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GEOLOGICAL SURVEY OF CANADA OPEN FILE 4085

# Focal mechanisms for eastern canadian earthquakes of 1999 

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#### Abstract

Although earthquake focal mechanism solutions provide information about the faults on which the earthquakes occur and on local stress field orientations and seismotectonics, they were not routinely determined in the past. We have begun to systematically attempt to determine focal mechanisms for all eastern Canadian earthquakes of magnitude 4.0 or greater and intend to publish the results annually in an Open File. This report covers the year 1999. We analyzed nine earthquakes. Two did not strictly meet our selection criteria but are included because they are of interest. We were able to determine focal mechanisms for six of the events. As has been the case in previous efforts, those events for which we could not obtain solutions were in offshore or northern regions where seismograph coverage is sparse.


#### Abstract

RÉSUMÉ Les mécanismes au foyer des tremblements de terre nous fournissent de l'information sur les failles sur lesquelles les tremblements de terre se produisent, et sur l'orientation des contraintes régionales et séismotectoniques mais auparavant on ne les déterminait pas régulièrement dans l'est du Canada. Nous avons commencé à essayer systématiquement de déterminer les mécanismes au foyer pour tous les tremblements de terre de magnitude 4,0 ou plus dans l'est du Canada. On a l'intention de publier les résultats tous les ans. Ce rapport-ci traite l'année 1999. On a analysé neuf tremblements de terre dont deux n'ont pas satisfait parfaitement nos critères mais ont attiré notre attention. On a obtenu des solutions pour six de ces tremblement de terre. Comme c'était le cas dans nos études précédents, les séismes pour lesquels on n'a pas pu obtenir des solutions se sont produits dans le nord ou des régions côtières où la couverture du réseau séismographique n'est pas suffisante.


## INTRODUCTION

Earthquake focal mechanism solutions provide information about the faults on which the earthquakes occur and on local stress field orientations. If focal mechanisms can be determined for many earthquakes in a given region then we also gain insight into regional seismotectonic precesses and stress fields. In eastern Canada earthquake focal mechanisms were not routinely determined in the past, although solutions may exist for individual earthquakes that were of particular interest as well as for some time periods or regions that have been systematically studied. We have begun to routinely analyze first motion data for all eastern Canadian (roughly defined as east of $110^{\circ} \mathrm{W}$ ) earthquakes of magnitude 4.0 or greater in an effort to determine their focal mechanisms. The data will be tabulated and mechanisms summarized annually in Geological Survey of Canada (GSC) Open Files. The years 1994 to 1998 have been completed (Bent and Perry, 1999, 2000; Bent and Drysdale, 2000). The current report covers the year 1999. Seven events met our selection criteria. We also examined two other earthquakes that were of interest although they did not strictly satisfy the criteria: one that was slightly smaller (magnitude 3.8) than our minimum magnitude but that was felt in the Toronto area and one that was just outside our area of interest (magnitude 5.7 in the western Arctic islands). Of the nine events studied we were able to determine focal mechanisms for six.

## Figure Captions:

Figure 1. Map of southeastern Canadian events studied. Symbol size is scaled to magnitude. Magnitude range 3.8-5.1.

Figure 2. Map of northeastern Canadian events studied. Symbol size is scaled to magnitude. Magnitude range 4.0-5.7.


Figure 1


## METHOD

The focal mechanisms were determined primarily from $P$ wave first motions. We usually only attempted to read $S$ polarities if the $P$ data did not provide a well-constrained solution. We looked at the vertical component of all digital seismograms of the Canadian National Seismogram Network or CNSN received in Ottawa either by continuous transmission or dial-up request. Rotated horizontal data were used as much as possible for $S$ polarities. Note that data from some western Canadian stations, particularly from the dense array on Vancouver Island, were not routinely archived in Ottawa. However, data from enough stations in British Columbia are routinely received that we can determine whether data from the additional stations are likely to be useful, in which case the waveform data can be requested.

We also include any first motion data contained in the Canadian Earthquake Epicenter File (CEEF) database for which the seismograms were not archived. Generally these polarities were supplied by operators of regional, usually U. S., networks or single stations within a few hundred kilometers of the epicenter. In the discussions of individual earthquakes, we note any polarities not determined directly by the authors.

We attempted to obtain non-CNSN seismograms or polarity data if they were likely to better constrain the solution, but did not routinely request external data for southern Canadian events except for earthquakes in southern Ontario. For the southern Ontario events, we tried to obtain data from the Southern Ontario Network (SONT) operated by the University of Western Ontario. Data from these stations improved the azimuthal coverage and provided more data at short distances.

For events from Baffin Bay to Ellesmere Island, we routinely requested data from two stations operated by the Incorporated Research Institutions for Seismology (IRIS) : Alert (ALE) on Ellesmere Island and Sondre Stromfjord (SFJ) on Greenland. Data from SFJ were also routinely requested for earthquakes in the Labrador Sea. For events in the western Arctic or the Boothia-Ungava seismic zone, we requested data from the IRIS station at Flin Flon, Manitoba (FFC). For earthquakes occurring offshore along the southeastern margin, we consulted analog seismograms from stations in Halifax (HAL) and Guysborough, Nova Scotia (GBN) and St. John's, Newfoundland (STJN), all three of which are archived in Ottawa.

Most of the earthquakes we studied were relatively small (magnitude less than 5.0). Therefore we did not routinely attempt to acquire teleseismic data although it was obtained for some events. For readers not familiar with station locations or codes, a map of all seismograph stations operating in Canada, including those that are not part of the CNSN, may be found on the internet. The current address is
http://www.seismo.nrcan.gc.ca/cnsn/stn_map.e.html.
As much as possible we read the polarities from unfiltered data. However, given that the
earthquakes studied were generally small, it was occasionally necessary to filter some seismograms, particularly those from more distant broadband stations. Because high frequencies are attenuated more than low frequencies, there is always some danger that a first motion read from filtered data will not be the true first motion. By restricting the filtering of the broadband data to the frequency range of the short-period instruments, we should not adversely affect the data any more than a short-period instrument response would do. Unless stated otherwise, all filters were two-pole, single-pass Butterworth filters. Any filtering is noted on the seismogram traces where "hp" is used to designate a highpass filter, "lp" a lowpass filter and "bp" a bandpass filter. Corner frequencies are also noted. All filtering of data was performed using the Seismic Analysis Code (SAC) routines (Tapley and Tull, 1991), which were also used to rotate horizontal records into their radial and tangential components.

There was also some concern, particularly for broadband seismograms, that the first motions could be affected by the FIR filters applied to the data. In any case where it was suspected that what appeared to be the first motion was actually a FIR filter artifact, the data were run through a program to deconvolve the FIR filter (Scherbaum and Bouin, 1997). A small apparent first motion followed by a significantly larger arrival of opposite polarity was the principal reason for de-FIR-ring the data. In the end, we found that these small first motions did not appear to be filter related.

As far as we have been able to ascertain polarities for CNSN instruments were correct during 1999. Instrument polarities are checked approximately once a year using large teleseismic events with known focal mechanisms. Polarities for stations of the Southern Ontario Network were confirmed by staff of the University of Western Ontario, who operate the network.

The first motions were read independently by the two authors. In case of disagreement, the seismograms in question were re-evaluated. If we could not reach a consensus, data from that particular station were not used to determine the focal mechanism.

To determine the focal mechanism, the grid search algorithm of Snoke et al. (1984) was employed. The search is based on the orientation of the B axis. Normal procedure was to initially search the focal sphere at $5^{\circ}$ intervals. If no solutions were found to satisfy all the data, the search increment was gradually decreased. If with a $1^{\circ}$ search no solutions were found, we returned to the $5^{\circ}$ search, but allowed 1 polarity error. The procedure of alternately decreasing the search interval and increasing the number of allowed polarity misfits was repeated until at least one solution was found. If the initial $5^{\circ}$ no error search resulted in a large number of solutions or if there were large differences among the solutions, we attempted to find more polarity data by reading $S$ wave polarities and/or looking for data from additional seismograph stations. For many of the offshore and northern events, the solutions could not always be further constrained.

The inversion procedure normally ignores emergent polarities. If the inversion resulted in a large number of potential solutions or a small number of very different solutions, we reran the program using a weighting scheme by which emergent stations were treated
as impulsive and each impulsive arrival was entered twice, effectively half-weighting the emergent phases. Focal mechanisms resulting from the weighted inversion were assumed to be more probable solutions than those which satisfied the impulsive data only.

Earthquake epicenters were taken from the CEEF. With a few exceptions, depths for eastern Canadian earthquakes are fixed, usually at 18 km (roughly the midpoint of the crust). Only one of the events selected (991126 in southern Ontario) had a calculated depth in the database. Bent and Perry (2000) determined the depths of two additional earthquakes as part of a study of earthquake depths. Since take-off angle is depth dependent, using an assumed depth may influence on the focal mechanism to some degree. Having said that, in most cases the effect will be minimal. As long as the hypocenter is within the crust, the take-off angle for Pn and teleseismic P will not be greatly influenced by the depth. At closer distances, the depth will have a greater effect on take-off angle and can also affect the Pg-Pn cross-over distance. However, the close stations are usually few in number, except in the Charlevoix region, where the depths are usually reliably calculated and not fixed. In any case, it is possible to determine how dependent the solution is on these stations by running the inversion using take-off angles appropriate for different depths.

## FOCAL MECHANISMS

The epicenters for events included in this study are shown in Figures 1 (southeastern Canada) and 2 (northeastern Canada) where they are identified by date of occurrence. In the tables that follow for each earthquake, we include a list of polarities used to determine the fault plane solution(s). C's and D's indicate compressional and dilatational $P$ polarities respectively. If the P polarity is emergent, $\mathrm{a}+$ indicates compression and a - indicates dilatation. SV waves are indicated by an F for motion toward the station or a B for motion toward the source. For SH waves, a < indicates counterclockwise motion (or to the left if one is facing the station with one's back to the source) and > indicates clockwise motion. Emergent $S$ waves were not used as their polarities are difficult to pick with confidence. We aiso include portions of all seismograms from which we picked the first motions as well as focal mechanism solutions plotted in terms of both stress axes and P nodal planes. Uniess it is explicitly stated that data from a particular station could not be obtained (based on the station selection criteria discussed in the previous section), it should be assumed that the authors evaluated the data but could not determine the first motion.

In the seismogram figures that follow, the velocity records are shown. In an effort to be consistent, we have plotted a constant time window ( 4 sec ) for each seismogram shown. This window is normally adequate to clearly illustrate the preceding noise and first motion, but we note that there may be occasions where showing a longer time segment or the displacement record might have provided stronger evidence for our choice of first motion.

All azimuths and strikes are measured in degrees clockwise from north. Take-off or incidence angles are measured from the vertical. Dip angles are measured from the horizontal. Rake is also known as slip angle. A rake of $0^{\circ}$ implies pure left lateral strikeslip motion, $90^{\circ}$ pure thrust motion, $180^{\circ}$ pure right lateral strike-slip motion, and $-90^{\circ}$ (or $270^{\circ}$ ) pure normal faulting. Seismic source zones are as defined by Basham et al. (1982)

Quality rankings of A, B, C or D are assigned to each preferred focal mechanism. The quality should be taken not as an indication of how well the solution fits the data but rather as a comment on the quality and quantity of the data used to determine the solution. Azimuthal distribution, redundancy of coverage, number of misfits and total number of observations are the principal criteria considered. The availability of $S$ polarity data can help compensate for poor azimuthal coverage as the $S$ and $P$ radiation patterns are different.

The Northwest Territories were partitioned to become two territories - Nunavut and the Northwest Territories - on 1 April 1999. For simplification, we use the new names for the entire calendar year.

## 1 JANUARY 1999: ARCTIC OCEAN

Seismic Zone: outside zones defined in 1982
Latitude: $\quad 79.79^{\circ} \mathrm{N}$
Longitude: $\quad 110.24^{\circ} \mathrm{W}$
Depth: 18 km (fixed)
Magnitude:
Origin Time:
Comments:

## Polarity Data

| Station | Distance (km) | Azimuth $\left(^{\circ}\right)$ | Take-off <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | :---: | ---: | :---: | :---: |
| RES | 680 | 138.7 | 49.1 | + |
| ALE | 847 | 47.2 | 49.1 | D |
| IGL | 1407 | 129.6 | 49.1 | D |
| INK | 1446 | 220.8 | 49.1 | + |
| YKW3 | 1931 | 186.8 | 49.6 | D |
| DAWY | 1973 | 244.6 | 43.8 | C |
| FRB | 2208 | 119.4 | 42.8 | + |
| WHY | 2287 | 215.8 | 40.5 | C |
| HYT | 2297 | 219.3 | 38.9 | C |
| FCC | 2413 | 156.9 | 38.7 | C |
| DLBC | 2476 | 208.0 | 35.0 | C |
| KUQ | 2807 | 123.9 | 34.5 | D |
| EDM | 2964 | 183.9 | 30.8 | C |
| LLLB | 3283 | 195.1 | 30.0 | C |
| ULM | 3344 | 161.5 | 32.7 | C |
| SOLO | 3399 | 3416 | 196.7 | 32.6 |


| Station | Distance (km) | Azimuth ( $\left.{ }^{\circ}\right)$ | Take-off <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | :---: | :---: | :---: | :---: |
| PGC | 3509 | 196.8 | 29.6 | C |
| TBO | 3570 | 153.7 | 32.4 | C |
| EEO | 3901 | 141.7 | 31.8 | - |
| LMQ | 3915 | 131.2 | 28.9 | - |
| CRLO | 3988 | 139.8 | 31.5 | C |
| SADO | 4105 | 142.2 | 31.2 | + |

## Focal Mechanism Solutions

| Strike | Dip | Rake |
| :---: | ---: | ---: |
| 101.65 | 42.06 | 31.11 |
| 107.73 | 44.81 | 35.53 |
| 107.75 | 42.06 | 31.11 |
| 113.83 | 44.81 | 35.53 |
| 119.93 | 44.81 | 35.53 |
| 125.42 | 47.85 | 39.32 |
| 131.52 | 47.85 | 39.32 |
| 76.63 | 42.27 | 17.14 |
| 83.53 | 43.96 | 22.18 |
| 89.96 | 46.03 | 26.73 |
| 90.07 | 43.96 | 22.18 |
| 96.5 | 46.03 | 26.73 |
| 102.48 | 48.44 | 30.79 |
| 109.02 | 48.44 | 30.79 |
| 84.58 | 50.14 | 22.91 |
| 90.41 | 52.24 | 26.57 |

Strike Dip ..... Rake
85.47 ..... 48.3618.88
91.64 22.91 ..... 50.14
97.47 ..... 52.24 ..... 26.57
total number solutions: ..... 19
grid search: ..... $5^{\circ}$
\# misfits: ..... 0
comments:
Preferred solution: take solution from above list that fits highest number of emergent stations using weighting scheme discussed previously; solution below fits all except FRB and EEO
Plane 1:
Strike: ..... $85^{\circ}$
Dip: ..... $48^{\circ}$
Rake: ..... $19^{\circ}$
Plane 2
Strike: ..... $343^{\circ}$
Dip: ..... $76^{\circ}$
Rake: ..... $137^{\circ}$
P-axis:
Trend: ..... $40^{\circ}$
Plunge: ..... $17^{\circ}$
T-axis:
Trend: ..... $295^{\circ}$
Plunge: ..... $40^{\circ}$
B-axis:
Trend: ..... $148^{\circ}$
Plunge: ..... $45^{\circ}$
Quality: B ; poor azimuthal coverage but large number ofobservations and high degree of redundancy atazimuths covered

## Figure Captions:

Figure 3a-f. Seismograms from which polarity data were read by the authors.
Figure 4. Focal mechanism solutions. Lower hemisphere projection. The P and T axes are shown on the left and the $P$ nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations.


Figure 3 a
Arctic Ocean: 1 January 1999


Figure 3b



Figure $3 c$

Arctic Ocean: 1



Figure 3d



Figure 3 e
Arctic Ocean: 1 January 1999





Figure 3 f
Arctic Ocean: 1 January 1999

Figure 4

## 21 FEBRUARY 1999: BAFFIN BAY



## Focal Mechanism Solutions

## total number solutions: 565

grid search: ..... $5^{\circ}$
\# misfits: ..... 1
comments: solutions based on weighted data; solutions misfit one ofYKW3, RES, SCHQ or ALE (listed in decreasing order ofnumber of solutions that misfit station)
Preferred solution: insufficient data to determine
Fault Plane:Strike: undeterminedDip: undeterminedRake: undetermined

```
P-axis:
    Trend: most solution SW-NW auadrant but others
        possible; not in S-SW octant
    Plunge: undetermined
T-axis:
    Trend: NW-SSE; not in S-NW range
    Plunge: undetermined
B-axis:
    Trend: undetermined
    Plunge: undetermined
```


## Figure Captions:

Figure 5a-b. Seismograms from which polarity data were read by the authors.
Figure 6. Focal mechanism solutions. Lower hemisphere projection. The $P$ axes are on the left, T axes on the right and B axes below. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations. Solutions based on weighted data.
Baffin Bay: 21 February 1999



Baffin Bay: 21 February 1999

Baffin Bay: 21 February 1999

B axes

## 2 MARCH 1999: DEVON ISLAND, NUNAVUT

| Seismic Zone: | Eastern Arctic Background |
| :--- | :--- |
| Latitude: | $76.32^{\circ} \mathrm{N}$ |
| Longitude: | $91.96^{\circ} \mathrm{W}$ |
| Depth: | 18 km (fixed) |
| Magnitude: | 4.0 ( $\mathrm{m}_{N}$ ) |
| Origin Time: | $07: 11: 13$ (UT) |
| Comments: |  |

## Polarity Data

| Station | Distance $(\mathrm{km})$ | Azimuth $\left(^{\circ}\right)$ | Take-off Angie <br> $\left(^{\circ}\right)$ | First Motion |
| :--- | :---: | :---: | :---: | :---: |
| RES | 201 | 205.7 | 49.1 | D |
| IGL | 843 | 151.8 | 49.1 | D |
| INK | 1620 | 258.7 | 49.1 | - |
| FRB | 1640 | 136.0 | 49.1 | + |
| SFJ | 1731 | 104.1 | 49.1 | + |
| YKW3 | 1750 | 221.0 | 49.1 | D |

## Focal Mechanism Solutions

| total number solutions: | 7477 impulsive only; 1021 using weighting scheme |
| :--- | :--- |
| grid search: | $5^{\circ}$ |
| \# misfits: | 0 |
| comments: | ALE unavailable |

Preferred solution: insufficient data to determine; comments based on solutions from weighted data

Nodal Planes: cannot be determined
P-axis: majority of solutions in SW quadrant; some solutions in NW quadrant and SSE octant

T-axis: majority of solutions in NE quadrant; some nearhorizontal solutions in ESE and NNW octants

| B-axis: | not near-vertical; not near-horizontal if at NE-SE |
| :--- | :--- |
| or SSW-W azimuths; aimost any other |  |
| orientation is possible |  |

## Figure Captions:

Figure 7a-b. Seismograms from which polarity data were read by the authors.
Figure 8. Focal mechanism solutions. Lower hemisphere projection. The $P$ axes are shown on the left and the $T$ axes on the right. Data are also plotted (lower). Solid symbols indicate compressional first motions and open symbols dilatations. Solutions shown are based on weighted data.
DEVON ISLAND, NUNAVUT: 2 MARCH 1999



DEVON ISLAND, NUNAVUT: 2 MARCH 1999


Figure 7b


## 16 MARCH 1999: CÔTE-NORD, QUEBEC

| Seismic Zone: | Lower St. Lawrence |
| :--- | :--- |
| Latitude: | $49.61^{\circ} \mathrm{N}$ |
| Longitude: | $66.32^{\circ} \mathrm{W}$ |
| Depth: | 18 km (fixed), 19.2 km (Bent and Perry, 2000) <br> Magnitude: <br> Origin Time: <br> Comments: <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br>  <br> felt in eastern Quebec and northern New Brunswick; largest <br> instrumentally recorded earthquake in Lower St. Lawrence seismic <br> zone |

## Polarity Data

| Station | Distance (km) | Azimuth $\left(^{\circ}\right)$ | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | :---: | :---: | :---: | :---: |
| ICQ | 70 | 261.9 | -75.5 | D |
| SMQ | 73 | 338.0 | -76.2 | D |
| GSQ | 97 | 216.8 | -79.5 | D |
| CNQ | 132 | 255.4 | -82.2 | D |
| MNQ | 203 | 301.1 | 49.1 | D |
| A21 | 327 | 230.7 | 49.1 | C |
| A64 | 330 | 234.3 | 49.1 | C |
| A61 | 351 | 233.9 | 49.1 | C |
| A16 | 362 | 230.2 | 49.1 | - |
| LMQ | 375 | 233.7 | 49.1 | C |
| A54 | 386 | 233.1 | 49.1 | - |
| A11 | 390 | 228.7 | 49.1 | C |
| DAQ | 406 | 245.0 | 49.1 | + |
| LMN | 434 | 164.3 | 49.1 | + |
| DPQ | 581 | 238.3 | 49.1 | + |
| SCHQ | 582 | 225.3 | 49.1 | C |
| MOQ | 655 | 49.1 | D |  |


| Station | Distance (km) | Azimuth ( ${ }^{\circ}$ ) | Take-off Angle $\left({ }^{\circ}\right)$ | First Motion |
| :---: | :---: | :---: | :---: | :---: |
| LG4Q | 700 | 312.6 | 49.1 | D |
| MNT | 715 | 233.0 | 49.1 | + |
| OTT | 849 | 240.0 | 49.1 | - |
| WBO | 849 | 236.2 | 49.1 | D |
| CRLO | 918 | 248.5 | 49.1 | - |
| KUQ | 951 | 352.6 | 49.1 | D |
| EEO | 1005 | 255.6 | 49.1 | - |
| SADO | 1110 | 245.8 | 49.1 | D |
| KAPO | 1171 | 275.3 | 49.1 | - |
| FRB | 1581 | 356.0 | 49.1 | + |
| TBO | 1683 | 275.2 | 49.1 | C |
| SOLO | 1847 | 281.3 | 49.6 | + |
| FCC | 2060 | 310.2 | 45.1 | D |
| ULM | 2111 | 283.2 | 45.1 | + |
| RES | 3089 | 344.2 | 30.4 | D |
| EDM | 3215 | 295.3 | 30.1 | - |
| YKW1 | 3222 | 314.6 | 30.1 | D |
| YKW2 | 3226 | 314.5 | 30.1 | D |
| YKW3 | 3228 | 314.8 | 30.1 | D |
| INK | 4100 | 325.3 | 28.5 | + |
| FINES | 5410 | 39.3 | 25.3 | C |

## Focal Mechanism Solutions

| Strike | Dip | Rake |
| :---: | :---: | :---: |
| 320.26 | 58.23 | 25.7 |

total number solutions: 1
grid search: ..... $5^{\circ}$
\# misfits: ..... 3
comments: ..... misfits SMQ, MOQ and TBO
Preferred solution:
Plane 1:
Strike: ..... $320^{\circ}$
Dip: ..... $58^{\circ}$
Rake: ..... $26^{\circ}$
Plane 2:
Strike: ..... $216^{\circ}$
Dip: ..... $68^{\circ}$
Rake: ..... $146^{\circ}$
P-axis:
Trend: ..... $270^{\circ}$
Plunge: ..... $6^{\circ}$
T-axis:
Trend: ..... $175^{\circ}$
Plunge: ..... $39^{\circ}$
B-axis:
Trend: ..... $8^{\circ}$
Plunge: ..... $50^{\circ}$
Quality: B; good azimuthal coverage with redundancy; wouldhave been an " $A$ " except for the fact that it misfits SMQ
Figure Captions:

Figure 9a-j. Seismograms from which polarity data were read by the authors.
Figure 10. Focal mechanism solutions. Lower hemisphere projection. The P and T axes are shown on the left and the P nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations.








CÔTE-NORD, QUEBEC: 16 MARCH 1999




CÔTE-NORD, QUEBEC: 16 MARCH 1999

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Côte-Nord, Quebec: 16 March 1999

Figure 10

# 29 JULY 1999: LABRADOR SEA 

| Seismic Zone: | Labrador Sea |
| :--- | :--- |
| Latitude: | $60.90^{\circ} \mathrm{N}$ |
| Longitude: | $58.03^{\circ} \mathrm{W}$ |
| Depth: | 18 km (fixed) |
| Magnitude: | $5.3\left(\mathrm{M}_{\mathrm{L}}\right) 4.8\left(\mathrm{~m}_{\mathrm{b}}\right)$ |
| Origin Time: | $12: 50: 59(\mathrm{UT})$ |
| Comments: |  |

## Polarity Data

| Station | Distance (km) | Azimuth ( ${ }^{\circ}$ ) | Take-off Angle ( ${ }^{\circ}$ | First Motion |
| :---: | :---: | :---: | :---: | :---: |
| FRB | 631 | 304.8 | 49.1 | - |
| KUQ | 665 | 246.7 | 49.1 | + |
| SCHQ | 854 | 221.5 | 49.1 | - |
| LG4Q | 1259 | 237.1 | 49.1 | C |
| DRLN | 1298 | 178.3 | 49.1 | - |
| SMQ | 1307 | 208.4 | 49.1 | C |
| MNQ | 1335 | 214.9 | 49.1 | + |
| ICQ | 1395 | 208.8 | 49.1 | C |
| CNQ | 1439 | 210.7 | 49.1 | C |
| IGL | 1449 | 320.7 | 49.1 | + |
| GSQ | 1454 | 207.4 | 49.1 | + |
| A64 | 1643 | 212.9 | 49.1 | + |
| A61 | 1662 | 213.2 | 49.1 | D |
| A16 | 1682 | 212.7 | 49.1 | D |
| LMQ | 1684 | 213.5 | 49.1 | C |
| A54 | 1696 | 213.6 | 49.1 | D |
| A11 | 1711 | 212.8 | 49.1 | D |
| LMN | 1734 | 197.9 | 49.1 | + |
| TRQ | 1956 | 220.8 | 47.4 | - |
| GRQ | 1967 | 224.0 | 47.4 | + |


| Station | Distance (km) | Azimuth $\left(^{\circ}\right.$ ) | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :---: | :---: | :---: | :---: | :---: |
| KAPO | 1996 | 214.4 | 47.4 | + |
| CRLO | 2080 | 226.0 | 45.1 | - |
| EEO | 2091 | 230.1 | 45.1 | C |
| RES | 2125 | 330.8 | 45.1 | - |
| SADO | 2271 | 227.3 | 42.8 | + |
| TBO | 2403 | 249.7 | 38.7 | - |
| SOLO | 2435 | 255.6 | 38.7 | C |
| ULM | 2619 | 260.3 | 35.5 | D |
| EDM | 3341 | 280.5 | 29.9 | + |
| INK | 3463 | 316.0 | 29.7 | C |
| WALA | 3679 | 274.9 | 29.3 | C |
| DLBC | 3861 | 298.8 | 29.0 | + |
| DAWY | 3923 | 311.3 | 28.9 | D |
| PNT | 3967 | 279.7 | 28.8 | C |
| LLLB | 3991 | 283.0 | 28.8 | C |
| HYT | 4048 | 305.9 | 28.6 | C |

Comments: FCC, DAQ and A21 not functioning
Focal Mechanism Solutions

| Strike | Dip | Rake |
| ---: | ---: | ---: |
| 133.30 | 41.41 | 40.89 |
| 96.48 | 43.96 | -22.18 |
| 103.37 | 42.27 | -17.14 |
| 117.21 | 41.03 | -11.69 |
| 107.76 | 45.86 | -9.85 |
| 114.71 | 45.22 | -4.98 |
| 86.70 | 58.23 | -25.70 |

total number solutions: ..... 7
grid search: ..... $5^{\circ}$
\# misfits: ..... 2
comments: all solutions misfit LMQ and ULM
Preferred solution: use weighting scheme to determine whether thrust or normalcomponent; solution 4 most consistent with emergent dataPlane 1:
Strike: ..... $117^{\circ}$
Dip: ..... $41^{\circ}$
Rake: ..... $-12^{\circ}$
Plane 2:
Strike: ..... $216^{\circ}$
Dip: ..... $41^{\circ}$
Rake: ..... $-130^{\circ}$
P-axis:
Trend: ..... $90^{\circ}$
Plunge: ..... $39^{\circ}$
T-axis:
Trend: ..... $337^{\circ}$
Plunge: ..... $26^{\circ}$
B-axis:
Trend: ..... $223^{\circ}$
Plunge: ..... $40^{\circ}$Quality: $\quad \mathrm{D}$; no stations in eastern hemisphere but large number of stationsand good azimuthal redundancy in western hemisphere; emergentdata used to determine sign of dip-slip component

## Figure Captions:

Figure 11a-i. Seismograms from which polarity data were read by the authors.
Figure 12. Focal mechanism solutions. Lower hemisphere projection. The $P$ and $T$ axes are shown on the left and the $P$ nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations.

LABRADOR SEA: 29 JULY 1999




Figure 11b





6เト อกถธ!
LABRADOR SEA: 29 JULY 1999





LABRADOR SEA: 29 JULY 1999

P nodal planes

## 31 OCTOBER 1999: MORIN HEIGHTS, QUEBEC

Seismic Zone: western Quebec<br>Latitude: $\quad 45.84^{\circ} \mathrm{N}$<br>Longitude: $\quad 74.31^{\circ} \mathrm{W}$<br>Depth:<br>Magnitude:<br>Origin Time:<br>Comments:<br>18 km (fixed), 20.8 km (Bent and Perry, 2000)<br>$4.2\left(\mathrm{~m}_{\mathrm{bLg}}\right)$<br>20:14:10 (UT)<br>felt in western Quebec and eastern Ontario

Polarity Data

| Station | Distance (km) | Azimuth $\left({ }^{\circ}\right)$ | Take-off Angle $\left(^{\circ}\right)$ | First Motion |
| :---: | :---: | :---: | :---: | :---: |
| TRQ | 45 | 334.5 | -68.3 | D |
| MNT | 66 | 126.3 | -74.7 | C |
| GAC | 93 | 259.9 | -79.0 | D |
| GAC | 93 | 259.9 | -79.0 | B |
| GAC | 93 | 259.9 | -79.0 | $>$ |
| BGR | 114 | 182.8 | -81.0 | C |
| OTT | 122 | 245.5 | -81.6 | D |
| WBO | 122 | 219.0 | -81.6 | D |
| GRQ | 146 | 305.4 | -83.0 | C |
| DPQ | 149 | 51.5 | -83.1 | D |
| PTN | 153 | 200.7 | -83.3 | 0 |
| MOQ | 171 | 110.0 | 49.1 | + |
| BRC | 188 | 212.8 | 49.1 | D |
| FINE | 190 | 201.3 | 49.1 | D |
| CRLO | 240 | 275.9 | 49.1 | C |
| QCQ | 255 | 65.2 | 49.1 | + |
| A54 | 347 | 57.7 | 49.1 | - |
| LMQ | 358 | 56.8 | 49.1 | D |


| Station | Distance $(\mathrm{km})$ | Azimuth $\left({ }^{\circ}\right)$ | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | :---: | ---: | :---: | :---: |
| EEO | 378 | 285.1 | 49.1 | + |
| WLVO | 388 | 237.9 | 49.1 | + |
| SADO | 398 | 254.1 | 49.1 | - |
| A64 | 402 | 55.4 | 49.1 | D |
| PKRO | 431 | 242.5 | 49.1 | - |
| ACTO | 520 | 243.3 | 49.1 | - |
| GPD | 538 | 181.4 | 49.1 | - |
| PAL | 540 | 176.5 | 49.4 | $<$ |
| GSQ | 641 | 55.4 | 49.1 | D |
| MNQ | 663 | 36.3 | 49.1 | + |
| KAPO | 734 | 305.9 | 49.1 | C |
| SMQ | 746 | 46.6 | 49.1 | + |
| LG4Q | 865 | 0.9 | 49.1 | + |

## Focal Mechanism Solutions



> Lamont-Doherty (LDEO) Network stations were examined; those from which polarities could be picked are noted in the data table; we were unable to pick polarities from the SONT stations BRCO, LDN, STCO ,TYNO, RD01, RD02, RD03 and RD04 or the LDEO stations ARNY, CHIP, CRNY, MANY, MSNY, TBR and WCC

$1^{\text {st }}$ solution has misfit at GRQ; GRQ is near Pg-Pn crossover distance; if if is actually a Pn, it would fit solution but we believe it has been correctly identified as a Pg; other 4 solutions misfit SH at GAC
same 5 -solutions result if we include emergent polarities using weighting scheme discussed in text so can't use emergent data to distinguish between 2 families of solutions
note that all solutions have one nodal plane with similar orientation (see Figure 14); use that plane to calculate mean and then take closest actual solution (solution 5)

## Preferred solution:

Plane 1:
Strike: ..... $95^{\circ}$
Dip: ..... $79^{\circ}$
Rake: ..... $44^{\circ}$
Plane 2:
Strike: ..... $355^{\circ}$
Dip: ..... $47^{\circ}$
Rake: ..... $165^{\circ}$
P-axis:
Trend: ..... $218^{\circ}$
Plunge: ..... $21^{\circ}$
T-axis:
Trend: ..... $325^{\circ}$
Plunge: ..... $38^{\circ}$
B-axis:
Trend: ..... $106^{\circ}$
Plunge: ..... $45^{\circ}$
Quality: B: good azimuthal coverage and redundancy but some uncertainty in strike

Figure Captions:

## Figure Captions:

Figure 13a-h. Seismograms from which polarity data were read by the authors.
Figure 14. Focal mechanism solutions. Lower hemisphere projection. The $P$ and $T$ axes are shown on the left and the P nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations. S data and nodal planes are plotted at the bottom. The arrows indicate the first motion direction. The solid lines represent the SH nodal planes and the dashed lines the SV nodal planes.



Morin Heights, Quebec: 31 October 1999




Morin Heights, Quebec: 31 October 1999



Figure $13 f$
Morin Heights, Quebec: 31 October 1999





Morin Heights, Quebec: 31 October 1999


## 2 NOVEMBER 1999: LAURENTIAN SLOPE

| Seismic Zone: | Laurentian Slope |
| :--- | :--- |
| Latitude: | $45.20^{\circ} \mathrm{N}$ |
| Longitude: | $56.08^{\circ} \mathrm{W}$ |
| Depth: | $18 \mathrm{~km}($ fixed $)$ |
| Magnitude: | $4.5\left(\mathrm{M}_{\mathrm{L}}\right)$ |
| Origin Time: | $08: 45: 21(\mathrm{UT})$ |
| Comments: |  |

## Polarity Data

| Station | Distance $(\mathrm{km})$ | Azimuth $\left({ }^{\circ}\right)$ | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | :---: | ---: | :---: | :---: |
| STJN | 368 | 43.1 | 49.1 | C |
| GBN | 427 | 275.0 | 49.1 | D |
| DRLN | 464 | 347.1 | 49.1 | C |
| LMN | 686 | 279.2 | 49.1 | 0 |
| ICQ | 972 | 303.6 | 49.1 | + |
| SMQ | 973 | 308.8 | 49.1 | C |
| CNQ | 1015 | 301.0 | 49.1 | 0 |
| A21 | 1081 | 289.8 | 49.1 | C |
| A64 | 1099 | 290.3 | 49.1 | + |
| A54 | 1131 | 287.9 | 49.1 | - |
| SCHQ | 1318 | 328.3 | 49.1 | + |
| GAC | 1514 | 279.0 | 49.1 | - |

## Focal Mechanism Solutions

total number solutions: ..... 950
grid search: ..... $5^{\circ}$
\# misfits: ..... 0
comments:

| Fault Planes: | insufficient data to determine |
| :---: | :--- |
| P-axis: |  |
| Trend: <br> Plunge: | southern hemisphere <br> undetermined |
| T-axis: |  |
| Trend: | probably NNW-NNE; anything in the NW to SE <br> hemisphere is possible |
| Plunge: | undetermined |
| B-axis: |  |
| Trend: | undetermined |
| Plunge: | not near vertical; not near horizontal if trend is <br>  |

## Figure Captions:

Figure 15a-c. Digital seismograms from which polarity data were read by the authors.
Figure 16. Focal mechanism solutions. Lower hemisphere projection. The $P$ axes are shown on the left and the $T$ axes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations.



Laurentian Slope: 2


Figure 15b

2 November 1999: Laurentian Slope

Figure 16

Seismic Zone: Latitude: Longitude: Depth: Magnitude: Origin Time: Comments:
none
43.71 N
$78.99^{\circ} \mathrm{W}$
10 km
$3.8\left(\mathrm{~m}_{\mathrm{N}}\right)$
22:33:01 (UT)
included (despite magnitude $<4$ ) because felt in the Toronto area Polarity Data

| Station | Distance $(\mathrm{km})$ | Azimuth $\left({ }^{\circ}\right)$ | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | ---: | ---: | ---: | :---: |
| PKRO | 29 | 347.5 | -71.2 | C |
| WLVO | 54 | 63.0 | -79.4 | D |
| EFO | 73 | 200.9 | -82.2 | C |
| RD01 | 76 | 244.6 | -82.5 | - |
| RD02 | 84 | 239.0 | -83.2 | D |
| RD04 | 84 | 223.2 | -83.2 | D |
| ACTO | 87 | 263.2 | -83.4 | D |
| RD03 | 97 | 233.3 | -84.1 | D |
| TYNO | 98 | 226.6 | -84.2 | D |
| SADO | 119 | 354.3 | -85.2 | C |
| SADO | 119 | 354.3 | -85.2 | F |
| SADO | 119 | 354.3 | -85.2 | $<$ |
| BRCO | 206 | 287.7 | 49.1 | C |
| CRLO | 289 | 25.6 | 49.1 | D |
| TRQ | 448 | 49.8 | 49.1 | D |

## Focal Mechanism Solutions

| Strike | Dip | Rake |
| ---: | ---: | ---: |
| 120.65 | 60.50 | -28.34 |
| 122.74 | 66.07 | -26.34 |

total number solutions: ..... 2
grid search: ..... $5^{\circ}$
\# misfits: ..... 0
comments: looked at data from USGS stations BINY, AAM and LBNH butwere unable to pick polarities; examined data from entireSONT network but could pick polarities only from thosestations noted in list
SADO S polarities from rotated horizontal records
Preferred solution: solutions similar; use solution 1
Plane 1
Strike: ..... $120^{\circ}$
Dip: ..... $61^{\circ}$
Rake: ..... $-28^{\circ}$
Plane 2
Strike: ..... $226^{\circ}$
Dip: ..... $66^{\circ}$
Rake: ..... $-147^{\circ}$
P-axis:
Trend: ..... $85^{\circ}$
Plunge: ..... $40^{\circ}$
T-axis:
Trend: ..... $352^{\circ}$
Plunge: ..... $3^{\circ}$
B-axis:
Trend: ..... $258^{\circ}$
Plunge: ..... $50^{\circ}$
Quality: B good azimuthal coverage but a few gaps; redundancy at some azimuths

## Figure Captions:

Figure 17a-d. Seismograms from which polarity data were read by the authors.
Figure 18. Focal mechanism solutions. Lower hemisphere projection. The $P$ and $T$ axes are shown on the left and the $P$ nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations. The $S$ data are plotted on the lower figure with the arrows pointing in the direction of the first motion. Solid lines represent SH nodal planes and dashed lines the SV nodal planes.


Figure 17a


Lake Ontario: 26 November 1999



## 7 DECEMBER 1999: McCLURE STRAIT, NORTHWEST TERRITORIES

Seismic Zone: none
Latitude: $\quad 75.38^{\circ} \mathrm{N}$
Longitude: $\quad 120.44^{\circ} \mathrm{W}$
Depth:
Magnitude:
Origin Time:
18 km (fixed)
$5.7\left(M_{L}\right) 4.7\left(m_{b}\right.$ NEIS $)$
05:35:58 (UT)
Comments:
a magnitude $4.0\left(\mathrm{M}_{\mathrm{L}}\right)$ aftershock followed at 06:02:49 (UT); focal mechanism not determined for aftershock due to insufficient data; first motion at RES same as mainshock; both are technically outside our area of interest

## Polarity Data

| Station | Distance (km) | Azimuth $\left({ }^{\circ}\right)$ | Take-off Angle <br> $\left({ }^{\circ}\right)$ | First Motion |
| :--- | ---: | ---: | ---: | :--- |


| Station | Distance (km) | Azimuth ( ${ }^{\circ}$ ) | Take-off Angle $\left({ }^{\circ}\right)$ | First Motion |
| :---: | :---: | :---: | :---: | :---: |
| LG4Q | 3135 | 114.2 | 33.0 | C |
| TBO | 3303 | 133.9 | 33.0 | C |
| SCHQ | 3206 | 150.0 | 32.9 | + |
| KAPO | 3367 | 127.3 | 32.7 | D |
| MNQ | 3593 | 110.5 | 32.4 | C |
| EEO | 3738 | 122.2 | 29.4 | C |
| CNQ | 3761 | 107.9 | 29.3 | + |
| DAQ | 3790 | 112.2 | 29.3 | - |
| GRQ | 3817 | 118.4 | 29.3 | + |
| A64 | 3839 | 110.8 | 29.2 | C |
| CRLO | 3841 | 120.5 | 29.2 | $+$ |
| A61 | 3848 | 111.1 | 29.2 | C |
| A21 | 3857 | 110.6 | 29.2 | C |
| A54 | 3864 | 111.6 | 29.2 | + |
| DPQ | 3885 | 114.8 | 29.1 | D |
| TRQ | 3890 | 117.1 | 29.1 | C |
| A11 | 3892 | 111.5 | 29.1 | C |
| GAC | 3922 | 118.5 | 29.1 | C |
| SADO | 3935 | 123.4 | 29.1 | C |
| DRLN | 4025 | 96.6 | 28.7 | C |
| MOQ | 4042 | 115.1 | 28.8 | C |
| LMN | 4178 | 106.6 | 28.5 | C |

## Focal Mechanism Solutions

Strike Dip Rake
168.02 ..... 60.5042.39
130.45 ..... 79.45 ..... 44.01
134.07 82.95 ..... 44.56
164.25 66.07 ..... 39.32
141.00 83.59 ..... 39.57
148.77 77.30 ..... 38.26
156.14 ..... 71.25 ..... 36.01
160.34 74.24 ..... 37.25
163.97 71.2536.01
total number solutions: ..... 9
grid search: ..... $5^{\circ}$
\# misfits: ..... 2
comments: ALE not availablemisfits for all solutions at KUQ and KAPO
Preferred solution: solutions similar; use ciosest to mean
Plane 1:
Strike: ..... $149^{\circ}$
Dip: ..... $77^{\circ}$
Rake: ..... $38^{\circ}$
Plane 2:
Strike: ..... $49^{\circ}$
Dip: ..... $53^{\circ}$
Rake: ..... $164^{\circ}$
P-axis:
Trend: ..... $274^{\circ}$
Plunge: ..... $16^{\circ}$
T-axis:
Trend: ..... $16^{\circ}$
Plunge: ..... $36^{\circ}$
B-axis:
Trend: ..... $164^{\circ}$
Plunge: ..... $50^{\circ}$
Quality: B: many observations and good azimuthal redundancy but pooroverall azimuthal coverage

## Figure Captions:

Figure 19a-j. Seismograms from which polarity data were read by the authors.
Figure 20. Focal mechanism solutions. Lower hemisphere projection. The P and T axes are shown on the left and the $P$ nodal planes on the right. Data are also plotted. Solid symbols indicate compressional first motions and open symbols dilatations.


Figure 19a



Figure 19c

 McClure Strait, NWT:



McClure Strait, NWT: 7 December 1999





McClure Strait, NWT: 7 December 1999





Figure 19i

McClure Strait, NWT:


Figure 19j
McClure Strait, NWT: 7 December 1999

Figure 20

## DISCUSSION

Southeast Mainland: As was the case in previous focal mechanism studies, the southeastern mainland was the region where it was most consistently possible to obtain well constrained focal mechanisms. We obtained solutions for all three events. Event 990316 is one of the few earthquakes in the Lower St. Lawrence seismic zone for which a focal mechanism has been determined. The mechanism implies a combination of strikeslip and thrust motion. One of the nodal planes is roughly parallel to the trend of the St. Lawrence River. The existing mechanisms for lower St. Lawrence earthquakes vary in detail but most have one nodal plane striking approximately parallel to the St. Lawrence River and have a thrust component of slip (Adams et al., 1988, 1989; Bent and Drysdale, 2000).

Event 991031 is typical of western Quebec earthquakes in that it has a primarily thrust mechanism with one nodal plane striking roughly northwest (for example, Wahlström, 1987; Adams et al., 1988, 1989; Adams, 1991; Bent, 1996a, b; Bent and Perry, 1999; Bent et al., 1999; Bent and Drysdale, 2000). This earthquake was of particular interest as it was the fourth earthquake of magnitude greater than 4.0 to occur in the region of Ste-Agathe-des-Monts during the last four years. All have typical western Quebec focal mechanisms but are not identical to each other (Bent et al., 1999). The focal mechanism of 991031 most closely resembles that of 970524, which is also the closest event to it spatially.

Event 991126 beneath Lake Ontario is the smallest of the events studied. We obtained an oblique-normal mechanism with faulting on either a SE-NW or SW-NE striking plane. The normal component of faulting is somewhat unusual for the southeastern mainland, which is dominated by thrust faulting. However, the P axis orientation is within the range expected for southeastern Canada. A preliminary focal mechanism for another event beneath Lake Ontario (20000524, Bent and Drysdale, unpublished data) indicates either thrust motion on a steeply dipping northwest-striking plane or oblique thrust on a shallowly dipping northeast striking plane.

Southeast Offshore: Because of poor azimuthal coverage it is difficult to obtain well constrained first motion focal mechanisms for small earthquakes in the Laurentian Slope region. We were unable to determine a focal mechanism for event 991102. Constraints that could be placed on the stress axes were minimal.

Labrador Sea- Baffin Bay: It is rarely possible to determine focal mechanisms from first motion data for small to moderate Baffin Bay earthquakes due to poor azimuthal coverage. We were not able to obtain a focal mechanism for the Baffin Bay event 990221. Azimuthal coverage for the Labrador Sea is somewhat better although the larger magnitude (5.3) of the Labrador Sea event 990729 was also a factor in being able to obtain a solution. Both nodal planes indicate a combination of strike-slip and normal faulting. Previous studies of Labrador Sea earthquakes have found evidence for all types of faulting (Bent and Hasegawa, 1992; Bent and Perry, 1999, 2000; Bent and Drysdale, 1999). The P-axis trend of $90^{\circ}$ is in the range of typical values for eastern Canada.

Far North: We were able to determine focal mechanisms for two of the three events in this region. As was seen in previous studies, the likelihood of obtaining a solution is a combination of earthquake magnitude and location. Not surprisingly, it becomes easier to determine focal mechanisms as the magnitude increases. It is also generally easier to obtain solutions for events in the central and western part of this region than for the
east due to differences in azimuthal coverage. A mechanism could not be determined for the magnitude 4.0 event 990302 on Devon Island.

Event 990101 in the Arctic Ocean was the result of either strike-slip faulting along an eastwest striking plane or oblique thrust motion along a near north-south striking plane. There are no focal mechanisms for events very near the epicenter of 990101 but those for events in adjacent regions exhibit many similarities to that of event 990101. Event 96022 of Bent and Perry (1999) near Melville Island (to the south of 990101) also had one nodal plane indicative of strike-slip faulting on an east-west plane. Hasegawa (1977) obtained primarily strike-slip mechanisms for four events in the Byam Martin Channel (southeast of 990101). Similarly, Bent and Perry (2000) found evidence for strike-slip (or oblique thrust) faulting in the Byam Martin Channel.

Our focal mechanism for event 991207 in the McClure Strait is indicative of oblique thrust faulting on a plane striking near east-west or northeast-southwest. This mechanism is similar to those of the events discussed in the previous paragraph, all of which occurred in the regions just to the east of 991207.

Figure 21. Focal mechanism solutions for earthquakes in the southeastern mainland. Lower hemisphere projection with shaded regions representing compressional quadrants. In this and subsequent figures black shading is used for $A$ and $B$ quality solutions, gray shading for $C$ and $D$ quality and a question mark (?) Indicates that a solution could not be determined. Symbol size is scaled to magnitude. Earthquakes are identified by date of occurrence.

Figure 22. Focal mechanism solutions for earthquakes in the southeastern offshore region. Format is the same as for Figure 21.

Figure 23. Focal mechanism solutions for earthquakes in the Labrador Sea to Baffin Bay region. Format is the same as for Figure 21.

Figure 24. Focal mechanism solutions for earthquakes in the far north. Format is the same as for Figure 21.


Figure 21


Figure 22


Figure 23


Figure 24

## CONCLUSIONS

We determined focal mechanism solutions from first motions for six of the nine earthquakes studied. The solutions are summarized in the Table. As was the case in previous such efforts, the events for which we could not obtain solutions occurred in offshore and northern regions where the azimuthal coverage of the CNSN and other seismograph networks is poor and station density is low. Alternate methods of focal mechanism determination, which may prove more useful than first motions in these regions, are under investigation. In contrast, in southeastern Canada both the azimuthal coverage and station density are better and we can obtain solutions for events of magnitude less than 4.0. Focal mechanisms determined in this study are consistent with those of many earlier studies in that we find evidence for primarily thrust faulting in the seismic zones adjacent to the St. Lawrence River, a mix of fault types in the Labrador Sea, and a significant component of strike-slip motion in the western Arctic.

## ACKNOWLEDGMENTS

We thank our non-GSC colleagues who supplied us with waveforms and/or polarity information and John Adams for reviewing the manuscript. All maps in this paper were generated using GMT (Wessel and Smith, 1991).

$$
\begin{gathered}
\text { Event } \\
990101 \\
990221 \\
990302 \\
990316 \\
990729 \\
991031 \\
991102 \\
991126 \\
991207
\end{gathered}
$$

$$
\begin{array}{ll}
* & \\
\mathrm{SS}=\text { strike-slip } & \text { (rake within } 30^{\circ} \text { of pure strike-slip) } \\
\mathrm{T}=\text { thrust } & \text { (rake within } 30^{\circ} \text { of pure thrust) }
\end{array}
$$

TABLE
sUMMARY OF SOLUTIONS
oblique-thrust)
data (abstract), Earthquake Notes, 55 (3), 15.
Tapley, W. C. and J. E. Tull (1991). SAC- Seismic Analysis Code Revision 4, Lawrence Livermore National Laboratory, Livermore, California, non-paginated.

Wahlström, R. (1987). Focal mechanisms of earthquakes in southern Quebec, southeastern Ontario and northeastern New York with implications for regional seismotectonics and stress field characteristics, Bull. Seism. Soc. Am., 77, 891924.

Wessel, P. and W. H. F. Smith (1991). Free software helps map and display data, EOS Trans. Am. Geophys. U., 72, 441, 445-446.

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$$

