## **DESCRIPTIVE NOTES**

The Dawson map area is composed of three major physiographic divisions: the Southern Ogilvie Mountains to the north and northeast; the Klondike Plateau to the south and southwest; and the Tintina Trench, which separates the first two divisions. The mountains are composed mainly of Tintina Trench, which separates the first two divisions. Precambrian and Paleozoic sedimentary rocks which are locally intruded by Cretaceous porphyritic syenite (Green, 1972). Intrusions form two of the three highest peaks in the region and served as the main source areas for local glaciers and ice caps. As a result of glaciation, the Southern Ogilvie Mountains are deeply dissected by "U" shaped valleys which cut across the main divide with a dominant trend from southwest to northeast. The Klondike Plateau is part of the Yukon Plateaus (Matthews, 1986) and is mainly developed on quartzite and schist of the middle and upper Paleozoic Nasina Series and Klondike Schist (GSC, 1988). The Klondike Plateau is largely dissected by dendritic "V" shaped vallies which reflect its unglaciated state. Notable exceptions are Sixtymi. River, which was glaciated prior to the last glaciation, and Fortymile River that features distal glaciofluvial terraces related to at least two glacial advances from the Yukon-Tanana Upland. The Tintina Trench is a graben of late Tertiary age developed along a Cretaceous - early Tertiary fault (Tempelman-Kluit, 1980). It is developed on unconsolidated to poorly consolidated Tertiary fluvial sand and gravel which include organic sediments with minor lignite.

Surficial deposits of this area can be divided into two major groups: I) colluvial deposits, which occupy approximately 70% of the area; and II) glacial deposits, which occupy about 25% of the area, primarily within the glaciated valleys of Southern Oglivie Mountains and the Tintina Trench. The emaining 5% includes organic, alluvial and lacustrine deposits.

Colluvial deposits consist of diamicton, rubble, and organic-rich silt and sand derived from bedrock and surficial materials by a variety of slope processes. Colluvium forms a blanket or veneer on the Klondike Plateau with rare bedrock outcrops, while bedrock and colluvium veneer dominate the interfluve areas in Southern Oglivie Mountains. In the southeast part of the Dawson area (Klondike interfluve areas in Southern Oglivie Mountains. In the southeast part of the Dawson area (Klondike interflude). Sollwish deposits commonly include organic-rich silt with massive ground ice. Slope goldfields), colluvial deposits commonly include organic-rich silt with massive ground ice. Slope complexes (Cx) occupy valley sides that may also include small alluvial fans, landslides and rock glaciers (in particular in cirques associated with modern glaciers and neoglacial moraines). Landslides (Cz) are very important in this region, particularly along the Tintina Trench were they cover about 35% of the total surface area. This total includes areas mapped as Me; moraine plains modified by landsliding. The Landslide category is usually included in the "slope deposits" category but because of the wide range of ages of this category it has been included in the Pilo-Pleistocene undifferentiated. Apparently landsliding was common in pre-glacial time, as shown by tilted Tertiary unamerenuated. Apparently landslitating was common in pre-glacial time, as shown by tilted Fertially sand and gravel unconformably overlain by glacial sediments east of Fitteenmile River (site 7). Postglacial downcut associated with the diversion and incision of the Yukon River on the Klondike plateau most likely triggered a new episode of landsliding following each period of glaciation, and continues today. Neotectonic movements may also be involved in landsliding.

Glacial deposits are grouped into three categories: glaciofluvial, morainal (till) and glaciolacustrine.

These deposits are the products of at least five major glaciations in addition to neoglacial activity.

Distinctly different degrees of preservation reflect ages from late Pliocene to late Pleistocene. The oldest deposits have been grouped as pre-Reid (pre-Illinoian), and includes three periods of glaciation. Reid deposits are of Illinoian age, whereas McConnell deposits relate to Late Wisconsinan glaciation. These formal names have been adopted because of their regional usage, in contrast with the terms "oldest, intermediate and last" used by Vernon and Hughes (1966) for the in contrast with the terms "oldest, intermediate and last" used by verticil and rughles (1960) to the glaciations in the Ogilvie Mountains. Pre-Reid deposits (glaciofluvial gravels and morainal) are generally oxidized and have highly weathered clasts in comparison with the better preserved younger deposits. As well, the thickest paleosols are associated with the older deposits (app. 2 m thick, deposits. As well, the thickest paleosols are associated with the older deposits (app. 2 m thick, deposits. As well, the thickest paleosols are associated with the older deposits (app. 2 in thick, contrasting 90 cm for Reid and 10-15 cm for McConnell deposits). Glaciofluvial deposits include proglacial and ice-contact sediments. They are generally well stratified to massive, poorly to well sorted, sand and gravel. Morainal deposits (till) consist of a nonsorted silt, sand and clay matrix with abundant pebbles, cobbles, and boulders. They occur in a variety of different and clay matrix with abundant pebbles, cobbles, and boulders. Glaciolacustring deposits consist of silt landscapes. landforms, including plains, hills, hummocks and ridges. Glaciolacustrine deposits consist of silt and clay with minor sand, commonly overlain by a blanket or veneer of organic deposits. Though lacustrine deposits appear in only a few places in the map area, they are an important component of Cz and glacial deposits (Me, Ge) of the Tintina Trench.

In the study area, older glaciations were generally more extensive than younger glacial events. The older glaciations reached the Tintina Trench from two different sources, the Cordilleran Ice Sheet (South Klondike and McQuesten Valleys) and the Ogilvie Mountains (from east to west; Brewery (South Klondike and McQuesten Valleys) and the Ogilvie Mountains (From east to west; Brewery (South Klondike and McQuesten Valleys) and the Ogilvie Mountains (From east to west; Brewery) Creek; North Klondike River and its northern tributaries, Chandindu River and tributary, Fifteenmile Creek; North Mondike River and its northern triputaries, Chandingu River and triputary, Fifteenmile River and Coal Creek). These source areas account for the mixed lithologies found in the three levels of pre-Reid glaciofluvial terraces mapped along the south part of the Tintina Trench (lower Klondike River). This is the only morphologic evidence for three pre-Reid glaciations in the study area, although stratigraphic evidence is found at several sites on the north side of the Tintina Trench (sites 3,4,5,6,7).

On the north slope of the Oglivie Mountains, pre-Reid moralnes form blankets and veneers of highly colluviated material with occasional erratics and extend to the foothills and plains. They were deposited by piedmont glaciers. These piedmont glaciers seem to have reached as far north as the southern slopes of Mount Skookum Jim, about 25 km north of the Dawson map-area's boundary. In the study area, however, Reid-age deposits dominate the plains and cover most of the areas of pre-Reid deposits along the Tintina Trench. In some areas (e.g. Chandindu and Fifteenmile rivers) it appears that Reid glaciers reached almost to the limit of glaciation. McConnell glaciation was very restricted, and its deposits are found within short distances of their sources (maximum 25 km).

Straugraphy

A preliminary correlation of deposits was done based on stratigraphic sections exposed at sites 1, 2, 3, 4, 5, 6, 7, along with paleomagnetic assessment and pollen analysis from site 3. A package of glacial sediments overlies Late Tertiary unconsolidated sand and gravel (Rs; sites 1, 3 and 7) of probable Pilo-Pleistocene age (White, 1996; Morison and Hein, 1987). A normal -reverse -normal probable Pilo-Pleistocene age (White, 1996; Morison and Hein, 1987). A normal -reverse -normal probable Pilo-Pleistocene age (White, 1996; Morison and Hein, 1987). sequence was found for the glacial sediments, Indicating a Gauss or Matuyama age for at least the first two glaciations to affect the area. The polarity signal as well as the high degree of clast weathering and oxidation of the three oldest glacial sediments distinguishes these deposits from the better preserved younger deposits. Because paleomagnetic assessment is preliminary, no polarity has been indicated in sections.

The pre-glacial landscape of the Yukon Territory and in particular the Klondike Plateau was very different than it is today. Drainage pattern analysis of central Yukon indicates that the Yukon River different than it is today. Drainage patiern analysis of central Tukon indicates that the Tukon invertible is an anomalous feature and is probably the result of glacial diversion. This idea of a southward flowing pre-glacial Yukon River was proposed by Tempelman-Kluit (1980). Pre-glacial drainage from Southern Ogilvie Mountains likely consisted of alluvial fans extending across the Tintina Trench area. This pattern was altered with the formation of the Tintina graben in late Pliocene.

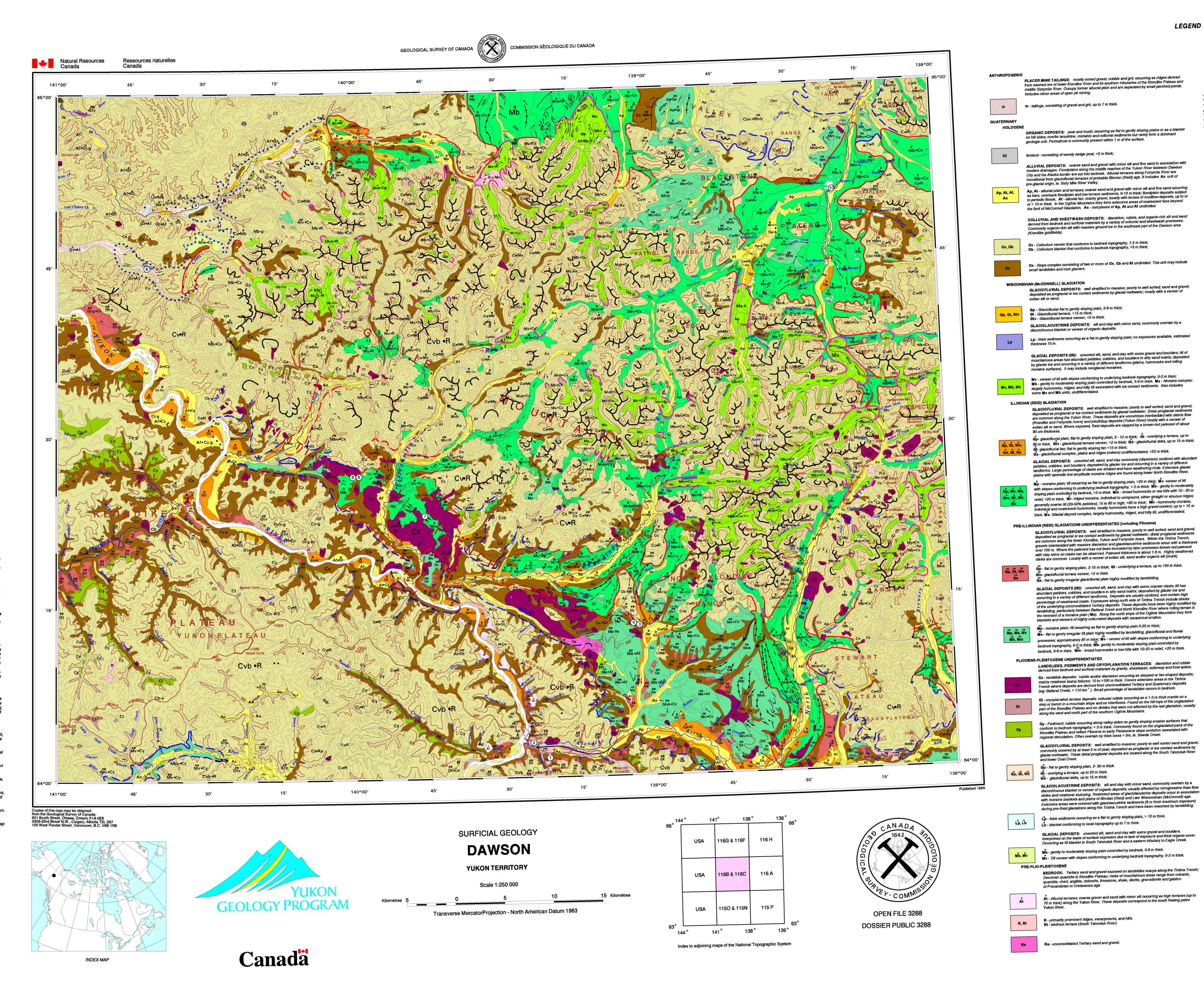
Paleo-topographic reconstruction of the Yukon Plateau indicates that the Yukon River drains along the northeastern paleo-foothills of the upland. This anomaly starts near Fort Selkirk, where the river crosses a bedrock spur before continuing along the paleo-foothills to the northwest. The river is crosses a bedrock spur before continuing along the paleo-foothills to the northwest. incised into the Klondike Plateau and crosses a paleo drainage divide between the Klondike and incised into the Klondike Plateau and crosses a paleo drainage divide between the Klondike and Fortymile rivers. Glaciation appears to be the only agent that could have determined the present-day course of the Yukon River. If this is correct it was the first glaciation, during which the Tintina Trench was occupied by glaciers, that diverted the river to the northwest and caused incision across the higher and more resistant terrain of the Klondike Plateau. Floodplains cut in bedrock can be observed along this stretch of the river.

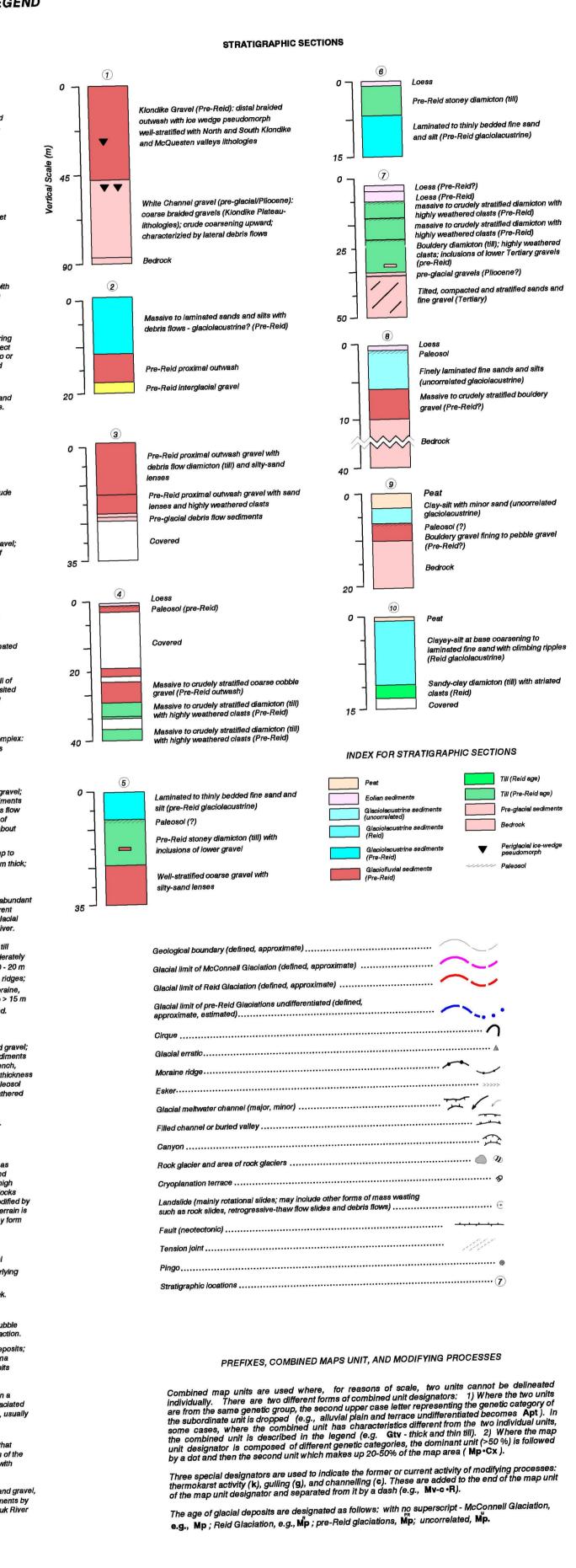
During the first glaciation of the central Yukon, Cordilleran and local glaciers coalesced and extended along the Tintina Trench (see limit of glaciation on map) as pledmont glaciers. The Grouse Mountain area marks the maximum western extent of Cordilleran ice. The ice moved between Moosehide and Klondike hills, diverting meltwater across the hills and establishing a permanent route for the Klondike River as an eastern tributary of the newly diverted Yukon River. Glaciers also caused westward drainage diversion of parts of Rock Creek and Chandindu River, and created a new drainage (Ballarat Creek) into a glacial lake impounded by Chandindu, Fifteenmile and Thane valley glaciers. The probable initial outlet of the lake was the lower Klondike meltwater channel, followed glaciers. The probable initial outlet of the lake was the lower Klondike meltwater channel, followed by the Chandindu-Fifteenmile-Yukon drainages. Another drainage diversion is found in the headwater of the southern-most tributary to Coal Creek, where part of Fifteenmile River was re-rerouted into Coal Creek. The subsequent pre-Reid glaciations seem to have had a similar pattern as the first glaciation, including the formation of glacial lakes. Reid glaciation, however, reached almost to the limit of glaciation of local glacier provenance, although Cordilleran ice did not reach the Tintina Trench at this time. The Fifteenmile valley glacier blocked the Yukon River and probably caused the formation of short-lived glacial lakes. McConnell Glaciation was very restricted and neoglacial formation of short-lived glacial lakes. McConnell Glaciation was very restricted and neoglacial moraines are present in cirques in the highest peaks of eastern Southern Ogilvie Mountains.

A great variety of placer deposits are found in the Dawson area, including pre-glacial gravels (White Channel, Klondike Plateau), interglacial gravels (Midnight Dome, Dawson town site), alluvial plains (lower Klondike River) and colluvium (Moose Creek). Sedimentological study of a gravel terrace (lower Klondike River) and colluvium (Moose Creek). Sedimentological study of a gravel terrace (site 2, mapped as outwash terrace) included a basal unit of interglacial gravels with a significant economic concentration of placer gold (Froese and Hein, 1996). Glacially derived outwash gravels overlying interglacial pay gravels are thick, crudely imbricate, containing large clasts, and sharply contrast with the gravel in the interglacial system. This raises the possibility of other outwashigh level terraces containing interglacial gravels with an economic concentration of placer gold in central Yukon (Froese, Msc. thesis in prep.).

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Geology by A. Duk-Rodkin (1995-96)

Geological digitization and cartography by: S.J. Hinds, Geodynamics Consulting

Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada

Base map at the same scale published by the Surveys and Mapping Branch in 1991

Copies of the topographical edition of this map may be obtained from the Canada Map Office, Department of Energy, Mines and Resources, Ottawa, Ontario, K1A 0E9

Mean magnetic declination 1990, 30°45′ E, decreasing 11.3′ annually. Readings vary from 30° 17′ E at centre of west edge to 31° 15′ E

Duk-Floukii, A. 1996: Surficial geology, Dawson, Yukon Territory; Geological Survey of Canada. Open File 3288, scale 1:250,000

at centre of east edge of the map

Elevations in metres above mean sea level