroxene spinifex zones, which contain pseudomorphs of blades (olivine up to 10 cm long, Sublimitic pelitic rocks (pyrite and pyrrhotite-bearing) are broadly interbedded with this komatiite. The mafic volcanic sequence is cut by late-tectonic, ~1821 Ma biotite - magnetite ± fluorite monzogranite (Sanborn-Barrie et al., 2001b; Skulski et al., 2003; Fig. 2). Plutonic rocks of the Walker Lake intrusive complex intrude the Prince Albert group subdomain or are in fault contact with it along the central Walker Lake shear zone (Fig. 2).

NORTHERN MIGMATITE SUBDOMAIN AND THE UPPER PRINCE ALBERT GROUP

The Northern Migmatite subdomain is dominated by high-grade metasedimentary rocks (unit Sm) including metabasite (metabasaltic/komatiitic migmatite with 10 to 40% interbanded granitoid melt/dykes, heterogeneous and homogeneous metatexite and diatexite, and local, garnet - cordierite ± anthophyllite - bearing, pelitic gneiss. Local indications of metatexite that are not coloured, migmatite-free quartz - biotite ± garnet ± sillimanite ± K-feldspar ± cordierite ± melt rocks that have either a migmatitic texture and reflect in situ partial melting and extraction, or contain <10% granitoid melt that is interpreted to have been derived from partial melting of the host rock. Local, garnet ± sillimanite - plagioclase ± garnet ± K-feldspar ± cordierite-anatectic (= migmatitic) pelitic gneiss (unit Ams) is interbedded with metatexite. Within the metasedimentary domain are rare wacke enclaves (paleosomes) that locally preserve bedding. One such enclave, located 1.5 km south of the Meadowbank gold deposit, yielded a detrital zircon age of 2.9 Ga (SHRIMP U/Pb, interpreted to be ~2931 ± 16 Ma (map number 3, Table 1; Fig. 2), in northeastern portion of map sheet 56-D, has yielded a SHRIMP U/Pb zircon age of 2800 ± 4 Ma (sample 7, Table 1), whereas a deformed, K-feldspar megacrystic, biotite granodiorite sheet in the Arrowsmith River shear zone, the granite has a weak to moderate foliation and is texturally heterogeneous (map number 4, sample number 8, Table 1; Fig. 2). The granite is interbedded with paramictite deposited between 2687 ± 10 Ma (youngest detrital zircon) and 2677 ± 18 Ma (cross-cutting pegmatoumus granite; SHRIMP U/Pb zircon; Skulski et al., 2003; T. Skulski, unpublished, Fig. 2).

PLUTONIC ROCKS

Intermediate to felsic plutonic rocks in the Prince Albert group subdomain include (from oldest to youngest): biotite tonalite, hornblende - biotite quartz diorite/tonalite, biotite - K-feldspar megacrystic granodiorite, biotite granodiorite monzogranite, and syn- to postkinematic pink syenogranite. Tonalite and granodiorite in the southern part of the Prince Albert group subdomain have yielded crystallization ages of 2690 ± 2 Ma (map number 5, Table 1; Fig. 2). Undeformed, late- to post-tectonic pink syenogranite, which is also a monzodiorite quartz diorite from the northern shore of Walker Lake (Fig. 2) yielded an age of 2630 ± 3 Ma, which is coeval with the Hoarding Wolf biotite tonalite (map number 6, Table 1; Fig. 2). Pegmatoumus granite (unit Agg) and pegmatite (unit Ap) are common throughout the map area and may be genetically related to the homogeneous, biotite - K-feldspar megacrystic (up to 50% by volume and 5 cm in length) granodiorite (unit Agg) bone sheet, which is exposed in the northeastern part of map sheet 56-D. Unit Agg is interpreted to be a syn-tectonic, shear zone. A K-feldspar augen granodiorite, which occurs in southern portion of map sheet 56-D, has yielded a SHRIMP U/Pb zircon age of 2692 ± 4 Ma (sample 7, Table 1), whereas a deformed, K-feldspar megacrystic, biotite granodiorite sheet in the Arrowsmith River shear zone, the granite has a weak to moderate foliation and is texturally heterogeneous (map number 4, sample number 8, Table 1; Fig. 2). The northern half of the Prince Albert group subdomain is dominated by volcanic rocks (unit Av) that are metamorphosed to mainly coarse-grained, foliated, augigranular biotite (± magnetite ± allanite) granodiorite (unit Agg). North of the Ellis, these granodiorite sheets are up to 1 km thick and have yielded crystallization ages of 2690 ± 2 Ma (map number 5, Table 1; Fig. 2) and 2692 ± 4 Ma (sample 7, Table 1; Fig. 2). The Hoarding Wolf stock (2630 ± 3 Ma; map number 6, Table 1; Fig. 2) has yielded a SHRIMP U/Pb zircon age of 2630 ± 3 Ma (sample 9, Table 1). Coarse-grained porphyritic granodiorite to monzogranite (unit Agp) is interpreted to be a subvolcanic intrusion. Unit Agp is spatially associated with the Hoarding Wolf stock (Fig. 2) and is interpreted to be a subvolcanic intrusion. Unit Agp is spatially associated with the Hoarding Wolf stock (Fig. 2) and is interbedded with pyroclastic rocks (unit Apy). Unit Agp contains 1 cm to 5 cm, pink - white K-feldspar megacrysts (up to 10 cm) and 5% to 10% biotite ± hornblende. Unit Agp is commonly foliated and is interpreted to be a subvolcanic intrusion. Unit Agp is spatially associated with the Hoarding Wolf stock (Fig. 2) and is interbedded with pyroclastic rocks (unit Apy). Unit Agp contains 1 cm to 5 cm, pink - white K-feldspar megacrysts (up to 10 cm) and 5% to 10% biotite ± hornblende. Unit Agp is commonly foliated and is interpreted to be a subvolcanic intrusion.

REFERENCES

Apex Geoscience Ltd.
1994: Committee Bay Joint Venture Exploration 1994, Laughland Lake – Curtis River Area, Prospecting permits 1295 to 1324, 1477, 1481, 1486, 1490, 1495, and 1498, District of Keewatin, N.W.T. Assessment Report submitted to Indian and Northern Affairs, Canada, No. 063436.

Ashton, K.A.,
1998: Precambrian geology of the southeastern Amer Lake area (66H/1), near Baker Lake, N.W.T, Ph.D. thesis, Queen's university, Kingston, Ontario, 551p.

Committee Bay Joint Venture
1993: Committee Bay Joint Venture, Gold Exploration 1993, Prospecting permits 1330 to 1334, District of Keewatin, N.W.T. Assessment Report submitted to Indian and Northern Affairs, Canada, No. 062501.

Cyprus Canada Inc.
1995: Committee Bay Project, 1995 Exploration Report on the Laughland Lake – Curtis River Area, District of Keewatin, N.W.T. NTS 56J, K, O, P. Assessment Report submitted to Indian and Northern Affairs, Canada, No. 063045.

Davis, H.J. and Zaleski, E.
1998: Geochronological investigations of the Archean Woodburn Lake Group, eastern Churchill Province, Northwest Territories; in Radiogenic Age and Isotopic Studies: Report 11, Geological Survey of Canada, Current Research 1998-F, 13p.

Dayell, C. and Sherlock, R.L.
2003: Iron Formation-hosted Gold Occurrences in the Ellice Hills area (NTS 56P), Committee Bay Belt, Nunavut, Geological Survey of Canada, Current Research, 2003-C10.

Frisch, T.
1982: Precambrian geology of the Prince Albert Hills, western Melville Peninsula, Northwest Territories, Geological Survey of Canada, Memoir 395, 57p.
2000: Precambrian geology of Boothia Peninsula and Lord Mayor Bay, Cape Barrow, part of Darby Lake map area, southcentral Nunavut, Geological Survey of Canada, Bulletin 540, 51 p.

Glengoyne, M., Little, E.C., Peterson, T.D., Hynes, A. and Scott, J.M.J.
2003: Quaternary glaciomarine environments west of Committee Bay, central mainland Nunavut; Geological Survey of Canada, Current Research 2003-C23.

Hanchar, S.R., Hanson, L.M., Ashton, K.E., Davies, T.
2001: The Murmac Bay group, Rae Province: record of a giant Archean rift? in 4th International Archean Symposium 2001, Extended Abstracts, A.P. Cassidy, J.M. Dunphy, and M.J. Van Kranendonk (eds.), Australian Geological Survey Organization–Geoscience Australia, Record 2001/37, p. 371-374.

Heuzenn, M. and Sumeza, S.
1986: Report of the investigations of rock geochemistry of permit 1090 and neighboring areas, Keewatin District, Northwest Territories, Assessment Report submitted to Indian and Northern Affairs, Canada, No. 063134.

Henderson, J.R.
1985: Structure and metamorphism of the Aphebian Penrhyn Group and the Archean basement complex in the Lyon Inlet area, Melville Peninsula, District of Keewatin, Geological Survey of Canada, Bulletin 354, 30p.

Heywood, W.W.
1981: Geological notes, northern District of Keewatin, Geological Survey of Canada, Paper 81-1B, 9p.

Hyde, D. J., Jenner, G. A., and Sherlock, R. L.
2002: Aphyric boninites from the Committee Bay River area, central mainland Nunavut, Geological Survey of Canada, Current Research, 2002-E7.

Johnstone, S.A.
2002: Structural and geochronological constraints on the tectonic evolution of the Walker Lake area, Committee Bay belt, Nunavut, unpublished M.Sc.thesis, University of Waterloo, Waterloo, ON., 188 pages.

Johnstone, S. A., Little, E.C., and Peterson, T.D.
2002: Significance of the Walker Lake shear zone with respect to regional deformation in the Committee Bay belt, central mainland, Nunavut; Geological Survey of Canada, Current Research, 2002-C15.

King Resources
2001: Report on Project Wagner, District of Keewatin, Northwest Territories, Assessment Report submitted to Indian and Northern Affairs, Canada, No. 063466.

Kiss E. Clyde M., and Dumont R.
2002a: Aeromagnetic Total Field Map, Nunavut, NTS 56 K/SE-I/SW, Geological Survey of Canada, Open File 4248, scale 1:50,000.
2002b: Aeromagnetic Total Field Map, Nunavut, NTS 56 K/NE-I/NW, Geological Survey of Canada, Open File 4249, scale 1:50,000.
2002c: Aeromagnetic Total Field Map, Nunavut, NTS 56 J/SE-I/SW, Geological Survey of Canada, Open File 4247, scale 1:50,000.
2002d: Aeromagnetic Total Field Map, Nunavut, NTS 56 J/NE-I/NW, Geological Survey of Canada, Open File 4246, scale 1:50,000.
2002e: Aeromagnetic Total Field Map, Nunavut, NTS 56 K/SE-I/SE, Geological Survey of Canada, Open File 4251, scale 1:50,000.
2002g: Aeromagnetic Total Field Map, Nunavut, NTS 56 P/NE-I/NE, Geological Survey of Canada, Open File 4253, scale 1:50,000.

LaChemiant, A.N. and Sauvé, J.C.
1991: LUPIN zircon evidence for widespread 2.6 Ga felsic magmatism in the central District of Keewatin, N.W.T.; in Radiogenic Age and Isotopic Studies: Report 4, Geological Survey of Canada, Paper 91-1A, p. 773-782.

Little, E.
2001: Preliminary results of relative ore-movement chronology of the Laughland Lake map area, Nunavut; Geological Survey of Canada, Current Research, 2001-C14.

Little, E.C., Peterson, T.D., Dyer, D.R., and Utting, D.J.
2003: Bedrock mapping in the Committee Bay Project, central Nunavut; Geological Survey of Canada, Current Research, 2003-C1.

MacHattie, T.
2002: Physical volcanology of komatiite in the Laughland and Walker Lake areas, Committee Bay belt, Nunavut, Geological Survey of Canada, Current Research, 2002-A14.

McMartin, I., Little, E.C., Peterson, T., Dyer, D.R., Utting, D.J., and Knight, R.D.
2002: Surficial geology, Committee Bay, central mainland, Nunavut; Geological Survey of Canada, Open File 4377, North and Ellice Islands, Geological Survey of Canada Open File 4377.

McMartin, I., Little, E.C., Peterley, T., Dyer, C.A., and Utting, D.J.
2003: Ice flow history and drift prospecting in the Committee Bay belt, central Nunavut: results from the Targeted Geoscience Initiative, Geological Survey of Canada, Current Research 2003-C4.

Oxyer, C. A. and Hirsch, S. I.
2002: Glacial landforms and preliminary chronology of ice-movement in the Arrowsmith River map area, Nunavut; Geological Survey of Canada, Current Research, 2002-C16.

Sanborn-Barrie, M., Skulski, T., Sanderson, H., Berman, R., Johnstone, S., MacHattie, T., and Hyde, D.
2002: Structural and metamorphic geology of the Walker Lake area, Committee Bay belt, Nunavut, Geological Survey of Canada, Current Research, no. 2002-C17.

Sanborn-Barrie, M., Sanderson, H., Skulski, T., Brown, J., Young, M., MacHattie, T., Dayell, C., Carson, C., Panagapko, D., and Hyde, D.
2003: Structural geology of the northeastern Committee Bay belt, Ellice Hills area, central Nunavut, Geological Survey of Canada, Current Research 2003-C23.

Sanderson, H.A., Brown, J.L., Gwinner, E., Hyde, D., Johnstone, S., MacHattie, T.G., Skulski-Gilbert, C., and Pleass, D.
2001a: Drilling, Laughland Lake, Nunavut (NTS 56K), Geological Survey of Canada, Open File 4190, scale 1:100 000, CD-ROM.

Sanderson, H.A., Brown, J., Skulski-Gilbert, C., MacHattie, T., Hyde, D., Johnstone, S., Gwinner, E., and Pleass, D.
2001b: Bedrock mapping in the Committee Bay Belt, Laughland Lake area (NTS 56K) central mainland, Nunavut, Geological Survey of Canada, Current Research 2001-C15.

Sanderson, H.A., Skulski-Gilbert, C., Brown, J., and Johnstone, S.A.
2001c: Regional structural and metamorphic geology of the Committee Bay Belt, Laughland Lake area, central mainland Nunavut, in Current Research, Geological Survey of Canada, Open File 4077.

Sherlock, R.L., and Dayell, C.
2002: Iron formation-hosted assessment report compilation for the Committee Bay region, central mainland, Nunavut (NTS 56-N,O,P), Geological Survey of Canada Open File 4329, 2 CD-ROMs.
2003: Geochemical and assessment report compilation for the Committee Bay region, central mainland, Nunavut, Geological Survey of Canada, Open File 1545, 2 CD-ROMs.

Skulski, T., Sanderson, H., Sanborn-Barrie, M., MacHattie, T., Hyde, D., Johnstone, S., Panagapko, D., and Dayell, C.
2003: Compilation, crustal evolution and mineral potential of the Committee Bay belt, Nunavut; in Current Research, Geological Survey of Canada, Open File 1545, CD-ROM.

Skulski, T., Sanderson, H., Sanborn-Barrie, M., Panagapko, D., Carson, C., Brown, J., Panagapko, D., Byrne, D., and Dayell, C.
2003: Geology, Committee Bay belt, Nunavut (NTS 56P), and new constraints on the regional geology of the Committee Bay area; in Current Research, Geological Survey of Canada, Current Research 2003-C21.

Utting, D.J. and Wood, B.C.
2003: Surficial interpretation of Quaternary glaciofluvial landforms identified during mapping of Walker Lake (NTS 56J,K, 19), Nunavut, Geological Survey of Canada, Paper 2002-C16.

Tella, S.
1994: Geology, Amer Lake (56H), Deep Rose Lake (56G), and parts of Nelly Lake (56F), District of Keewatin, Northwest Territories, Geological Survey of Canada, Open File 2945, scale 1:250,000.

Zaleski, E., Davis, W.J. and Sanderson, H.A.
2001: Continental extension, mantle magmas and basement cover relationships, in 4th International Archean Symposium 2001, Extended Abstracts, A.P. Cassidy, J.M. Dunphy, and M.J. Van Kranendonk (eds.), Australian Geological Survey Organization–Geoscience Australia, Record 2001/37, p. 374-376.

STRUCTURAL GEOLOGY

The Walker Lake – Arrowsmith River area structural geology as presented in Sanborn-Barrie et al. (2002) is summarized below.

D1 STRUCTURES

In the supracrustal rocks, the planar tectonic fabrics that predate the regional northeast-striking foliation are interpreted as S1 and occur predominantly in komatiitic and plutonic rocks. In komatiite, the northeastward striking, bedding-parallel, spinifex-textured flow tops are commonly transposed into S1. In plutonic rocks, S1 is defined by a moderately southeast, N-plunging mineral lineation. S1 is recognized in the orientation of >1-mm-long axes of FY-folds where it is deformed by aligned biotite and quartz at a high angle to S0. In the Northern Migmatite subdomain, S1 is represented by a migmatitic layering that is parallel to the regional northeast-striking foliation and is axial planar to isoclinal, same-facing (to the south) F1 folds which occur at the western end of the NHFB and trend north to northeast.

D2 STRUCTURES

The dominant fabric throughout the Committee Bay region is a moderate to strong, northeast-striking foliation. This fabric is axial planar to tight to isoclinal, F2 folds that are commonly rootless. In the Prince Albert group subdomain, composite S1 ≥ S2 fabrics, which are commonly shallow, but moderately, to steeply developed subparallel to bedding. In the Northern Migmatite subdomain, S1 and S2 are parallel and compositional layering is commonly a single foliation of higher strain.

S2 is the dominant, regional, penetrative, ductile fabric throughout the map area. It defines variations in both dip direction and orientation within the supracrustal belts. S2 plunges moderately to steeply (0 to 60°) to the southeast, intersecting, northeast-striking, bedding-parallel, S1 fabrics at a high angle to produce a mineral lineation. The mineral lineation is southeast to steeply northeast dipping, whereas in the northern part of the Prince Albert group subdomain, the S2 ≥ S1 composite fabric is southeast dipping (Sanborn-Barrie et al., 2002; Sanborn-Barrie et al., 2003; Fig. 2). S2 is a locally-confident, strain-gradient fabric that is defined by a moderately to strongly developed, northeast-striking, subvertical foliation. The shallower dip of the S2 ≥ S1 foliation in the northern part of the Prince Albert group subdomain results in the development of a composite, northeast-striking, S2 ≥ S1 foliation.

In the Northern Migmatite subdomain, S2 is locally identified as a weak planar to folds that modify the orientation of S1, and is typically a northeast- to east-striking schistosity that is defined by biotite, sillimanite and garnet, and commonly contains anchor axes. In contrast to S1, S2 is typically steeply dipping (>70°) with northeast dips dominating.

Numerous regional northeast-trending F2 folds within the plutonic rocks of the Prince Albert group subdomain are defined by the orientation of the S1 foliation. F2 folds are tight to isoclinal, and plunge moderately (30 to 60°) to the NE. They are upright to overturned (northwest vergent), noncylindrical, antiforms and synforms that tend to die out over a strike length of 4 to 12 km. On the basis of closure configuration, regional F2 folds appear to plunge westward in the eastern part of the belt (Fig. 2). F2 folds are commonly rootless in the Northern Migmatite subdomain. Axial planes to F2 folds are typically subvertical, northeast-striking, and commonly dip to the northwest, whereas F2 folds are commonly overturned to the southeast.

A reversal in younging has been determined along the length of the CBs, from southward younging at Kaltet River (Sanborn-Barrie et al., 2002) to northward younging at Ellice Hills (Sanborn-Barrie et al., 2003). The CBs is interpreted to represent a south-facing, homoclinal sequence that, through oroclinal asymmetric Z-folding, has reversed its segments (i.e., Ten Peaks quartzite ridge) of north-younging strata. Pervasive development of structures with dextral shear sense in the Walker Lake area (Sanborn-Barrie et al., 2002; Sanborn-Barrie et al., 2003; Fig. 2) is interpreted to implicate inversion of significant southeast-directed transpressional strain during, or subsequent to, D2 folding. Northward younging at one locality of the TB and NHHR may be a manifestation of regional-scale fold lineages between the belts.

ARROWSMITH RIVER SHEAR ZONE

An arcuate corridor of annealed porphyroclastic mylonitic rocks, designated the Arrowsmith River shear zone, extends for >100 km from the Ellice River to the Arrowsmith River (Fig. 2). The Arrowsmith River shear zone is an arcuate strands of very strongly foliated to mylonitic (plutonic rocks, 4 to 2 km wide east-striking high-strain corridor extends for ≥ 40 km across the northern part of the map area. Within this corridor, shallow (< 10°) south-dipping mylonitic fabrics are axial planar to tight, F2 folds that plunge moderately to the southeast (Sanborn-Barrie et al., 2002). For the shear plane, to the west, lineations generally plunge shallowly southeast (< 30°), whereas east of the Arrowsmith River they generally plunge eastward. Asymmetrical structures (mainly sigma-type wrapped porphyroclasts) indicate a dextral sense of shear. The Arrowsmith River shear zone is interpreted to be the result of dextral, oblique-slip, ductile reactivation of the east-plunging mineral lineations. The general displacement vector for the northern block may therefore be one of oblique-normal displacement (hanging-wall down-to-the-south).

The Arrowsmith River shear zone is characterized by high-strain intensity, fabric development and proximity to that of the northern shear zone. This shear zone is moderately shallow south-dipping (10 to 30°) structure with strong ductility. Kinematics indicate that plunge to the east to southeast (20 to 35°). Shear sense indicators (winged porphyroclasts and asymmetrical boudins) are consistent with dextral shear sense and dextral-oblique displacement. Kinematics are the result of reactivation of the east-plunging mineral lineations in the transition and perpendicular to foliation and strike-up-down to the southeast, southwest.

D3 STRUCTURES AND THE WALKER LAKE SHEAR ZONE

The moderately to strongly developed regional northeast-striking, S2 composite foliation typical of the Prince Albert group subdomain progressively changes its character and intensity in the southwest portion of the map area where it is overprinted by the Walker Lake shear zone. The Walker Lake shear zone is a 1 to 2 km wide, moderately to steeply striking and moderately northwest dipping (60 to 80°) to west striking and steeply north dipping zones. Mineral lineations progressively decrease in pitch with proximity to the west striking zone, which contains nearly horizontally plunging, mineral lineations. The Walker Lake shear zone is a zone of dextral, oblique-slip, ductile shear (Fig. 2; Sanborn-Barrie et al., 2002). Kinematic indicators, such as winged porphyroclasts and asymmetrical boudins, are consistent with dextral shear sense and dextral-oblique displacement. Kinematics are the result of reactivation of the east-plunging mineral lineations in the transition and D1 structures at the southwest band in the Arrowsmith River may be due to an open northeast-trending F3 folding.

METAMORPHISM

Metamorphic grades vary from lower amphibolite facies in the southeast to middle amphibolite facies in the eastern and northern parts of the Prince Albert group subdomain. Metamorphic assemblages at lower grade include staurolite – muscovite – garnet, cordierite – andalusite – muscovite and cordierite–garnet–biotite in metapelitic rocks, andalusite and cordierite in metapsammitic rocks, and hornblende – plagioclase in metabasite. Metamorphic grade appears to be quite symmetrical with respect to the dominant D2 ≥ S2 foliation, with the exception of some postkinematic garnet. Higher grade semipelitic rocks contain sillimanite + garnet + muscovite + biotite, with pink to sylvinebearing, garnet – cordierite ± anthophyllite gneiss in the Northern Migmatite subdomain. Metamorphic assemblages in post-kinematic plutonic rocks in the Prince Albert group subdomain yielded SHRIMP U/Pb ages of around 2580 Ma (map numbers 6, 7 and 9; Table 1), it is not clear how this metamorphic event is related to structural fabric development.

The Northern Migmatite subdomain contains domains of upper amphibolite facies rocks, with migmatitic paragenes exhibiting the assemblages garnet + sillimanite + biotite + quartz ± K-feldspar and garnet + sillimanite + cordierite + biotite + quartz ± K-feldspar in metapelitic rocks. Leucosomes in these rocks contain garnet, sillimanite, perthitic, anemogranite and pyrite are associated with monzonitic, garnet-bearing metasedimentary units containing 10 to 40% interbanded granitoid melt/dykes