

INTRODUCTION

New 1:50,000-scale mapping in the Mackenzie-Victory lakes area (Figure 1) indicates that: 1) the belt of allochthonous rocks extending for ca. 40 km from south of Mackenzie Lake to northwestern Victory Lake and previously mapped as "Mackenzie Lake metamorphism" (Bell, 1968, 1971; Davidson, 1970; Hamner et al., 1998) consists of the lower part of the Hurwitz Group; 2) Archean volcanic rocks (Kaminak Group) which host base metal massive sulphide prospects are unconformably beneath the Hurwitz Group; 3) both basement and cover experienced pervasive deformation (including light to isoclinal folding) and upper greenschist facies metamorphism before ca. 1.83 Ga (presented age of unroofing and unmetamorphosed lamprophyre dykes); and 4) structures related to intense Paleoproterozoic deformation, and the effects of Archean strain are not obvious.

Kaminak Group (units Am; A1)
The oldest lithostratigraphic unit in the area comprises a succession of bimodal mafic/felsic volcanic rocks and volcanogenic sedimentary intervals. The volcanic rocks form a continuum that ranges from mafic flows with local felsic intercalations (unit Am) to felsic flows and tuffs with minor mafic layers (unit A1). The mafic volcanic rocks are typically massive but locally contain pillow and brecciated beds. Similarly, quartz-felsic volcanic rocks are typically massive and contain only rare proximal-eye volcanic breccias. Highly deformed tuffs and tuffaceous sandstones and shales from metacarbonate beds within mixed mafic-felsic units. Mafic volcanic and tuffaceous units commonly contain coarse hornblende-plagioclase ± garnet, some pelitic beds with tuffaceous units are cordierite-bearing.

Granite, granodiorite, diorite (unit Ag)
On the northern and northern margins of the map area, granitic to dioritic rocks (unit Ag) are in fault contact with the Kaminak Group. Intrusive relationships are inferred because, adjacent to the contact, the granitic bodies contain numerous supracrustal enclaves and the supracrustal rocks are cut by granitic satellite dykes.

Gabbro dykes (Kaminak swarm; unit Kdy)
Rare north and northeast-trending gabbro dykes (unit Kdy), locally containing felsic megacrysts, cross cut the Archean supracrustal and granitic rocks. Significantly (see Structure below), the dykes contain a penetrative foliation that is concordant with foliation in the host rocks. They are likely part of the ca. 2.45 Ga (Hamner, 1998) Kaminak swarm.

Hurwitz Group (units PHk; PHm; PHp; PHm ± p; PHm; PHm ± p)
A sequence of allochthonous rocks originally mapped as "Mackenzie Lake metamorphism" outcrops in a belt from southwest of Mackenzie Lake to northwest of Victory Lake. Contacts with the Kaminak Group are inferred but a basal unconformity is inferred because: 1) conglomerate-bearing beds in the sequence contain clasts derived from supracrustal granitic and volcanic rocks; 2) granitic dykes and megacrystic gabbro dykes which cut Archean volcanic strata are lacking; and 3) the change from exclusively volcanogenic rocks in the Kaminak Group to exclusively allochthonous rocks is abrupt and signifies a dramatic change in depositional regime. Three main units are defined. As described below, individually and collectively these units bear strong resemblance to lower Hurwitz Group strata preserved elsewhere (see Aspler and Chalmers, 1997).

The lowest unit (unit PHk) consists of massive, parallel stratified and locally cross-stratified subarkose to quartz arkose, local quartz-pebble conglomerates and some pebbles. The maximum thickness of this unit is ca. 600 m, but locally only a few metres are preserved because of fault cut-out adjacent to basement. This unit, considered equivalent to the Nomat Formation (Aspler and Chalmers, 1998-9), signifies fluvial sedimentation during initial stages of Hurwitz Basin subsidence.

The middle unit (unit PHm; PHm ± p) defines a thick (up to 1000 m) section consisting predominantly of cobble-boulder conglomerates (see Figure 5 in Aspler et al., 2000). These conglomerates are typically massive and block-bedded, but locally define channels and graded sheets, and contain parallel and trough cross-stratified wedges and sheets of subarkose and arkosa. Clasts (up to 1 m in diameter) are generally subrounded to well rounded and are well-sorted in a coarse arkose matrix. Monomictic granitic megacrystic conglomerates (unit PHm) and polymictic conglomerates (unit PHm ± p) locally define discrete mappable sub-units. The monomictic conglomerates consist almost exclusively of different types of granitic clasts (fine to coarse grained, locally gabbroic) with rare amphibole and gabbroic fragments, whereas the polymictic conglomerates contain a diverse suite including granitic, quartzitic, felsic volcanic and mafic volcanic clasts. Locally, sections of monomictic conglomerates contain polymictic conglomerate sheets (unit PHm ± p). Significantly, the clasts lack a pre-existing tectonic fabric (see Structure below). Diastrophic zones of intense alteration occur within the granitic clasts. The appearance of stratified conglomerates and immature arkoses above a section of relatively mature arkoses containing quartz pebble conglomerates (see Figure 5 in Aspler et al., 2000) is interpreted as evidence of a sequence of erosion and deposition. The upper unit (unit PHp) is a 100 m thick and consists of grey to white subarkose to quartz arkose, local quartz-pebble conglomerates, and some pebbles. These conglomerates are generally well-sorted and concentrated in a scale framework-triad beds, occur sporadically throughout the unit. One ca. 100 m-thick horizon, consisting of thick bedded (20 to 10 m) sheets of heavy mineral-bearing quartz-pebble conglomerates (unit PHp) proved to be a useful marker. The upper unit is virtually identical to parts of the Maguag Member (King Formation; see Davidson, 1970) exposed elsewhere in the Hurwitz Basin. It is interpreted to represent local deposition on a plain in a wet paleoclimate. The paucity of coarse-grained detritus and restriction of clast types to diabase lithologies are indicators of low topographic relief and a long residence time for detritus. The absence of coarse-grained detritus and the lack of diastrophic zones are consistent with a long residence time for detritus. Xenite likely resulted from metamorphism of aluminous bifluorides found elsewhere in the Maguag Member and considered the product of tropical weathering (Bell, 1970b; Young, 1973).

Lamprophyre dykes (unit Pld)
At one locality between Mackenzie Lake and Victory Lake, an east-trending undeformed and unmetamorphosed lamprophyre dyke cuts penetratively through Hurwitz Group polymictic conglomerates and sandstones. Presumably part of the Churchill Island Formation magmatism, currently estimated at 1850 ± 300 Ma (L.P.F. zircon, Talla et al., 1980; 1825 ± 22 Ma (FAR) zircon, Roddick and Miller 1994) and 1832 ± 39 Ma (Pb/Pb isochron, apatite, MacRae et al., 1996).

Post-tectonic granites (unit Pg)
In the southwestern part of the map area, a non-foliated fluorite/tourmaline-bearing granite pluton (and associated satellite dykes) cuts penetratively through Hurwitz Group, likely one of the post-tectonic granitic intrusions within the western Churchill Province defined by Peterson and van Breemen (1999, ca. 1.83 Ga Hudson suite, ca. 1.76 Ga Nulivut suite).

STRUCTURE

South of Mackenzie Lake, basement and Hurwitz Group define a northeast-trending synclinorium that tapers to the southwest between two basement-cover faults. On the eastern side of the synclinorium, a high-angle embasement fault juxtaposes granitic rocks against the Kaminak Group. Near the "Marce prospect", the granitic rocks are structurally above a relatively thick Kaminak Group section (see cross-section B-B'). To the northeast they are structurally beneath, and the Kaminak Group is cut out against a high-angle basement-cover fault (see axial projection A-A'). The top of this basement-cover fault also charges through the Hurwitz Group, resulting in a complex relationship over outcrop in the south, and cover-over-basement relationships in the north. This fault lies near the Nomat Formation basement contact in the south (see cross-section B-B'), but cuts up-section northeast and eliminates the Nomat Formation near axial projection A-A'. In the interior of the synclinorium, light to isoclinal folds in Hurwitz Group strata display synclinal axial surfaces and penetrative planar foliation and shallowly plunging hinge lines and steeping lineations. In the south, near the "Marce prospect", the synclinorium comprises a narrow southeast-vergent anticline caught between two faults (see cross-section B-B'). To the northeast, the synclinorium broadens markedly and exposes higher stratigraphic levels as a basement-cover fault shifts east up-section and is folded together with the Hurwitz Group. Rocks on the eastern boundary, Mackenzie and Victory lakes are separated from the southern synclinorium by a northeast-trending cross fault adjacent to which relatively low-strain Hurwitz Group rocks are located. Faults at the "Victoria prospect" and "Marce prospect" are highly eroded Archean granites and Kaminak Group rocks are faulted above the Hurwitz Group, and all three map units define a south-vergent overturned package.

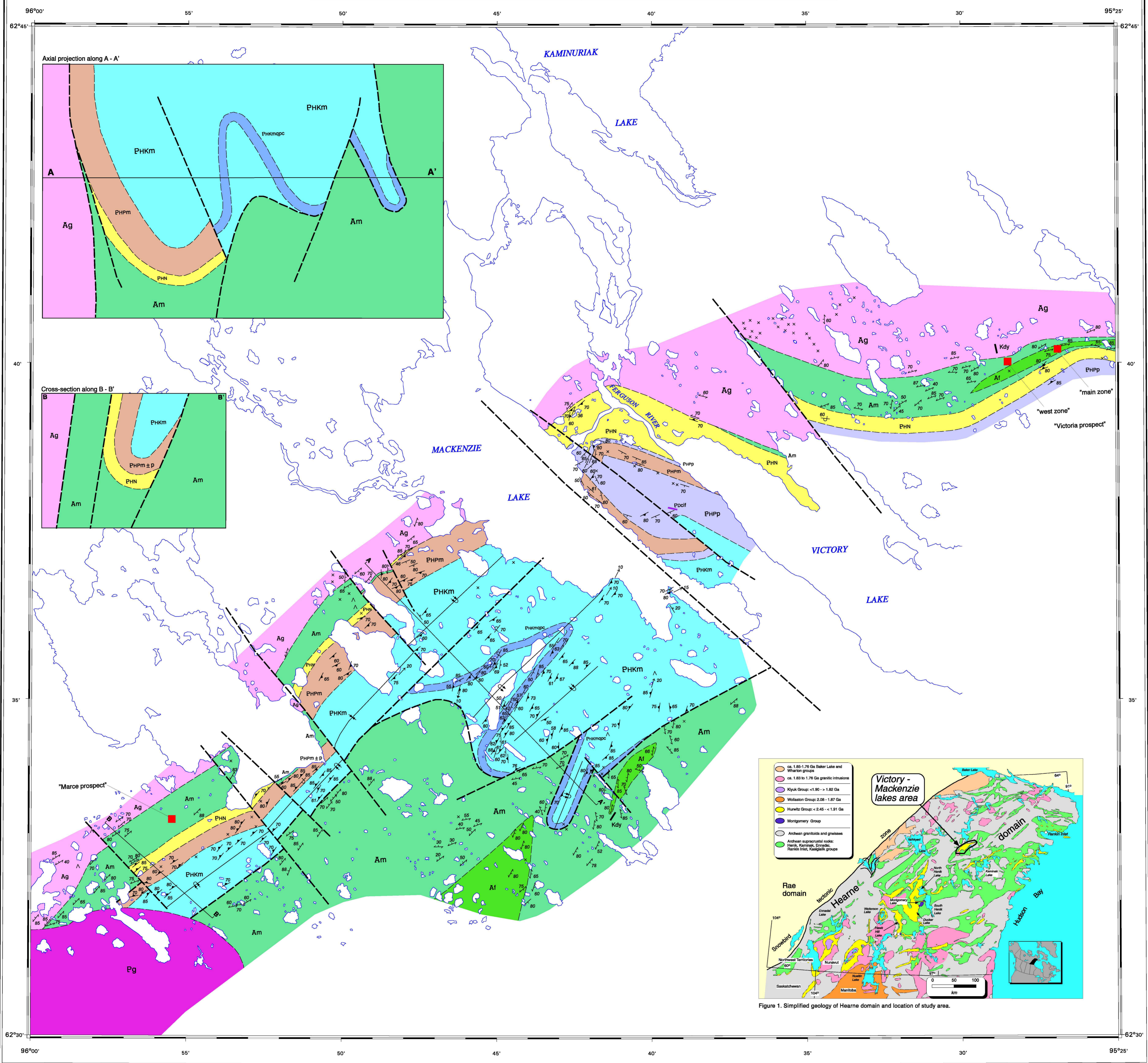
We interpret the post-tectonic Hurwitz Group structures in the Mackenzie-Victory lakes area to represent an extreme example of a style of deformation found elsewhere in the Hearne domain, in which complex structures arise from extension in the core of the map area. Deformation in the Hearne domain is inferred to have occurred pre-1.83 Ga because an undeformed and unmetamorphosed lamprophyre dyke cuts structures in the Hurwitz Group, and post-1.83 Ga because Kaminak dykes contain foliation related to a strike-slip fault. Archean basement rocks beneath the Hurwitz Group generally bear a penetrative fabric. Despite this, basement clasts within voluminous Hurwitz Group conglomerates lack a pre-depositional tectonic fabric. Hence, we conclude that, in contrast to the Kaminak Lake area to the southeast (Hamner et al., 1998) structures in the Mackenzie-Victory lakes area reflect intense Paleoproterozoic deformation, and that the effects of Archean strain are not obvious.

BASE METAL MASSIVE SULPHIDE PROSPECTS

Base metal massive sulphides hosted by Kaminak Group volcanogenic rocks have recently been discovered by Comex Minerals Corp. on the Victory Lake property. Values reported below are from Auriferous sulphide data, Comex Minerals Corp. 1998, 1999. At the "west zone" of the "Victoria prospect", massive and stringer pyrite, galena, chalcopyrite and garnet are hosted by a strike length of ca. 400 metres. They are hosted by silicified and quartz-veined amphibole-garnet schists (up to 25 m thick) thought to be metamorphic equivalents of hydrothermally altered (chloritic and silicified) mafic to intermediate volcanic rocks. Average chip samples from trenching across a 4 metre thick lens of massive mineralization assayed: 6.71% Zn, 0.2% Cu, 0.2% Pb, 24 g/t Ag and 0.04 g/t Au. Diamond drilling passed this zone at an interval of 5.12 metres. Stringer to disseminated chalcopyrite has assayed as high as 0.3% Cu over 7.1 metres. The "main zone" of the "Victoria prospect" is a 2000 metre long horizon containing massive and stringer pyrite, sphalerite and galena in felsic to intermediate volcanogenic strata. Metre-scale zones of strongly pyritic muscovite-quartz ± andalusite ± garnet ± graphite are interpreted to have been derived from metamorphism of tuffaceous beds that previously underwent intense acidic hydrothermal alteration. Samples from a massive sulphide area yielded up to 0.19% Zn, 0.04% Pb, and 240 g/t Ag. These values from stringer mineralization yielded up to 0.82% Zn, 0.26% Pb, and 64 g/t Ag. At the "Marce prospect" southeast of Mackenzie Lake, surface samples from sulphide cores hosted by mafic to intermediate volcanic rocks averaged 0.2%, 1.4% Cu, and 250g/t Ag. Gold assays as high as 7.2 g/t. The geological and mineralogical characteristics of base metal mineralization in the Mackenzie-Victory lakes area are typical of volcanic massive sulphide deposits. The Marce and Victoria prospects may represent a single horizon that was disrupted by Archean plutonism and Paleoproterozoic deformation.

REFERENCES

Aspler, L.B., and Chalmers, J.R.
1996a. Stratigraphy, sedimentology and physical volcanology of the Emeke-Bankin generations belt, Northwest Territories, Canada. Late Archean paleogeography of the Hearne Province and tectonic implications. *Precambrian Research*, v. 77, p. 59-88.
1996b. Relationships between the Monargony Lake and Hurwitz Group, and stratigraphic revision of the lower Hurwitz Group, District of Kaminak, Canadian Journal of Earth Sciences, v. 33, p. 1243-1256.
1997. Initial of ca. 2.45-2.1 Ga intracratonic basin sedimentation of the Hurwitz Group, Kaminak-Hinterland, Northwest Territories, Canada. *Precambrian Research*, v. 81, p. 265-297.
Aspler, L.B., Armitage, A.E., Ryan, J.J., Hausheu, M., Sarmacz, S., and Harvey, B.J.A.
2000. Precambrian Geology, Victory and Mackenzie lakes, Nunavut Territory and adjacent parts of "Mackenzie Lake metamorphism", Paleoproterozoic, Hurwitz Group. In *Current Research 2000-01*, Geological Survey of Canada, 10p.
Bell, N.T.
1968. Preliminary notes on the Proterozoic Hurwitz Group, (Tasari) (ESR) and Kaminak Lake (ESL) areas, District of Kaminak, Geological Survey of Canada, Paper 69-36, 17p.
1970a. Preliminary notes on the Hurwitz Group, Pichee map area, Northwest Territories, Geological Survey of Canada Paper 69-50, 15p.
1970b. The Hurwitz Group - a prototype for deposition on metastable cratons. In *Symposium on Basins and Geosynclines of the Canadian Shield*, A.J. Baer (ed.), Geological Survey of Canada Paper 70-40, p.159-199.
1971. Geology of Honik lakes (east half) and Ferguson Lake (east half) map-areas, District of Kaminak, Geological Survey of Canada Paper 70-61, 31p.
Hamner, S., Rastbali, R. H., Sandeman, K. A. L., Peterson, T. D., and Ryan, J. J.
1998. Field contributions to thematic studies related to the Kaminak greenstone belt, Kwilung Region, Northwest Territories. In *Current Research 1998-0*, Geological Survey of Canada, p. 77-84.
Haasmak, L.M.
1994. 2.45 Ga global mafic magmatism: Earth's oldest superplume? In *Abstracts of the Eighth International Conference on Geochronology, Cosmochronology and Isotope Geology*, M.A. LaSpere, G.G. Dalrymple, S.D. Turner (ed.), United States Geological Survey Circular 103, p. 102.
Davidson, A.
1970. Precambrian geology, Kaminak Lake map-area, District of Kaminak, Geological Survey of Canada Paper 69-91, 27p.
MacRae, N.N., Armitage, A.E., Miller, A.R., Roddick, J.C., Jones, A.L., and Mudry, M.P.
1996. The diamoniferous Athlaku lamprophyre dyke, Gibson Lake area, N.W.T., in LeCheminant, A.N., Richardson, D.G., Gualdoni, R.W., and Richardson, J.A., eds., *Searching for diamonds in Canada*, Geological Survey of Canada Open File 3228, p. 10'-107'.
Peterson, T.D., and van Breemen, G.
1999. Review and progress report of Proterozoic granitoid rocks of the western Churchill Province, Northwest Territories (Nunavut). In *Current Research 1998-0*, Geological Survey of Canada, p. 119-128.
Roddick, J.C., and Miller, A.R.
1994. "A" - 700 Ma age from the REE-enriched Emeke-Bankin ultrapotassic intrusive suite and implications for timing of ultrapotassic magmatism in the central Churchill Province, Northwest Territories. In *Current Research 1994-0*, Geological Survey of Canada, p. 69-74.
Talla, S., Heywood, W.W., and Lovelidge, W.D.
1985. A U-Pb age from a quartz syenite intrusion, Amer Lake, District of Kaminak, a part of the Churchill structural province. Geological Survey of Canada Paper 85-18, p. 307-310.
Young, G.M.
1973. Tuffs and aluminous quartzites as possible time markers for Middle Precambrian (Alphabetic) rocks of North America. In *Nunavut: Stratigraphy and Sedimentation* (G.M. Young (ed)), Geological Association of Canada Special Publication 12, p. 97-127.



LEGEND

Pg	Non-foliated fourite and tourmaline-bearing granite
Pld	Unmetamorphosed, non-foliated lamprophyre dyke
--- Intrusive contact ---	
PHk	HURWITZ GROUP
PHkmpc	KINGA FORMATION: Maguag Member: Grey and white subarkose to quartz arkose. Local block parallel and cross-stratified heavy mineral beds. Framework intact quartz-pebble conglomerates containing clasts of well rounded spherical white quartz, blue and grey chert in sheets 20 cm to 10 m thick
PHp	PADLE FORMATION: Massive cobble-boulder conglomerate; local channels, graded sheets; subarkose with parallel and trough cross stratification. Fine polymictic clast suite (granitic, mafic and felsic volcanic, quartzite, chert); PHm; monomictic conglomerates with different types of granitic clasts; PHm ± p: monomictic conglomerates with polymictic interbeds
PHN	NOMAT FORMATION: Parallel-stratified subarkose to quartz arkose, local quartz-pebble conglomerate, semipelite
--- unconformity ---	
Kdy	Kaminak dykes. Northeast- and northwest-trending, well foliated gabbro dykes; with feldspar megacrysts
--- Intrusive contact ---	
Ag	ARCHEAN
Ag	Well-foliated granodiorite, granite, diorite
--- Intrusive contact ---	
KAMINAK GROUP	
A1	Felsic to intermediate flows; tuffs; local mafic volcanic layers
Am	Mafic flows, locally pillowed; local felsic volcanic layers. Meta-scale tuffaceous sandstones, siltstones; local pebbles

Geological boundary (approximate)	---
Fault (approximate)	---
Axial trace fold (syncline)	---
Axial trace fold (anticline, syncline overturned)	---
Outcrop	---
Front have: probable subjacent outcrop	---
Bedding (inclined, overturned, tops unknown)	---
Bedding from pillow lavas (overturned)	---
Foliation (Hurwitz Group) (inclined)	---
Foliation (Kaminak Group and related plutonic rocks; likely in part Paleoproterozoic) (inclined, vertical)	---
Minor fold (aximuth and plunge)	---
Stretching lineation (aximuth and plunge)	---
Base metal prospect	---
Axial projection line	---
Cross-section line	---

Geology by L.B. Aspler, A.E. Armitage, M. Hausheu, S. Sarmacz and J.J. Ryan
Geological compilation by L.B. Aspler, 2000
Co-ordinated through the auspices of the Western Churchill NATMAP Project
Digital cartography by R.L. Allard, Earth Sciences Sector Information Division (ESS Info)
Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada
Digital base map supplied by the Author modified from data compiled by Geomatics Canada
Magnetic declination 2000, 0° 38' W, increasing 0.4' annually

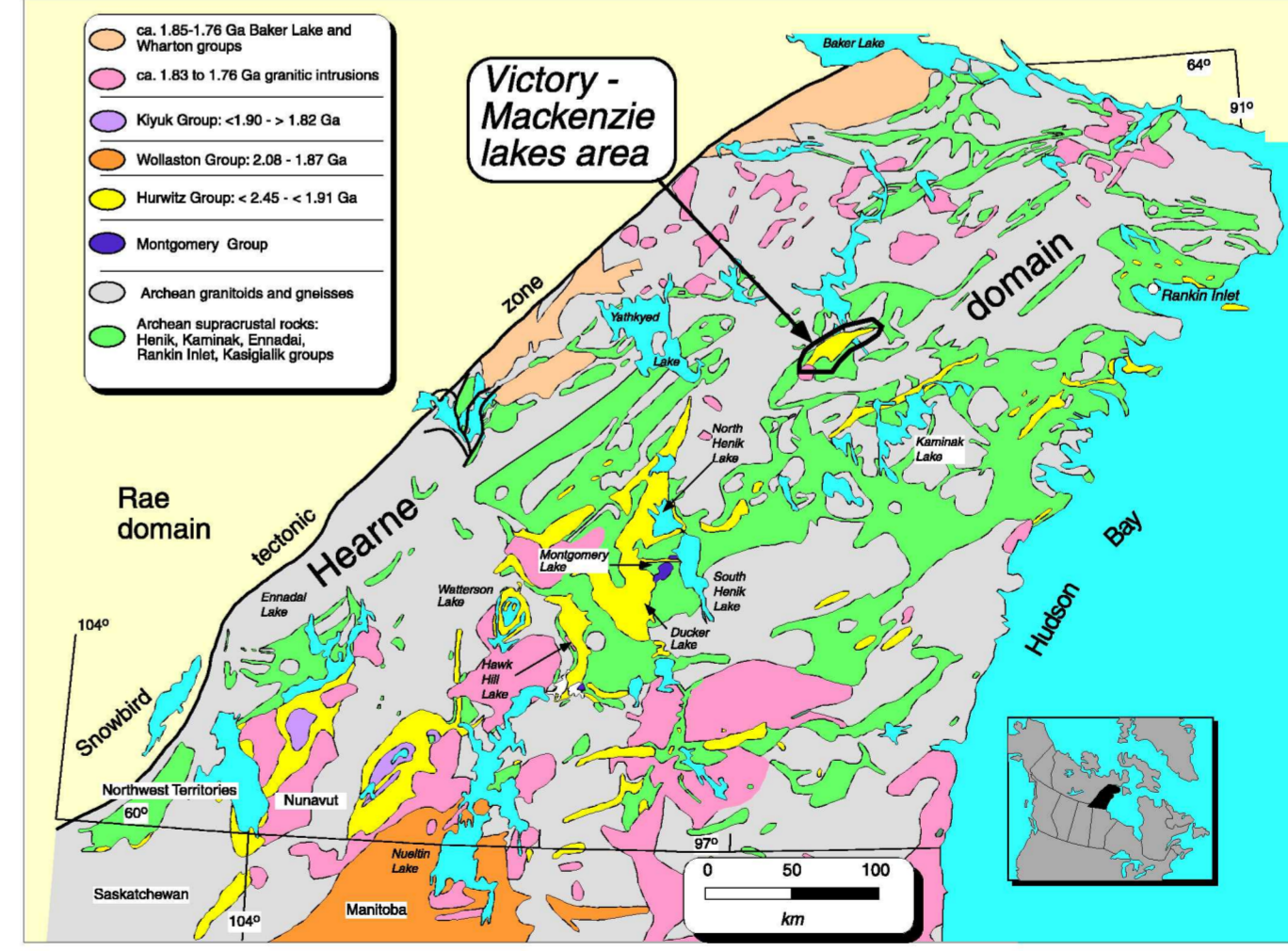


Figure 1. Simplified geology of Hearne domain and location of study area.

EGS 2000-13
OPEN FILE 3752
GEOLOGY
VICTORY AND MACKENZIE LAKES
NUNAVUT
Scale 1:50 000/Echelle 1/50 000
Universal Transverse Mercator Projection / Projection transversale universelle de Mercator
North American Datum 1927 / Système de référence géodésique nord-américain, 1927
© Her Majesty the Queen in Right of Canada, 2000 / © Sa Majesté la Reine du chef du Canada, 2000

66-98	66-99	66-100	66-101
55 L13	55 L12	55 L11	55 L10
66-98	66-99	66-100	66-101

OPEN FILE
DOSSIER PUBLIC
3752
GEOLOGICAL SURVEY OF CANADA
COMMISSION GÉOLOGIQUE DU CANADA
OTTAWA
12/2000
EGS 2000-13

Recommended citation:
Aspler, L.B., Armitage, A.E., Hausheu, M., Sarmacz, S., and Ryan, J.J.
2000. Geology, Victory and Mackenzie Lakes, Nunavut, Geological Survey of Canada, Open File 3752, scale 1:50 000.