

# DETAILED OUTCROP MEASURED SECTION OF THE ST. MARY RIVER FORMATION, OLDMAN RIVER, WEST OF MONARCH, SOUTHERN ALBERTA

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**JULY 1998** 

Although every effort has been made to ensure accuracy, this Open File Report has not been edited for conformity with Geological Survey of Canada standards.

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#### INTRODUCTION

The enclosed measured outcrop description section represents a partially complete succession through the Maastrichtian-age St. Mary River Formation. The nearly continuous outcrop, located about 200 km south-southeast of Calgary (Fig. 1), is one of the most complete surface records of these strata in the Plains portion of the basin. It includes about 530 m of the St. Mary River Formation. Beds exposed at this location were briefly examined by Williams and Dyer (1930), Russell and Landes (1940) and Tozer (1952), concentrating particularly on the St. Mary River/Willow Creek contact. Nadon (1991) provided detailed descriptions and paleocurrent data for individual channel sandstones throughout this outcrop area, but no complete measured section or direct stratigraphic correlation to surrounding surface and subsurface occurrences.

Background geological information on the St. Mary River Formation, Edmonton Group and Bearpaw Formation, both in surface and subsurface, were summarized in Hamblin (1998a, 1998b, 1998c). In addition, information pertinent to the petroleum geology, discovered reserves and hydrocarbon potential of the strata of the Edmonton Group was detailed by Hamblin and Lee (1997). Further outcrop and subsurface studies of the Edmonton, and related St. Mary River Formation, are currently in progress.

The exposures are good to excellent, are scattered along 20 km of the Oldman River (Fig. 1) in a line oblique to the strike of the bedding and reveal portions of most of the ~530 m of stratigraphic thickness. The outcrops are located in Twp. 9-10, Rge 23-25W4. Together, the exposed sections cover about 410 m of the succession, with a large gap only in the lower 1/3 of the unit, here estimated at about 90 m (Fig. 2). Paleocurrent data, suggesting dominant direction of sediment dispersal, is summarized in Figure 3. A subsurface Gamma Ray-Sonic log from a nearby well to the west is also included for comparison to the outcrops (Fig. 4).

### SUMMARY OF SEDIMENTOLOGY AND STRATIGRAPHY

The exposed section of Maastrichtian rocks detailed here covers about 410 metres of the succession, missing only a 90 m interval in the lower third, and several very minor intervals elsewhere (Fig. 2). The basal contact of the St. Mary River Formation is exposed at the Monarch Fault Zone, 3 km south of the village of Monarch, where it overlies marine shale of the Bearpaw Formation. The upper contact is exposed at the northward bend of the river, 6 km north of the village of Pearce, where the lowest beds of the overlying Willow Creek Formation appear.

# Bearpaw Formation

The Bearpaw Formation, as exposed here, consists of grey to dark grey silty mudstone, slightly coarsening-upward, with abundant bioturbation. Only the uppermost few metres is present.

# St. Mary River Formation

The St. Mary River Formation is generally characterized by thinly interbedded siltstone, very fine to medium sandstone and minor carbonaceous shale to coal. The lower 60 m exposed here (Fig. 2a) includes alternating very fine to fine sandstone, grey burrowed siltstone with gastropod shells, minor greenish grey pedogenic siltstone and numerous seams of coal and carbonaceous mudstone. These facies are similar to those commonly associated with the lower

Horseshoe Canyon Formation to the north (see Hamblin, 1998b) and the lower portion of the St. Mary River Formation as exposed near Little Bow Provincial Park (see Hamblin, 1998a). This is interpreted as a shoreline-related transitional succession, with significant marine influence at the 30 m and the 40 m levels. No significant shoreface deposits are preserved here, although this section is approximately correlative to the Blood Reserve Formation of the southernmost Foothills (see Hamblin, 1998b). This exposure is followed by a covered interval, estimated at about 90 m in thickness (Fig. 1, 2a,b).

The succeeding section, from about 150 m to about 455 m (fig. 2b-e), consists of a monotonous succession of thinly interbedded greenish grey pedogenic siltstone, sharp-based fining-upward fine to medium sandstone, with sideritic concretion horizons and minor thin bentonite beds. Roots and dinosaur footprints are common, but thin carbonaceous mudstones are rare. These facies are typical of the St. Mary River Formation throughout its area of distribution in the southern Plains and Foothills (see Hamblin, 1998b). These sediments are interpreted as the result of floodplain deposition in a non-marine setting. Some sandstone units are thick, multi storied, have erosional bases and are characterised by abundant trough and ripple cross bedding. These are interpreted as significant channel deposits. Abundant paleocurrent data suggests dominant sediment dispersal toward the northeast (Fig. 3), a conclusion similar to that of Nadon (1991). Indications of some possible marine (or open lacustrine?) influence are present in the 305-330 m level: presence of *Ophiomorpha* and other vertical burrows, hummocky cross stratification, and thin coarsening-upward sequences. These features, positioned at approximately 2/3 from the base of the formation, are in a similar position to that of the Drumheller Marine Tongue of the Horseshoe Canyon Formation to the north.

From about 455 m to 533 m represents the upper 78 m of the St. Mary River Formation. It includes interbedded thick fine to medium sandstones, lesser greenish grey pedogenic siltstone, and abundant thin carbonaceous mudstones. Roots are ubiquitous. Sideritic concretion horizons are absent. These facies are more reminiscent of the Horseshoe Canyon Formation to the north. Paleocurrent data indicate sediment dispersal was much more variable than in the previous unit including transport to the northeast, southeast, northwest and southwest. These deposits are interpreted to represent floodplain deposition, with re-establishment of conditions favourable to preservation of peat accumulation. Ten metres below the upper contact a white, trough cross bedded sandstone with large vertical tree roots is present, approximately correlative to the well-known Whitemud Formation of southeastern and central Alberta. In the upper 10 m are two distinct dark grey to brownish, laminated, bentonitic mudstone units which are very similar to the well-known Battle Formation of southeastern and central Alberta.

## Willow Creek Formation

Only about 10 m of the overlying Willow Creek Formation is poorly exposed at this location. Bright green shale alternates with whitish very fine to fine sandstone.

### INTERPRETATION

Comparison of the surface and subsurface sections (Fig. 4), and with the section near Carmangay (Hamblin 1998a), suggests that the intertonguing of St. Mary River-type facies of the south with Horseshoe Canyon-type facies of the north occurs in a definable and stratigraphically organized fashion at this latitude. Horseshoe Canyon-type facies (pale grey thick channel

sandstones interbedded with burrowed marine-brackish mudstone, abundant coals and carbonaceous shales, and lesser pedogenic siltstone) are prevalent in the lower ~70 m and the upper ~80 m. However, strata of the intervening ~350 m are dominated by St. Mary River-type facies (greenish grey pedogenic siltstone with abundant caliche/paleosol horizons and roots, interbedded with thin grey to buff sandstones and rare thin carbonaceous shales). Here, at the latitude of Twp 10, the Horseshoe Canyon-like units are thinner and less organic-rich than those at the latitude of Twp 14 near Carmangay (see Hamblin 1998a). Conversely, the St. Mary Riverlike facies are more dominant.

This stratigraphic arrangement may be interpreted as representing a N-S climatic trend (more humid to the north, more arid to the south; first expounded by Jerzykiewicz and Sweet, 1988) which shifted first northward, then southward through the time of Maastrichtian deposition. Conversely, it may be the result of the large-scale intertonguing of two different clastic wedges in this area which had different depositional styles, sediment sources and patterns of sediment dispersal. A third possibility is a relation to transgressive-regressive patterns. A combination of all three is possible.

### HYDROCARBON POTENTIAL

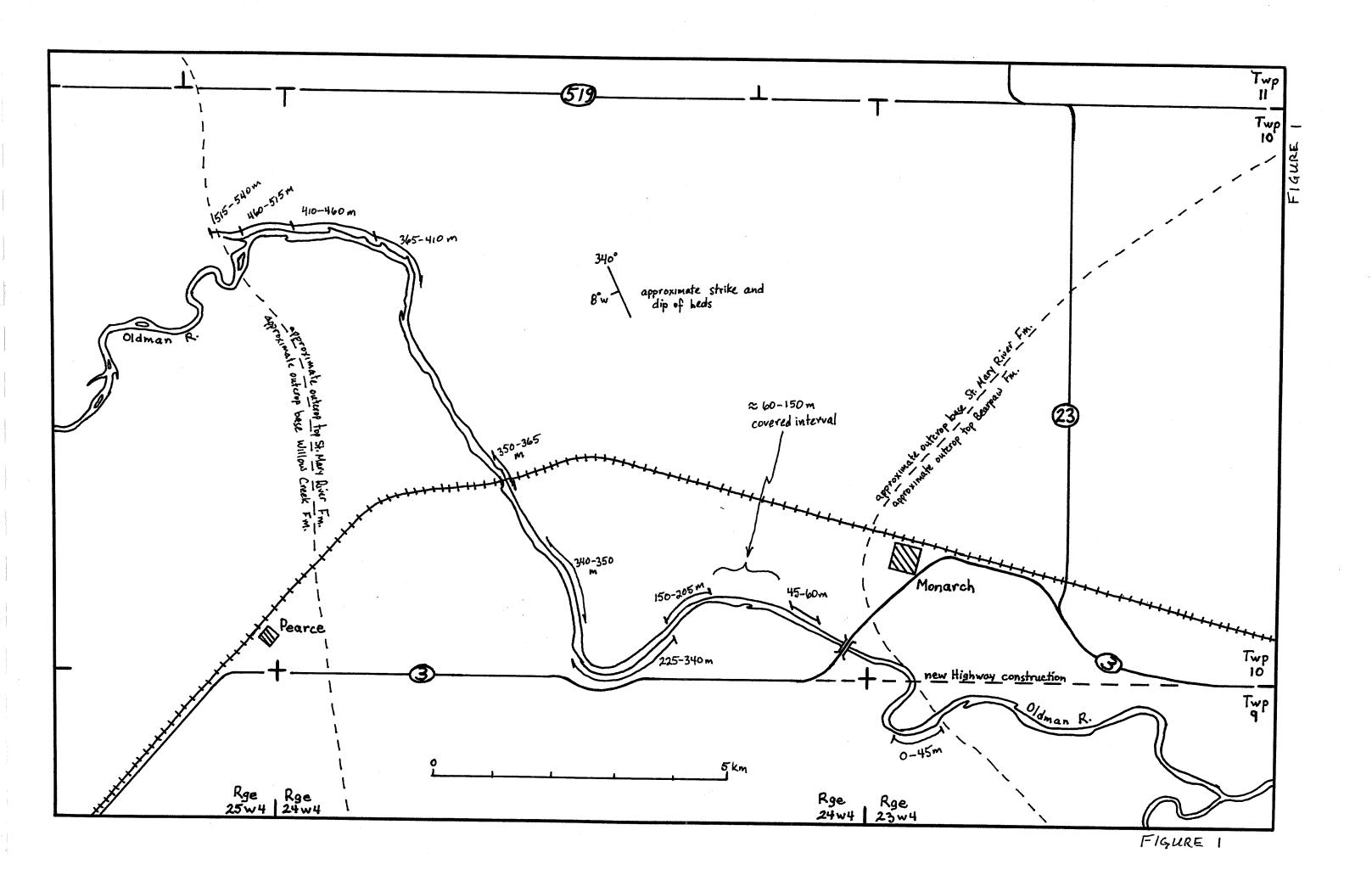
Further work is required to resolve the regional-scale stratigraphic architecture of the Maastrichtian deposits of southern Alberta. This may be important because a limited number of gas pools are known in the Horseshoe Canyon to the north, although relatively little exploration has ever been targeted directly toward these units. However, vast areas of Alberta shallowly underlain by these units are essentially unexplored at this stratigraphic level, and may offer undiscovered pools of modest size, low pressure, but inexpensive exploration costs.

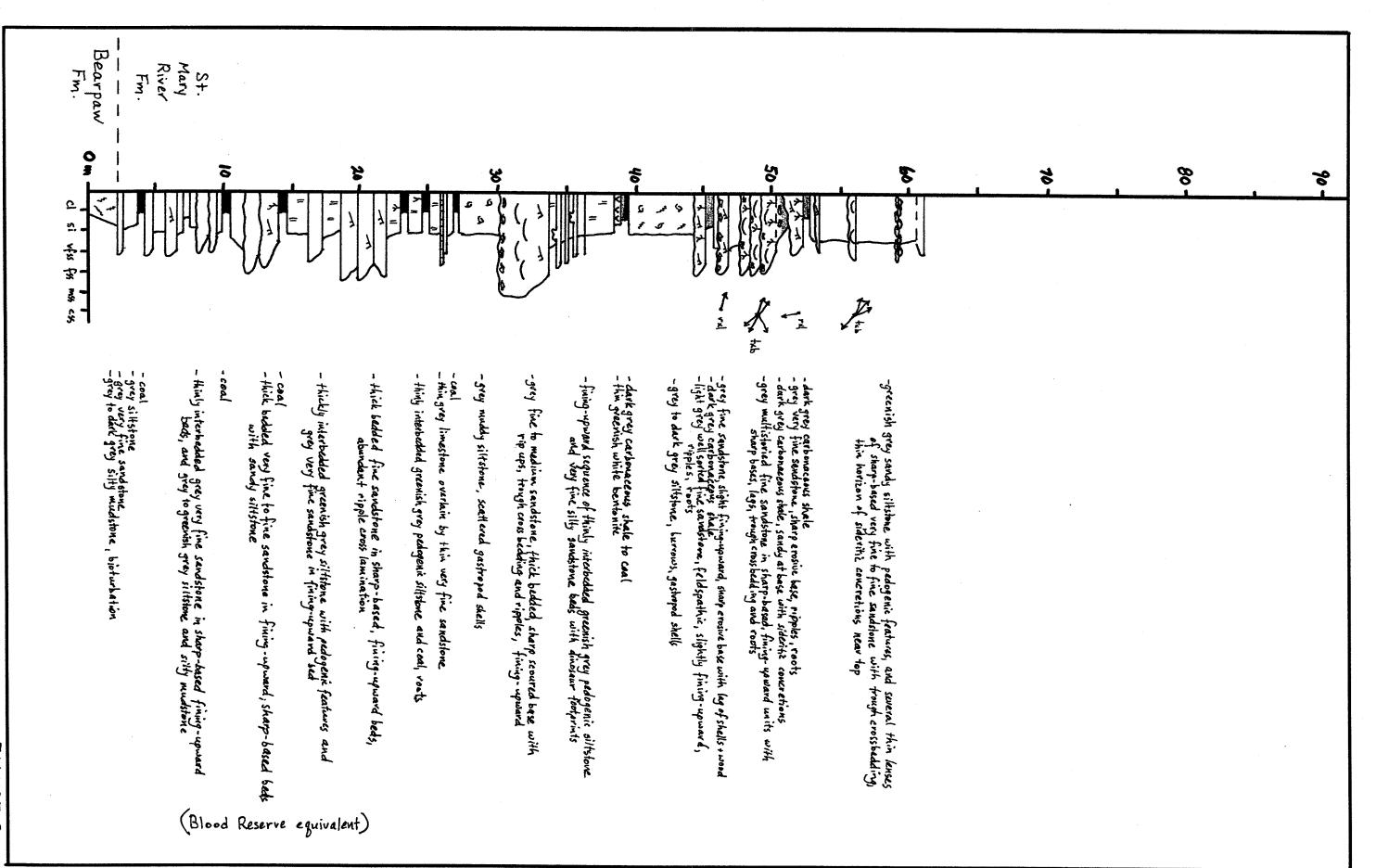
# LIST OF FIGURES

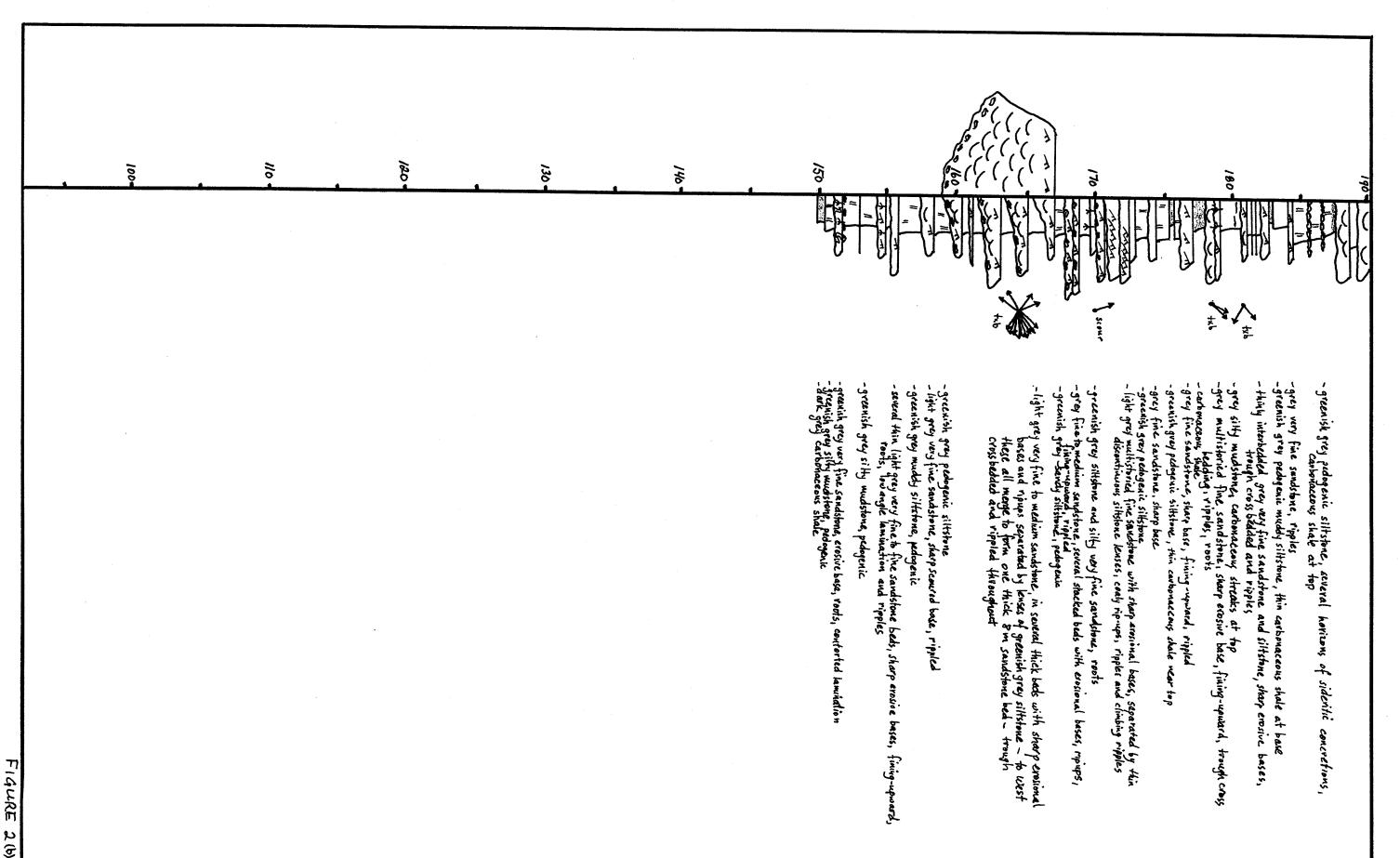
- 1. Location map for measured outcrop sections along Oldman River near Monarch. Thicknesses refer to those of measured section in Figure 2.
- 2 (a-f). Measured outcrop sections in stratigraphic context.
- 3. Summary of paleocurrent data from outcrops. A. Summary of measurements of trough cross bedding, ripple cross lamination and linear scours, with vector means of dispersal directions. B. Total of all paleocurrent measurements from these outcrops, with vector mean of dispersal direction.
- 4. Comparison of Gamma Ray Sonic log from well 12-4-10-27W4 with measured outcrop sections from Oldman River.

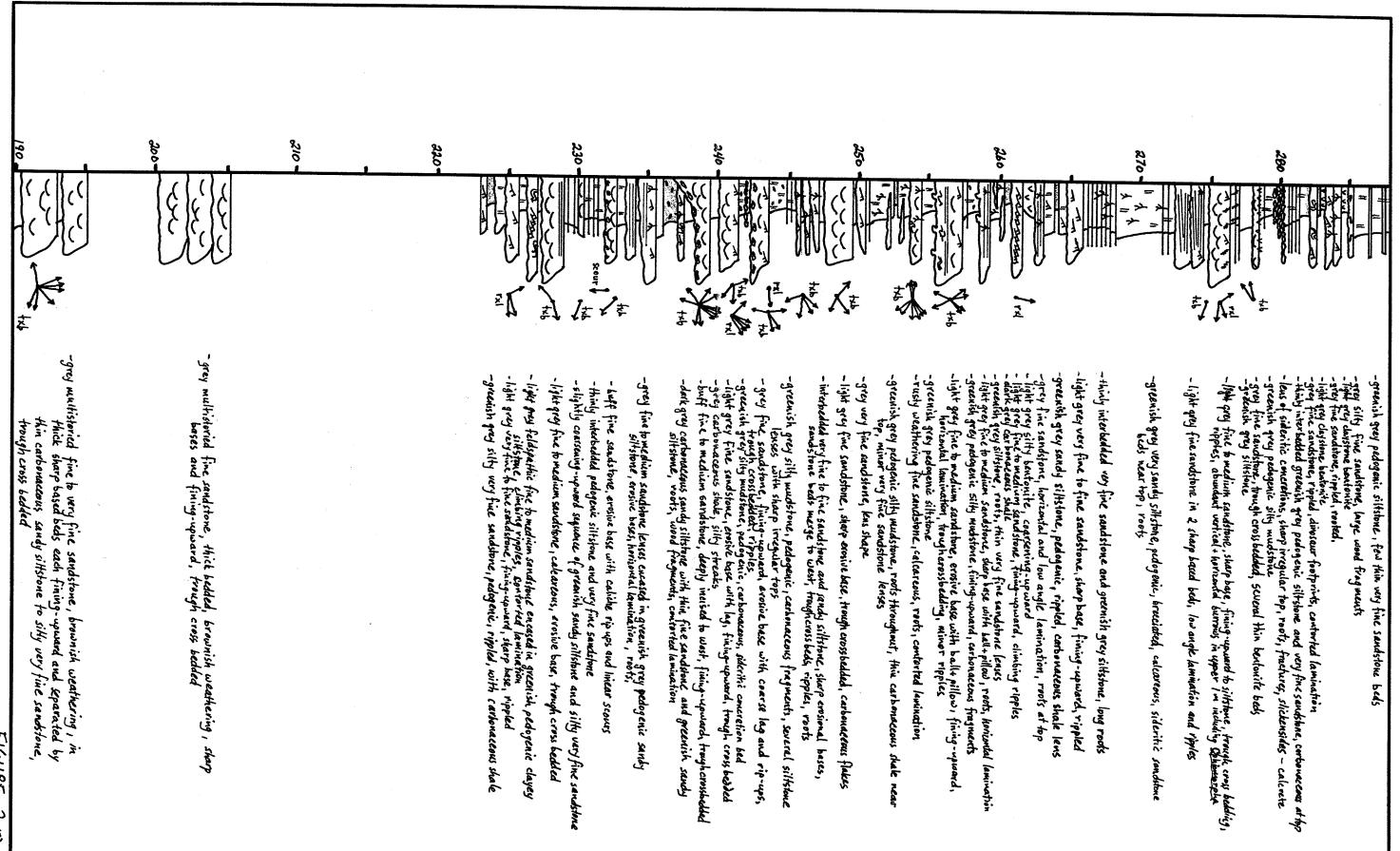
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FIGURE

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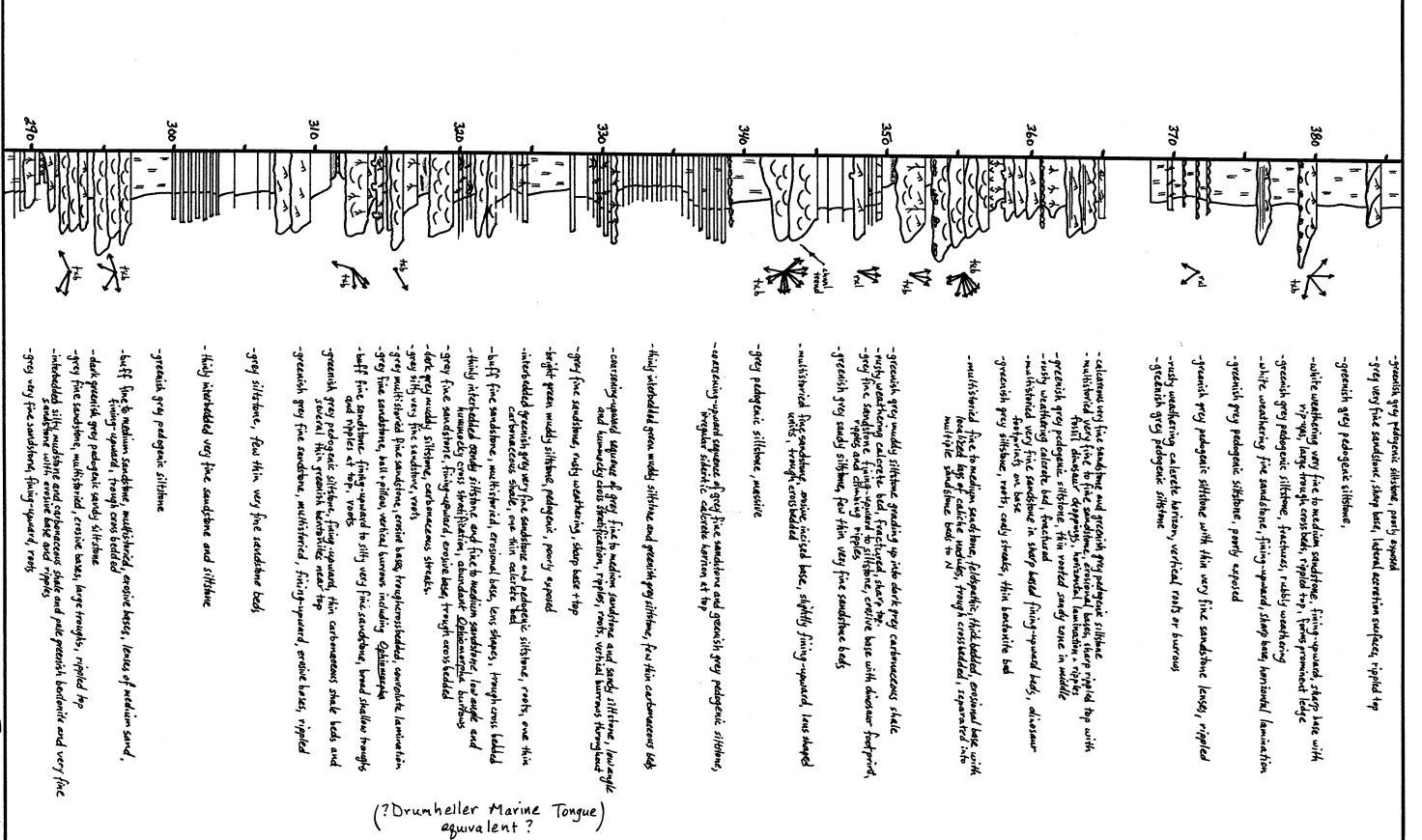
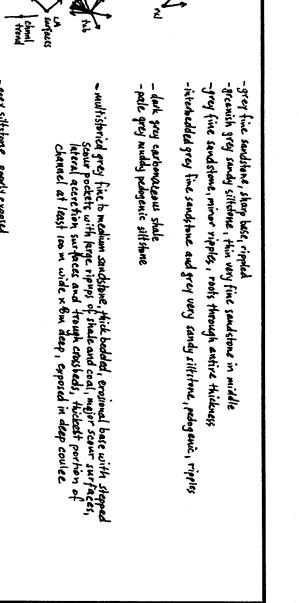


FIGURE 2(d)



FIGURE

2(0)

480

grey siltstone, poorly exposed

grey fine to medium sandstone, sharp flat base, horizontal Amination and trough cross beds, roots through antire thickness

-grey fine to medium sandstone, leaticular, fining-upward, large trough cross beds, ripples -greenish grey muddy siltstone, sideritic horizons near top - 2 storey fine sandstone, fining-upward, evosional bases, dinosaur footprints, wood fragments, roots, desication cracks at top, lower nippled, upper trough cross hedded

8

-darkgrey carbonaceous shale laminested roots; shell fragments -grey very fine sandstone, gradational base, sharp top -greenish grey silfstone, pedogenic, few thin sandy horrooms, shell fragments at top

-grey fine to medium sandstone, thick bedded, erosional base, rippled, abundant long roots through entire thickness

-greenish grey rooted pedogenic sillstone with many very thin rippled sandy beds and several sidesitic horizons

450

批

-several thin goog fine sandshine beds separated by thin sillshine

-greenish grey pedogenic siltsione, few very thin sandy horizons

multistoried grey five to medium sandstone, thick bedded, well sorted, exosional base with deep scour pockets, few sifty lenses to E, horizontal lamination, minor ripples + trough cross bads, long roots at top

440

grey coarsening-upward sendy sittstune with abundant burrows at base, rooted calcrete cape-brown sity bentwrite, micaceous
-greenish grey pedogonic sittsbane with grey rappled fine sandstone in middle
-coarsening-upward sequence of greenish grey sandy siltstone to silty very fine
sandstone, padogonic, calcareous cap -grey to greexish grey pedageric sandy sillstone with very fine to fine sandstone in middle, easily base with ripys, trough crossbeds and ripples

-grey very fine sandstone, sharp flat bese + tap -coarsening-upward sequence of greenish gray sandy siltstone, pedagenic, very calcareau

430

<u>-</u>

-grex sandy siltstone, thin very fine sandstone in middle
-grex fine to medium sandstone, faldspathic, erosional base, slightly fining-uptroughcross bedding and climbing ripple
-greenish grex pedogenic siltstone, sandy, long rods cut off by overlying unit -several stacked thick grey fine to medium sandstone, feldspathic, well sorted, erosponal bases, divosaur footprints, trough cross bedded, separated by thin siltstone lenses to w -greenish grey pedogenic silfstone

420

=

4

-3 stacked grey fine sandstone lenses separated by thin sandy sillstone, erosional bases, horizontal lamination, vinor trough cross-bedding, ripples

-greenish grey pedogenic siltstane with several thin very fine sandstane beds with sharp bases and ripples

410

-interbedded very fine to fine sands tone and grey siltstone, roots at top dinosaw-footprints, horizontal lamination, ripples, trough crossbads

-light grey silty mudstone, beminated, ?bentonitie?
- grey multistoried very fine to fine sandstone, crosive bases, horizontal lamination and ripples, siltstone lens in middle

-greenish grey pedogenic silfstone

-light grey fine to medium sendstone, well sorted, thick bedded, self-papper, pinches out to S, exosional base with scour pockets full of ripups/caliche/wood/dinosaur bone, trough cross bedded, long roots in middle -greenish grey pedagenic siltstane -multistaried grey fine sandstane separated by shale lenses, arasianal bases, trough cross budding and ripples, calcanous cap bed

400

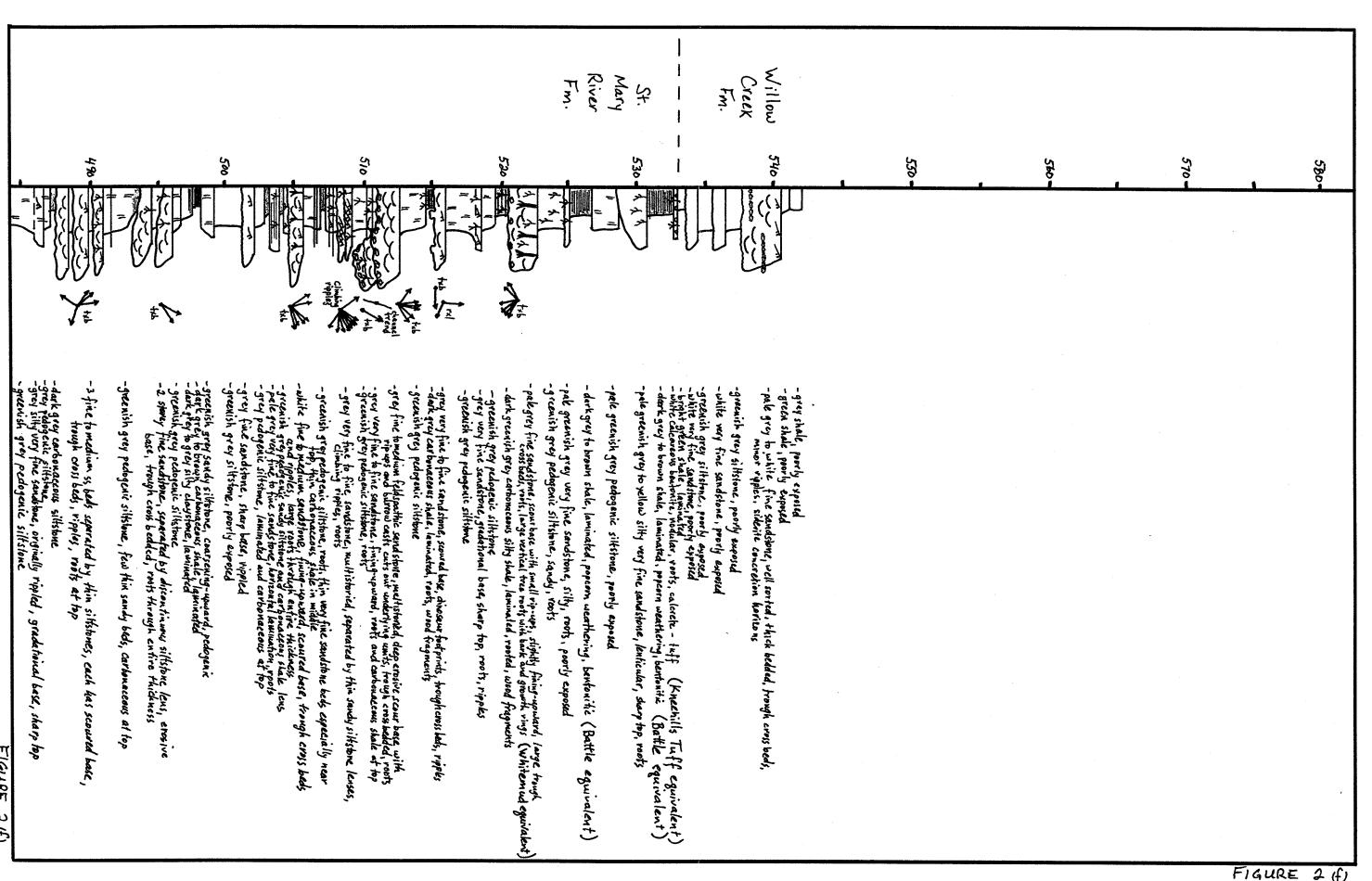
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interbedded greenish grey sillstone and very fine to fine sandstone, ripples

390

FIGURE શ

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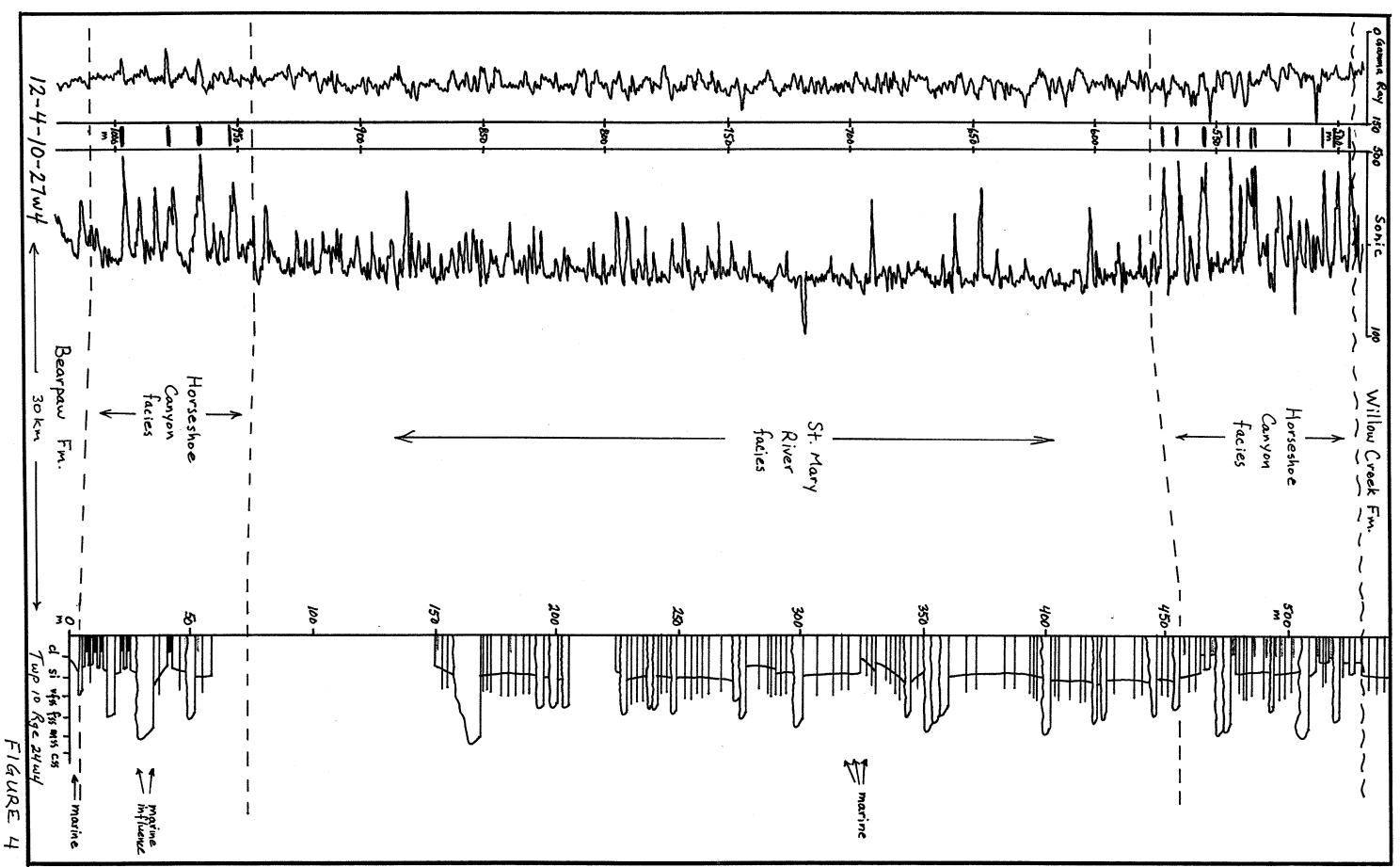


FIGURE 4

