

LITHOLOGICAL MAP OF THE HIGH-SULFIDATION HOPE BROOK GOLD DEPOSIT, NEWFOUNDLAND^{1-2,3}

Dubé, B., Lauzière, K., Boisvert, E., Geological Survey of Canada, GSC-Québec, 2535 Laurier Boulevard, P.O. 7500, Sainte-Foy, Québec, G1V 4C7

LEGEND

EARLY DEVONIAN

- CH Chetwynd Granite (390 ± 3 Ma)
- HO Hornfels (metamorphic contact zone)

SILURIAN

- LA POILE INTRUSIVE SUITE
- HN Hawks Nest Pond Porphyry (418 ± 2/1.5 Ma) fine-grained biotite-bearing plagioclase-quartz-potassic feldspar porphyritic intrusion
- LA POILE GROUP (429 ± 2 to 422 ± 2 Ma)
- CC Cinq Cerf Brook Conglomerate
- CCb Polymictic boulder-pebble conglomerate
- CCv Volcaniclastic tuffaceous sandstone and tuff
- GB Mylonitized Grand Bruit Gull Pond Tuff (429 ± 2 Ma)
- GBqp Quartz pebble conglomerate
- GBs Sandstone
- GBv Felsic volcanic tuff, chert and volcaniclastic rocks intruded by local quartz-feldspar porphyry dikes
- GBpc Polymictic conglomerate

LATE PROTEROZOIC AND YOUNGER

- MD Undistinguished fine- to medium-grained mafic to intermediate dikes and sills. Commonly highly magnetic. Unstrained to highly strained. Intermixed with local Whittle Hill Sandstone units.

LATE PROTEROZOIC SEQUENCE

- DS Sill-dike complex
- DSap Sill-dike complex with up to 70-80% felsic to intermediate blue quartz-bearing quartz-feldspar porphyritic rocks (OFP) (based on the interpretation that most of the protolith of the advanced argillic alteration zone is OFP), mafic intrusions (Late Proterozoic and younger) and includes some volcaniclastic rocks
- DSqm Sill-dike complex with similar proportion of felsic to intermediate blue quartz-bearing quartz-feldspar porphyry and mafic intrusions (Late Proterozoic and younger) and includes some volcaniclastic rocks
- DSm Sill-dike complex mainly composed of mafic to intermediate intrusions of various ages (Late Proterozoic and younger) commonly feldspar-phryic and magnetite-bearing, with some quartz-feldspar porphyry and some volcaniclastic rocks. Mainly based on drill holes intersections.
- DSx Undivided zone of felsic to intermediate blue quartz-bearing quartz-feldspar porphyry and mafic intrusions. The latter could be younger than Late Proterozoic. Mainly compiled from Stewart (1992).

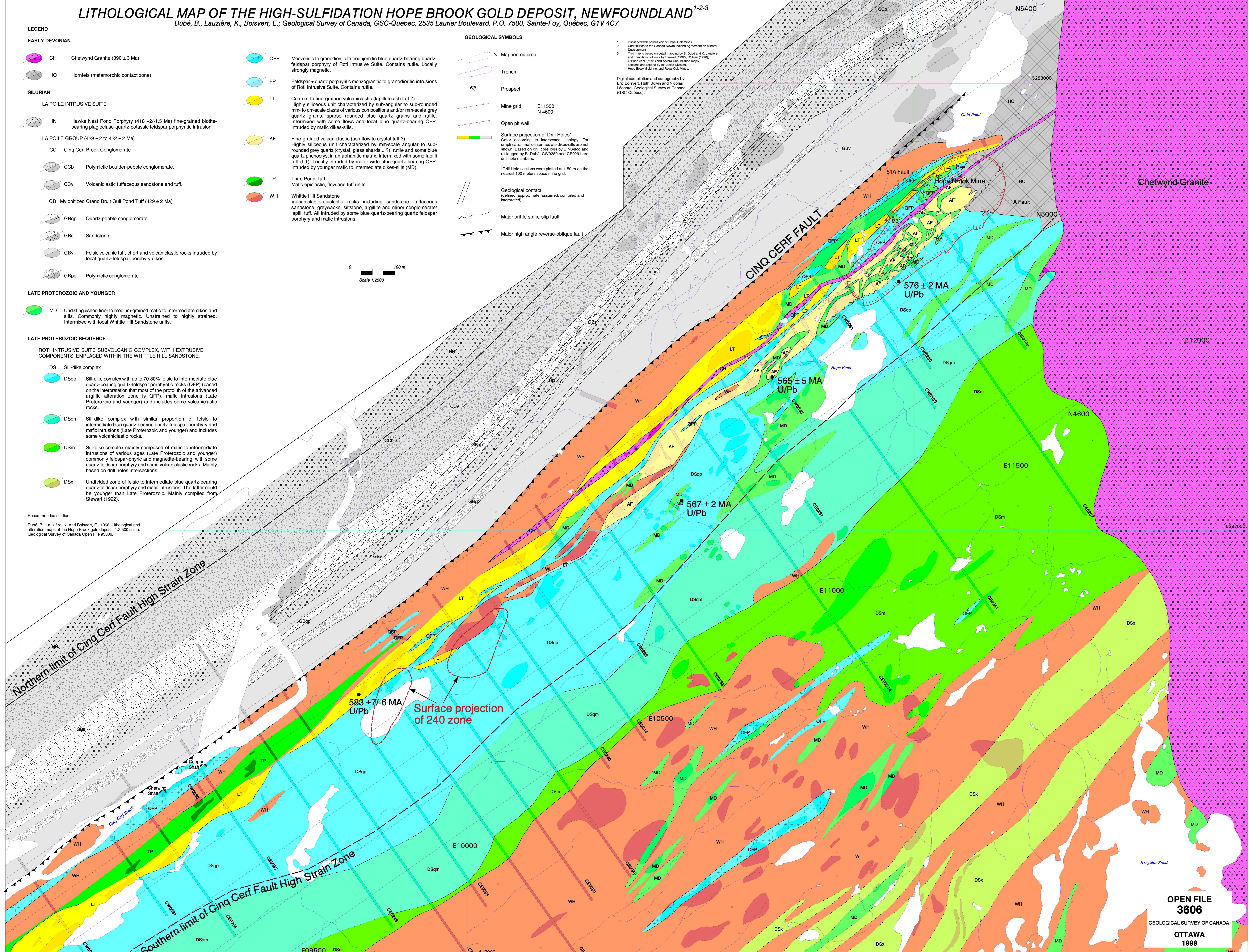
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GEOLOGICAL SYMBOLS

- × Mapped outcrop
- Trench
- Prospect
- Mine grid E11500 N4600
- Open pit wall
- Surface projection of Drill Holes*
Color according to intersected lithology. For simplification mafic-intermediate dikes-sills are not shown. Based on drill core logs by BP-Selco and relogged by B. Dubé, CW6050 and CC501 are drill hole numbers.
- *Drill Hole sections were plotted at a 50 m on the nearest 100 meters space mine grid.
- Geological contact (defined, approximate, assumed, compiled and interpreted)
- Major brittle strike-slip fault
- Major high angle reverse-oblique fault

- 1 Published with permission of Royal Oak Mines
- 2 Contribution to the Canada-Newfoundland Agreement on Mineral Development
- 3 The map is based on aerial mapping by B. Dubé and K. Lauzière and compilation of work by Stewart (1992), O'Brien (1990), O'Brien et al. (1991) and several unpublished maps, sections and reports by BP-Selco Division, Hope Brook Gold Inc. and Royal Oak Mines.

Digital compilation and cartography by
Eric Boisvert, Ruth Bowen and Nicolas Léonard, Geological Survey of Canada (GSC-Québec).



ABSTRACT

The Hope Brook gold mine is a high-sulfidation epithermal gold deposit hosted by Late Proterozoic rocks of the Northern Appalachian Avalon Zone, in southwestern Newfoundland. The deposit (451 Au) is enclosed within a zone of hydrothermal alteration > 3 km long and up to 400 m wide.

REGIONAL GEOLOGICAL SETTING
The area around the deposit comprises two main lithological sequences, from the southeast to the northwest: 1) a greenschist grade Late Proterozoic sequence and 2) the La Poile Group, a Silurian (429-422 Ma) greenschist facies cover sequence of subvolcanic and associated volcanic rocks, unconformably overlying the Late Proterozoic cover sequence and now separated from it by the Cinq Cerf Fault Zone (O'Brien et al., 1991). The Late Proterozoic cover sequence consists of the Whittle Hill Sandstone and overlying Third Pond Tuff (O'Brien et al., 1991). Both units are related by the Roti Intrusive Suite (Dunning and O'Brien, 1986) a multi-phase calc-alkaline suite distinguished by the presence of blue quartz veins. The different phases of the Roti Intrusive Suite include granodiorite (576 ± 2 Ma), the Betty's Pond Tonalite (565 ± 4 Ma) and quartz-feldspar porphyry intrusions (O'Brien et al., 1991; Stewart, 1992; Dubé et al., 1998).

The Whittle Hill Sandstone and the overlying Third Pond Tuff constitute a conformable, regionally north-facing succession that includes, in its basal parts, polymictic conglomerates, grading upward into quartz-rich sandstone and its upper parts into local clay siltstones with rare felsic volcanic layers (O'Brien et al., 1993). The Third Pond Tuff is dominated by water-lain mafic and felsic tuffs and agglomerates with local volcanogenic sedimentary rocks formed late in the tectonic evolution of the basin (O'Brien et al., 1993). The nature and distribution of these Late Proterozoic units indicate that mafic and felsic volcanism took place during sedimentation (O'Brien, 1989).

The Chetwynd Granite is a post-tectonic, high-level biotite granite dated at 390 ± 3 Ma (O'Brien et al., 1991). It cuts across all ductile fabrics and has a well developed thermal aureole (McKenzie, 1986; Yule et al., 1990; O'Brien et al., 1991; Stewart, 1992).

The Silurian evolution of the area comprises two main stages: 1) Mid Silurian opening of the La Poile basin and deposition of the La Poile Group overlying the Late Proterozoic cover sequence, and 2) the Late Silurian closing of the basin and cover deformation of the composite Late Proterozoic - Early Devonian sequence (O'Brien et al., 1991). The Cinq Cerf Fault Zone, which in most areas juxtaposes the Whittle Hill-Third Pond succession against the La Poile Group cover sequence, represents the most important Silurian structure. The fault dips steeply to the SE and is interpreted as a Silurian strike-slip zone (O'Brien, 1989). The Cinq Cerf fault zone is also thought to be regionally transgressive relative to the La Poile Group stratigraphy (O'Brien, 1989), although it is interpreted to coincide overall with the unconformity originally located at the base of the La Poile Group.

GEOLOGY OF THE DEPOSIT
The Hope Brook deposit occurs within the Whittle Hill Sandstone-Third Pond Tuff succession, in the hanging wall of the Cinq Cerf Fault Zone (Yule et al., 1990; Stewart, 1992; Dubé et al., 1998). The deposit occurs in a zone of highly strained, metamorphosed mafic and felsic volcanic rocks in which it is difficult to recognize the original host lithologies. Nevertheless, several descriptive lithological units or domains can be mapped in the deposit area. From NW to SE, the mine sequence is composed of Whittle Hill Sandstone which includes sandstone, tuffaceous sandstone, greywacke, siltstone, argillite and minor conglomerate/lapilli tuff. Further to the northwest, the Whittle Hill Sandstone is in tectonic contact with highly strained felsic volcanic and volcaniclastic rocks of the La Poile Group across the Cinq Cerf Fault Zone. Locally, mafic volcaniclastic rocks of the Third Pond Tuff and blue quartz-bearing quartz-feldspar porphyries (OFP) of the Roti Intrusive Suite are intercalated within the Whittle Hill Sandstone. Other OFP dikes with rutile and blue quartz that are strongly magnetic are located immediately to the northwest of the deposit.

Farther SE, there is a sequence of coarse to fine-grained felsic volcaniclastic rocks (LT), up to 70 m thick and more than 3 km long. This sequence is characterized by cm- to mm-scale, sub-angular to sub-rounded clasts of various compositions (porphyry, sedimentary rocks and reworked, sub-angular quartz as well as minor blue quartz and rutile in a fine grained matrix). These volcaniclastic rocks are interpreted as comprising mainly lapilli to ash tuff and are now highly siliceous due to intense hydrothermal alteration. They are most highly siliceous in the zone containing similar unaltered tuff (Third Pond Tuff) present at the same stratigraphic position within the highly altered mafic zone to the southwest (Stewart et al., 1993). This volcaniclastic unit is in strike with a highly altered mafic zone of coarse volcaniclastic rock, occurring immediately to the north of the ore zone, which is variously interpreted as a debris flow, a volcano-hydrothermal breccia or mixed with possible conglomerate. It is difficult to definitively establish the primary nature of the lithologies due to the intense hydrothermal alteration and deformation. All these rocks are locally intruded by blue-quartz-bearing OFP's.

The host rocks to the ore zone are interpreted to consist mainly of fine-grained felsic volcaniclastic rocks (ash flow to pyroclastic tuff) and/or very fine grained intrusions, intermixed with some coarse-grained volcaniclastic rocks (lapilli tuff). All these units are now highly siliceous due to the intense younger hydrothermal alteration. The SW into a homogeneous fine- to medium grained unit characterized by quartz, white mica, albite and pyrite, which is interpreted as a strongly altered Whittle Hill Sandstone horizon preserved mainly to the SW of the deposit.

The structural hanging wall to the ore zone consists of a sill-dike complex more than 400m thick. It corresponds to the dike swarm described by Stewart (1992). This complex can be subdivided into three sub-zones, from northwest to southeast: 1) a zone (DSap) up to 200 m wide dominated by blue quartz-bearing OFP's and minor feldspar-quartz porphyries (QFP) of iron/yttrium-bearing to monzonitic to granodiorite compositions with a subordinate proportion of mafic intrusions and volcaniclastic units; 2) an approximately 150m wide zone (DSqm) containing similar proportions of blue quartz-bearing OFP's and intermixed mafic intrusions with some volcaniclastic rocks; 3) a 40-50 m wide zone (DSm) composed mainly of mafic to intermediate intrusions of various ages, commonly feldspar-phryic and magnetic, with some OFP's/QFP and volcaniclastic rocks. Immediately southeast of the deposit, this zone is more than 200 m wide and represents an almost massive mafic-intermediate intrusive complex. Southeast of the sill-dike complex, the area is underlain by Whittle Hill Sandstone epithermal rocks intruded by mafic and some FOP's sills and dikes which give way southward to the Betty's Pond Tonalite.

Note: For a detailed geological description of the deposit and U-Pb zircon dating see Stewart (1992) and Dubé et al. (1998).

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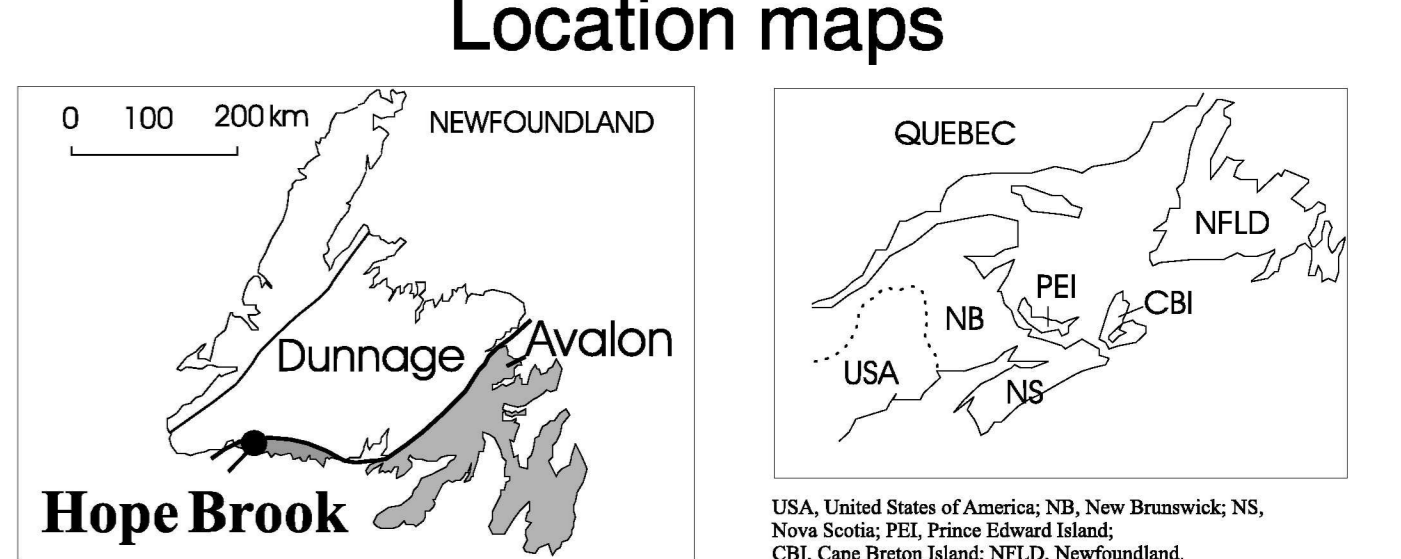
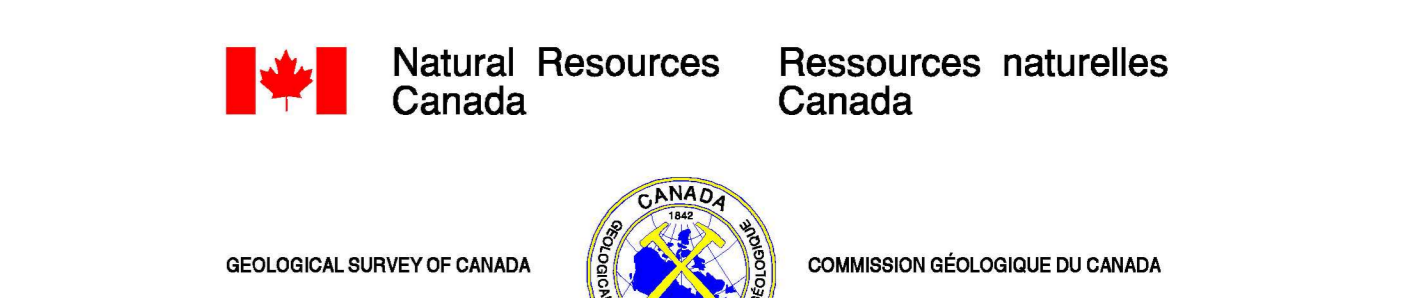
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