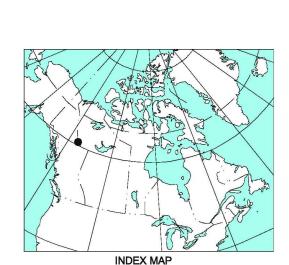






UNIVERSAL TRANSVERSE MERCATOR GRID, ZONE 10

444000m. E.



419000m. E.

**CONTOUR INTERVAL 100 FEET** 

Elevations in Feet above Mean Sea Level

**OPEN FILE 1615** 

SURFICIAL GEOLOGY

# **CHINKEH CREEK**

NORTHWEST TERRITORIES - YUKON TERRITORY

Scale 1:50 000/Échelle 1/50 000 2

Universal Transverse Mercator Projection North American Datum 1983 © Her Majesty the Queen in Right of Canada, 2003

Projection transverse universelle de Mercator Système de référence géodésique nord-américain, 1983

© Sa Majesté la Reine du chef du Canada, 2003

3 Kilomètres

#### Compilation by I.R. Smith based on fieldwork and studies of vertical air photographs 2000-2002. THIS MAP IS A PRODUCT OF THE CENTRAL FORELAND NATMAP PROJECT

Surficial geology from field work by I.R. Smith 2000-2002.

Digital cartography by I.R. Smith.

Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada.

Base map at the same scale published by Surveys and Mapping Branch in 1971.

95C/15	95C/16	95B/13	
Dendale Lake	Etanda Lakes	Sawmill Mountain	
95C/10	95C/9	95B/12	
Tika Creek	Chinkeh Creek	Mount Flett	
	GSC OF 1615		
95C/7	95C/8	95B/5	
Brown Lake	Babiche Mountain	Fisherman Lake	
	GSC OF 1558	GSC OF 4360	
NATIONAL TOPOGE	NATIONAL TOPOGRAPHIC SYSTEM REFERENCE AND INDEX		

TO ADJOINING GEOLOGICAL SURVEY OF CANADA MAPS

**LEGEND** 

Coloured legend blocks indicate units that appear on this map

## **QUATERNARY**

**SURFICIAL DEPOSITS** 

POST LAST GLACIATION

50 m thick; forming aprons and fans below cliffs

## NONGLACIAL ENVIRONMENTS

ORGANIC DEPOSITS: organic matter; >1 m thick; formed by the accumulation of vegetation in poorly drained depressions (swamps and bogs); usually forms flat terrain

COLLUVIAL DEPOSITS: block accumulations and mass wasting debris, 1-50 m thick

Talus (scree): accumulations of blocks; commonly exceeding 2 m in diameter; as much as

Rock Glaciers (relict): rock debris deformed by the down-slope flow of buried or interstitial ice, forming pronounced transverse and longitudinal ridges and furrows

Debris slump deposits: unconsolidated material; generally smaller blocks or more localized masses, but may include larger masses (>10 m thick) where associated with thick till, glaciolacustrine or glaciofluvial deposits; internal structure of material may be retained; often traceable upslope to active scarps; where sufficient moisture is present the slump may become a flow, producing characteristic levees along its lateral margins and a spatulate form at the base of slope.

Bedrock slump deposits: large rotational blocks in bedrock, shallow to 10's of metres thick; internal structure of material may be retained; often traceable upslope to active scarps; where sufficient moisture is present the slump may produce a flow at its base, forming a characteristic spatulate form; prominent in areas of moderate to shallow-dipping Cretaceous shale, siltstone and sandstone beds of the Fort St. John Group; associated with the largest mass movements in the region

Rock slide deposits: chaotic landscape of irregular and stacked bedrock blocks; prominent in areas of moderately dipping, poorly-indurated sandstone and shale-rich beds in the Mattson and Fantasque formations

ALLUVIAL DEPOSITS: gravel, sand, and organic detritus >1 m thick

Fluvial deposits: well sorted gravel and sand with detrital organic beds, including concentrations of logs, >1 m thick; Ap, floodplains and mantling valley floors, forming meander scars and point bars; At, terraces along valley wall sides

Alluvial fan: poorly sorted gravel and sand with organic detritus and buried soils; fans are commonly crossed by debris flow channels and levees and subject to shifting stream courses; >1 m thick

#### POSTGLACIAL OR LATE WISCONSINAN

## PROGLACIAL AND GLACIAL ENVIRONMENTS

GLACIOLACUSTRINE DEPOSITS: coarse to fine sand, silt and clay, with gravel debris flow layers and dropstones; deposited in glacier-dammed lakes; level topography; Lp, thin discontinous veneers, <1 m thick; Lt, forming terraces, often deeply dissected by postglacial

GLACIOFLUVIAL DEPOSITS: gravel, sand, minor sandy diamict, usually >1 m thick; deposited on, beneath, or in front of glacier margins

Proglacial outwash: Gd, braided outwash deltas; Gdt, delta terraces; Gf, fans; Gp, outwash plains and mantling valley floors; Gt, level outwash terraces

Ice contact stratified drift: deposited behind or at the ice magin; topography is undulating, irregular, or ridged; It, lateral kame terraces; Idt, delta terraces; Ik, kettle holes; Ih, hummocky moulin kame fields, or ice block disintegration terrain; Ir, eskers or crevasse fillings TILL: nonsorted diamict deposited directly by glacial ice; matrix is sandy to clayey and contains striated clasts of various lithologies

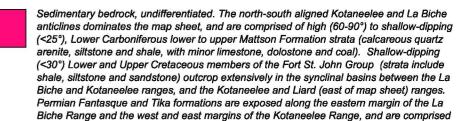
Till blanket: > 2 m thick; forming undulating topography that obscures underlying bedrock

Till veneer: < 2 m thick and discontinuous; surface mimics underlying bedrock structure

structure; Tbk, distinctly kettled

## PRE-QUATERNARY

### **BEDROCK**



Geological boundary (defined, gradational)

arenite, siltstone and shale, with minor limestone, dolostone and coal). Shallow-dipping (<30°) Lower and Upper Cretaceous members of the Fort St. John Group (strata include shale, siltstone and sandstone) outcrop extensively in the synclinal basins between the La Biche and Kotaneelee ranges, and the Kotaneelee and Liard (east of map sheet) ranges. Permian Fantasque and Tika formations are exposed along the eastern margin of the La Biche Range and the west and east margins of the Kotaneelee Range, and are comprised of diverse strata that includes chert, siltstone, limestone, dolostone, and sandstone. Devonian and Carboniferous Besa River Formation (mostly shale with some sandstone) is exposed along the south-central Kotaneelee Range and the southern La Biche Range (see Lane and Fallas, 2003).

NOTE: In areas where the surficial cover forms a complex mosaic, the area is coloured according to the predominant unit and labelled with hyphenated letters in descending order of cover

## MAP SYMBOLS

Cirque; peaks and sharp ridges formed by glacial erosion Striae (glacial flow direction known, unknown) Proglacial meltwater channel; abandoned or occupied by small underfit stream (wide, narrow with direction of flow inferred) Lateral meltwater channel (barb points upslope and down flow)

Drift geochemistry sample site

Canadian Shield erratic

Mass Wasting is the collective term given to the range of processes and resultant landforms that relate to the gravitational downslope movement of rock and/or unconsolidated material without the direct conveyance by water, air or ice. Water and ice are, however, often key components in initiating and perpetuating mass wasting by reducing the strength of materials and in their plastic and fluid behaviour. Different types of mass wasting are distinguished by the type of materials involved (e.g., bedrock, talus, till), the mode of deformation (e.g., creep, slide, slump, flow), speed of movement, morphology of the moving mass, and water content.

Creep is the slow (mm's to cm's per year), often imperceptible, downslope movement of soil, talus or other unconsolidated material. Creep occurs episodically in response to solutional weathering, seasonal wetting and drying, or freeze-thaw cycles and may include the plastic deformation of clay-rich soils. While more prevalent on steep slopes, creep can occur on slopes <5°. Evidence of creep is seen where tree trunks or structures (e.g., hydro poles) are tilted downslope, soil accumulates upsolope of retaining walls, and cracks develop in the soil perpendicular to the slope. Creep is also responsible for the formation of gelifluction lobes, prominent, small-scale (metres in length, centimetres thick), periglacial landforms found along the upper reaches of local mountain ranges (but not included in the regional surficial geology

Slides are the rapid, downslope movement of bedrock or unconsolidated material. Failure occurs along bedding and/or fracture planes in bedrock, and along bedrock contacts, or structural and sedimentological boundaries within unconsolidated material. Slides can be initiated at shallow or considerable depths. Slumps involve the rotational movement of bedrock and/or unconsolidated material along failure planes. Slumps may occur as individual blocks or amorphous masses (reflecting water content and structural integrity of the failing material). Slumps often extend progressively up-slope through time, and can be associated with active scarp or headwall retreat. Slumps can be initiated by failure along bedding, fracture, or sedimentological planes, by infiltration of surface water, through lateral incision and undercutting of slopes by streams, or excavation activities (e.g., road building, pipeline trenching). Slumps are prominent in areas of moderately dipping, poorly-indurated sandstone and shale-rich beds in the Mattson and Fantasque formations, and in moderate to shallow-dipping shale, siltstone and sandstone beds of the Cretaceous Fort St. John Group. Slumps are associated with the largest mass movements in

While different earth surface materials and geological settings are often strongly associated with various types of mass wasting, predicting their occurrence, magnitude and rate of deformation is often not possible. Some areas that are prone to mass wasting include regions of steeply dipping bedrock, poorly indurated and shale-rich bedrock, and along stream courses and meandering river channels. Human activities such as road building, pipeline trenching, logging and seismic exploration can also initiate mass wasting, particularly where they undercut slopes, or act to destabilize surficial materials.

Glacial History: The Chinkeh Creek map area was glaciated during the last (late Wisconsinan) glaciation (ca. 25-10 000 years ago) by the continental Laurentide Ice Sheet flowing from the northeast (Keewatin Sector) and by the Cordilleran Ice Sheet flowing from the west. The Laurentide Ice Sheet dispersed distinctive granite erratics, originating from the Canadian Shield. These granite erratics were found atop Kotaneelee Range (1390 m above sea level (asl)), and throughout the Chinkeh and Kotaneelee river valleys. Immediately south of this map sheet, a granite erratic was also found atop the La Biche Range (1620 m asl). Sandstone erratics of unknown provenance were found at the crest of the Kotaneelee Range (1850 m asl), establishing a minimum upper limit of glaciation for the region. Cross-cutting ice flow directional indicators (striae and flutings) indicate that glaciers first moved westward across the region, followed by an eastward flow. It can thus be concluded that this map area was first occupied by Laurentide ice which inundated the entire landscape. Subsequently, Cordilleran ice advanced eastward, displacing the Laurentide ice. Cordilleran ice did not extend to the highest summits of the Kotaneelee and northern

La Biche ranges. Cirque basins are prominent along the northern Kotaneelee and southern La Biche ranges. Small, arcuate moraines within the cirques and immediately down-valley indicate a period of alpine/cirque glaciation subsequent to retreat of the Cordilleran and Laurentide ice sheets. The fact that the moraines are well vegetated, have a subdued morphology, and that no cirque glaciers were found anywhere in the 95C map sheet, suggests that these moraines were not formed during the Little Ice Age (1400 - 1900 AD), but instead relate to post-late Wisconsinan (possibly mid-Holocene) glacial activity. Deglacial landforms and deposits are prominent in the Chinkeh Creek map area. Many of these are associated with the impoundment of drainage between westward retreating Cordilleran ice and eastward

retreating Laurentide ice, including a glacial lake in the upper Kotaneelee River valley where well-sorted

# deposits of sand and silt >90 m thick were found.

Open File map 1674, scale 1:50 000.

Lane, L.S., and Fallas, K.M. (compilers) 2003: Geology, Chinkeh Creek (95C/9), Yukon and Northwest Territories; Geological Survey of Canada,

References

OPEN FILE DOSSIER PUBLIC	Open files are products that have not gone through the GSC formal publication process.	
1615		
GEOLOGICAL SURVEY OF CANADA COMMISSION GÉOLOGIQUE DU CANADA	Les dossiers publics sont des produits qui n'ont pas été soumis au processus officiel de	
2003	publication de la CGC.	