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## Introduction

A generalized representation of the EM fields due to an arbitrary source in a layered earth is presented in terms of toroidal and poloidal modes from an arbitrary dipole source situated in a whole space. Section 2 presents a discussion on how the modal representation can be extended to account for stratified media and it is shown how propagator matrices (see, for example, Kennett 1983) can be used to express the modal potentials at any point in a layered earth (Boerner \& West, 1989; note that this paper contains some typographical errors). This result is then used to derive an algorithm useful for computing the EM fields from an arbitrarily oriented point source (electric or magnetic dipole), located at any position in a stratified earth.

## 1 A Modal Description of EM Fields

Consider the standard form of Maxwell's equations and a single Fourier component proportional to $e^{i \omega t}$,

$$
\begin{gather*}
\nabla \cdot \mathbf{B}=0  \tag{1}\\
\nabla \times \mathbf{E}+i \omega \mathbf{B}=0  \tag{2}\\
\nabla \times \mathbf{B}-\mu \alpha \mathbf{E}=\mu \mathbf{J}^{\prime} \tag{3}
\end{gather*}
$$

E is the electric field, $\mathbf{B}$ is the magnetic induction and $\mathbf{J}^{\prime}$ is the applied source current which is assumed to be unaffected by $\mathbf{E}$ and $\mathbf{B}$. The electrical properties of the uniform and isotropic medium are represented by the magnetic permeability $\mu$ and the admittivity $\alpha=\sigma+i \omega \epsilon$, where $\sigma$ is the electrical conductivity and $\epsilon$ is the permittivity. For the development of the layered earth resposne, it is assumed that the spatial variation of these parameters is confined to the $z$ direction. Cylindrical and Cartesian coordinates will be used in the development with a common origin and $z$-axis and with $r, x$ and $y$ in directions parallel to layering.

Our goal is to derive a solution to Maxwell's equations (1)-(3) in terms of two independent modes. The axis of separation for these modes is chosen to be the direction in which the model parameters vary, i.e., the $z$-axis. The "Toroidal Magnetic" or TM mode is characterized by current loops in the $r-z$ planes and a toroidal magnetic field (i.e., a solenoidal field which has no z-component). Since the electric currents associated with this mode cut across the changing medium properties, one can expect that

TM modes are sensitive to the concentration of charge on conductivity gradients. The "Poloidal Magnetic" or PM mode consists of current loops lying perpendicular to the z-axis which generate a poloidal magnetic field (i.e., a solenoidal field whose curl has no z-component). PM mode currents are coupled by induction and are sensitive to layers of high conductivity in the medium, rather than to conductivity gradients.

The modal solution to Maxwell's equations can be obtained by employing a standard theorem of vector analysis to decompose any vector field into a combination of three scalar fields (cf. Morse \& Feshbach 1953, Chapter 13),

$$
\begin{equation*}
\mathbf{F}=\nabla \phi+\nabla \times(\psi \hat{\mathbf{z}})+\nabla \times \nabla \times(\chi \hat{\mathbf{z}}) \tag{4}
\end{equation*}
$$

where $\phi, \psi$, and $\chi$ are eigenfunction solutions of the scalar equations $\nabla^{2} \phi+k^{2} \phi=0$, etc. When this decomposition is applied to the magnetic induction, the fact that B must be solenoidal requires a representation of the form

$$
\begin{equation*}
\mathbf{B}=\nabla \times(\Pi \hat{\mathbf{z}})+\nabla \times \nabla \times(\Psi \hat{\mathbf{z}}) \tag{5}
\end{equation*}
$$

where $\Pi$ represents a scalar potential which generates $T M$ modes and $\Psi$ is the scalar potential associated with the PM modes. The common Hertz potential separation used by Weaver (1970), among others, is obtained by replacing $\Pi$ with $\alpha \Pi$. In fact, $\Pi$ and $\Psi$ are not uniquely determined by this representation until a further constraint is applied (see Backus $1986 \S 1$ ). In our case, an appropriate constraint is applied when the source current is specified by vertical, horizontally irrotational and divergenceless components,

$$
\begin{equation*}
\mathrm{J}^{\prime}=J_{z}^{\prime} \hat{\mathbf{z}}+\nabla_{h} T+\nabla \times(\Upsilon \hat{\mathbf{z}}) \tag{6}
\end{equation*}
$$

where the subscript $h$ refers to the horizontal components, and the functions $T$ and $\Upsilon$ must satisfy the Poisson equations

$$
\begin{gather*}
\nabla_{h}^{2} T=\nabla_{h} \cdot \mathbf{J}_{h}^{\prime}  \tag{7}\\
\nabla_{h}^{2} \Upsilon=-\left(\nabla_{h} \times \mathbf{J}_{h}^{\prime}\right) \cdot \hat{\mathbf{z}} \tag{8}
\end{gather*}
$$

It is important to point out that there are conditions on $T$ and $\Upsilon$ (see Backus 1986 §5.2), but these do not usually introduce problems when dealing with controlled sources. From Maxwell's equations (1)-(3) and the vector representations in (5)-(6), we can obtain differential equations for $\Pi$ and $\Psi$.

$$
\begin{align*}
\nabla_{h}^{2} \Pi+\alpha \partial_{z}\left(\partial_{z} \Pi / \alpha\right)-i \omega \mu \alpha \Pi & =-\mu J_{z}^{\prime}+\mu \alpha \partial_{z}(T / \alpha)  \tag{9}\\
\nabla^{2} \Psi-i \omega \mu \alpha \Psi & =-\mu \Upsilon \tag{10}
\end{align*}
$$

Also, the electric field is

$$
\begin{equation*}
\mathbf{E}=\frac{1}{\alpha} \nabla_{h}\left[\partial_{z} \Pi / \mu-T\right]-\frac{1}{\alpha}\left(\nabla_{h}^{2} \Pi / \mu+J_{z}^{\prime \dot{\prime}}\right) \hat{\mathbf{z}}-i \omega \nabla \times(\Psi \hat{\mathbf{z}}) \tag{11}
\end{equation*}
$$

Equations (9) and (10) illustrate the assertions made in the introduction to this section concerning the current distributions which produce the TM and PM modes. Namely, that divergenceless, horizontally circulating source currents $\nabla \times(\Upsilon \hat{\mathbf{z}})$ produce the PM modes, while $J_{z}^{\prime}$ and $\nabla_{h} T$ generate the TM modes.

The next step is to find solutions for $\Pi$ and $\Psi$ which satisfy the differential equations given above. In particular, we are concerned with point sources which suggests that (9) and (10) may be conveniently solved using the Hankel transform pair,

$$
\begin{align*}
& \tilde{f}(\lambda, z)=\int_{0}^{\infty} r J_{0}(\lambda r) f(r, z) d r  \tag{12}\\
& f(r, z)=\int_{0}^{\infty} \lambda J_{0}(\lambda r) \tilde{f}(\lambda, z) d \lambda
\end{align*}
$$

$J_{0}(\lambda r)$ is the zeroth order Bessel function of the first kind, $r$ is the horizontal separation and $\lambda$ is the horizontal wavenumber. There is a close relationship between (12) and a two dimensional (2D) Fourier transform where $\lambda^{2}=p^{2}+q^{2}$. The 1D Hankel transforms are preferred for problems involving point sources to reduce the numerical effort of transformation to the space domain by exploiting known cylindrical symmetries of the resulting fields.

Applying the Hankel transform to the governing differential equations yields

$$
\begin{equation*}
\alpha \partial_{z}\left(\partial_{z} \tilde{\Pi} / \alpha\right)-u^{2} \tilde{\Pi}=-\mu \tilde{J}_{z}^{\prime}+\mu \alpha \partial_{z}(\tilde{T} / \alpha) \tag{13}
\end{equation*}
$$

and

$$
\begin{equation*}
\partial_{z}\left(\partial_{z} \tilde{\Psi}\right)-u^{2} \tilde{\Psi}=-\mu \tilde{\Upsilon} \tag{14}
\end{equation*}
$$

where $u^{2}=\lambda^{2}+i \omega \mu \alpha$, and a tilde denotes variables expressed in the Hankel transform domain. To solve these differential equations requires a knowledge of the specific source term, and for the purposes of this discussion we shall consider the source to be an arbitrarily oriented electric dipole of moment $J=I\left(d x^{\prime}, d y^{\prime}, d z^{\prime}\right)$ at an arbitrary point in the model. By solving (13) and (14) using the usual variation of parameters method for each source orientation (see also, Chave \& Cox 1982), we find the potential on either side of the source as

$$
\left[\begin{array}{l}
\Pi\left(\left.z\right|_{z \leq z^{\prime}}\right)  \tag{15}\\
\Pi\left(\left.z\right|_{z \geq z^{\prime}}\right)
\end{array}\right]=\frac{\mu I}{4 \pi} \int_{0}^{\infty} \underset{\sim}{\widetilde{C}}\left[\begin{array}{rrr}
d x^{\prime} & d y^{\prime} & d z^{\prime} \\
-d x^{\prime} & -d y^{\prime} & d z^{\prime}
\end{array}\right]\left[\begin{array}{l}
\partial_{x} \\
\partial_{y} \\
\frac{\lambda^{2}}{u}
\end{array}\right] \frac{J_{0}(\lambda \xi)}{\lambda} d \lambda
$$

and

$$
\left[\begin{array}{l}
\Psi\left(\left.z\right|_{z \leq z^{\prime}}\right)  \tag{16}\\
\Psi\left(\left.z\right|_{z \geq z^{\prime}}\right)
\end{array}\right]=\frac{\mu I}{4 \pi} \int_{0}^{\infty} \underset{\sim}{\widetilde{\mathrm{C}}}\left[\begin{array}{lll}
d x^{\prime} & -d y^{\prime} & d z^{\prime} \\
d x^{\prime} & -d y^{\prime} & d z^{\prime}
\end{array}\right]\left[\begin{array}{c}
\partial_{y} \\
\partial_{x} \\
0
\end{array}\right] \frac{J_{0}(\lambda \xi)}{u \lambda} d \lambda,
$$

where $\underset{\sim}{\mathbf{C}}$ is a propagator matrix (see, for example, Gilbert \& Backus 1969, Kennett 1983, and Ursin 1983) describing the continuation of the potentials away from the source level,

$$
\underset{\sim}{\widetilde{\mathbf{C}}}=\left[\begin{array}{cc}
\mathrm{e}^{u\left(z-z^{\prime}\right)} & 0  \tag{17}\\
0 & \mathrm{e}^{-u\left(z-z^{\prime}\right)}
\end{array}\right]
$$

This matrix is based on the fact that the potentials must obey the scalar homogeneous Helmholtz equation everywhere except at the source. Although the concept of propagator matrices is likely to be more familar to seismologists than EM researchers, the technique is useful to keep the expressions for the potentials manageable in the case of complicated layering. This formulation also simplifies programming due to modularity.

The horizontal separation between the receiver location $(r, \phi, z)$ and the source point ( $r^{\prime}, \phi^{\prime}, z^{\prime}$ ) is given by the law of cosines

$$
\begin{equation*}
\xi=\sqrt{r^{2}+r^{\prime 2}-2 r r^{\prime} \cos \left(\phi-\phi^{\prime}\right)} \tag{18}
\end{equation*}
$$

Notice that the EM fields in a 1D earth exhibit translational and rotational invariance, i.e., only the relative separation and orientation between the source and receiver is important and the absolute horizontal reference position is unnecessary. In fact, the only lateral reference point in a layered earth is at infinity where the potentials must approach zero.

Expressions (15) and (16) can be used to find the modal potentials at ( $r, \phi, z$ ) in a uniform whole space from a source at $\left(r^{\prime}, \phi^{\prime}, z^{\prime}\right)$. We next extend the theory to account for layering in the earth model.

## 2 Propagator Matrices for Stratified Media

Finding the potentials in a stratified space requires relatively simple modifications of the theory outlined in section 1 in that only the propagator matrices must he changed. To begin, we adopt a right handed coordinate system with the positive \% axis directed downwards. Imaginary boundaries are inserted in the model at the source and receiver level and layers are numbered in increasing order away from the source level. This potentially confusing notation is resolved by denoting any quantities on the same side of the source as the receiver with a superscript asterisk (*).

The modal potentials in a stratified halfspace are found by combining the potentials for a source in a whole space in (15) (16) with solutions of homogeneous forms of the differential equations (13) and (14) to satisfy the boundary conditions at interfaces. Essentially, we add to the whole space potential from the source upward and downward "secondary" potentials whose amplitudes are determined by the property contrasts at boundaries in the media. The net effect is to introduce inward travelling components to the outward components (15)-(16) and alter the outward travelling amplitudes. A convenient manner of representing the effect is with a reflection ratio, which at any level in the medium is the ratio of the potential which is travelling towards the source (purely reflected potentials) to the potential travelling away from the source (primary and reflected potentials). With this definition and noting that the secondary potentials must propagate unchanged through the source layer, the total potential at the source level can be written in terms of the reflection ratio on either side of the source. The result expressed in the Hankel domain is

$$
\left[\begin{array}{l}
\widetilde{\Pi}_{o u t}^{*}\left(z^{\prime}\right) \\
\widetilde{\Pi}_{i n}^{*}\left(z^{\prime}\right)
\end{array}\right]=\widetilde{\mathcal{R}}^{\Pi}\left[\begin{array}{l}
\widetilde{\Pi}\left(\left.z\right|_{z=z^{\prime}+\varepsilon}\right) \\
\widetilde{\Pi}\left(\left.z\right|_{z=z^{\prime}-\varepsilon}\right)
\end{array}\right] \quad \text { and } \quad\left[\begin{array}{l}
\widetilde{\Psi}_{o u t}^{*}\left(z^{\prime}\right) \\
\widetilde{\Psi}_{i n}^{*}\left(z^{\prime}\right)
\end{array}\right]=\widetilde{\mathcal{R}}^{\Psi}\left[\begin{array}{l}
\widetilde{\Psi}\left(\left.z\right|_{z=z^{\prime}+\varepsilon}\right) \\
\widetilde{\Psi}\left(\left.z\right|_{z=z^{\prime}-\varepsilon}\right)
\end{array}\right]
$$

where $\varepsilon$ is infinitesimally small and

$$
\widetilde{\mathcal{R}}^{\Pi}=\frac{1}{1-\widetilde{R}_{1}^{\Pi} \widetilde{R}_{1}^{\Pi *}}\left[\begin{array}{cc}
1 & \widetilde{R}_{1}^{\Pi}  \tag{19}\\
\widetilde{R}_{1}^{\Pi *} & \widetilde{R}_{1}^{\Pi} \widetilde{R}_{1}^{\Pi *}
\end{array}\right],
$$

and

$$
\widetilde{\mathcal{R}}^{\Psi}=\frac{1}{1-\widetilde{R}_{1}^{\Psi} \widetilde{R}_{1}^{\Psi *}}\left[\begin{array}{cc}
1 & \widetilde{R}_{1}^{\Psi}  \tag{20}\\
\widetilde{R}_{1}^{\Psi *} & \widetilde{R}_{1}^{\Psi} \widetilde{R}_{1}^{\Psi *}
\end{array}\right] .
$$

With no stratification, there are no reflected potentials and thus the reflection ratios are all zero. In this case, (19) and (20) reduce to unity and the whole space potentials are unaltered by this component of the propagation matrices.

Expressions for the reflection ratios can then be found by translating the boundary conditions on the EM fields to boundary conditions on the modal potentials. The relationship between the ratios defined in adjacent layers for the TM mode is given by

$$
\begin{equation*}
\widetilde{R}_{i}^{\Pi}=\left[\frac{\widetilde{X}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}+1\right)+\widetilde{Y}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}-1\right)}{\widetilde{X}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}+1\right)-\widetilde{Y}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}-1\right)}\right] \mathrm{e}^{-2 u_{i} t_{i}} \tag{21}
\end{equation*}
$$

where the $i^{\text {th }}$ reflection ratio is evaluated at the boundary between the $i^{\text {th }}-1$ and $i^{\text {th }}$ layer, $t_{i}$ is the thickness of the $i^{t h}$ layer and

$$
\tilde{X}_{i}^{\Pi}=\frac{u_{i}}{u_{i+1}}, \quad \widetilde{Y}_{i}^{\Pi}=\frac{\alpha_{i}}{\alpha_{i+1}}
$$

The PM mode reflection ratio has an identical form to (21),

$$
\begin{equation*}
\widetilde{R}_{i}^{\Psi}=\left[\frac{\widetilde{X}_{i}^{\Psi}\left(\widetilde{R}_{i+1}^{\Psi}+1\right)+\widetilde{Y}_{i}^{\Psi}\left(\widetilde{R}_{i+1}^{\Psi}-1\right)}{\widetilde{X}_{i}^{\Psi}\left(\widetilde{R}_{i+1}^{\Psi}+1\right)-\widetilde{Y}_{i}^{\Psi}\left(\widetilde{R}_{i+1}^{\Psi}-1\right)}\right] \mathrm{e}^{-2 u_{i} t_{i}} \tag{22}
\end{equation*}
$$

except that

$$
\widetilde{X}_{i}^{\Psi}=\frac{u_{i}}{u_{i+1}}, \quad \widetilde{Y}_{i}^{\Psi}=\frac{\beta_{i}}{\beta_{i+1}}
$$

where $\beta_{i}=i \omega \mu_{i}$ is the impedivity of the $i^{\text {th }}$ layer.
The form of (21) and (22) indicates that reflection ratios can be evaluated recursively. The starting point for the recursion is anywhere beyond the stratification since then there are no subsequent layers to reflect the potential back towards the source and $R_{n}=0$. Thus a calculation is performed on each side of the source, starting at the terminating halfspace and continuing in to the source horizon until $\tilde{R}_{1}$ and $\tilde{R}_{1}^{*}$ (the reflection coefficients on both sides of the source level) are each known.

Once the total potential at the source level is known, the potentials at some other depth are found by using a propagator matrix similar to that described for the whole space model. The form of this matrix for the $i^{t h}$ layer is

$$
{\underset{\sim}{\underset{\sim}{C}}}_{i}=\left[\begin{array}{cc}
\mathrm{e}^{-u_{i} t_{i}} & 0  \tag{23}\\
0 & \mathrm{e}^{+u_{i} t_{i}}
\end{array}\right]
$$

and it describes the attenuation and amplification of the outward and inward travelling potentials, respectively, as well as their respective phase shifts. However, as this matrix only represents a continuation of the potentials through homogeneous material, and it is necessary to enforce the boundary conditions on the potentials at the interfaces with an appropriate boundary condition matrix. These are

$$
\widetilde{\sim}_{i}^{\Pi}=\frac{1}{2}\left[\begin{array}{rr}
\left(\tilde{Y}_{i}^{\Pi}+\tilde{X}_{i}^{\Pi}\right) & \left(\tilde{Y}_{i}^{\Pi}-\widetilde{X}_{i}^{\Pi}\right)  \tag{24}\\
\widetilde{R}_{i+1}^{\Pi}\left(\widetilde{Y}_{i}^{\Pi}+\widetilde{X}_{i}^{\Pi}\right) & \widetilde{R}_{i+1}^{\Pi}\left(\widetilde{Y}_{i}^{\Pi}-\widetilde{X}_{i}^{\Pi}\right)
\end{array}\right],
$$

and

$$
\widetilde{\sim}_{\sim}^{\widetilde{\mathbb{P}}^{\Psi}}=\frac{1}{2}\left[\begin{array}{rr}
\left(\widetilde{Y}_{i}^{\Psi}+\widetilde{X}_{i}^{\Psi}\right) & \left(\widetilde{Y}_{i}^{\Psi}-\widetilde{X}_{i}^{\Psi}\right)  \tag{25}\\
\widetilde{R}_{i+1}^{\Psi}\left(\widetilde{Y}_{i}^{\Psi}+\widetilde{X}_{i}^{\Psi}\right) & \widetilde{R}_{i+1}^{\Psi}\left(\widetilde{Y}_{i}^{\Psi}-\widetilde{X}_{i}^{\Psi}\right)
\end{array}\right] .
$$

when propagating the potentials from the $i^{t h}$ to the $i^{t h}+1$ layer.
The final response matrices for converting the primary source potential at the source level to inward and outward propagating potentials at the receiver level are denoted $\underset{\sim}{\mathbb{T}}$ and $\underset{\sim}{\widetilde{P}}$ for the toroidal and poloidal magnetic modes and have the form
and

$$
\begin{equation*}
\widetilde{\sim} \underset{\sim}{\widetilde{\mathbf{T}}}=\widetilde{\sim}_{R}^{\Psi} \times \widetilde{\sim}_{\widetilde{\mathbf{B}}_{R-1}}^{\Psi} \times \widetilde{\sim}_{R-1}^{\Psi} \times \ldots \times \widetilde{\sim}_{S}^{\widetilde{\mathbf{B}}_{S}^{\Psi}} \times \widetilde{\sim}_{S}^{\widetilde{\mathbf{C}}^{\Psi}} \times \widetilde{\mathcal{R}}^{\Psi}, \tag{27}
\end{equation*}
$$

where the layers containing the source and receiver are designated by the subscripts $S$ and $R$ respectively.

Summarizing, the potentials at any receiver location can be calculated from

$$
\left[\begin{array}{l}
\Pi_{o u t}^{*}(z)  \tag{28}\\
\Pi_{i n}^{*}(z)
\end{array}\right]=\frac{\mu I}{4 \pi} \int_{0}^{\infty} \underset{\sim}{\widetilde{\mathbf{P}}}\left[\begin{array}{rrr}
d x^{\prime} & d y^{\prime} & d z^{\prime} \\
-d x^{\prime} & -d y^{\prime} & d z^{\prime}
\end{array}\right]\left[\begin{array}{c}
\partial_{x} \\
\partial_{y} \\
\frac{\lambda^{2}}{u}
\end{array}\right] \frac{J_{0}(\lambda \xi)}{\lambda} d \lambda
$$

and

$$
\left[\begin{array}{l}
\Psi_{\text {out }}^{*}(z)  \tag{29}\\
\Psi_{\text {in }}^{*}(z)
\end{array}\right]=\frac{\mu I}{4 \pi} \int_{0}^{\infty} \underset{\sim}{\widetilde{T}}\left[\begin{array}{lll}
d x^{\prime} & -d y^{\prime} & d z^{\prime} \\
d x^{\prime} & -d y^{\prime} & d z^{\prime} \cdot
\end{array}\right] \cdot\left[\begin{array}{c}
\partial_{y} \\
\partial_{x} \\
0
\end{array}\right] \frac{J_{0}(\lambda \xi)}{u \lambda} d \lambda .
$$

The sum of the outgoing and ingoing modal potentials can be substituted into the definitions of the EM fields (11) and (5) to obtain expressions for the electric field and magnetic induction from a dipole source. After some algebra and collecting terms according to the the spatial derivatives, a general representation of the EM fields from
an arbitrary electric dipole source is written

$$
\left[\begin{array}{c}
E_{x}  \tag{30}\\
E_{y} \\
E_{z} \\
B_{x} \\
B_{y} \\
B_{z}
\end{array}\right]=\frac{I}{4 \pi} \int_{0}^{\infty} \widetilde{\mathcal{F}}\left[\begin{array}{c}
\lambda^{-2} \partial_{x x} \\
\lambda^{-2} \partial_{x y} \\
\partial_{x} \\
\partial_{y} \\
1
\end{array}\right] \lambda J_{0}(\lambda \xi) d \lambda
$$

where $\widetilde{\mathcal{F}}$ is a $(6 \times 5)$ matrix provided in Table 1 (which is actually for a transversely isotropic medium; see below). Notice that $\underset{\sim}{\mathcal{F}}$ contains all the information about the electrical properties of the layered earth for an arbitrary dipole source excitation and the geometrical variations of the fields are contained in the $(1 \times 5)$ matrix of partial differentials.

A useful alternative representation is to employ Bessel function relationships to define the geometrical operators in terms of sines and cosines instead of differentials. For example,

$$
\left[\begin{array}{c}
\partial_{x x}  \tag{31}\\
\partial_{x y} \\
\partial_{x} \\
\partial_{y} \\
1
\end{array}\right] J_{0}(\lambda \xi) \Longleftrightarrow\left[\begin{array}{cc}
-\frac{\cos 2 \theta+1}{2} \lambda^{2} & \cos 2 \theta \frac{\lambda}{\xi} \\
-\frac{\sin 2 \theta}{2} \lambda^{2} & \sin 2 \theta \frac{\lambda}{\xi} \\
0 & -\lambda \cos \theta \\
0 & -\lambda \sin \theta \\
1 & 0
\end{array}\right]\left[\begin{array}{c}
J_{0}(\lambda \xi) \\
J_{1}(\lambda \xi)
\end{array}\right]
$$

and $\theta$ is measured from the $x$-axis.
The final expression for the fields in terms of the geometrical functions can be
written

$$
\left[\begin{array}{l}
E_{x}  \tag{32}\\
E_{y} \\
E_{z} \\
B_{x} \\
B_{y} \\
B_{z}
\end{array}\right]=\frac{I}{4 \pi} \int_{0}^{\infty} \tilde{\mathcal{\mathcal { F }}}\left[\begin{array}{cc}
-\frac{\cos 2 \theta+1}{2} & \frac{\cos 2 \theta}{\lambda \xi} \\
-\frac{\sin 2 \theta}{2} & \frac{\sin 2 \theta}{\lambda \xi} \\
0 & -\lambda \cos \theta \\
0 & -\lambda \sin \theta \\
1 & 0
\end{array}\right] \lambda\left[\begin{array}{l}
J_{0}(\lambda \xi) \\
J_{1}(\lambda \xi)
\end{array}\right] d \lambda
$$

Table 1 contains the matrix $\underset{\mathcal{F}}{\widetilde{\mathcal{F}}}$ for two types of sources, the electric dipole source developed above, and the magnetic dipole source $\mathbf{M}$, (i.e., a divergenceless source current). Provided the source is infinitesimally small, altering the theory to account for a magnetic dipole source is relatively simple since the source terms only alter the differential equations (13)-(14) and whole space potentials (15)-(16), but not the propagation matrices. The subscripts $S$ and $R$ in Table 1 refer to the properties of the layers containing the source and receiver, respectively, while the terms $\widetilde{P}_{n}$ and $\widetilde{T}_{n}$ refer to linear combinations of the elements of the propagation matrix $\underset{\sim}{\widetilde{P}}$ and $\underset{\sim}{\widetilde{T}}$, viz.,

$$
\begin{align*}
& \widetilde{T}_{1}=\left(\widetilde{T}_{1,1}+\widetilde{T}_{1,2}+\widetilde{T}_{2,1}+\widetilde{T}_{2,2}\right) \\
& \widetilde{T}_{2}=\left(\widetilde{T}_{1,1}+\widetilde{T}_{1,2}-\widetilde{T}_{2,1}-\widetilde{T}_{2,2}\right) \\
& \widetilde{T}_{3}=\left(\widetilde{T}_{1,1}-\widetilde{T}_{1,2}-\widetilde{T}_{2,1}+\widetilde{T}_{2,2}\right) \\
& \widetilde{T}_{4}=\left(\widetilde{T}_{1,1}-\widetilde{T}_{1,2}+\widetilde{T}_{2,1}-\widetilde{T}_{2,2}\right)  \tag{33}\\
& \widetilde{P}_{1}=\left(\widetilde{P}_{1,1}+\widetilde{P}_{1,2}+\widetilde{P}_{2,1}+\widetilde{P}_{2,2}\right) \\
& \widetilde{P}_{2}=\left(\widetilde{P}_{1,1}+\widetilde{P}_{1,2}-\widetilde{P}_{2,1}-\widetilde{P}_{2,2}\right) \\
& \widetilde{P}_{3}=\left(\widetilde{P}_{1,1}-\widetilde{P}_{1,2}-\widetilde{P}_{2,1}+\widetilde{P}_{2,2}\right) \\
& \widetilde{P}_{4}=\left(\widetilde{P}_{1,1}-\widetilde{P}_{1,2}+\widetilde{P}_{2,1}-\widetilde{P}_{2,2}\right) .
\end{align*}
$$

Examining Table 1 indicates that once the source orientation is specified, a maximum of only four of these scalar functions are necessary to represent all the components of the EM field. Furthermore, if the source and receiver are on the same horizontal plane this number reduces to two and there is substantial redundancy in how EM field components sample a stratified earth structure.

Equation (30) is a complete representation of the EM Green's function multiplied by the source distribution and no approximations or assumptions (other than choosing
the specific nature of the dipole source) have been incorporated into the development. Because no quasi-static approximation (e.g., Wait 1982) has been made, this form of the EM fields is equally valid for wave and diffusion problems (assuming a known source current that is unaffected by $\mathbf{E}$ and $\mathbf{B}$ ).

When the source and receiver are both located at the earth/air interface, the propagation matrices reduce to

$$
\underset{\sim}{\widetilde{\mathbf{P}}}=\left[\begin{array}{cc}
1 & 0  \tag{34}\\
\widetilde{R}_{1}^{\Pi} & 0
\end{array}\right]
$$

and

$$
\underset{\sim}{\widetilde{T}}=\left[\begin{array}{cc}
1 & 0  \tag{35}\\
\widetilde{R}_{1}^{\Psi} & 0
\end{array}\right] .
$$

Further simplification can be achieved when the source and receiyer are located at the air/earth contact and the earth is a uniform halfspace. In this case we write

$$
\begin{align*}
& \widetilde{R}_{0}^{\Pi}=\left[\begin{array}{c}
\frac{\lambda}{u}-\frac{\alpha_{0}}{\alpha} \\
\frac{\lambda}{u}+\frac{\alpha_{0}}{\alpha}
\end{array}\right]  \tag{36}\\
& \widetilde{R}_{0}^{\Psi}=\left[\frac{\frac{\lambda}{u}-\frac{\beta_{0}}{\beta}}{\frac{\lambda}{u}+\frac{\beta_{0}}{\beta}}\right] \tag{37}
\end{align*}
$$

since $\widetilde{R}_{2}^{\Pi}=0$ and $\widetilde{R}_{2}^{\Psi}=0$. For this particular case, it is a simple matter to write down the expressions for $\underset{\sim}{\widetilde{\mathbf{P}}}$ and $\underset{\sim}{\widetilde{\mathbf{T}}}$.

$$
\begin{align*}
& \widetilde{P}_{1}=\widetilde{P}_{4}=1+\widetilde{R}_{0}^{\mathrm{II}}=\frac{2 \lambda \alpha}{\lambda \alpha+u \alpha_{0}}  \tag{38}\\
& \widetilde{P}_{2}=\widetilde{P}_{3}=1-\widetilde{R}_{0}^{\Pi}=\frac{2 u \alpha_{0}}{\lambda \alpha+u \alpha_{0}}  \tag{39}\\
& \widetilde{T}_{1}=\widetilde{T}_{4}=1+\widetilde{R}_{0}^{\Psi}=\frac{2 \lambda \beta}{\lambda \beta+u \beta_{0}}  \tag{40}\\
& \widetilde{T}_{2}=\widetilde{T}_{3}=1-\widetilde{R}_{0}^{\Psi}=\frac{2 u \beta_{0}}{\lambda \beta+u \beta_{0}} \tag{41}
\end{align*}
$$

In evaluating closed forms of these expressions, it is useful to recall that $\alpha_{0} \approx 0$.
A useful extension to the theory outlined here is to consider transversely isotropic media (see, for example, Wait 1982). That is, the medium is characterized by an admittivity tensor having the principal axis of anisotropy coincident with the axis of separation
for the modal potentials.

$$
\underset{\sim}{\alpha}=\left[\begin{array}{ccc}
\sigma_{h}+i \omega \epsilon_{h} & 0 & 0  \tag{42}\\
0 & \sigma_{h}+i \omega \epsilon_{h} & 0 \\
0 & 0 & \sigma_{v}+i \omega \epsilon_{v}
\end{array}\right]=\left[\begin{array}{ccc}
\alpha_{h} & 0 & 0 \\
0 & \alpha_{h} & 0 \\
0 & 0 & \alpha_{v}
\end{array}\right]
$$

Thus, the horizontal admittivity may be different than the vertical admittivity.
The development of the modal potentials and EM fields in anisotropic media is more involved than the theory presented here but contains the same essential elements (see the discussion by Nobes 1984). Rather than present a complete theoretical exposition, we simply quote the changes to the theory presented in this paper which are required to represent the EM fields in an anisotropic media.

First, the coefficient of anisotropy is defined to be

$$
\begin{equation*}
K^{2}=\alpha_{h} / \alpha_{v} \tag{43}
\end{equation*}
$$

and it is found that

$$
\begin{equation*}
v=\sqrt{K^{2} \lambda^{2}+i \omega \mu \alpha_{h}} \tag{44}
\end{equation*}
$$

describes the vertical continuation operator for the TM potential in the transversely isotropic media. Therefore the continuation operator for this mode is simply

$$
\widetilde{\mathrm{C}}_{i}^{\Pi}=\left[\begin{array}{cc}
\mathrm{e}^{-v_{i} t_{i}} & 0  \tag{45}\\
0 & \mathrm{e}^{+v_{i} t_{i}}
\end{array}\right]
$$

The boundary conditions at layer interfaces must be modified to account for the anisotropic case and therefore yield a new definition of the reflection ratio and the boundary condition matrix elements $\widetilde{X}_{i}^{\Pi}$ and $\widetilde{Y}_{i}^{\Pi}$.

$$
\begin{equation*}
\widetilde{R}_{i}^{\Pi}=\left[\frac{\widetilde{X}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}+1\right)+\widetilde{Y}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}-1\right)}{\widetilde{X}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}+1\right)-\widetilde{Y}_{i}^{\Pi}\left(\widetilde{R}_{i+1}^{\Pi}-1\right)}\right] \mathrm{e}^{-2 v_{i} t_{i}} \tag{46}
\end{equation*}
$$

where

$$
\widetilde{X}_{i}^{\Pi}=\frac{v_{i}}{v_{i+1}} \frac{K_{i+1}^{2}}{K_{i}^{2}}, \quad \widetilde{Y}_{i}^{\Pi}=\frac{\alpha_{v, i}}{\alpha_{v, i+1}} .
$$

Thus the changes introduced by the anisotropic model'alter the continuation operator and boundary conditions of the TM modal potential. The other change enters in the actual differential equations governing the modal potentials. Notice that all expressions presented in Table 1 reduce to the isotropic case when $\alpha_{v}=\alpha_{h}=\alpha$ so that $K^{2}=1$ and $u=v$. Also, one should expect that fields generated by horizontal circulations of current (PM mode) should be unaffected by transverse isotropy, which is indeed the case.

Table 1 a
where

$$
\begin{array}{rlrl}
A=\left(\frac{v_{R}}{K_{R}} \frac{\widetilde{P}_{3}}{\alpha_{v, S}}-i \omega \mu_{R} \frac{\widetilde{T}_{1}}{u_{S}}\right), & B & =\left(\mu_{R} \frac{\alpha_{v, R}}{\alpha_{v, S}} \widetilde{P}_{4}-\mu_{R} \frac{u_{R}}{u_{S}} \widetilde{T}_{2}\right) \\
C & =\frac{v_{R}}{v_{S}} \frac{K_{S}}{K_{R}} \frac{\widetilde{P}_{2}}{\alpha_{v, S}}, & D & =\mu_{R} \frac{\alpha_{v, R}}{\alpha_{v, S}} \frac{K_{S} \widetilde{P}_{1}}{v_{S}}
\end{array}
$$

and

$$
E=\frac{\widetilde{P}_{4}}{\alpha_{v, S}}, \quad F=\mu_{R} \frac{\widetilde{T}_{1}}{u_{S}}
$$

## TABLE 1b

|  | $\left[-A I_{y, m}^{\prime}\right.$ | $A I_{x, m}^{\prime}$ | 0 | $C I_{z, m}^{\prime}$ | $-\frac{\mu_{R}}{\mu_{S}} \widetilde{T}_{4} I_{y, m}^{\prime}$ |
| :---: | :---: | :---: | :---: | :---: | :---: |
| ${\underset{\mathcal{F}}{\sim}}_{\sim}^{m}=i \omega \mu_{S} d A$ | $-A I_{x, m}^{\prime}$ | $-A I_{y, m}^{\prime}$ | $-C I_{z, m}^{\prime}$ | 0 | $\frac{K_{S}}{K_{R}} \frac{v_{R}}{v_{S}} \widetilde{P}_{2} I_{x, m}^{\prime}$ |
|  | 0 | 0 | $-E I_{y, m}^{\prime}$ | $-E I_{x, m}^{\prime}$ | 0 |
|  | $B I_{x, m}^{\prime}$ | $B I_{y, m}^{\prime}$ | $D I_{z, m}^{\prime}$ | 0 | $\mu_{R} \alpha_{v, R} \frac{K_{S} \widetilde{P}_{1}}{v_{S}} I_{x, m}^{\prime}$ |
|  | $-B I_{y, m}^{\prime}$ | $B I_{x, m}^{\prime}$ | 0 | $D I_{z, m}^{\prime}$ | $-\mu_{R} \frac{u_{R} \widetilde{T}_{3}}{\beta_{S}} I_{y, m}^{\prime}$ |
|  | 0 | 0 | $F I_{x, m}^{\prime}$ | $F I_{y, m}^{\prime}$ | $-\mu_{R} \lambda^{2} \frac{\widetilde{T}_{1}}{u_{S} \beta_{S}} I_{z, m}^{\prime}$ |

where

$$
\begin{array}{cl}
A=\left(\frac{\mu_{R}}{\mu_{S}} \widetilde{T}_{4}-\frac{K_{S}}{K_{R}} \frac{v_{R}}{v_{S}} \widetilde{P}_{2}\right) . & B=\mu_{R}\left(\frac{u_{R} \widetilde{T}_{3}}{\beta_{S}}-\alpha_{v, R} \frac{K_{S} \widetilde{P}_{1}}{v_{S}}\right), \\
C=\frac{\mu_{R}}{\mu_{S}} \frac{\widetilde{T}_{1}}{u_{S}}, & D=\mu_{R} \frac{u_{R}}{u_{S}} \frac{\widetilde{T}_{2}}{\beta_{S}}
\end{array}
$$

and

$$
E=\frac{K_{S} \widetilde{P}_{1}}{v_{S}}, \quad F=\mu_{R} \frac{\widetilde{T}_{4}}{\beta_{S}}
$$

## Special Cases

One useful generalization of the above derivation is to find the electric potential of a pole source (i.e., the DC resistivity case). The potential can be obtained by integrating the electric field along the source and over the receiver dipole. For example, it is easy to show from equation (30) that the electric potential generated by a pole source (i.e. after integrating the source term from 0 to $\infty$ ) is

$$
\begin{equation*}
V_{\text {pole }}=\frac{I d l}{4 \pi} \int_{0}^{\infty}\left(\frac{v_{R}}{K_{R}} \frac{\widetilde{P}_{3}}{\alpha_{v, S}}-i \omega \mu_{R} \frac{\widetilde{T}_{1}}{u_{S}}\right) \frac{J_{0}(\lambda \xi)}{\lambda} d \lambda . \tag{47}
\end{equation*}
$$

For an isotropic earth at zero frequency, this term reduces to

$$
\begin{equation*}
V_{\text {pole }}=\frac{I d l}{4 \pi} \int_{0}^{\infty}\left(\frac{\widetilde{P}_{3}}{\sigma_{v, S}}\right) J_{0}(\lambda \xi) d \lambda \tag{48}
\end{equation*}
$$

Based on the above derivation of $P_{3}$ for a uniform halfspace, direct substitution yields

$$
\begin{equation*}
V_{\text {pole }}=\frac{I d l}{2 \pi \sigma r} . \tag{49}
\end{equation*}
$$

Thus equation (48) can be used to find the electrical response of any electrode array by superposition.

Another relevant generalization is to consider a finite source having arbitrary topology. That is, the source could be an extended bipole that follows a curvilinear path, but which is grounded at both ends. Alternatively, the source could be a polygonal loop of wire that closes on itself, i.e. is not grounded at any point. Such arbitrary sources can always be constructed from integrating the fields of an electric dipole along the path of the wire (since the electric fields are conservative). However, care must be taken in the integration since the terms associated with the double spatial derivatives of the EM kernel functions (see equation (30)) must be evaluated only at the grounding points. To be more specific, we rewrite equation (30) to indicate the integration over the source and specifically isolate the path dependent terms

$$
\left[\begin{array}{c}
E_{x}  \tag{50}\\
E_{y} \\
E_{z} \\
B_{x} \\
B_{y} \\
B_{z}
\end{array}\right]=\frac{I}{4 \pi} \int_{0}^{\infty} \stackrel{\mathcal{F}}{\sim}\left[\begin{array}{c}
\lambda^{-2} \partial_{x x} \\
\lambda^{-2} \partial_{x y} \\
0 \\
0 \\
0
\end{array}\right] \lambda J_{0}(\lambda \xi) d \lambda+\frac{I}{4 \pi} \int_{0}^{\infty} \underset{\sim}{\mathcal{F}}\left[\begin{array}{c}
0 \\
0 \\
\partial_{x} \\
\partial_{y} \\
1
\end{array}\right] \lambda J_{0}(\lambda \xi) d \lambda
$$

The first integral is ignored when the source closes upon itself (i.e., is a loop). If the wire bends between the grounding points, the first contribution of the first integral is only dependent on the location of the grounding points, while the second integral is evaluated over the path taken by the wire.

## 3 Routines

The algorithm outlined above has been coded into a general purpose FORTRAN routine that is able to compute the fields from any orientation of source and receiver, in a completely arbitrary earth. The numerical evaluation of the Hankel transforms is done using the FHT method, or direct integration, and is thus limited to the diffusive realm. No attempt has been made to optimize the code and there are many means of doing this. The primary purpose of this code is simply to illustrate the somewhat abstract theory described above in an algorithmic form.

The following provides listings of the various software routines used in the package.
The main program is emdipole and it calls the various input/output subroutines to define the layered earth model. For each frequency, emdipole computes all six electromagnetic field components using the prescribed source/receiver geometry. At the final step, the computed fields in the space domain are written to output files.

## Listing of emdipole.f

```
c EMDIPOLE drives the subroutine package that calculates the
c electric and magnetic fields in a transversely isotropic layered
c earth.
c x - north \> Right hand coordinate system.
c z - down /
c include 'dipole.inc'
        include mapfre
        parameter(maxfre=250)
c
        real*8 sigmah(maxlay), sigmav(maxlay), epsiln(maxlay)
        real*8 freq1, den, theta, radii(maxrad), freqmx
        real*8 tol, freq(maxfre), tinc
        complex*16 field(maxrad,6,maxfre)
        integer numrad, ndec, i, nfreq, ifre, irad, n1, n2
        integer fht_type, ntheta
        character fc(6)*2, aa*2, ab*2
        data fc/'Ex','Ey','Ez','Hx','Hy','Hz'/
c
        call reader(maxlay,layers,sigmah,sigmav,epsiln,mu,
    > thckns,freq1,ndec,den,maxrad,numrad,radii,ntheta,
    > theta,tinc,js,ms,sz,rz,tol,fht_type)
c
        nfreq = nint(den*ndec)
        freqmx=freq1*(10.0**float(ndec))
        call look(maxlay,layers,sigmah,sigmav,epsiln,mu,
        > thckns,freq1,freqmx,den,maxrad,numrad,radii,
            ntheta,theta,tinc,js,ms,sz,rz)
        call orient(wc,theta)
        call initd
        do i = 1, nfreq
            freq(i) = freq1*10**((i-1)/den)
            if(mod(i-1,den).eq.0) write(6,*) 'Frequency = ',freq(i)
            call dipole(sigmah,sigmav,epsiln,freq(i),numrad,
        > radii(1),theta,tol,radii,field(1,1,i),
        > fht_type)
        end do
c
```

```
    open(unit=11,file='ex.fld',status='unknown')
```

    open(unit=11,file='ex.fld',status='unknown')
        open(unit=12,file='ey.fld',status='unknown')
        open(unit=12,file='ey.fld',status='unknown')
        open(unit=13,file='ez.fld',status='unknown')
        open(unit=13,file='ez.fld',status='unknown')
        open(unit=14,file='hx.fld',status='unknown')
        open(unit=14,file='hx.fld',status='unknown')
        open(unit=15,file='hy.fld',status='unknown')
        open(unit=15,file='hy.fld',status='unknown')
        open(unit=16,file='hz.fld',status='unknown')
        open(unit=16,file='hz.fld',status='unknown')
        open(unit=16,file='hz.fld',status='unknown')
        open(unit=16,file='hz.fld',status='unknown')
        do j = 1, 6
        do j = 1, 6
            write(i+10,'(a)') '@xaxis label "Frequency (Hz)"'
            write(i+10,'(a)') '@xaxis label "Frequency (Hz)"'
            write(i+10,'(a)') '@xaxis ticklabel format power'
            write(i+10,'(a)') '@xaxis ticklabel format power'
        write(i+10,'(a)') '@yaxis label "'//fc(i)//' Amplitude"'
        write(i+10,'(a)') '@yaxis label "'//fc(i)//' Amplitude"'
        end do
        end do
        do irad = 1, numrad
        do irad = 1, numrad
        call itoa(aa,2*irad-2,n1)
        call itoa(aa,2*irad-2,n1)
        call itaa(ab,2*irad-1,n2)
        call itaa(ab,2*irad-1,n2)
        do i = 1, 6
        do i = 1, 6
            write(i+10,'(a)') '@s'//aa(1:n1)//' color 2'
            write(i+10,'(a)') '@s'//aa(1:n1)//' color 2'
            write(i+10',(a)') '@s'//ab(1:n2)//' color 1'
            write(i+10',(a)') '@s'//ab(1:n2)//' color 1'
            write(i+10,'(a1,2x,f15.6)') '# ',radii(irad)
            write(i+10,'(a1,2x,f15.6)') '# ',radii(irad)
            do ifre = 1, nfreq
            do ifre = 1, nfreq
                write(i+10,*) freq(ifre),
                write(i+10,*) freq(ifre),
    > (dreal(field(irad,i,ifre))),
> (dreal(field(irad,i,ifre))),
> (dimag(field(irad,i,ifre)))
> (dimag(field(irad,i,ifre)))
end do
end do
write(i+10,*) '>'
write(i+10,*) '>'
end do
end do
end do
end do
close(unit=11)
close(unit=11)
close(unit=12)
close(unit=12)
close(unit=13)
close(unit=13)
close(unit=14)
close(unit=14)
close(unit=15)
close(unit=15)
close(unit=16)
close(unit=16)
stop
stop
format(13(1pe14.7,1x))
format(13(1pe14.7,1x))
end

```
    end
```

The following file is included in almost all other routines and provides the common parameter definitions and array dimensions.

## Listing of dipole.inc

```
sz = depth of the source
c s = layer containing the source
sdepup = distance to nearest interface above the source
sdepdn = distance to nearest interface below the source
sdist = distance to interface nearest the source, but between
c
crz = depth of the receiver
    r = layer containing the receiver
    rdepup = distance to nearest interface above the receiver
    crepdn = distance to nearest interface below the receiver
    cdist = distance to interface nearest the receiver, but between
    c the source and receiver
    c alphah = horizontal admittivity vector
    c alphav = vertical admittivity vector
    beta = impedivity vector
    c kappa = transverse isotropy vector
    c
        integer maxlay, maxrad
        parameter (maxlay = 100, maxrad = 50)
```

```
complex*16 alphah(maxlay), alphav(maxlay)
complex*16 beta(maxlay), kappa(maxlay)
real*8 mu(maxlay), pi/3.14159265358979/
real*8 thckns(maxlay), rdist, sdist, rz, sz
real*8 sdepup, sdepdn, wc(2,5), js(3), ms(3)
integer r, s, layers
common /properties/ alphah, alphav, beta, kappa, mu
common /structure/ thckns, layers
common /source/ sdist, sdepup, sdepdn, sz, js, ms, s
common /receiver/ rdist, rz, r
common /geometry/ wc
```

dipole computes the EM fields for the input geometry and layered earth model at a single frequency. This routine essentially computes equation (32).

## Listing of dipole.f

```
c---
```

c---
> ri,theta,tol,radii,tfield,fht_type)
> ri,theta,tol,radii,tfield,fht_type)
dipole
dipole
c computes the EM fields due to any electric or magnetic dipole source
c computes the EM fields due to any electric or magnetic dipole source
c within any arbitrarily layered anisotropic earth.
c within any arbitrarily layered anisotropic earth.
c the routine works in the frequency domain using lagged convolution to
c the routine works in the frequency domain using lagged convolution to
c generate the fields at various radii, along a single angle.
c generate the fields at various radii, along a single angle.
c input variables:
c input variables:
sigmah ---- a vector containing the horizontal conductivities of
sigmah ---- a vector containing the horizontal conductivities of
each layer. the air halfspace is a layer.
each layer. the air halfspace is a layer.
sigmav _--- a vector containing the vertical conductivity of
sigmav _--- a vector containing the vertical conductivity of
each layer.
each layer.
epsiln ---- a vector containing the values of epsilon_0 for each
epsiln ---- a vector containing the values of epsilon_0 for each
layer.
layer.
freq ------ frequency (in hz) for which the response is
freq ------ frequency (in hz) for which the response is
calculated.
calculated.
numrad ---- number of radii at which to calculate the fields.
numrad ---- number of radii at which to calculate the fields.
r1-------- initial radius value
r1-------- initial radius value
theta ---- the angle which the receiver makes with the source.
theta ---- the angle which the receiver makes with the source.
tol ----- FHT tolerance
tol ----- FHT tolerance
fht_type -- select FHT method or Pade approximate method
fht_type -- select FHT method or Pade approximate method
common variables:
common variables:
layers ---- the number of layers in the model.
layers ---- the number of layers in the model.
thckns ---- model thickness vector
thckns ---- model thickness vector
sz ------- z coordinate of the transmitter ( z positive down)
sz ------- z coordinate of the transmitter ( z positive down)
rz ------- z coordinate of the receiver (z positive down)
rz ------- z coordinate of the receiver (z positive down)
js(3) ----- electric source vector
js(3) ----- electric source vector
ms(3) ----- magnetic source vector
ms(3) ----- magnetic source vector
output variables:
output variables:
radii ----- a vector of exponentially spaced radius values.
radii ----- a vector of exponentially spaced radius values.
the first element is r1, and the vector is filled
the first element is r1, and the vector is filled
by the routine.
by the routine.
tfield ---- a vector of the field values as a function radius
tfield ---- a vector of the field values as a function radius
external response
external response
include 'dipole.inc'
include 'dipole.inc'
real*8 delta, tol, height
real*8 delta, tol, height
common /density/ delta
common /density/ delta
integer i, j, numrad, nr, fht_type

```
            integer i, j, numrad, nr, fht_type
```

```
    real*8 radii(maxrad), theta, freq, r1
    real*8 sigmah(maxlay), sigmav(maxlay), epsiln(maxlay)
    complex*16 tfield(maxrad,6), rfield(maxrad,6)
    logical newh
c
    nr = numrad
    do i = 1, nr
        radii(i) = ri*exp(delta*(i-1))
    end do
    height = abs(sz-rz)
C
c set up the frequency dependent variables
    call initw(sigmah,sigmav,epsiln,freq)
c Begin the hankel transform routines
    newh = .true.
c J0
    do i = 1, 6
        if(fht_type.eq.1) then
            call fht(newh,nr,radii,height,0,tol,response,i,tfield(1,i))
            else
                    call hantrn(tfield(j,i),0,radii(j),i,tol)
            end do
        end if
    end do
c J1 terms with a 1/r dependance
c
    do i = 1, 6
        if(fht_type.eq.1) then
            call fht(newh,nr,radii,height,1,tol,response,i+6,rfield(1,i))
            else
            do j = 1, numrad
                    call hantrn(rfield(j,i),1,radii(j),i+6,tol)
            end do
        end if
        do j = 1, numrad
            tfield(j,i) = tfield(j,i) + rfield(j,i)/radii(j)
        end do
    end do
c J1 terms without the 1/r dependance
    do i = 1, 6
        if(fht_type.eq.1) then
            call fht(newh,nr,radii,height,1,tol,response,i+12,rfield(1,i))
            else do j = 1, numrad
                    call hantrn(rfield(j,i),1,radii(j),i+12,tol)
            end do
        end if
        do j = 1, numrad
            tfield(j,i) = tfield(j,i) + rfield(j,i)
        end do
    end do
    return
    end
    subroutine hantrn(field,order,rad, n,tol)
    external funct
    integer order, nker, ierr, n, ic, ij, ir
    complex*16 field
    real*8 tol, rad
    logical new/.true./
    character*2 ls(6)/'Ex','Ey','Ez','Hx','Hy','Hz'/
```

```
    character*2 bs(2)/'J0','J1'/, rs(2)/'/r',' '/
    common /fun/ nker
c
c call besautz(field,order,1,7,rad,funct,tol,0.1*tol,1,new,ierr)
    write(6,*) 'BESAUTZ - routine missing !'
    stop
    if(ierr.eq.1) then
        if(n.le.6) then
            ic=n
            ij=1
            else if (n.gt.6.and.n.le.12) then
                ic=n-6
            ir = 1
            elve
                ic}=n-1
                ij =2
            ir =
        write(6,*) 'BESAUT failed to converge; ',ls(ic),
    >
        end if
        return
    complex*16 function funct(lambda)
    external response
    real*8 lambda
    integer iker, nker
    complex*16 kern(-45:301,18), temp, response
c
    common /kernels/ kern, iker
    common /fun/ nker
    iker = 1
    temp = response(lambda)
    funct = kern(iker,nker)
c
    return
    end
```

dipole provides the option of not using the FHT algorithm and computing the Hankel transform by direct integration (Chave 1983). This approach is generally more accurate than the FHT convolution method, but is also much more time consuming. This routine requires the subroutine besautz and associated subroutines available from A.D. Chave. For this release, a besautz stub routine is provided, but it just returns an error code. To exploit this option, obtain besautz, compile and link it with the emdipole code.
reflctn computes the reflection coefficients at the source level, both above and below the source. The subroutine uses recurrsion based on an initial value of $R_{N}=0$ and equations (22) and (46).

Listing of reflctn.f :

```
c-----
    subroutine reflctn(u,v,rtor,rpol,lamsq,rs_pol,rs_tor)
c calculate the reflection coefficents for the source
c calculate the reflection coefficents for the source
    include 'dipole.inc'
    integer i
```

```
    complex*16 u(maxlay),v(maxlay),rtor(maxlay),rpol(maxlay)
    complex*16 gamma2,a,b,expont,temp,rs_pol,rs_tor
    real*8 lamsq,distan
c
c do i = 1, layers
    gamma2 = alphah(i)*beta(i)
            u(i) = casqrt( lamsq + gamma2)
            v(i) = cdsqrt(kappa(i)*lamsq + gamma2)
    end do
c}find the reflection coefficent above the source
    rpol(1) = dcmplx(0.0d0,0.0d0)
    rtor(1) = dcmplx(0.0d0,0.0d0)
c
    do i = 2, s
        if(i.eq.s) then
            distan = -2.0*sdepup
            else
                distan = -2.0*thckns(i)
            end if
c
c
    a = v(i)*alphah(i-1) + v(i-1)*alphah(i)
    b = v(i)*alphah(i-1) - v(i-1)*alphah(i)
    rpol(i) = (rpol(i-1)*a + b)*expont/(rpol(i-1)*b + a)
C
    expont = cdexp(u(i)*distan)
        a = u(i)*beta(i-1) +u(i-1)*beta(i)
            b = u(i)*beta(i-1) - u(i-1)*beta(i)
            rtor(i) = (rtor(i-1)*a + b)*expont/(rtor(i-1)*b +a)
        end do
        rs_pol = rpol(s)
    rs_tor = rtor(s)
C
Find the reflection coefficent below the source.
    rpol(layers) = dcmplx(0.0d0,0.0d0)
    rtor(layers) = dcmplx(0.0d0,0.0d0)
c
    do i = layers-1, s, -1
    if(i.eq.s) then
        distan = -2.0*sdepdn
        else
            distan = -2.0*thckns(i)
    end if
    expont = cdexp(v(i)*distan)
    a = v(i)*alphah(i+1) + v(i+1)*alphah(i)
    b}=v(i)*alphah(i+1) - v(i+1)*alphah(i
    rpol(i) = (rpol(i+1)*a + b)*expont/(rpol(i+1)*b + a)
c
    expont = cdexp(u(i)*distan)
    a = u(i)*beta(i+1) +u(i+1)*beta(i)
    b = u(i)*beta(i+1) - u(i+1)*beta(i)
    rtor(i) = (rtor(i+1)*a + b)*expont/(rtor(i+1)*b + a)
    end do
C
c The ordering of the reflection coefficents is important. rs_pol
c and rs_tor should be the reflection coefficients on the
c side of the source away from the receiver.
    if(sz.gt.rz) then
    temp = rpol(s)
    rpol(s) = rs_pol
    rs_pol = temp
```



$$
\begin{aligned}
& \text { temp }=\text { rtor }(s) \\
& \text { rtor }(s)=\text { rs_tor } \\
& \text { rs_tor }=\text { temp } \\
& \text { end if } \\
& \text { return } \\
& \text { end }
\end{aligned}
$$

proptn computes the propagation matrices that use the reflection coefficients to take the primary potentials at the source level to the receiver level. The relevant equations are (19) and (20), as well as (24) and (25).

## Listing of proptn.f

```
subroutine proptn(u,v,rtor,rpol,rs_pol,rs_tor,t,p)
c generate the propagation matrices fromc the source to the receiver
c
    include 'dipole.inc'
c
        integer i, step
        real*8 distan
        complex*16 proppm(2,2), proptm(2,2), pmprop (2,2), tmprop (2,2)
        complex*16 contpm(2,2), conttm(2,2), fact, rs_pol, rs_tor
        complex*16 u(maxlay), v(maxlay), rpol(maxlay), rtor(maxlay)
        complex*16 p(4),t(4)
c
c calculate the total potentials at the source level.
c - note the extra factor for this matrix added to the summation
    definitions below
    pmprop(1,1) = 1.0d0
    pmprop(1,2) = rs_pol
    pmprop(2,1) = rpol(s)
    pmprop (2,2) = rpol(s)*rs_pol
c
    tmprop}(1,1)=1.0d
    tmprop(1,2) = rs_tor
    tmprop(2,1) = rtor(s)
    tmprop(2,2) = rtor(s)*rs_tor
c continue the source level total potentials to the nearest interface
c
    step = 1
    if(r.lt.s) step = -1
    do i = s, r-step, step
c
            distan = thckns(i)
            if(i.eq.s) distan = sdist
            call cont(contpm,conttm,u(i),v(i),distan)
c
            call matmult(contpm,pmprop,pmprop)
            call matmult(conttm,tmprop,tmprop)
c
c the potential that is propagating away from the source level is then
c continued across the interface, and the potential,propagating towards
c the source is found from the potential propagating away from the source.
c this is done using the reflection coefficents to help eliminate
c round-off errors.
    fact = 0.5 /(alphav(i+step)*kappa(i)*v(i+step))
    proppm(1,1) = fact*(alphah(i)*v(i+step) + alphah(i+step)*v(i))
    proppm(1,2) = fact*(alphah(i)*v(i+step) - alphah(i+step)*v(i))
    proppm(2,1) = proppm(1,1)*rpol(i+step)
```

```
c
    proppm(2,2) = proppm(1,2)*rpol(i+step)
    fact = 0.5 /(u(i+step)*beta(i+step))
    proptm(1,1) = fact*(beta(i)*u(i+step) + beta(i+step)*u(i))
    proptm(1,2) = fact*(beta(i)*u(i+step) - beta(i+step)*u(i))
    proptm(2,1) = proptm(1,1)*rtor(i+step)
    proptm(2,2) = proptm(1,2)*rtor(i+step)
c
c the contribution from this interface is multiplied into the
c propagation matrix.
c
    call matmult(proppm,pmprop,pmprop)
    call matmult(proptm,tmprop,tmprop)
        end do
c
c boundary to the receiver level
c
    call cont(contpm,conttm,u(r),v(r),rdist)
    call matmult(contpm,pmprop,pmprop)
    call matmult(conttm,tmprop,tmprop)
c
    fact = 1.0d0 / ( 1.0d0 - rpol(s)*rs_pol )
    p(1)=(pmprop (1,1)+pmprop (1,2)+(pmprop (2,1)+pmprop (2, 2)))*fact
    p(2)=(pmprop (1,1)+pmprop(1,2)-(pmprop(2,1)+pmprop (2, 2)))*fact
    p(3)=(pmprop (1,1)-pmprop (1,2)-(pmprop (2,1)-pmprop (2,2)))*fact
    p(4)=(pmprop(1,1)-pmprop (1,2)+(pmprop (2,1)-pmprop (2,2)))*fact
c
    fact = 1.0d0 / ( 1.0d0 - rtor(s)*rs_tor )
    t(1)=(tmprop (1,1)+tmprop (1,2)+(tmprop (2,1)+tmprop (2, 2)))*fact
    t(2)=(tmprop (1,1)+tmprop (1,2)-(tmprop (2,1)+tmprop (2, 2)))*fact
    t (3) =(tmprop (1,1)-tmprop (1,2)-(tmprop (2,1)-tmprop (2,2)))*fact
    t(4)=(tmprop (1,1)-tmprop (1,2)+(tmprop (2,1)-tmprop (2,2)))*fact
    c
    return
    end
```

cont is the routine to continue the potentials vertically across a uniform media. The equations used are (23) and (45). As the two non-zero elements of this matrix are inverses of each other, only the term with the negative expoential is calculated and inverted. This saves expensive complex exponential evaluations.

## Listing of cont.f

```
subroutine cont(contpm,conttm,u,v,distan)
c-------m--------------------------------------------------------------------------
c cont generates an continuation matrix for pi and gamma. u and v are
c the attenuation factors for the layer to be continued across, and
c distan is the distance through the layer.
c
    complex*16 contpm(2,2), conttm(2,2), u, v
    real*8 distan
c if(abs(u*distan).gt.80) then
        conttm(1,1) = dcmplx(0.0d0,0.0d0)
        conttm(2,2) = dcmplx (0.0d0,0.0d0)
        else
        conttm(1,1) = cdexp(-u*distan)
        conttm(2,2) = 1.0d0/conttm(1,1)
    end if
C
    conttm(1,2) = dcmplx(0.0d0,0.0d0)
    conttm(2,1) = dcmplx(0.0d0,0.0d0)
```

```
c
```

if(abs(v*distan).gt.80) then
contpm(1,1) $=\operatorname{dcmplx}(0.0 \mathrm{~d} 0,0.0 \mathrm{dO})$
contpm(2,2) $=$ dcmpix(0.0d0,0.0d0)
else
contpm (1,1) $=\operatorname{cdexp}(-v * d i s t a n)$
contpm $(2,2)=1.0 \mathrm{~d} 0 /$ contpm $(1,1)$
end if
contpm $(1,2)=\operatorname{dcmplx}(0.0 \mathrm{~d} 0,0.0 \mathrm{~d} 0)$
contpm $(2,1)=$ dcmplx (0.0d0,0.0d0)
return
end
response clculates the wavenumber domain response of the layered earth to a single wavenumber $\lambda$ of source excitation. This routine is called as a function by the Fast Hankel Transform routine fht.

## Listing of response.f



```
    complex function response(lambda)
c Calculates the response in the Hankel domain of the
c layered earth to a given dipole excitation.
c
    include 'dipole.inc'
c
    integer iker, i, j
    complex*16 u(maxlay), v(maxlay), rs_pol, rs_tor
    complex*16 rpol(maxlay), rtor(maxlay), p(4), t(4)
    complex*16 hfield(6,5), kern(-45:301,18)
    real*8 lambda, lamsq
C
    common /kernels/ kern, iker
    lamsq = lambda*lambda
c compute the reflection coefficients
    call reflctn(u,v,rtor,rpol,lamsq,rs_pol,rs_tor)
c compute the propagation matrices
    call proptn (u,v,rtor,rpol,rs_pol,rs_tor,t,p)
c introduce the electric dipole source normalization
    if(js(1).ne.0.or.js(2).ne.0.or.js(3).ne.0) then
        call jsrch(u,v,lamsq,t,p,hfield)
c}\mathrm{ c introduce the magnetic dipole source normalization
        else if(ms(1).ne.0.or.ms(2).ne.0.or.ms(3).ne.0) then
        call msrch(u,v,lamsq,t,p,hfield)
    end if
c
    do j = 1, 6
        kern(iker,j) = mmplx(0.0,0.0)
        do i = 1, 5
            kern(iker,j) = kern(iker,j) + wc(1,i)*hfield(j,i)*lambda
    end do
    end do
c J1 kernels with an r dependance
```

```
c
            do j = 1, 6
            kern(iker,j+6) = cmplx(0.0,0.0)
            do i = 1, 2
                kern(iker,j+6) = kern(iker,j+6) + wc(2,i)*hfield(j,i)
            end do
        end do
c
    do j = 1, 6
        kern(iker,j+12) = cmplx(0.0,0.0)
        do i = 3,5
            kern(iker,j+12) = kern(iker,j+12) +
    >
                wc(2,i)*hfield(j,i)*lambda**2
            end do
        end do
        return
        end
```

hantrn provides the option of not using the FHT algorithm and computing the Hankel transform by direct integration (Chave 1983). This approach is generally more accurate than the FHT convolution method, but is also much more time consuming. This routine requires the subroutine besautz and associated subroutines available from A.D. Chave. For this release, a besautz stub routine is providedi, but it just returns an error code. To exploit this option, obtain besautz, compile and link it with the emdipole code.
fht is the primary computational routine. It calculates the lagged convolution of the layered earth kernel function (computed in response) with the Fast. Hankel Transform filter coefficients. All kernels are calculated for each wavenumber. but the FHT convergence checking is done sequentially for each kernel to ensure accuracy. The unused kernels are saved for later evaluation.

## Listing of fht.f



```
            subroutine fht(new, nrad,radii,hgt,order, tol, kernel, nker,field)
c-------------------------------------------------------------------
c this subroutine computes the zeroth or first order hankel transform
c for any diffusive kernel function using lagged convolution.
c new........is a logical variable set to .true. each time a new lagged
        convolution is performed (ie for each new frequency).
c nrad.......is the number radii at which the field is to be calculated.
    when using lagged convolution, it is advantageous to evaluate
    the field at many radial points.
radii......is a vector of length nrad containing the radii at
    which the field is to be evaluated. the user must generate
    this array from delta in the common block /density/ using
    the expression;
                            radii(i) = radii(1)*exp(delta*(i-1)) i=1,nrad
    where radii(1) is the minimum radius. delta is the sampling
    density of the filter weights for fht.
hgt........is the vertical separation between the source and receiver
c order......is the order of the hankel transform (0 or 1).
tol........is the desired tolerence of the field caculated. essentially
```

```
c determines the number of significant figures in the final
c answer. tests for convergence are made by comparing the
magnitude of the next term in the sum to tol*(summed total).
if the next term in the sum is smaller than this value, the
    series is assumed to have converged.
kernel.....is a user defined complex function of lambda, ie.
    complex function kernel(lambda). the results of the
    evaluation of the kernel function are stored in the
    common block /kernels/ in the array kern. the variable
        iker points to the storage location for a given wavenumber
        (lambda), and the variable nker points to the hankel
        transform kernel. allows the user to store several kernel
    functions that can be computed simultaneously and then
    Hankel transform each kernel. the array kern must be
    dimensioned to match the number of kernels to be evaluated.
nker.......is the array indicator for the kernel function. ie determines
        which kernel is to be transformed.
field......are the fields variations with radius for each of the kernel
    vectors and is dimensioned fld(nrad)
    external kernel
this include statement inputs the filter weights
    include 'fhtwts.inc'
    integer irad, iker, ifilt, nker, nrad, order
    integer istep, kstart
    complex*16 field(nrad), f, temp, kern(-45:301,18), kernel
    real*8 radii(nrad), tr, ti, tmaxr, tmaxi, tol, coef
    real*8 hgt, rho, error1, error2, lambda(-45:301)
    logical converged, new, stored(-45:301)
c
c user supplied common block /kernels/.
c
    common /kernels/ kern, iker
    common /density/ delta
C c initialize the lambda vector
    if (new) then
        do irad = -nrad, nfilt
            lambda(irad) = lambda1*exp(delta*(irad-1))/radii(1)
        end do
c find the starting position in lambda space which is near the peak of
c the kernel function. it appears when rho = 1, this peak is at 1.0e-03.
    rho = sqrt( radii(1)*radii(1) + hgt*hgt )
    kstart = dlog(1.0e-03/(lambda1*rho))/delta
c
c
    do irad = -nrad, nfilt
        stored(irad) = .false.
        end do
    new =.false.
    end if
C
evaluate the transformation over the radii values..
    do irad = 1, nrad
    field(irad) = cmplx(0.0,0.0)
c chankel transform loop over lambda
c
    converged = .false.
    istep = 1
    iker = kstart
    ifilt = kstart + irad - 1
```

```
                    do while(.not.converged)
if the kernel functions have not been evaluated at the required value of
lambda, then they must be found and stored in the common block kernels.
c
    if(.not.stored(iker)) then
                        temp = kernel(lambda(iker))
        stored(iker) = .true.
    end if
C
evaluate next term in the summation in real and imaginary
parts to improve the computational speed.
c
    if(order.eq.0) then
        coef = dble(wt0(ifilt))
        else if(order.eq.1) then
        coef = dble(wti(ifilt))
        end if
        f = coef*kern(iker,nker)
        field(irad) = field(irad) + f
c
    if(iker.gt.kstart+20.or.istep.lt.0) then
        tr = dabs(dreal(f))
        tmaxr = dabs(dreal(field(irad)))*tol
        ti = dabs(dimag(f))
        tmaxi = dabs(dimag(field(irad)))*tol
c
c
    if(tr.le.tmaxr.and.ti.le.tmaxi) then
    if converged on the rhs, start summation on lhs.
        if(istep.gt.0) then
            iker = kstart
            istep = -1
            else if(istep.lt.0) then
C if converged on both sides, return normalized field value.
c
            field(irad) = field(irad) / radii(irad)
            converged = .true.
                end if
            end if
c
    iker = iker + istep
    ifilt = iker + irad - 1
c
    if((ifilt.gt.nfilt.or.ifilt.lt.1).and..not.converged) then
        error1 = 100.0*dreal(f)/dreal(field(irad))
        error2 = 100.0*dimag(f)/dimag(field(irad))
        if(ifilt.gt.nfilt) write(6,*)'+fht overflow: kernel ',nker
        if(ifilt.lt.1) write(6,*) '-fht overflow: kernel ',nker
        write(6,*) '% error : real ',error1,'; imag ',error2
C
continue the convolution if the lhs has not been done
        if(ifilt.lt.1) then
            field(irad)=field(irad)/radii(irad)
            converged = .true.
            else
            iker = kstart - 1
            ifilt = iker + irad - 1
            istep = -1
        end if
        end if
```

```
    end do
    end do
    return
    end
```

The FHT filter coefficients were calculated by the method of Johansen and Sørensen (1979), as implemented by Christensen (1990).

## Listing of fhtwts.inc




| 130 | + 1.8246721E-06, 2.8919071E-06, | 4.5833553E-06, 7.2641076E-06, |
| :---: | :---: | :---: |
| 131 | + 1.1512781E-05, 1.8246396E-05, | $2.8918252 \mathrm{E}-05,4.5831503 \mathrm{E}-05$, |
| 132 | + 7.2635914E-05, 1.1511488E-04, | $1.8243139 \mathrm{E}-04,2.8910090 \mathrm{E}-04$, |
| 133 | + 4.5810955E-04, $7.2584422 \mathrm{E}-04$, | $1.1498526 \mathrm{E}-03,1.8210664 \mathrm{E}-03$, |
| 134 | + 2.8828366E-03, 4.5606276E-03, | 7.2069676E-03, 1.1369739E-02, |
| 135 | + 1.7887297E-02, $2.8021393 \mathrm{E}-02$, | 4.3588251E-02, 6.7066119E-02, |
| 136 | + 1.0131146E-01, 1.4866030E-01, | $2.0732317 \mathrm{E}-01,2.6530137 \mathrm{E}-01$, |
| 137 | + $2.8524462 \mathrm{E}-01,2.0254407 \mathrm{E}-01$, | $5.7375154 \mathrm{E}-02,-3.7135384 \mathrm{E}-01$, |
| 138 | +-3.1963734E-01, 4.6955122E-01, | -2.7699099E-02, -2.0391968E-01, |
| 139 | + 1.8174329E-01,-1.0118230E-01, | $4.4856879 \mathrm{E}-02,-1.8088601 \mathrm{E}-02$, |
| 140 | + 7.1995987E-03,-2.8661435E-03, | 1.1410288E-03, $4.5425151 \mathrm{E}-04$, |
| 141 | + 1.8084077E-04,-7.1994008E-05, | 2.8661331E-05,-1.1410281E-05, |
| 142 | + 4.5425148E-06,-1.8084077E-06, | 7.1994007E-07, -2.8661331E-07, |
| 143 | + 1.1410281E-07, -4.5425148E-08, | 1.8084077E-08,-7.1994007E-09, |
| 144 | + 2.8661331E-09,-1.1410281E-09, | 4.5425148E-10, -1.8084077E-10, |
| 145 | + 7.1994007E-11, $-2.8661331 \mathrm{E}-11$, | 1.1410281E-11, $4.5425148 \mathrm{E}-12$, |
| 146 | + 1.8084077E-12,-7.1994007E-13, | $2.8661331 E-13,-1.1410281 E-13$, |
| 147 | + 4.5425148E-14,-1.8084077E-14, | 7.1994007E-15, $-2.8661331 \mathrm{E}-15$, |
| 148 | + 1.1410281E-15, $4.5425148 \mathrm{E}-16$, | 1.8084077E-16, -7.1994007E-17, |
| 149 | + 2.8661331E-17, -1.1410281E-17, | $4.5425148 \mathrm{E}-18,-1.8084077 \mathrm{E}-18$, |
| 150 | + 7.1994007E-19,-2.8661331E-19, | 1.1410281E-19, -4.5425148E-20, |
| 151 | + 1.8084077E-20,-7.1994007E-21, | 2.8661331E-21,-1.1410281E-21, |
| 152 | + 4.5425148E-22,-1.8084077E-22, | $7.1994007 \mathrm{E}-23,-2.8661331 \mathrm{E}-23$, |
| 153 | + 1.1410281E-23,-4.5425148E-24, | $1.8084077 \mathrm{E}-24,-7.1994007 \mathrm{E}-25$, |
| 154 | + 2.8661331E-25,-1.1410281E-25, | 4.5425148E-26,-1.8084077E-26, |
| 155 | + 7.1994007E-27, $-2.8661331 \mathrm{E}-27$, | 1.1410281E-27, $4.5425148 \mathrm{E}-28$, |
| 156 | + 1.8084077E-28,-7.1994007E-29, | 2.8661331E-29,-1.1410281E-29, |
| 157 | + 4.5425148E-30,-1.8084077E-30, | 7.1994007E-31,-2.8661331E-31, |
| 158 | + 1.1410281E-31,-4.5425148E-32, | $1.8084077 \mathrm{E}-32,-7.1994007 \mathrm{E}-33$, |
| 159 | + 2.8661331E-33,-1.1410281E-33, | 4.5425148E-34, -1.8084077E-34, |
| 160 | + 7.1994007E-35,-2.8661331E-35, | 1.1410281E-35,-4.5425148E-36, |
| 161 | + 1.8084077E-36/ |  |

initw is an initiation routine, called for each frequency, that computes the admittivity, impedivity and coefficient of anisotropy for each layer of interest.

## Listing of initw.f

```
c----------------------------------------------------------------------------------------
    subroutine initw(sigmah,sigmav,epsiln,freq)
c-------------------------------------------------------------
c sets up the parameters for the common blocks
c
    include 'dipole.inc'
    integer i
    real*8 epsiln(maxlay), sigmah(maxlay), sigmav(maxlay)
    real*8 omega, freq
    omega = freq * 2.0d0 * pi
    do i = 1, layers
        alphah(i) = dcmplx(sigmah(i),omega*epsiln(i))
        alphav(i) = dcmplx(sigmav(i),omega*epsiln(i))
        kappa(i) = alphah(i)/alphav(i)
        beta(i) = dcmplx(0.0d0,omega*mu(i))
        end do
        return
        end
```

initd is called only when the layered earth model changes and computes the interface depths and distances between the source and receiver and the nearest interfaces.

## Listing of initd.f



```
    integer i
    real*8 depth(maxlay)
c find the relative positions of the layers in the earth
c also locate the layers containing the source and receiver
c
    r = 1
    depth(1) = 0.0d0
    do i = 2, layers-1
        depth(i) = depth(i-1) + thckns(i)
            if(rz.le.depth(i).and.rz.gt.depth(i-1)) r = i
            if(sz.le.depth(i).and.sz.gt.depth(i-1)) s = i
    end do
    if(rz.gt.depth(layers-1)) r = layers
    if(sz.gt.depth(layers-1)) s = layers
    write(6,*) 'Source layer = ',s,'; Receiver layer = ',r
c
c find the distance from the source to the nearset interface
c above and below the source level ( required for the reflection
c coefficients
    if(s.ne.1) sdepup = abs(sz - depth(s-1))
    if(s.ne.layers) sdepdn = abs(sz - depth(s))
c
    if(r.ne.s) then
        if(rz.ge.sz) then
            sdist = sdepdn
            rdist = abs(rz - depth(r-1))
            else if(rz.lt.sz) then
            sdist = sdepup
            rdist = abs(rz - depth(r))
            end if
            else if(r.eq.s) then
            sdist = abs(sz - rz)
            rdist = sdist
    end if
    return
    rend
```

orient finds the geometrical orientation factors for an arbitrary dipole at an angle of $\theta$ with respect to the x-axis.

## Listing of orient.f

```
c-------------------------------------------------------------------------------------
    subroutine orient(wc,theta)
c------------------------------------------------------------------------------------------
c This subroutine generates the geometrical orientation factors for the
c arbitrary dipole problem with an angle theta.
c FACTOR is 1/(4 pi) and theta is measured from the axis of the dipole.
c Boerner & West, 1989, Geophysical Journal, 97, 529-547.
C
    real*8 factor, costt, sintt, theta, wc(2,5)
    parameter (factor = 0.07957747154594768)
C
    costt = cosd(2.0d0 * theta )
    sintt = sind( 2.0d0 * theta )
    wc(1,1) = -factor * ( costt + 1.0d0 ) / 2.0d0
    wc(1,2) = -factor * sintt / 2.0d0
    wc(1,3) = 0.0d0
    wc}(1,4)=0.0\textrm{dO
```


jsrch finds the total field kernels due to electric dipole sources as expressed in the Hankel domain. The components calculated are essentially those given in Table 1a.

## Listing of jsrch.f

```
c------------------>\mp@code{subroutine jsrch(u,v,lambda2,t,p,hfield)}
c total fields due to electric sources in Hankel Domain.
    include 'dipole.inc'
c
    complex*16 u(maxlay), v(maxlay), p(4), t(4), hfield(6,5)
    complex*16 ej(5), hj(4), rat
    real*8 lambda2
C
    rat = kappa(s) * v(r) / ( kappa(r) * v(s) )
    ej(1) = v(r) * p(3)/(alphav(s) * kappa(r) )
    ej(2) = beta(r) * t(1) / u(s)
    ej(3) = rat * p(2) / alphav(s)
    ej(4) = lambda2 * v(r) * p(i) / (alphav(s) * kappa(r) )
    ej(5) = p(4) / alphav(s)
c
    hfield(1,1) = (ej(1)-ej(2))*js(1)
    hfield(1,2) = (ej(1)-ej(2))*js(2)
    hfield(1,3) = ej(3)*js(3)
    hfield(1,4) = cmplx(0.0,0.0)
    hfield(1,5) = -ej(2)*js(1)
c
    hfield(2,1) =-(ej(1)-ej(2))*js(2)
    hfield(2,2) = (ej(1)-ej(2))*js(1)
    hfield(2,3) = cmplx(0.0,0.0)
    hfield(2,4) = ej(3)*js(3)
    hfield(2,5) = -ej(1)*js(2)
c
    hfield(3,1) = cmplx(0.0,0.0)
    hfield(3,2) = cmplx(0.0,0.0)
    hfield(3,3) = ej(5)*js(1)
    hfield(3,4) = ej(5)*js(2)
    hfield(3,5) = -ej(4)*js(3)
c Magnetic field components
    rat = alphav(r)/alphav(s)
    hj(1) = u(r) * t(2) / u(s)
    hj(2) = rat * p(4)
    hj(3) = rat * kappa(s) * p(1) / v(s)
    hj(4)= t(1) /u(s)
c
    hfield(4,1) = (hj(2)-hj(1))*js(2)
    hfield(4,2) =-(hj(2)-hj(1))*js(1)
    hfield(4,3) = cmplx(0.0,0.0)
    hfield(4,4) = hj(3)*js(3)
```

```
hfield(4,5) = hj(2)*js(2)
c hfield(5,1) = (hj(2)-hj(1))*js(1)
    hfield(5,2) = (hj(2)-hj(1))*js(2)
    hfield(5,3) = -hj(3)*js(3)
    hfield(5,4) = cmplx(0.0,0.0)
    hfield(5,5) = -hj(1)*js(1)
c
    hfield(6,1) = cmplx(0.0,0.0)
    hfield(6,2) = cmplx(0.0,0.0)
    hfield(6,3) = -hj(4)*js(2)
    hfield(6,4) = hj(4)*js(1)
    hfield(6,5) = cmplx(0.0,0.0)
    return
    end
```

msrch finds the total field kernels due to magnetic dipole sources as expressed in the Hankel domain. The components calculated are essentially those given in Table 1b.

Listing of msrch.f

```
c-------------------------- subroutine msrch(u,v,lambda2,t,p,hfield)
c total fields due to magnetic sources in Hankel Domain.
c include 'dipole.inc'
c complex*16 u(maxlay), v(maxlay), p(4), t(4), hfield(6,5)
    complex*16 rat, ej(4), hj(5)
    real*8 lambda2
c
    rat = kappa(s) / kappa(r)
    ej(1) =-rat * v(r) * p(2) * beta(s) / v(s)
    ej(2) = t(4) * beta(r)
    ej(3)= t(1) * beta(r)/u(s)
    ej(4) = kappa(s) * p(1) * beta(s) / v(s)
c
    hfield(1,1) =-(ej(1)+ej(2))*ms(2)
    hfield(1,2) = (ej(1)+ej(2))*ms(1)
    hfield(1,3) = cmplx(0.0,0.0)
    hfield(1,4) = ej(3)*ms(3)
    hfield(1,5) = -ej(2)*ms(2)
c
    hfield(2,1) =-(ej(1)+ej(2))*ms(1)
    hfield(2,2) =-(ej(1)+ej(2))*ms(2)
    hfield(2,3) = -ej(3)*ms(3)
    hfield(2,4) = cmplx(0.0,0.0)
    hfield(2,5) = -ej(1)*ms(1)
c
    hfield(3,1) = cmplx(0.0,0.0)
    hfield(3,2) = cmplx(0.0,0.0)
    hfield(3,3) = -ej(4)*ms(2)
    hfield(3,4) = ej(4)*ms(1)
    hfield(3,5) = cmplx(0.0,0.0)
c
c Magnetic field components
    hj(1) = u(r)*t(3)
    hj(2) =-alphav(r) * kappa(s) * p(1) * beta(s) / v(s)
    hj(3) = u(r)* t(2)/u(s)
    hj(4) = t(4)
```


look is a simple I/O routine that echos the input information out to the screen to check the veracity of the input data file.

## Listing of look.f

```
c-----------------------------------------------------------
    > thckns,freq1,freqmx,den,maxrad,numrad,radii,
    > ntheta,theta,tinc,js,ms,sz,rz)
c look at the input parameters
c
    character*2 chars(6)
    integer i, ij, layers, numrad, maxlay, maxrad, ntheta
    real*8 freq1,freqmx,sigmah(maxlay),siigmav(maxlay), tinc
    real*8 epsiln(maxlay), js(3), ms(3), mu(maxlay), jms(3)
    real*8 theta, radii(maxrad), den, thckns(maxlay), sz, rz
    data chars/'Jx','Jy','Jz','Mx','My','Mz'/
c
    ij = 1
    jms(1) = js(1)
    jms(2) = js(2)
    jms(3) = js(3)
    if(ms(1).ne.0.or.ms(2).ne.0.or.ms(3).ne.0) then
            ij = 4
            jms(1) = ms(1)
            jms(2) = ms(2)
            jms(3) = ms(3)
            end if
            write(6,1) freqi, freqmx, den, chars(ij), chars(ij+1),
    > chars(ij+2), jms(1), jms(2), jms(3), sz, rz,
    > numrad, radii(1), ntheta, theta, tinc
    format(' frequency range ; ',1pe9.3,' to ',1pe9.3,' Hz',/,
    > ' ; ',Opf5.2,' points per decade',/,
    > , source component ; ',2x,3(5x,a2,5x),/,
    > , ; ',2x,3(1pe12.4),/,
    > , Tx/Rx height (m) ; ',1pe11.3,'/',1pe11.3,/,
    > '# and initial r ; ',i6, ' / ',1pe11.3,/,
```

```
> , # and initial theta; ',i6, , / ',Opf7.2,/,
> , theta increment ; ',Opf7.2,/)
c
2
    format(/, 2x,'#', 3x,'sigma_h', 4x,'sigma_v',
> 3x,'epsilon',7x,'mu',9x,'tau' ,7x,'thckns'/,
> 6x,'(S/m) ', 4x,' (S/m) ',
> 4x,'(F/m)',7x,'(H/m)',7x,' (m) '/)
    i=0
    do i=1,layers
        if(i.ne.layers.and.i.ne.1) then
            if(sigmav(i).eq.sigmah(i)) then
                write(6,5) i,sigmah(i),sigmav(i),epsiln(i),mu(i),thckns(i)
                else
                write(6,7) i,sigmah(i),sigmav(i),epsiln(i),mu(i),thckns(i)
            end if
        else
            if(sigmav(i).eq.sigmah(i)) then
            write(6,3) i,sigmah(i),sigmav(i), epsiln(i),mu(i)
            else
                write(6,8) i,sigmah(i),sigmav(i),epsiln(i),mu(i)
            end if
        end if
        if(i.ne.layers) write(6,4)
    end do
    write(6,6)
    format(1x,i2,5(1x,1pe10.3),1x,' Infinite ')
    format(1x,72('-'))
    format(1x ,i2,6(1x,1pe10.3))
    format (//)
    format('K',i2,5(1x,1pe10.3))
    format('K',i2,4(1x,1pe10.3))
    return
    end
```

reader reads the input data file.

## Listing of reader.f

```
c-------------------------------------------------------------------------------------
    subroutine reader(maxlay,layers,sigmah,sigmav,epsiln,mu,
                                    thckns,freq1,ndec, den,maxrad, numrad, radii,
                                    ntheta,theta,tinc,js,ms,sz,rz,tol,fht_type)
c this file reads the required data from an existing data file
c
    integer i, ndec, layers, numrad, maxlay, maxrad, ntheta
    integer st1, fht_type
    real*8 thckns(maxlay), den, epsiln(maxlay)
    real*8 mu(maxlay), freqn, freq1, sx, sy, idl
    real*8 sigmav(maxlay), sigmah(maxlay)
    real*8 radii(maxrad), theta, sz, rz, eps, js(3), ms(3)
    real*8 saz, sin, ds, eps_rel
    real*8 mu_rel, mufs, tol, rden, rmax, tmax, tinc
    character*80 name, line
c
    parameter( eps = 8.8542090d-12,
    > mufs = 1.25663706143591729d-06 ) ,
c
    name='dipole.dat'
    open(unit=1,file=name,status='old')
    call get_line(line,*10)
    read(line,*) tol, fht_type
    call get_line(line,*10)
```

```
    read(line,*) st1
    call get_line(line,*10)
    read(line,*) sx, sy, sz
    call get_line(line,*10)
    read(line,*) idl, saz, sin, ds
    if (sti.eq.0) then
        js(1) = idl * cosd( saz ) * cosd( sin )
        js(2) = idl * sind( saz ) * cosd( sin )
        js(3) = idl * sind( sin )
        else
        ms(1) = ds * cosd( saz ) * cosd( sin )
        ms(2) = ds * sind(saz ) * cosd( sin )
        ms(3) = ds * sind( sin )
    end if
    call get_line(line,*10)
    read(line,*) freq1, freqn, den
        ndec = nint(dlog10(freqn/freq1))
        call get_line(line,*10)
        read(line,*) radii(1), rmax, rden, rz, theta, tmax, tinc
        numrad = max(int(dlog10(rmax/radii(1))*rden),1)
        ntheta = max(int((tmax - theta)/tinc),1)
c
    call get_line(line,*10)
    read(line,*) layers
    layers = layers + 2
    call get_line(line,*10)
    read(line,*) sigmah(1), sigmav(1), eps_rel, mu_rel
    epsiln(1) = eps * eps_rel
    mu(1) = mufs * mu_rel
    do i = 2, layers-1
        call get_line(line,*10)
        read(line,*) thckns(i), sigmah(i), sigmav(i),
    >
        epsiln(i) = eps * eps_rel
        mu(i) = mufs * mu_rel
        end do
        call get_line(line,*10)
        read(line,*) sigmah(layers), sigmav(layers),
    > eps_rel, mu_rel
        epsiln(layers) = eps * eps_rel
        mu(layers) = mufs * mu_rel
c
C
c
        eps_rel, mu_rel
return
end
```

matmult multiplies two $2 \times 2$ matrices together.
Listing of matmult.f

```
c----------------------------------------------------
    subroutine matmult (a,b,c)
c multiply a and b (2x2 matrices) together and store the
c result in c.
            complex*16 a(2,2), b(2,2), c(2,2), temp(2,2)
c
    temp(1,1) =a(1,1) * b(1,1) +a(1,2) * b (2,1)
    temp(1,2) =a(1,1) * b(1,2) +a(1,2) * b(2,2)
    temp(2,1) =a(2,1) * b(1,1) +a(2,2) * b(2,1)
    temp(2,2) =a(2,1)*b(1,2) +a(2,2)*b(2,2)
```

```
c
    c(1,1) = temp(1,1)
        c(1,2) = temp (1,2)
        c(2,1) = temp (2,1)
        c(2,2) = temp (2,2)
    return
    end
```

itoa converts an integer to an ascii character.
Listing of itoa.f


```
    character char*(*)
    integer num, i
    if(num.le.9) then
        write(char(1:1),'(i1)') num
        i = if 1
    if(num.gt.9.and.num.1t.100) then
        write(char(1:2),'(i2)') num
        i =2
    end if
    return
    end
```

get_line reads a line from the input control file and ignores it if the line begins with a comment character (\#).

## Listing of getline.f

```
c-----------------------------------------------------------------------------------
    subroutine get_line(line,*)
    character*(*) line
100 read(1,'(a)',end=10) line
    if(line(1:1).eq.'#') goto 100
    if(line.eq.' ') goto 100
    return
    return 1
    end
```

The following listing shows an example of the input data file format.
Listing of dipole.dat.example

```
# - all lines begining with # are ignored
# FHT definitions, tolerance = relative error
# (type: 1 = Convolution, 2 = Direct integration )
# tolerance type
    1.0e-08
        1
# 1.0e-08
# 0 = Electric Dipole, 1 = Magnetic Dipole
1
# source location (meters)
src_x src_y src_z
```



## 4 Testing the Software

One of the most reliable checks on the software being compiled properly and producing valid result is to demonstrate reciprocity. This involves computing the fields from a given source and then interchanging the source and receiver and recomputing the fields. Reciprocity says that in a linear medium (i.e., a medium exhibiting no dispersive properties), the source and receiver are exactly reciprocal. A tcl script file has been provided along with the code to check automatically that reciprocity is valid for each source-receiver pair.
tcl , tool command language, is a multi-purpose scripting language available for many computer systems (e.g. UNIX, Windows95, etc. It is freely available in both source and binary distributions from http://sunscript.sun.com/ . The path of the emdipole executable must be specified in the script files and then run recip.tcl to output the real and imaginary parts of the EM fields as well as the error in the reciprocity calculation. Reciprocity is quite a stringent test since it exercises all parts of the propagation matrix calculation, the reflection coefficients and the source matrices. Furthermore, depending upon the parameters of the layered earth model for which it is run, reciprocity also tests the implementation of the transverse anisotropy derivation.

The following is an example of the recip.tcl script file.

## Listing of recip.tcl

[^0]```
# - note that to check reciprocity between electric and magnetic
# fields you need to normalize by i omega mu_R. This script only
# normalizes by i omega mu_freespace.
set PROGRAM ./emdipole
#
    set test parameters -
    # FRQ = test frequency
    # RAD = source/receiver separation
    # ANG = angle between the }x\mathrm{ axis and the receiver
# SDP = source depth
# RDP = receiver depth
#
set FRQ 1.00e-05
set RAD 100.00
set ANG 60.00
set RDP 0.01
proc main {} {
    global FRQ FRH RAD ANG SDP RDP COPY
    puts "\nPercentage difference in real and imaginary parts\
            as determined from reciprocity calculations.\n"
    puts " REAL IMAGINARY Pair %Error Re %Error Im"
    set FRH [expr $FRQ * 5.0]
#
# Compute only non-redundant source-receiver pairs.
    set com { exex exey exez exhx exhy exhz
            eyey eyez eyhx eyhy eyhz
                                    ezez ezhx ezhy ezhz
                                    hxhx hxhy hxhz
                                    hyhy hyhz
                                    hzhz }
    foreach pair $com {
        set src [string range $pair 0 1]
        set fld [string range $pair 2 3]
        set dir [string index $pair 1]
        set typ [string index $pair 0]
        set SIN 0.0
        set SAZ 0.0
        if { $dir == "y" } { set SAZ 90.0 }
        if {$dir == "z" } { set SIN 90.0 }
        if { $typ == "h" } { set DIP 1 }
        run $SDP $RDP $DIP $SAZ $SIN
        read $fld a b
# Reverse the source and receivex parameters
    set dir [string index $pair 3]
    set typ [string index $pair 2]
    set SIN 0.0
    set SAZ 0.0
    set DIP 0
    if { $dir == "y" } { set SAZ 90.0 }
    if { $dir == "z" } { set SIN 90.0 }
    if { $typ == "h" } { set DIP 1 }
    run $RDP $SDP $DIP $SAZ $SIN
    read $src c d
# # print out comparison
    puts [format "(%10.3e %10.3e) (%10.3e %10.3e) - $pair: %.3e %.3e " \
```

```
    $a $c $b $d [error $a $c] [error $b $d]]
}
proc error {a b} {
    if { abs($b) != 0.0} {
        return [expr 100*(abs($a-$b))/abs($b)]
    } else {
    return 0.0
    }
}
proc run { SDP RDP DIP SAZ SIN } {
    global ANG RAD FRQ FRH PROGRAM
#
# write a data file with the appropriate parameters
    set f [ open dipole.dat w ]
        puts $f "1.0e-10 1"
        puts $f "$DIP"
        puts $f "0.0 0.0 $SDP"
        puts $f "1.0 $SAZ $SIN 1.0"
        puts $f "$FRQ $FRH 2 0 0.0"
        puts $f "$RAD $RAD 10 $RDP $ANG $ANG 5"
# specify the layered earth model (layers + two terminating halfspaces)
    puts $f "1"
# -- air
# -- layer parameters (repeat as necessary) (sh, sv, epsilon, mu)
# -- lower halfspace
#
    puts $f " 1.0e-14 1.0e-14 1.0 1.0"
    puts $f "30 1.0e-01 1.0e-05 80.0 1.0"
    puts $f " 1.0e-05 1.0e-01 1.0 1.0"
    close $f
    exec $PROGRAM >& /dev/null
}
proc read { fld real imag } {
    upvar $real r
    upvar $imag i
    global FRQ
    set con 7.89568352087148689e-06
    set f [ open $fld.fld r ]
    while {[ [gets $f line ] >= 0 && $out == "" } {
        regsub D [lindex $line 0] e input
        if { $FRQ == $input } {
            set out $line
        }
    }
    close $f
#
# change the D in the double precision fortran output to an e
#
    regsub D [lindex $out 1] e re
    regsub D [lindex $out 2] e im
#
    If it is a magnetic field, we need to convert .
    if [string match *h* $fld] {
        set r [expr -$con * $im * $FRQ]
        set i [expr $con * $re * $FRQ]
    } else {
        set r $re
        set i $im
```

| 141 | $\}$ |
| :--- | :--- |
| 142 | $\}$ |
| 143 | main |
| 144 | exit |

recip.tcl produces screen output showing the calculated fields and a measure of the percent difference created by exchanging the source and receiver. An example output follows based on model contained in the recip.tcl script shown above.

## Listing of recip.out

Percentage difference in real and imaginary parts as determined from reciprocity calculations.

| PEAT |  |  | \%Error Re |  |
| :---: | :---: | :---: | :---: | :---: |
| (-5.880e-09 -5.88 |  |  |  |  |
| 12 8.663e-12) |  |  |  |  |
| 3.913e-11-3 |  |  |  | 1. |
| $1.527 e-24$ 1.527e-24) | ( $2.837 e-16$ 2.837e-16) |  | $4.991 \mathrm{e}-1$ |  |
| $29 e-24-2.629 e-24)$ | ( $7.121 e-17 \quad 7.121 e-17)$ | 迷 |  |  |
|  | (-4.571e-16-4.571e-16) |  |  |  |
|  | 3-1.065e-13) |  | 0 | 0 |
| $6.778 e-11-6.778 e-11)$ | (-1.047e-16-1.047e-16) |  | 1.623 | $9.550 e^{-12}$ |
| $393 e-24$ 4.393e-24) | $2.564 e-16$ 2.564e-16) |  | 6 |  |
| 24-1.527e-24) | $(-2.837 e-16-2.837 e-16)$ |  | 4. | 9. |
| 89e-24 2.489e-24) | ( $2.639 \mathrm{e}-16 \quad 2.639 \mathrm{e}-16$ ) |  | 8. | $1.849 e^{-12}$ |
| 99e-17-2.589e-17) | ( $-2.368 e-23-2.368 e-23$ ) | - ezez | -1 | 1.689e-1 |
| 1.057e-30-1.057e-30) | $(-4.019 e-25-4.019 e-25)$ | - ezh | $3.051 \mathrm{e}-1$ | 1 |
| $6.102 \mathrm{e}-316$ | $2.321 \mathrm{e}-25$ 2.321e-25) | - ezhy |  |  |
| 00e+00 0.000e+00) | $00 \quad 0.000 \mathrm{e}+00)$ |  | 0. | 0 |
| -3.034e-26-3.034e-26) | ( $-1.754 e-18-1.754 e-18$ ) | - hxh | $0.000 e+00$ | 00 |
| $3.358 e-26$ 3.358e-26) | ( 6.105e-18 6.105e-18) |  | $0.000 e+00$ | 00 |
| -3.125e-26-3.125e-26) | ( $-2.474 e-18-2.474 e-18$ ) |  |  |  |
| 8.432e-27 8.432e-27) | 5.296e-18 5.296e-18) |  |  |  |
| -5.413e-26-5.413e-26) | ( $-4.285 e-18-4.285 e-18$ ) |  | $0.000 \mathrm{e}+00$ |  |
| -2.195e-26-2.195e- | ( 3.543e-18 3.543e-18) | hzhz: | 0. | $0.000 e+00$ |

Reciprocity is a necessary, but not sufficient test of the algorithm. Another useful test is to examine the boundary conditions by placing a receiver just, above and just below every interface in the model to see if the continuity of tangential $\mathbf{E}$ and $\mathbf{H}$, and normal $\mathbf{J}$ and $\mathbf{B}$ is maintained.

The reciprocity and boundary condition tests described above are important in verifying the code is working properly, but they do not guarantee the source fields are correct. To test the calculation more fully, it is important to compare the results with analytical expressions, or results obtained by other authors. Two particularly useful test suites are the analytical expressions for a wholespace (e.g. Ward and Hohmann, 1991), and for a uniform halfspace as provided by Bannister (1966). A tcl script file for computing two of the examples shown in Ward and Hohmann (1991) is also included in this distribution. The script file is called ward.tcl and it contains the reference to the original paper and the figures that the examples should replicate. The output file of ward.tcl is called ward.out.

Listing of ward.tcl

```
#!/apps/local/bin/wish
# Test to match examples from Ward and Hohmann
# "Electromagnetic Theory for Geophysical Applications" Chapter 4, pages 131-311
# Electromagnetic Methods in Applied Geophysics - Theory
# Edited by Misac N. Nabighian, Society of Exploration Geophysicists
# 1991
set PROGRAM ./emdipole
# OUTPUT FILES = ward.out (containing real and imaginary parts)
# - You shouldn't change these constants for the comparison tests
#
set FRH
set
set
set SDP
set RDP
set SIGMA
set MEIHOD 1.
set DIP 1.0
set fld hx
# Comment out the tests you don't want
# fg2 -Hx equitorial wholespace: page 176 - fig 2-2
# fg3 -Hx coaxial wholespace :page 177 - fig 2-3
#set type fg3
set type fg2
if { $type == "fg2" } { set ANG 90.0 }
if { $type == "fg3" } { set ANG 0.0 }
proc main {} {
    global FRQ FRH RAD ANG SDP RDP COPY DIP SAZ SIN outfile fld
    global SIGMA
#
    set outfile [ open ward.out w ]
    run $SIGMA
    read $fld $SIGMA
}
proc run { SIGMA } {
    global ANG RAD FRQ FRH SDP RDP DIP SAZ SIN METHOD PROGRAM
##
# write a data file with the appropriate parameters
    set f [ open dipole.dat w ]
        puts $f "1.0e-10 $METHOD"
        puts $f "$DIP"
        puts $f "0.0 0.0 $SDP"
        puts $f "1.0 $SAZ $SIN 1"
        puts $f "$FRQ $FRH 10 0 0.0"
        puts $f "$RAD $RAD 10 $RDP $ANG $ANG 5"
#
    # specify the layered earth model (two terminating halfspaces)
        puts $f "O"
##
    -- air
    -- lower halfspace
        puts $f " $SIGMA $SIGMA 1.0 1.0 0.0"
        puts $f " $SIGMA $SIGMA 1.0 1.0 0.0"
    close $f
    exec $PRDGRAM >& /dev/null
}
```

```
proc read { fld SIGMA } {
    global DIP RAD outfile con
    set f [ open $fld.fld r ]
    while { [gets $f line] >= 0} {
        if { [string match Q* $line] || \
                [string match # [lindex $line 0]] || \
                [string match > [lindex $line 0]] } {
            puts $outfile $line
        } else {
            regsub D [lindex $line 0] e freq
            regsub D [lindex $line 1] e re
            regsub D [lindex $line 2] e im
            set re [expr abs($re)]
            set im [expr abs($im)]
            puts $outfile [format "%15.8e %15.8e %15.8e" $freq $re $im]
        }
    }
    close $f
}
main
exit
```

Another suite of results involving layered models appears in the paper by Spies and Frischknecht (1991). The tcl script spies.tcl creates output files spies.amp and spies.pha using the program emdipole to compute the model examples presented by Spies and Frischknecht (1991) for comparison purposes.

## Listing of spies.tcl

```
#!/apps/local/bin/wish
##
to match examples from Spies and Fischknecht:
# "Electromagnetic Sounding" Chapter 5, pages 285-425
# Electromagnetic Methods in Applied Geophysics - Applications Part A
# Edited by Misac N. Nabighian, Society of Exploration Geophysicists
# 1991
set PROGRAM ./emdipole
#
# OUTPUT FILES = spies.amp spies.pha (containing amplitude and phase)
#
# - You shouldn't change these constants for the comparison tests
#
set FRQ 0.01
set FRH r 100.00
set RAD 1000.00
set SAZ 100.00
set SDP rr 0.01
set SIGMA
set METHOD 1
# Uncomment the test you want
# hcp -horizontal coplanar :page 302 - fig 2-1 (a)' and (b)
# vcp -vertical coplanar :page 302 - fig 2-2 (a) and (b)
# vca -vertical coaxial :page 302 - fig 2-3 (a) and (b)
# ehz -electric equitorial Hz :page 304 - fig 2-8 (a) and (b)
# ehy -electric equitorial Hy :page 304 - fig 2-9 (a) and (b)
# eex -electric equitorial Ex :page 305 - fig 2-11(a) and (b)
# ihy -electric inline Hy :page 305 - fig 2-10(a) and (b)
# iex -electric inline Ex :page 305 - fig 2-12(a) and (b)
```

```
# ppl -perpendicular loops :page 303-fig 2-4 (a) and (b)
```


# ppl -perpendicular loops :page 303-fig 2-4 (a) and (b)

# -to do ppl calculation, you need to get the direct integration routine

# -to do ppl calculation, you need to get the direct integration routine

        by Alan Chave - see documentation
        by Alan Chave - see documentation
    \#set type ppl
\#set type ppl
set type hcp
set type hcp
\#set type vcp
\#set type vcp
\#set type vca
\#set type vca
\#set type ehz
\#set type ehz
\#set type ehy
\#set type ehy
\#set type ihy
\#set type ihy
\#set type eex
\#set type eex
\#set type iex
\#set type iex
if { \$type == "hcp" } {
if { \$type == "hcp" } {
set ANG
set ANG
set ANG
set ANG
set fld [expr 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]
set fld [expr 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]
}
}
if { \$type == "vcp" } {
if { \$type == "vcp" } {
set ANG 90.0
set ANG 90.0
set ANG
set ANG
set fld [expx 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]
set fld [expx 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]
}
}
if { \$type == "vca" } {
if { \$type == "vca" } {
set ANG
set ANG
set fld
set fld
}
}
if { \$type == "ppl" } {
if { \$type == "ppl" } {
set ANG
set ANG
set DIP 1 1
set DIP 1 1
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD * \$RAD]

# 

# 

# FHT seems to fail in this particular case (probably because

# FHT seems to fail in this particular case (probably because

# of poor convergence on the imaginary part) - use direct integration

# of poor convergence on the imaginary part) - use direct integration

    set METHOD 2
    set METHOD 2
    }
}
if { \$type == "ehz" } {
if { \$type == "ehz" } {
set ANG 90.0
set ANG 90.0
set SIN 0.0
set SIN 0.0
set DIP O
set DIP O
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
}
}
if { \$type == "ehy" } {
if { \$type == "ehy" } {
set ANG
set ANG
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
}
}
if { \$type == "ihy" } {
if { \$type == "ihy" } {
set ANGG 0.0
set ANGG 0.0
lol
lol
set fld hy
set fld hy
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
set con [expr 3.14159 * 4.0 * \$RAD * \$RAD]
}
}
if { \$type == "eex" } {

```
if { $type == "eex" } {
```

```
105
}
if { $type == "iex" } {
    set ANG 0.0
    set SIN 
    set con [expr 3.14159 * $RAD * $RAD * $RAD * $SIGMA]
}
proc main {} {
    global FRQ FRH RAD ANG SDP RDP COPY DIP SAZ SIN amp_file pha_file fld
    global SIGMA
#
    set com { 0 0.3 1.0 3.0 10.0 30.0 100.0 }
    set amp_file [ open spies.amp w ]
    set pha_file [ open spies.pha w ]
    foreach fact $com {
        run $SIGMA $fact
        read $fld $SIGMA
    }
}
proc run { SIGMA fact } {
    global ANG RAD FRQ FRH SDP RDP DIP SAZ SIN METHOD PRDGRAM
#
# write a data file with the appropriate parameters
#
    set SIGMA2 [expr $SIGMA*$fact]
    set f [ open dipole.dat w]
        puts $f "1.0e-10 $METHOD"
        puts $f "$DIP"
        puts $f "0.0 0.0 $SDP"
        puts $f "1.0 $SAZ $SIN 1.0"
        puts $f "$FRQ $FRH 5"
        puts $f "$RAD $RAD 10 $RDP $ANG $ANG 5"
# specify the layered earth model (layers + two terminating halfspaces)
# puts $f "1"
# _-- air
## -- air first layer (repeat as necessary)
# -- lower halfspace
        puts $f " 1.0e-14 1.0e-14 1.0 1.0 0.0"
        puts $f "250 $SIGMA $SIGMA 1.0 1.0 0.0"
        puts $f " $SIGMA2 $SIGMA2 1.0 1.0 0.0"
    close $f
    exec $PROGRAM >& /dev/null
}
proc read { fld SIGMA } {
    global DIP RAD amp_file pha_file con
    set f [ open $fld.fld r ]
    while { [gets $f line ] >= 0 } {
        if { [string match Q* $line] || \
                    [string match # [lindex $line 0]] || \
                    [string match > [lindex $line 0]] } {
            puts $amp_file $line
            puts $pha_file $line
            } else {
```

```
        regsub D [lindex $line 0] e freq
        regsub D [lindex $line 1.] e re
        regsub D [lindex $line 2] e im
        puts $amp_file [format "%10.3e %10.3e"\
            $freq [expr sqrt($re*$re+$im*$im)*$con]]
        puts $pha_file [format "%10.3e %10.3e"\
            $freq [expr atan2($im,$re)*180/3.14159]]
        }
    }
    close $f
}
main
exit
```


## 5 Numerical Considerations

Although the exact expressions for the EM fields in a layered earth from a dipole source can always be written explicitly, there are some numerical limitations in evaluating these expressions. Some care should be used in computing fields with this program. In particular, numerical problems arise when the source or receiver are situated exactly on the same horizontal plane. (with the air/earth boundary presenting the worst case). It is generally prudent to separate the source and receiver by a millimeter or more to help ensure stability of the numerical calculation. This vertical offset is only important at very high wavenumber (very short distances) where it damps the diverging kernel functions. Such damping is important for the FHT which implicitly assumes a band-limited input kernel function. As long as the vertical separation of the source and receiver is quite small relative to the horizontal offset, the vertical offset should not be deleterious. In some cases, the direct integration method of Chave (1983) is able to properly evaluate formally divergent Hankel transform integrals.

There are certainly many other cases where the fields may be calculated correctly, but result in the wrong answer because of numerical instabilities. For most purposes of sources on or below the air/earth interface there should be few problems. It is important to be careful, and to perform simple tests when uncertain abou the calculation. This code has certainly not been optimized for stability, it is only to show the method of calculation. By examining certain scenarios in details, it may be possible to develop more robust codings of the EM field calculation described here.

## 6 Summary

In this manuscript, we have presented both a description of, and means of calculating, the modal representation of Maxwell's equations for the EM fields in a layered earth. The purpose was to illustrate the underlying simplicity of the complicated stratified earth Green's functions and to present a generalized representation of all EM methods for layered earths. The specific problem of EM sounding was generalized to give the representation arbitrary source/receiver configurations. Two scalar kernel functions
are sufficient to describe the electric kernels of the halfspace sounding problem, a notion consistent with the theoretical development in terms of toroidal and poloidal magnetic modes. A further advantage of the factorization is that it isolates the geometrical portion of the electromagnetic response, facilitating studies of the source-receiver geometry of the observed fields.

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[^0]:    1 \#!/apps/local/bin/wish
    2
    3
    \# Test Reciprocity

