

Codes of this map may be obtained from the Geological Survey of Canada, 601 Booth Street, Ottawa, Ontario K1A 0E8, 353-262 Street, V-111, Québec, Alberta T2L 2A7

LEGEND

Coloured legend blocks indicate map units that appear on this map

QUATERNARY HOLOCENE

- O ORGANIC DEPOSITS: peat and muck up to 2 m thick, formed predominantly by the accumulation of vegetative material in bogs, occur in depressions and along valley bottoms; permafrost is commonly present; contains small pebbles, ice-wedge polygons, and thermoclastic collapse structures. Small unmetamorphosed organic deposits occur in most terrain units.
- A ALLOWED DEPOSITS: silt and sand, deposited by modern streams and rivers; deposits generally are stratified and moderately sorted; 1 to 5 m thick, occur as floodplains in places covered by glacial till.

PLEISTOCENE (WISCONSIN GLACIATION)

- L GLACIOLACUSTRINE DEPOSITS: silt and sand, cross-stratified to planar bedded; 1 to 8 m thick, deposition into temporary glacial-dammed lakes and ponds.
- G2 Outwash: rounded gravel and sand, massive to cross-stratified, probably less than 5 m thick, occurs as braided fans.
- G1 Esker sediments: sand, silt and gravel in planar, cross-stratified and massive beds; 1 to 40 m thick, forms ridges with both sharp-topped and flat-topped segments, mounds, and flanking aprons, deposited at or behind the ice margin; formed subglacially or in subglacially exposed ice-walled channels. Zones of washed rock, small transverse gravel ridges associated with this unit, isolated kame deposits, and circular rim ridges are shown by symbols.
- T3 Hummocky till: from 5 to 30 m thick, forms irregular to rolling terrain with relief up to 15 m; some areas have abundant small meltwater channels and lag concentrations of boulders in depressions.
- T2 Till blanket: from 2 to 10 m thick, occurs as till plains mimicking bedrock topography or as drumlins; small rock outcrops in this unit are shown by symbols.
- T1 Till veneer: less than 2 m thick, rock structure is generally visible on airphoto; unit includes patches of bedrock and till blanket.

PRE-QUATERNARY

- R3 Granitoid rocks
- R2 Metasedimentary rocks
- R1 Volcanic rocks

Geological boundary

- Ice-wedge polygon
- Frost heaved and shattered rock
- Geological date
- Radiocarbon date
- Fossil locality
- Raised beach
- Lag concentration of glacially abraded boulders
- Subglacial or proglacial meltwater channel
- Esker (direction of flow known, unknown)
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- Arenas and till plateaus
- Moraine
- Drumlinoid till form
- Rock crag-and-fall till form
- Roche moutonnée or whaleback
- Striation (ice flow direction known, unknown; 1=deduced)
- Striation, poorly defined (ice flow direction known, unknown)
- Gossan
- Small rock outcrop
- Sample site

DESCRIPTIVE NOTES

Aylmer Lake map area lies in central District of Mackenzie, on the watershed of major drainage systems flowing northward (Black River, south to Great Slave Lake, northward (Climaxine River, and north via Caribou Lake). Elevations range between 380 m and 500 m. Local relief is low, commonly <10 m in areas of rock and till blanket, although relief >20 m is observed in areas of hummocky till and, on some slopes, over low till. Rugged rocky areas north of Muskeg Lake have about 50 m local relief. Numerous lakes occur on the bedrock, and smaller water bodies occur in depressions on the surface. Most drainages are enclosed streams that flow into bedrock or the till plain. The area lies beyond the treeline. Glacial deposits are covered by low shrubs, sedges, and ferns with vegetation typical of the Slave Province National Geoscientific Mapping Programme (NATMAP) was designed to provide a regional framework for geological interpretation, environmental management, and drift prospecting. The project involved helicopter-assisted ground work, including terrain mapping, till sampling, and measuring of ice flow indicators. Field traverses and airphoto interpretation provided information on the nature and distribution of surface materials shown on the map. About 170 samples from this area were collected for heavy mineral and trace element geochemistry (Dredge et al., 1996a), 55 additional 10 kg samples were collected for heavy mineral and kimberlite indicator analysis (Dredge et al., 1996a), and pedee samples were collected to assess loess, silt, and glacial transport distance. Maps were constructed using 1:50 000 scale airphoto and recompiled onto a 1:250 000 scale. Open File 2798 (Dredge et al., 1994) and this map are the first regional-scale geological studies of the area.

The area lies within the central Slave Province of the Canadian Shield. Lord and Barnes (1954; Fig. 1) mapped the bedrock and made causal observations on striations and eskers. The rocks consist primarily of Archean volcanics and related fine-grained turbidite sediments belonging to the Yellowknife Supergroup (McGlynn and Henderson, 1972; Plafghan and Aderson, 1991). Sedimentary rocks are the most abundant type in the area. These rocks have been folded, faulted, and metamorphosed to various degrees, and formed into phyllites, schists, gneisses, and migmatites. The metamorphisms were later introduced by granite. Slight metamorphism has occurred along faults near sedimentary contacts. Dike swarms trending northeast, east, and north-south cross the area, and some dykes are 30 m wide. Kimberlite pipes have introduced older rock types. Ontonagon and Early Tertiary metasediments, both remain in metadiagenesis, and the latter date of 52 Ma indicate the age of some of the slope (Northern Miner, 1993).

The metasedimentary rocks have been subjected to intensive fracturing in the postglacial period. Large frost-heaved blocks, heaved above the level of the surrounding rock surface, are common features in the Yellowknife Supergroup.

Bedrock

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Glacial deposits

Till is the most extensive deposit in the area. Only one stratigraphic unit of till has been recognized and it is attributed to the Wisconsin Glaciation (Dredge et al., 1994). The till consists of a matrix-supported sediment, with the matrix ranging from silt sand to sand. The stone content varies from 5 to 40% of the total volume of material, but is generally about 20%. Matrix size varies according to bedrock composition and the degree of weathering of the bedrock. Till derived from granitoid rocks has a sandy matrix, while that derived from the Yellowknife Supergroup is clayey. Where glacial meltwater was associated with till deposition the matrix is commonly matrix to coarse grained sand.

The till sheet has been divided into three subunits based on thickness and surface morphology: veneer, blanket, and hummocky till. Till veneer is generally <2 m thick, contains exposures of bedrock, and conforms to the underlying bedrock morphology. It is generally thin and is deposited in the lee of the main ice margin. Major topographic elements of local provenance. Bedrock structure is usually visible on airphoto. Till blanket is generally 2 to 10 m thick, but thickens up to 40 m in places. It is generally composed of a matrix of silt and sand, with rounded to subangular clasts of local provenance. Bedrock structure is usually visible on airphoto. Till blanket is generally 2 to 10 m thick, but thickens up to 40 m in places. It is generally composed of a matrix of silt and sand, with rounded to subangular clasts of local provenance. Bedrock structure is usually visible on airphoto. Till blanket is generally 2 to 10 m thick, but thickens up to 40 m in places. It is generally composed of a matrix of silt and sand, with rounded to subangular clasts of local provenance. Bedrock structure is usually visible on airphoto.

Glacial events

White eskers are fluted forms related to turbidite granules within the ice, they are indicators of ice flow in the glaciers they record commonly run towards the ice margin, that is, they parallel ice flow. The largest eskers have major ridge segments trending westward. Near Aylmer Lake they have additional northward trending segments and may reflect a later ice flow shift towards the northwest (Fig. 3).

Glacial overprints

- Striations indicate that an early ice flow towards the southwest crossed the map area (Fig. 5b). Since the striations and eskers they are found on are unweathered, they are assumed to post-date the last glaciation. The flow is thus thought to relate to the building phase of the last glaciation.
- Cross-cutting striations, streamlined rock, and drumlins indicate a later major flow towards the west (Fig. 5b). This flow is tentatively placed during the main part of the Lake Wisconsin Glaciation and may reflect flow from the Mackenzie Ice Divide (Dyke and Prest, 1987). Though it is directly east of the Aylmer Lake map area. Some of the west-trending eskers may have developed at this time.
- White eskers flow post-dated in the south, striation sequences and re-oriented drumlins suggest that the rest of the area experienced a progressive shift in flow towards the northwest (Fig. 5c). These shifts may be responses to recession and related ice lobing. Major esker segments and their tributaries in the central and northern parts of the area following this northward trend may have developed during the same late glacial phase.
- The occurrence of till and streamlined forms on top of many eskers suggests that some developed while the ice was active. The apparently subparallel grain size of the till sediments suggests that much material was deposited into a low energy environment. The angularity of the sediment and preservation of mica and other minerals, and suggest either that the grain size of the till sediments was controlled by a protective water slurry. Since there is no evidence of glacial lakes in the area, the sediment must have been deposited into subglacial or proglacial water. The low energy environment suggests that the till was deposited in a protective water slurry. Since there is no evidence of glacial lakes in the area, the sediment must have been deposited into subglacial or proglacial water. The low energy environment suggests that the till was deposited in a protective water slurry.
- Striations, drumlins, and rock moutonnées in the northward record a late Wisconsin ice flow (Fig. 5d). The record of these features suggest that they relate to a late ice lobe confined to the area north of Muskeg Lake. Striations drumlins suggest that the flow of this lobe may have been affected by ice in the area south of the Aylmer Lake.

The broad band of hummocky moraine in the central part of the map area, with associated outcrop logs, rim ridges, and till plateaus, suggests that part of the ice sheet advanced to the area north of Muskeg Lake. This ice sheet advanced to the area north of Muskeg Lake. This ice sheet advanced to the area north of Muskeg Lake.

Esker under till, cut into till, and till at the surface above both till and outcrop; therefore they span a considerable part of the glacial/interglacial cycle. Associated with the eskers, either adjacent to, or on, the washed zones, are small kame features formed primarily of ripple and ripple drift beds, commonly ripple and sand ridges, large scale cross-bedding in fine sand. Footwalling revealed sharp changes in surface sediment from fine sand to boulder gravel. Exposure of bare, washed rock up to 1 km wide, commonly of till ridges and connect esker large segments.

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The peat deposits are common in low areas along poorly defined watercourses. Most of the ice-wedge polygons in the map area occur within the organic unit.

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QUATERNARY HOLOCENE

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DESCRIPTIVE NOTES

The area lies within the central part of the Mackenzie sector of the Laurentide Ice Sheet (Dyke and Prest, 1987; Dyke and Dredge, 1996). Aylmer Lake map area lies directly west of the Mackenzie Ice Divide, which developed during the Lake Wisconsin Glaciation (Dredge et al., 1994). From 10 000 BP the ice divide shifted eastwards into the District of Keewatin and disappeared there about 7000 years ago. At its maximum, the northern margin of the Mackenzie ice sheet lay north of the map area and the area is Mackenzie Valley. Radiocarbon dates on marine shells beyond the boundaries of this map area suggest that ice still covered the region about 10 000 BP, but that the ice by east of Aylmer Lake by 9000 BP.

Ice flow indicators

Some aspects of the glacial history and assessment of glacial transport of material can be determined by the relative age and strength of striations, although many rock outcrops are later covered or have rough, weathered surfaces, small granite pegmatite outcrops, prevalent throughout the map area, have maintained their glacial polish and striations. Striations on other rock types were found on individual mineral outcrops and on fresh surfaces exposed by peeling back the turf. Several techniques were used to determine relative ages of striations. At some sites, age relationships could be determined directly by examining dates of cross-cutting striations, although in many places, the cross-cutting relationship was ambiguous. Numerous outcrops have been revealed by more than one flow. The ice flow sequence on these can be deduced from the orientation of the striations, and the corresponding presence or absence of striations. Similarly, striations on planed faces (beaches) of rock moutonnées were assumed to post-date the flow that formed and slacked the outcrop. The striation date indicate an early southwestern regional ice flow, a westward flow, a shift to northwesterly flow, and a final southward flow in the Muskeg Lake area (Ward et al., 1994). All these striations are thought to represent the same glaciation because the surfaces they were found on are all relatively fresh (i.e., unweathered).

Streamlined bedrock forms are common throughout the map area, and conform in direction to ice flows determined from striations. A prominent esker of rock moutonnées and drumlins on the volcanic rocks north of Muskeg Lake indicate a sequence, but evenly limited southward ice flow. A first flow containing rounded drumlins forms an area from west to northward in the vicinity of Yon Lake.

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Geology by L.A. Dredge, B.C. Ward, and D.E. Kerr with assistance from M. Gingras, R. Paulsen, and B. Piens

Compilation by L.A. Dredge

Digital cartography by the Geological Survey of Canada

Contribution to the Slave Province Natmap Project

Logistical support was provided by the Polar Continental Shelf Project

Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada

Digital base map at the scale of 1:250 000 from the Surveys, Mapping and Remote Sensing Sector, Natural Resources Canada was assembled by the Geological Survey of Canada

Copies of the topographical outlines covering this map area may be obtained from the Canada Map Agency, Department of Natural Resources Canada, Ottawa, Ontario, K1A 0E8

Mean magnetic declination 1995, 24°20' E, decreasing 20.9" annually. Readings vary from 22°46' E in the SE corner to 25°01' E in the NW corner of the map

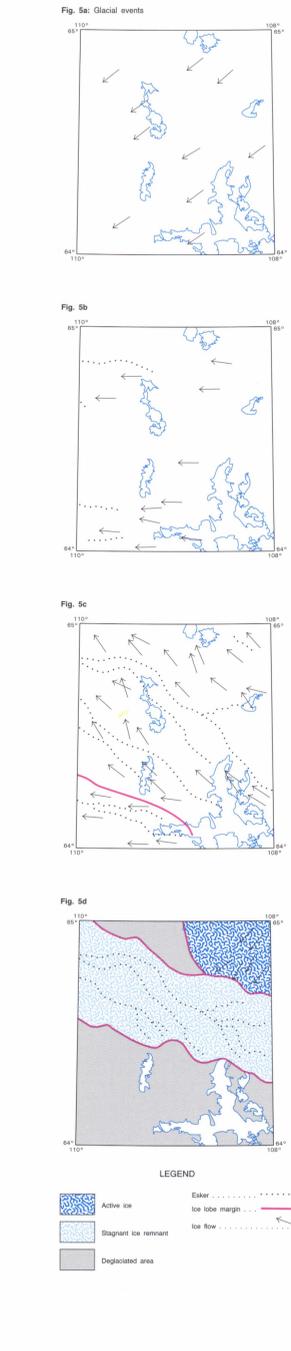
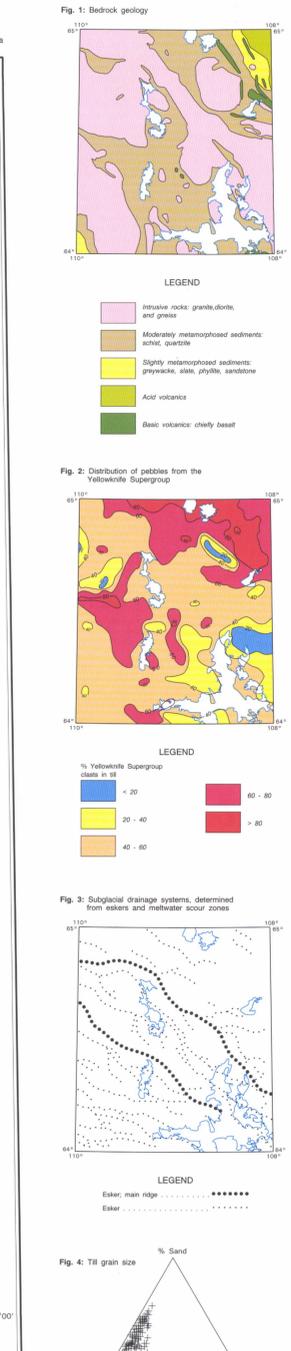
Elevations in feet above mean sea level

LEGEND

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1867A (Sb)

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