



GEOLOGICAL SURVEY OF CANADA

OPEN FILE 3043

Revised seismicity of the Grand Banks and offshore Newfoundland

J. Adams, R. Wahlström

1995



Natural Resources
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REVISED SEISMICITY OF THE GRAND BANKS AND OFFSHORE NEWFOUNDLAND

John Adams¹ and Rutger Wahlström^{1,2}

¹Geophysics Division, Geological Survey of Canada

1 Observatory Crescent, OTTAWA K1A 0Y3, CANADA

² now: Seismological Department, Uppsala University

Box 2101, S-750 02, UPPSALA, Sweden

Geological Survey of Canada Open File 3043

58 Pages including 5 Tables, 10 Figures and an Appendix

ABSTRACT

A comprehensive review of the catalogued seismicity of the Grand Banks of Newfoundland (1922–1994) has revised or added earthquakes in four time periods: 1922–1970, the two largest earthquakes ($M \approx 5$) relocated to Baffin Island (1922) and the northern Labrador Sea (1962); 1977–1982, added four new earthquakes; 1969–1984, revised the locations and magnitudes for five earthquakes; 1984–1994, added 15 earthquakes. Many earthquakes appear to be spatially associated with Mesozoic extensional structures that trend northeastwards across the Grand Banks. A cluster of six post-1970 earthquakes lies 40–80 km northwest of the Hibernia oilfield and represents the most seismically active region of the Grand Banks.

RÉSUMÉ

Un examen approfondi de la sismicité cataloguée antérieurement à la région des Grands Bancs de Terre-Neuve (1922–1994) a révélé quelques révisions et quelques séismes nouveaux dans les quatre périodes suivantes: 1922–1970, relocalisation des deux séismes les plus importants ($M \approx 5$) à l'île Baffin (1922) et dans le nord de la mer du Labrador (1962); 1977–1982, quatre séismes nouveaux; 1969–1984, révisions aux localisations et aux magnitudes de cinq séismes; 1984–1994, ajout de 15 séismes. Plusieurs séismes semblent associés spatialement à des structures extensionnelles du Mésozoïque qui tendent vers le nord-est à travers les Grands Bancs. Une grappe de six séismes postérieurs à 1970 se trouve de 40 à 80 km au nord-ouest du champ pétrolifère Hibernia et constitue la région la plus active des Grands Bancs.

INTRODUCTION

This open-file is the third of a series documenting our current knowledge of earthquakes off the southeastern Canadian continental margin (Fig. 1). The other reports in the series include: Adams and Staveley (1985) for the island of Newfoundland, and Adams and Simmons (1991) for the Labrador Sea, and are expected to include: Adams (in prep.) for the Laurentian Slope seismic zone, Adams (in prep.) for the Laurentian Channel and Fan, and Wahlström and Adams (in prep.) for the Scotian Margin. The current report covers a part of the region defined by the Grand Banks Basin Atlas sheet (a Frontier Geoscience project of the Geological Survey of Canada), and revises the epicentres and magnitudes of earthquakes on the Grand Banks and adjacent parts of the offshore.

As a side note, it should be noted that the “Grand Banks” earthquake of 1929, at magnitude 7.2 the largest on the southeastern Canadian continental margin, occurred in the Laurentian Channel (44.5°N 56°W), south of Newfoundland, and not on the Grand Banks proper. Thus it is not discussed in the present report.

Besides the more long-term scientific aspects, a recent interest in the seismicity of the Atlantic Margin has arisen from the need to produce seismic design levels (e.g. Foo and Crouse, 1986) for the Hibernia platform, near 46.7°N 48.9°W, east of St. John’s, and to refine estimates of seismic hazard to future offshore hydrocarbon production platforms such as may be emplaced off Sable Island, east of Halifax, or on the Grand Banks (Keen et al., 1990). Earlier hazard estimates (Basham et al., 1983; Adams, 1986) recognized the poorly determined nature of the offshore seismicity and the need to improve seismograph coverage to detect more of the smaller offshore earthquakes and to relocate the older earthquakes.

The current project was begun in 1982 by Adams, and many of the original earthquake phase arrivals were re-read during R. Wahlström’s study period in Ottawa in October 1984 - October 1985. Preliminary epicentres and magnitudes for the older earthquakes were determined at that time. Although only informal, these results have been available for use in seismic hazard analyses and have appeared in various reviews (e.g., Basham and Adams, 1983; Adams, 1986; Adams and Basham, 1989); it is our intention here to finally report them

in a formal manner, and to update the results to take account of more recent earthquakes, 1985-1994.

HISTORY OF EARTHQUAKE MONITORING

A brief, and probably still incomplete, history of earthquakes in Newfoundland is given by Adams and Staveley (1985). Prior to the advent of instrumental monitoring of earthquakes in Atlantic Canada, all earthquakes were detected by their felt effects. All of the earthquakes felt on the island of Newfoundland documented by Adams and Staveley (1985) appear to pertain to small, onshore earthquakes and not to large offshore earthquakes. By contrast, some earthquakes felt in coastal Labrador are likely to have been large offshore events (Basham and Adams, 1983; Adams and Staveley, 1985), though A. Stevens (pers. comm., 1991) considers they might have been local earthquakes along the coast. For example, an event felt widely (but not strongly in any one place) in Nova Scotia and New Brunswick in 1882 might have been an offshore earthquake (Ruffman and Peterson, 1988, p. 284). Another feature that might signify an offshore earthquake is the occurrence of a tsunami (e.g., Adams and Staveley, 1985; Ruffman et al., 1991); however no earthquake has yet been convincingly confirmed from the numerous reports of tsunami-like waves.

Seismograph monitoring of the Grand Banks has improved with time, so that small earthquakes that can currently be detected would have passed unnoticed in previous years. The following brief history is adapted in part from Stevens (1980). A seismograph was operational in Halifax (HAL) from 1915, but it operated at a low magnification (first a Bosch and then a Mainka instrument, with magnifications of 100 - 150x) and was able to detect only the largest earthquakes. The installation of the Wood-Anderson seismographs (peak magnification at 1 second = 2700x) at Seven Falls (SFA) and Shawinigan Falls (SHF) in 1927 improved the monitoring, as did the Benioff seismograph ($\approx 8,000x$ for 1937-1939 and $\approx 80,000x$ subsequently) at Ottawa in 1937 and the Willmore seismograph ($\approx 30,000x$) at Halifax in 1952.

All of these seismographs lay distant from and southwest of the Grand Banks, and it was not until 6-component Canadian Standard Seismographs were opened at Schefferville (SCH) in July 1962 and St. John's (STJ) in June 1964 that coverage improved significantly. STJ, while it was well-located with respect to Grand Banks earthquakes, operated at a magnification of only 35,000x for the short-period instruments because of the high microseismic and cultural noise of the site.

A further improvement in monitoring occurred when Memorial University of Newfoundland started to operate short-period vertical component seismographs in St. John's (in the same vault as STJ, but herein named 'MUNF' to distinguish it for the period when the STJ standard station was also in operation) in February 1975 and Corner Brook (CBK) in October 1976. These instruments operated at a higher gain ($\approx 100,000x$) than STJ because of their restricted low frequency response and proved very successful at detecting smaller offshore earthquakes. However, they were not always operated to modern standards with respect to timing, calibration, or record annotation, and the records were not routinely read with the other Canadian seismograms. In June 1983, Adams retroactively read the earlier seismograms, and the subsequent seismograms have been forwarded to Ottawa and have been systematically read by Adams as part of the national seismograph network.

The most recent changes affecting the earthquakes in this report are as follows. STJ was closed in March 1989, and the MUNF seismometer was relocated from a vault on the main campus to nearby Mount Scio and was renamed to STJN in September 1991. STJN has been operated under contract from Memorial University to the Geological Survey of Canada since April 1989, and CBK was operated under contract from April 1989 to March 1994. GSC's new digital seismograph at Deer Lake (DRLN) opened in December 1993 and replaced the function of CBK.

Although coverage has improved with time, it is still poor with respect to the rest of Canada. As will be seen (e.g., Fig. 4), very few earthquakes of magnitude less than 4 are detected more than 300 km distant from St. John's. The few larger earthquakes that are located (or only detected, see Table 2) must therefore be indicative of many smaller earthquakes that pass unnoticed in the same active areas. The last earthquakes mapped in this report occurred in 1992. There were no earthquakes located in the study area in 1993; the three that occurred in 1994 and one in early 1995 are added only to the Appendix and

the “Results” section as a “note in proof”, since the report was substantially completed in July 1992, even though its issue was delayed until March 1995.

PROCEDURES

Choice of events studied

Figure 1 delimits the area of the southeastern offshore covered by the Basin Atlas mapsheets and the sub-region treated in this report. The earthquakes studied in this paper represent: all known earthquakes thought to lie in the Grand Banks region according to the Canadian Earthquake Epicentre File (CEEF) in October 1985 (being the previous state of knowledge, see Fig. 2 and Table 1); all subsequent earthquakes located by current methods; and some older earthquakes (not in the CEEF) found by searching the older Memorial University seismograms.

In all, 23 pre-1993 earthquakes are studied in this paper: the two earliest are moved away from the area (having been grossly mislocated), 4 are new events added from the Memorial seismograms, and 17 are earthquakes which have either been relocated (pre-1985 events) or new events located during this project. In addition a number of unlocatable events were detected, usually only on STJ or MUNF, but occasionally on CBK alone if the high-gain MUN station in St. John’s was not operating (Table 2). These earthquakes were too small to register on other seismographs and are thus probably offshore.

Determination of Epicentres

For all the earthquakes studied here the epicentres and focal depths have been determined with the standard crustal model and location program used by the Geophysics Division. This model is a good average for the Canadian Shield and is used for offshore earthquakes in the absence of sufficient information to refine the average velocity structure. The crustal model has a 36-km-thick crust of P-velocity 6.2 km/s overlying a mantle of P-velocity 8.2 km/s. S-wave velocities are 3.62 km/s (revised from the previous crustal ‘Lg’ velocity of 3.57 km/s by Connors and Adams, 1988) and 4.70 km/s, respectively. It would

not be surprising if the use of this simple structure leads to some systematic bias in the epicentres (likely up to a few tens of kilometres, but possibly even larger), because with the distribution of land-based seismographs (all to the east of the earthquakes) the epicentres move to minimize the residuals created by an inappropriate crustal model. However, it is impossible to be sure because there have been no earthquakes large enough to give a reliable teleseismic epicentre for comparison, and no earthquakes located both by ocean-bottom seismometers and the seismograph network (contrast the situation on the west coast of Canada, e.g. Wahlström and Rogers, 1993; Wahlström et al., 1990). For these reasons the reader should be warned that the apparent precision of the results (e.g. the errors on Table 3 or in the the Appendix “pikfile” solution lines) is misleading. Furthermore, because of the distribution of stations, the true epicentres are often more uncertain in latitude than longitude, even where this is not indicated by the errors in Table 3.

For each earthquake, all existing Canadian seismograms were read for phase identification, phase onset times (Pn, Pg, Sn, Lg), maximum amplitudes and corresponding periods (Sn or Lg), and P-wave polarity (Pn or Pg). Many of the records for the earlier earthquakes were read by Wahlström at Lamont-Doherty Geological Observatory, New York, where the older Canadian seismograms were stored until the mid-80’s. Adams subsequently re-read selected phases in Ottawa. Phases from the Greenland stations were added from Danish bulletins when available. Timing was uncertain on the Memorial seismogram until about 1983; notes in the Appendix indicate where an arbitrary time correction was applied in order to use the S-P information.

The routine determination of epicentres by the Geophysics Division uses all available phases except those that obviously misfit. Hypocentral depths are often assigned to 18 km, being simply the mid-crustal depth according to the model. Epicentres in the present paper were recalculated from the earthquake phase arrival times according to the following guidelines:

- 1) STJ and MUNF were co-located in the same vault. Therefore where individual phases were read on both instruments, only one was used in the location, the other being zero-weighted (‘X-ed’ out). Usually it was the Pn from MUNF (its higher gain gave a sharper arrival) and the Sn from STJ (the Sn could be read best from the horizontal component; MUNF was vertical component only) that were used.

- 2) Pg and Sg arrivals for stations within about 100 km were fit well, even at the expense of larger residuals on the many distant stations (affects only events #69, #94b, #94c, and #95).
- 3) Pn and some Sn arrivals for the closest stations were fit well, at the expense the fit of arrivals at more distant stations.
- 4) Lg phases at distances beyond about 1000 km were given lesser weight because of their less sharp onsets and uncertainty in the Lg velocity.
- 5) In a few cases, a single station was chosen to replace a cluster of stations at a similar azimuth and distance to avoid undue bias.

The results of this data selection can be seen by examining the residuals for the earthquake solutions in the Appendix. While such recomputations of standard epicentres – which occasionally revise the routinely-determined epicentres by many kilometres – do not guarantee that an **accurate** epicentre has been found, we believe that they represent an improvement on the routinely determined epicentres. The epicentres are listed in Table 3 and plotted in Figure 3.

Focal depth

It is in general difficult to compute focal depth for small local earthquakes in the absence of a very dense local seismic network and a good knowledge of the crustal velocities. For large ($M > 5$) earthquakes that are recorded at teleseismic distances (> 2000 km) it is possible to compare arrivals of different waves to determine the depth of the earthquake, e.g. the delay between the direct, downward P-wave and the wave that propagates upwards before being reflected down at the earth’s surface as P (pP and sP). All the earthquakes studied here are too small for this method.

At close epicentral distances (conventionally, within twice the focal depth of the earthquake being studied), the depth can be computed by minimizing the combined residuals on the upward- and downward-propagating rays; however, the earthquakes in this report lie offshore and have poor seismograph coverage at small distances. Consequently all but one earthquake in this report are placed at the mid-crustal depth (18 km) of the standard model.

Determination of magnitudes

Magnitudes are determined according to current GSC practice, namely: a m_b or M_S from teleseismic P amplitudes is preferred for the larger events (say $M > 5$) if it is available from USGS or ISS; a m_{bLg} (also described as a m_N) determined from the Lg-wave amplitude is preferred where the Lg has propagated normally and not been severely attenuated; and finally a M_L determined from the Sn amplitude is used as a last resort. Supplementary magnitudes (e.g. m_b where a m_{bLg} is preferred) are retained in the CEEF. The M_L calculation with Sn amplitudes uses California attenuation and overestimates magnitudes from stations involving long paths across the high-Q Canadian Shield. Results reported by Adams and Simmons (1991), who compared teleseismic m_b with M_L for nine Labrador Sea earthquakes, suggest that the M_L 's for the magnitude range 4.5 to 5.4 might be higher than m_b by 0.4 magnitude units. However, their results are too poorly based to suggest that the M_L magnitudes reported here should be uniformly reduced by that estimate of bias.

Determination of polarities

Where possible, polarities were read from the first P (Pg or Pn) on all available records. These produced a sparse set of data, often only the polarity on the closest station being readable. Polarities were assigned either full or half weight. The polarities were read by both authors and readings on which we differed were either not used or a consensus was reached and the polarity was used at half-weight. Full weight polarities are impulsive and unambiguous. Half weight polarities are emergent, less strong, or may occur on noisy records.

RESULTS

The results of the relocations are summarized in Tables 3 and 4, and as a map of the revised seismicity in Figure 3. Details of the relocated solutions are in the Appendix, which also contains detailed comments on the phases and solutions. Figure 4 shows a map of the relocation vectors. The earthquakes are mapped at their revised locations

and the tails point back to their old epicentre. Figure 5 shows the new epicentres with diamonds showing the computed precision of the epicentre (large diamonds represent poorly determined epicentres), Figure 6 shows the detail of the area east of St John's, and Figure 7 shows the detail of the cluster of earthquakes near Hibernia.

In the following discussion, and in Tables 3 and 4 and the Appendix, earthquakes are denoted by their date, e.g., #89b, which is the second earthquake in 1989.

#22 was originally located by the International Seismological Summary (ISS) at 50°N , 50°W based on six stations, only OTT being in Canada. Although SAS was operating, the records were apparently not saved, and the records for VIC have still not been located. The OTT records were not found, and the microfilm of the Bosch and Milne-Shaw records show no clear signal. However, we did find the HAL Mainka record (one horizontal component). This had a time correction of -45 seconds, and gave clear phases $P=06\text{h}34\text{m}33\text{s}$, amplitude 0.10mm at a period of 2.6 s; $S=06\text{h}39\text{m}05\text{s}$; and surface wave maximum at 06h42m amplitude 0.15mm at a period of 4.0 s. This indicates a distance of 26.5° (irrespective of the exact time correction; that P and S are the recorded phases is confirmed by later arriving surface waves), which is more than twice the distance of HAL from the assigned ISS epicentre (the HAL readings were not included in the ISS solution).

The distance from HAL is also consistent with the recorded surface wave maximum at 06h42m. Using the arrival times given by ISS and a revised phase interpretation we find origin time 062855, OTTAWA distance= 25° (S and RM; the P misfits by about 1 min, which may be a 1 minute reading or reporting blunder because accurate timing was not a problem at the Dominion Observatory), CHICAGO distance= 28.5° (RM), WASHINGTON distance= 32° (RM), and ESKDALEMUIR distance= 32° (RM). 'RM' denotes the Rayleigh wave maximum, the time of which can be used for crude locations in the absence of phase onset arrival times (At Uppsala a set of RM travel time tables prepared by Båth are used). These readings give an approximate epicentre at 70°N , 75°W (Baffin Island). For this solution, the two other ISS readings give distances about 4° too long if the recorded phase is interpreted as RM for this solution (BIDSTON distance= 38.5° and DE BILT distance= 41°). The

Uppsala Weichert records were checked, but there was absolutely no trace of the earthquake. Thus, that two of the three European stations with data in the ISS are off by a few minutes should not be discouraging: the signals must be very weak.

Dr. Anne Stevens has suggested that the HAL, OTT, and WAS phases would fit neatly as Pn, Lg; Pn, Sn; and Lg respectively and give an epicentre near 59N 53W, off southwestern Greenland. We can not agree with this location, for: 1) if the epicentre is in eastern Labrador Sea, then the second phase at HAL is not Lg, because Lg does not propagate across this oceanic region (the same goes for the WAS reading); and 2) the clear 3rd phase at HAL is not explained. The “reasonable fit” of these three stations is worth little compared with our new solution which satisfies practically all stations.

Our solution, with identified and accurately timed phases at HAL, giving a reliable distance, and several fitting LR maximum arrivals for a Baffin Island location, is no doubt the best from the documented data. The greatest “obstacle” is the misfitting OTT P arrival and that this cannot be checked since the records are lost. Our solution is both the best and is very rough (see Discussion below).

Smith (1962), gave a fictitious intensity of MM VI-VII to this earthquake, which was not reported felt, and from this intensity the current M_L 5.3 in the CEEF was derived. From the horizontal component at HAL (orientation unknown) we read a trace amplitude of 0.15 mm at a period of 4.0 s. Assuming i) a magnification of 150 (the maximum magnification was 100-150 according to Stevens, 1980) for the found record, and ii) zero amplitude for the missing orthogonal component, we compute $M_s(\text{Prague-Moscow})=5.1$. Due to the two assumptions this is the minimum value.

In conclusion #22 is a magnitude 5 earthquake in Baffin Island (about 2850 km from its ISS epicentre) and not a Grand Banks earthquake.

#62 was originally located east of the Strait of Belle Isle, and prior to revision was a single large earthquake in an otherwise aseismic area. Both the original and revised locations are based on only three phases from two stations, a P and an S on SCH and one phase on HAL, which Smith (annotations on Dominion Observatory epicentre cards) took to be Sn. A inland epicentre west of Schefferville was ruled out by Smith because the earthquake was not detected on southern stations such as SFA or

MNT, so he located it on the southern Labrador Shelf. After re-examining the HAL seismogram and confirming that only one phase was recorded, we have re-interpreted Smith's Sn reading as the Pn phase. This, together with the absence of Lg energy at Schefferville, is consistent with an earthquake in the oceanic crust of the Labrador Ridge. A slightly larger Labrador Ridge event, 621202, close to the revised #62 epicentre, is well-located, and is also found to lack a Sn reading for HAL (Adams and Simmons, 1991). HAL and SCH were at similar distances from the old #62 epicentre, but the SCH record had a 30 mm amplitude, while HAL signal was "very small"; this suggests the earthquake was much further from HAL, as we now locate it. The epicentre of #62 has thus been revised by over 1000 km.

#69 occurred very close to St. John's, but was not reported in the newspapers as being felt (Adams and Staveley, 1985). The epicentre is revised 17 km to the SE.

#70 is a relatively large earthquake, M_L 4.3, north of the Grand Banks. It is located using only stations STJ and SCH, but the absence of phases on HAL excludes the southern alternative epicentre.

#71 is now the largest earthquake on the Grand Banks at m_b 4.8. It lies within 65 km of the Hibernia Oilfield, and has been moved 32 km SE of the CEEF epicentre. The maximum amplitude readings on Canadian seismographs for this earthquake represent Lg energy, and the m_N 4.8 so calculated is similar to the teleseismic m_b from the ISC. However, m_N is the preferred magnitude scale since it best represents the frequencies of engineering interest when this earthquake is used in eastern Canadian seismic hazard calculations.

#76 is confirmed as being located north of the Grand Banks, the revised epicentre being essentially the same as that in the CEEF.

#77 is a new earthquake found by searching the Memorial University seismograms, but was too small to be confirmed by GSC seismograph records. From the absence of phases at Halifax, the earthquake lies to the northeast of St. John's on the northern Grand Banks, rather than south of St. John's.

#78 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It lies far to the north of the Grand Banks, close to the similar-sized #70.

#81 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It is one of the most easterly of the Grand Banks earthquakes, but some Lg energy did propagate to STJ.

#82 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It is the first known earthquake from the Newfoundland Ridge – Tail of the Bank area. It is magnitude $M_L 4.0$, a relatively large earthquake to have been missed during the routine compilation of seismicity (it occurred during the Miramichi aftershock sequence and may simply have been overlooked). However, it is probably close to the detection threshold for this remote area, and should be taken as indicative of seismic activity along the southeastern margin of the Grand Banks that is currently undetectable.

#84a is a moderate earthquake on the northern Grand Banks, notable for having a $M_L 2.6$ #84b aftershock two and a half hours later. The mainshock epicentre is relocated 17 km to the SE of the CEEF epicentre. The aftershock was entered into the CEEF as a note, but here is proposed to have its own entry (#84b). The STJ seismograms for both mainshock and aftershock show significant Lg energy, indicating a source in continental rather than thinned continental or oceanic crust.

#85 is an earthquake found first by searching the Memorial University seismograms, and then confirmed by GSC seismograph records. By contrast to #84a and #84b, the Lg wave did not propagate efficiently to STJ, suggesting the earthquake occurred in oceanic or thinned continental crust. #85 lies farther offshore from St. John's than #84, and may reflect crustal thinning towards the edge of the Grand Banks.

#88a is similar to #85, occurs in the same general area, and also lacks Lg energy at STJ. The current epicentre is 10 km to the SW of the CEEF epicentre.

#88b occurred close (within 20 km) to another large earthquake (#71) in the vicinity of the Hibernia Oilfield. Lg-wave energy was not recorded at STJ, even though the STJ seismogram of #71 showed an Lg-wave.

#89a is a large, $M > 5$, earthquake near the Mid-Atlantic Ridge, which is included for completeness. Many earthquakes occur on the ridge to the east of the map area, and some of the larger ones give Pn and sometimes Sn energy on Atlantic Canada seismographs. These earthquakes are best located using world-wide data (we choose the ISC solution

instead of computing a GSC solution), and have no seismic hazard implications for Canada, because they are too distant to cause significant ground shaking and these mid-Atlantic Ridge earthquakes are not known to trigger tsunamis.

#89b is an earthquake on the Grand Banks. It appears to be close to the Hibernia Oilfield and is the largest earthquake in the vicinity since 1971.

#90 is only the second known earthquake from the southernmost Grand Banks. It is considerably larger (M_L 4.7) than #82, which is nearby. The m_b determined by the ISC from 8 stations is 4.3, consistent with the anticipated overestimation of magnitude by M_L .

#91 is an earthquake on the shelf east of St. John's. It is clearly closer to St. John's than the cluster of events (#71, #88b, #89b, #92a,b,d) that lie about 80 km farther east.

#92a is the third magnitude 4 earthquake to occur in a cluster about 270 km east of St. John's. This earthquake was well-recorded throughout eastern Canada, with Lg-wave energy being seen as far as Alberta. Relative arrivals at St. John's and Halifax suggest a similar location to the much larger 1971 earthquake (see Discussion).

#92b is a further earthquake in the cluster about 270 km east of St. John, the second of three in 1992.

#92c is the first earthquake known from the eastern margin of the Grand Banks, at the southern end of Flemish Pass. It is located 200 km southeast of the nearest cluster of earthquakes and has a distinctly different location from the other Grand Banks earthquakes in 1992. It lies under the continental slope, where the crustal thickness is about 10-15 km (see Discussion). Despite this, strong Lg-energy was observed on STJN.

#92d is a further earthquake in the cluster about 270 km east of St. John, the last of three in 1992.

“Note added in proof” The following four earthquakes occurred after the preparation of the figures, tables, and text of this report were completed in 1992. They are referenced only here and in the Appendix.

#94a is a M_L 3.2 earthquake on 940108 near the northeast trend of earthquakes discussed below under “Northeast Newfoundland Shelf and Basin”.

#94b is a m_{bLg} 3.1 earthquake 35 km north of St. John's on 940811. It occurred during daylight hours, but despite numerous enquiries, no blasting source was identified. #69 was a similar-sized earthquake at a similar epicentral distance, but the unclear Sg arrival at STJN and the poor distribution of seismograph stations prevents a decision whether or not both occurred in the effectively same place. A subsequent event occurred on 950122 (#95).

#94c is a m_{bLg} 2.6 earthquake on 941201, about 80 km east of St. John's, distinctly farther east of St. John's than #69 or #94b.

#95 is a small, m_{bLg} 2.4, event on 950122, 35 km from St. John's, and likely in the same place as #69 and #94b. It occurred at night, so blasting can probably be ruled out.

DISCUSSION

Revision of epicentres

Both of the two earliest earthquakes originally placed on the Grand Banks have proved to be mislocated northern events. #22 gives a poorly constrained location in central Baffin Island, but likely occurred in the northern of the two seismically-active regions along the coast of Baffin Island (and not actually in the interior where the crude epicentre places it). In that region, its magnitude is not exceptional. #62 is confirmed as a Labrador Ridge event, and is one of several magnitude 5 events on the ridge. It would likely not have been mislocated if it had occurred a year later, when FBC had begun operating.

Of the other five earthquakes in the pre-1985 CEEF (Table 1; Fig. 2), one (#71) moves 32 km, two move about 17 km, and two less than 6 km (Table 4). Fifteen new events have been added to the CEEF since 1984 (11 given in Table 3, plus 4 in 94/95). Their CEEF epicentres reflect the incorporation of our enhanced research effort into the ongoing compilation of seismicity. Past experience suggests that without this enhanced effort some of these earthquakes would have been missed, and the routinely-determined epicentres for others would have needed revision.

Four earthquakes between 1975 and 1982 are reported for the first time, having been found from examination of the Memorial seismograms. During the same period only one earthquake (#76) had been located in the study area during the routine analysis. It is disturbing that three of the new earthquakes are larger than M_L 4 (4.0 to 4.2) and the closest (#81) is only 380 km from STJ (but 1280 km from the next closest seismograph of the Canadian Network, HAL). Although the amplitude on STJ was easily detectable, #81 gave less than 0.8 mm amplitude on all the other GSC seismograms.

Detection and location thresholds, and rates of activity

Now that the re-examination of the Memorial seismograms is complete it seems likely that since 1977 almost all earthquakes on the Grand Banks have been **detected** if they exceed the magnitude thresholds for the following radial distances east of St. John's:

350 km	3.0	m_N
550 km	3.5	M_L
900 km	4.0	M_L

However, for a **location** requiring, say, HAL or SCH, the threshold has been considerably higher:

350 km	4.0	m_N
550 km	4.2	M_L
900 km	4.6	M_L

Note that these thresholds are subjective estimates of a 90% confidence ability, based on the appearance of past earthquakes on the seismograms together with the attenuation rate implied by the M_L formula. A more rigorous analysis is intended in the context of a Canada-wide assessment of detection capability.

The location capability improved in 1983 with the opening of station GBN in Nova Scotia, and improved slightly for the period 1981–1991 because of the operation of four ECTN stations in New Brunswick. However, it is clear that smaller earthquakes on the inner Grand Banks, and larger earthquakes along the outer continental margin are not all being detected, and it is difficult to estimate how many are being missed, since the known annual rate of earthquakes in a particular magnitude range and in any given region is highly variable.

In addition, for the Grand Banks, it is difficult to calculate the annual probability of earthquakes larger than a given size because the location capability varies so dramatically in space and has varied also with time. If, however most of the magnitude 4 and greater earthquakes that occurred within 400 km of St. John's since 1977 are assumed to have been located, then Table 4 shows four earthquakes in this category between 1977 and 1994, i.e. four earthquakes in 18 years, or an average rate of 0.2 earthquakes of magnitude 4 and greater per annum on the inner Grand Banks. Since the located number is so small, then the annual rate is only a very rough estimate. If even two earthquakes in this category had been missed, then the annual rate would be 50% higher. Since there are even fewer earthquakes located with magnitudes greater than 4.5, they provide no direct basis for an annual rate of higher magnitude earthquakes east of St. John's. To make such estimates requires assumptions about the relative rates of big and small earthquakes; such assumptions are commonly made for input to seismic hazard calculations.

Lg Attenuation and propagation

There are numerous references in the brief earthquake descriptions above regarding the extent to which Lg-wave energy has propagated from the earthquake to St. John's. Lg-wave energy is known to be very sensitive to crustal structure, propagating very efficiently in normal continental crust but very poorly in oceanic or thinned-continental crust.

Figure 8 uses thick, moderate, and thin lines to show the paths with excellent, weakened, and zero Lg-wave propagation to St. John's. There are still too few paths to fully delimit the region of normal crust, though the 20-km contour of crustal thickness (see Fig. 10) appears to divide the earthquakes correctly. The difference in Lg propagation between #88b and that from the nearby earthquakes could be due to local crustal variations or to earthquake depth, but must be considered unexplained at this time.

Earthquakes with significant Lg-wave energy at St. John's commonly also show Lg-wave energy at Schefferville, but almost never in Nova Scotia or southern New Brunswick. The later path contains oceanic-thickness crust south of Newfoundland that attenuates the Lg-wave.

Focal Mechanisms (J. Adams has sole responsibility for the contents of this section)

None of the earthquakes discussed in this report is large enough to determine a reliable focal mechanism at present. This is mostly due to the extremely poor station distribution for these offshore earthquakes, as #71 would have been easily large enough for a reliable mechanism, had it occurred onshore.

At m_{bLg} 4.8, #71 was large enough to be recorded by 84 seismographs reporting to the International Seismological Centre. Thirteen (3 being Canadian) reported polarities. Both Adams and Wahlström independently read all Canadian seismograms, confirmed the 3 Canadian readings reported to the ISC, and added 10 more (Table 4). Figure 9A shows the distribution of polarities. Of particular note is that STJ and HAL have opposite polarities, despite having similar azimuths from the epicentre. Figure 9B shows a range of mechanisms that fit the polarity data acceptably well (each misfits the equivalent of 4 full-weight polarities) without misfitting either STJ or HAL. These mechanisms, which represent almost pure strike-slip faulting on north-south or east-west planes, misfit our Canadian compressional readings at OTT, GWC, BLC, SES, and FFC and the ISC-reported polarity at ELT. Of these only the misfit of the clear polarity at FFC is of concern, but as the reported phase arrives 5 s late for GSC’s solution, it may well be a depth phase rather than the direct P-wave arrival.

Two points to note about this poor mechanism are:

- pure strike-slip faulting is very uncommon in onshore southeastern Canada.
- the P-axis is oriented NW-SE, at a high angle to the “regional” stress field orientation which is NE-SW (Adams and Bell, 1991, Figure 3). However, local stress orientations estimated from oil-well breakouts in the Hibernia oilfield (70 km to the southeast) include anomalous east-west and north-south compression directions (Adams and Bell, 1991, Figure 16).

I suggest that future modelling (for example, of the Pnl wave on STJ and the teleseismic P-waves at Canadian Arctic stations) should be carried out to test my tentative mechanism and also to establish the depth of the 1971 earthquake.

Polarities from other earthquakes in this report are given in the PIK files (Appendix). Worthy of note is that STJ and MUNF are dilatational for Pn arrivals and HAL and GBN are

usually compressional, and that both compressional arrivals at STJ are for upward-directed arrivals from the close events #69 and #94b.

Seismotectonics

Figure 10 shows the relocated earthquakes together with bathymetry and crustal thickness contours taken from Shih et al. (1988). We discuss the earthquakes in five groups.

Mid-Atlantic Ridge. Event #89a lies in the extreme northeast of the study area, near the Mid-Atlantic Ridge. These earthquakes have no seismic hazard implications for Canada.

Newfoundland Ridge. Two earthquakes (#82 and #90) lie south of the Tail of the Bank, on the Newfoundland Ridge. Little is known of the seismotectonics of this region, though the pair of earthquakes might be associated with the rifted continental margin (Adams and Basham, 1989) and might represent reactivation of the extensional faults formed during the opening of the Atlantic Ocean.

Newfoundland Rifted Margin. A single earthquake (#92c) is known to have occurred under the continental slope that marks the southeastern edge of the Grand Banks. However, few earthquakes smaller than M4 can be located in this area. It occurred in a region of continental crust somewhat thinned by the rifting that formed the Atlantic Ocean. Although this is the first known earthquake from the 1000-km-long rifted margin that bounds the Grand Banks to the southeast and southwest, seismicity of the margin was predicted by the ‘ESX’ seismicity model of Basham and Adams (1983) and Basham et al. (1983). Their model suggested that the entire rifted margin would be capable of generating large ($M \approx 7$) earthquakes. By implication, either the entire margin would need to be seismically-active at a low level (below the detection threshold) or earthquakes were occurring at long intervals, so that contemporary seismicity (which at the time was absent) might not reveal all active structures. Therefore #92c is important supporting evidence for the ESX model, which itself is important because of the seismic hazard large earthquakes on the margin would pose to facilities producing hydrocarbons from the Jeanne D’Arc Basin.

Grand Banks Proper. Seven earthquakes (#71, #88b, #89b, #91, #92a, #92b, and #92d) lie on the Grand Banks in the vicinity of the Hibernia Oil Field. #91 lies about 70 km to the west of the other six, but the azimuthal precision is such that the remaining six earthquakes (except, perhaps #89b) could well have occurred in the same place (Fig. 7). Although not conclusive, an analysis of relative arrival times for Nova Scotia and Quebec stations (representing the maximum azimuthal spread of the data) suggests a NNE-SSW to N-S alignment of the epicentres. This possible alignment roughly parallels the structural trends on the central Grand Banks.

The epicentres lie on the Bonavista Platform, about 20 km west of, but parallel to, the Mercury Fault (Adams and Bell, 1991, Figure 16). In view of the potential location bias (due to the unknown crustal velocity structure), it is not impossible that they actually occurred on the Mercury Fault. Alternatively, they may have occurred on a less-significant or an unknown fault within the Bonavista Platform. The Mercury Fault is an east-dipping listric detachment which bounds the northwest side of the Jeanne D'Arc Basin, the largest of a number of northeast-trending basins that cut across the Grand Banks. These basins were formed by normal faulting of the basement consequent on crustal stretching during the opening of the Atlantic Ocean (Ziegler, 1988).

This cluster of activity represents the most active region on the Grand Banks, having generated three M4, and four M3 earthquakes. Allowing for different periods of completeness, the recorded earthquakes suggest an annual rate for $M \geq 4.0$ earthquakes of about 0.1, which makes the cluster about a third as active as the Laurentian Slope zone south of Newfoundland (Keen et al., 1990; Adams, unpub.).

East Newfoundland Basin and Shelf. One earthquake (#81) lies south of Orphan Knoll and 200 km northeast of the cluster of earthquakes near Hibernia. The remaining nine earthquakes lie along a northeast trend extending from just off St. John's (#69) to 52°N, 47°W. The earthquakes in this trend lie in a band about 70-100 km wide, and might extend southwest to include earthquakes off Bonavista in 1965 and on the northern Avalon Peninsula in 1884 and 1956 (Adams and Staveley, 1985).

The earthquakes near the middle of the trend (#76, #77, #84a,b, #85, #88a) lie in a region where the basement surface has a local relief of 2-4 km and is formed into linear, northeast-trending ridges (Grant, 1988). These are interpreted as the result of

strong block faulting of the Paleozoic basement in Jurassic times (Proctor et al., 1984, p. 39).

The two most northeasterly earthquakes (#70 and #78) occurred on oceanic crust close to the inshore extension of the Charlie-Gibbs Fracture Zone. Smaller earthquakes could seldom be located in this area. In this region the fracture zone trends east-northeast. In the Labrador Sea to the north, Adams and Simmons (1991) have found that many of the earthquakes on the oceanic crust occur in the vicinity of the major fracture zones.

In summary, the seismicity discussed in this report may be due to fracture zones weakening the oceanic crust around the Grand Banks, or to northeast-trending normal faults and rift features that broke the integrity of the basement in the Mesozoic during the opening of the Atlantic Ocean. At this point in our understanding there are no correlations with individual faults, only spatial associations with broad zones of weakness. These conclusions are not very different from those of Adams and Basham (1989).

Seismic Hazard Implications

Earthquakes on the Grand Banks have significant seismic hazard implications for offshore hydrocarbon production platforms such as will be emplaced at Hibernia or elsewhere on the Grand Banks. Some earlier hazard estimates included the two largest earthquakes (1922 and 1962) in the hazard calculations, but these earthquakes have now been relocated away from the Grand Banks. The largest earthquake is now m_{bLg} 4.8, but more importantly it is now seen to lie within an active cluster, revealed chiefly because of activity in the past five years.

The problems with detection and location thresholds (discussed above) make it difficult to assess the rates of earthquakes and hence the hazard for the offshore. Previous hazard estimates (Basham et al., 1983; Adams, 1986) recognized the poorly determined nature of the offshore seismicity and the need to relocate the older earthquakes and to improve seismograph coverage to detect more of the small offshore earthquakes. Some relatively inexpensive additions to the onland seismograph network – for example a station near Bonavista, Cape Freels, or St. Anthony – would certainly improve the detection and location of offshore

earthquakes. Although expensive, and technically difficult to install and operate, a tethered ocean bottom seismometer in the vicinity of the Hibernia platform would present the best near-term opportunity to improve seismic monitoring on the outer Grand Banks. One further conclusion of the present report is that effort needs to be made to assess the earthquake location threshold in both time and space so that the seismicity record can be interpreted in terms of recurrence rates for the entire Grand Banks.

CONCLUSIONS

1. The two largest earthquakes once considered on the Grand Banks have been relocated out of the area. The largest earthquake is now m_{bLg} 4.8.
2. Four new earthquakes between 1977 and 1985 have been found, the epicentres of five previously-known earthquakes prior to 1985 have been revised slightly, and post-1984 records have been searched for earthquakes with more care, resulting in a further 15 earthquakes being added.
3. Most earthquakes appear to be spatially associated with Mesozoic extensional structures that trend northeastwards across the Grand Banks, although another decade of monitoring may be needed to test this conclusion.
4. A cluster of activity northwest of the Hibernia site represents the most active region on the Grand Banks in recent years.
5. The earthquakes catalogued in this report form the most reliable basis for seismic hazard calculations for the Grand Banks.

ACKNOWLEDGEMENTS

We thank Kurt Kennedy (Memorial University of Newfoundland Engineering COOP student, 1985) for the initial scanning of the MUN records and Kerry Fagan (Memorial University Engineering COOP student, 1992) for setting up some of the computer-generated maps. Professor Jim Wright at Memorial University loaned the early Memorial University seismograms and coordinated the excellent post-1983 operation of the STJN and CBK. R.W.'s stay at the Earth Physics Branch (subsequently the Geophysics Division, GSC)

was financed by the Swedish Natural Sciences Research Council, Contract G-PD 4012-100. Reviews by Allison Bent and Stephen Halchuk were appreciated, as were additional comments by Anne Stevens. Continued funding of the eastern offshore aspects of the GSC's seismology program by the Panel for Energy Research and Development made possible the collection of the earthquake records and their interpretation in this report.

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TABLE CAPTIONS

Table 1. 1922 to 1984 earthquake epicentres, as known in 1985.

Table 2. Probable earthquakes recorded on Memorial seismographs, but unlocatable. (possible Grand Banks earthquakes)

Table 3. A. New solutions for earthquakes no longer in the study area. B. Revised, new, & added solutions for earthquakes in the study area.

Table 4. Summary of revisions to Grand Banks earthquakes.

Table 5. Polarity data for the 1971 Grand Banks earthquake.

FIGURE CAPTIONS

Figure 1. Map of the Eastern Margin showing the geographic scope of the two published reports and the region of the present study.

Figure 2. Seismicity of the Grand Banks as known in 1985, taken from the then-current Canadian Earthquake Epicentre File. Only earthquakes within the study area (east of the heavy line) are shown. Symbols in this and subsequent figures are:— star, magnitude (M) ≥ 5.0 ; square, $4.0 \leq M \leq 4.9$; triangle, $3.0 \leq M \leq 3.9$; and cross, $M \leq 3.0$.

Figure 3. Relocated Grand Banks earthquakes. The label next to each earthquake symbol identifies the earthquake by year (76, 81, etc) and sequence in the year (84a, 84b, etc) and reference to Table 3.

Figure 4. Relocation vectors for the studied earthquakes. The heavy lines extend back from the revised epicentre (also marked by a symbol) to the pre-revision epicentre. For most earthquakes the revisions are small, less than the symbol size, but two events have moved north, off the map area (see text).

Figure 5. Computed uncertainty for the revised epicentres, as represented by precision diamond (the axes represent the standard errors in latitude and longitude). For almost all the earthquakes, the computed precision is less than the space occupied on the map by the earthquake symbol (see also Fig. 6 and 7), though the plotted epicentres may be inaccurate by an amount at least as large as the symbol size.

Figure 6. Detail of Figure 5, showing the seismicity of the Grand Banks proper, together with the precision diamonds.

Figure 7. Detail of Figure 5, showing the epicentres in the cluster of earthquakes near Hibernia, together with their precision diamonds.

Figure 8. Propagation of Lg-wave energy to St. John's from offshore earthquakes. Thick lines represent paths with efficient propagation, single moderate-thickness line represents partial attenuation of Lg, and thin lines represent paths that completely attenuate Lg.

Figure 9. Focal mechanism information for event #71. A) Data (see Table 5). C and D represent full-weight data, c and d represent half-weight data. B) Possible nodal planes, representing strike-slip faulting on a N-S or an E-W plane. P, B, and T represent the axes of the solution.

Figure 10. Relocated offshore earthquakes, crustal thickness (taken from Shih et al., 1988), and 500-m isobath. Compare with Figure 3 to determine the identity of the events. As in previous maps, earthquakes west of the heavy line are not plotted.

TABLE 1

1922 TO 1984 EARTHQUAKE EPICENTRES AS KNOWN IN 1985

REF	DATE	TIME (UT)	LAT	LONG	DEPTH (km)	MAGNITUDE	TYPE
ISC	19220726	062855	50.0 N	50.0 W		5.3	ML
GSC	19620803	013102	52. N	54.2 W		4.8	ML
GSC	19690805	215324	47.74 N	52.31 W	18	3.3	MN
GSC	19701031	174445	52.17 N	46.37 W	18	4.3	ML
ISC	19710815	061715	47.46 N	49.53 W		4.8	MB
GSC	19760828	192330	50.10 N	48.85 W	18	4.0	ML
GSC	19840526	192549	48.89 N	51.06 W	18	3.5	MN

REF: GSC - solution from GSC, ISC - solution from ISC

TABLE 2

PROBABLE EARTHQUAKES RECORDED ON MEMORIAL
SEISMOGRAPHS, BUT UNLOCATABLE.
(POSSIBLE GRAND BANKS EARTHQUAKES)

DATE	TIME UT	STATION	S-P (s)	DIST (km)	AMP (mm)	MAGNITUDE ¹ ML
750403	0941	MUNF	47	470	2.0	3.4
760217	0658	MUNF	45	450	3.0	2.9
760327	1101	MUNF	63	645	2.5	3.8
761124	2154	MUNF	160	1710	7.0	-
770811	1048	MUNF	42	420	2.5	3.5
780618	1316	CBK	69	710	-	-
780228	0756	MUNF	75	780	2.0	3.7
790408	2132	MUNF	43	425	1.5	3.4
790712	0314	MUNF	67	690	1.5	3.6
790725	0444	MUNF	62	635		
791014	1407	MUNF	24	220	4.0	2.9
791112	1429	MUNF	32	305	6.0	3.0
800731	1609	MUNF	82	850	-	-
810121	0426	CBK	61	625	-	-
821215	0654	MUNF	23	205	-	-
920218	0346	STJN	54	550	0.8	3.0
920925	1818	STJN	63	645	0.3	2.9

¹Magnitude assuming Sn amplitudes; '-' not determined

TABLE 3

A. NEW SOLUTIONS FOR EARTHQUAKES NO LONGER IN THE STUDY AREA.

ID ¹	ST	LAT	LONG	MAG	TIME	DATE	ERRORS	STN	PH	M	RMS	DEPTH
22	M	+70.000-	75.000	MS=5.1	0628550	26071922	00.0000.000	0.0	2	3	0	0.00N 18.00
		(Baffin Island)										
62	M	+61.027-	58.225	ML=4.8	0131042	03081962	00.0000.000	0.0	2	3	1	0.00 18.00
		(Labrador Ridge)										

B. REVISED, NEW, & ADDED SOLUTIONS FOR EARTHQUAKES IN THE STUDY AREA.

ID ¹	ST	LAT	LONG	MAG	TIME	DATE	ERRORS	STN	PH	M	RMS	DEPTH
69	R	+47.632-	52.156	MN=3.4	2153207	05081969	00.0860.110	0.2	5	9	4	0.94 18.00
70	R	+52.189-	46.333	ML=4.3	1744440	31101970	00.0310.110	0.3	2	4	2	0.29 18.00
71	R	+47.308-	49.160	MN=4.8	0617095	15081971	00.1310.215	0.2	10	17	6	2.50 18.00
76	R	+50.133-	48.910	ML=4.1	1923304	28081976	00.0260.108	0.3	5	8	3	0.54 18.00
77	N	+49.337-	49.586	ML=3.4	0805261	25031977	00.1230.155	0.0	2	4	1	0.12 18.00
78	N	+52.573-	46.752	ML=4.1	1435140	03011978	00.0280.073	0.1	3	6	3	0.36 18.00
81	N	+48.643-	47.820	ML=4.2	1306126	26091981	00.0670.170	0.2	5	10	6	1.16 18.00
82	N	+41.367-	47.282	ML=4.0	0407239	22071982	00.1040.124	0.1	6	8	4	1.05 18.00
84a	R	+48.852-	50.837	MN=3.6	1925475	26051984	00.0190.037	0.3	8	18	6	0.48 18.00
84b	A	+48.852-	50.837	MN=2.6	2204198	26051984	00.0000.000	0.0	1	1	0	0.00H 18.00
85	A	+49.787-	50.480	ML=2.8	1806035	03081985	00.0150.053	0.3	4	6	2	0.14 18.00
88a	A	+49.834-	49.982	ML=3.5	0715448	09011988	00.0180.058	0.2	8	11	6	0.34 18.00
88b	A	+47.463-	49.246	ML=3.1	0016451	09081988	00.1950.120	0.0	5	8	2	0.76 18.00
89a	A	+52.730-	35.200	MS=5.4	2253371	14051989	00.0390.030	0.0	-	-	25	0.00H 10.00
89b	A	+46.960-	49.247	ML=4.2	0515398	03121989	00.0500.088	0.4	6	9	4	0.71 18.00
90	A	+41.918-	48.341	ML=4.7	2140041	24041990	00.0270.062	0.1	14	23	13	0.67 18.00
91	A	+47.364-	50.167	MN=3.2	1123013	23071991	00.0820.146	0.2	6	10	3	0.79 18.00
92a	A	+47.245-	49.236	MN=4.0	0607283	13011992	00.0570.100	0.2	17	25	14	0.93 18.00
92b	A	+47.326-	49.346	ML=3.0	1658161	06071992	00.3710.268	0.2	3	6	3	0.74 18.00
92c	A	+46.118-	47.438	MN=3.9	0420227	17071992	00.1120.087	0.1	9	17	2	0.77 18.00
92d	A	+47.353-	49.114	MN=3.4	1131520	10081992	00.0830.070	0.0	5	9	1	0.39 18.00

¹This table is modified from GSC's CEEF format. A brief description is given by the column heads as follows: ID, see text; ST, status - M moved, R revised, N new, A added; LAT, LONG latitude and longitude in decimal degrees; MAG, magnitude; TIME, hour minute seconds in Universal Time; DATE; ERRORS, standard errors on latitude, longitude, and magnitude; STN, number of stations used for location; PH, number of phases used; M, number of stations used for average magnitude; RMS, root mean square residual (seconds); DEPTH, assigned depth (km).

TABLE 4

SUMMARY OF REVISIONS TO GRAND BANKS EARTHQUAKES

ID	SEQ	DATE	TIME	Δ LAT	Δ LONG	Δ MAG	Δ Z	STATUS	Δ KM
22	1	19220726.0628		20.000	-25.000	-0.2	18.00	moved	2852.8
62	2	19620803.0131		9.027	-4.025	0.0	18.00	moved	1040.9
69	3	19690805.2153		-0.108	0.154	0.1	0.00	revised	16.6
70	4	19701031.1744		0.019	0.037	0.0	0.00	revised	3.3
71	5	19710815.0617		-0.152	0.370	-0.1	18.00	revised	32.5
76	6	19760828.1923		0.033	-0.060	0.1	0.00	revised	5.6
77	7	19770325.0805		0.000	0.000	3.4	0.00	new	0.0
78	8	19780103.1435		0.000	0.000	4.1	0.00	new	0.0
81	9	19810926.1306		0.000	0.000	4.2	0.00	new	0.0
82	10	19820722.0407		0.000	0.000	4.0	0.00	new	0.0
84a	11	19840526.1925		-0.038	0.223	0.1	0.00	revised	16.8
84b	12	19840526.2204		0.000	0.000	2.6	0.00	added	0.0
85	13	19850803.1806		0.000	0.000	2.8	0.00	added	0.0
88a	14	19880109.0715		0.000	0.000	3.5	0.00	added	0.0
88b	15	19880809.0016		0.000	0.000	3.1	0.00	added	0.0
89a	16	19890514.2353		0.000	0.000	5.4	0.00	added	0.0
89b	17	19891203.0515		0.000	0.000	4.2	0.00	added	0.0
90	18	19900424.2140		0.000	0.000	4.7	0.00	added	0.0
91	19	19910723.1123		0.000	0.000	3.2	0.00	added	0.0
92a	20	19920113.0607		0.000	0.000	4.0	0.00	added	0.0
92b	21	19920706.1658		0.000	0.000	3.0	0.00	added	0.0
92c	22	19920717.0420		0.000	0.000	3.9	0.00	added	0.0
92d	23	19920810.1131		0.000	0.000	3.4	0.00	added	0.0

ID and SEQ represent the earthquake identification number and a simple counting sequence, respectively; Δ LAT, Δ LONG, Δ MAG and Δ Z represent the changes to these parameters established in this Open File; STATUS describes the reason for the change: moved - moved out of the area; revised - existing solution improved; new - event found on old MUN seismograms; added - event added from enhanced reading of MUN and GSC seismograms; Δ KM represents the distance moved in kilometres.

TABLE 5

POLARITY DATA FOR THE 1971 GRAND BANKS EARTHQUAKE

ID	AZ	TO	POL	SOURCE
STJ	274.000	49.000	D	ISC and Wahlstrom
SCH	311.000	49.000	-	Adams -, ISC D, Wahlstrom C
MNT	272.000	46.000	-	ISC and Adams
CLE	267.000	33.000	D	ISC
FAV	267.000	29.000	C	ISC
DOU	65.000	29.000	C	ISC
NOR	8.000	28.000	C	ISC
GOL	280.000	28.000	D	ISC
ILT	340.000	23.000	D	ISC
BNG	103.000	19.600	D	ISC
ELT	26.000	19.200	D	ISC
AAB	37.000	18.000	C	ISC
GAR	44.000	18.000	C	ISC
HAL	261.000	49.000	C	Wahlstrom and Adams good quality, not strong
SIC	290.000	49.000	-	Wahlstrom C, Adams -
OTT	274.000	42.000	+	very weak, noisy record
UNB	270.000	49.000	-	clear, not strong, Wahlstrom D
SFA	277.000	49.000	D	good
GWC	305.000	39.000	+	v weak on noisy record
SUD	280.000	35.000	E	low signal strength, emergent
BLC	321.000	33.000	+	Adams 900925
FFC	303.000	32.000	C	Adams 900925 good polarity but 5 sec late
SES	298.000	30.000	+	Adams 900925
INK	329.000	29.000	D	Adams 900925
MCE	303.000	29.000	-	ISC

ID, Station providing the reading; AZ azimuth from earthquake to station; TO take-off angle from earthquake to station; POL P-wave first motion:- C and D are compression and dilatation; + and - are half-weight C and D.

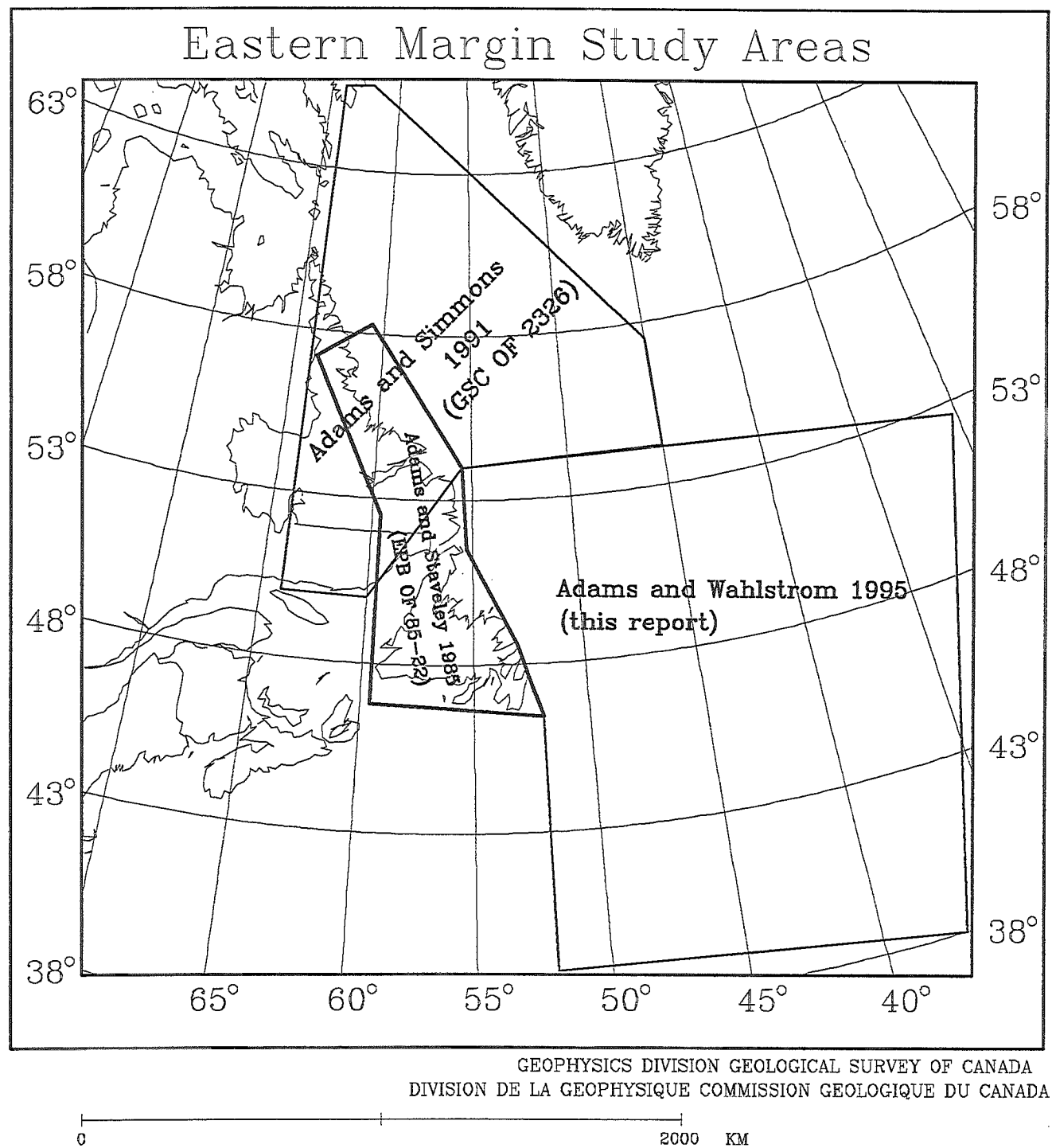
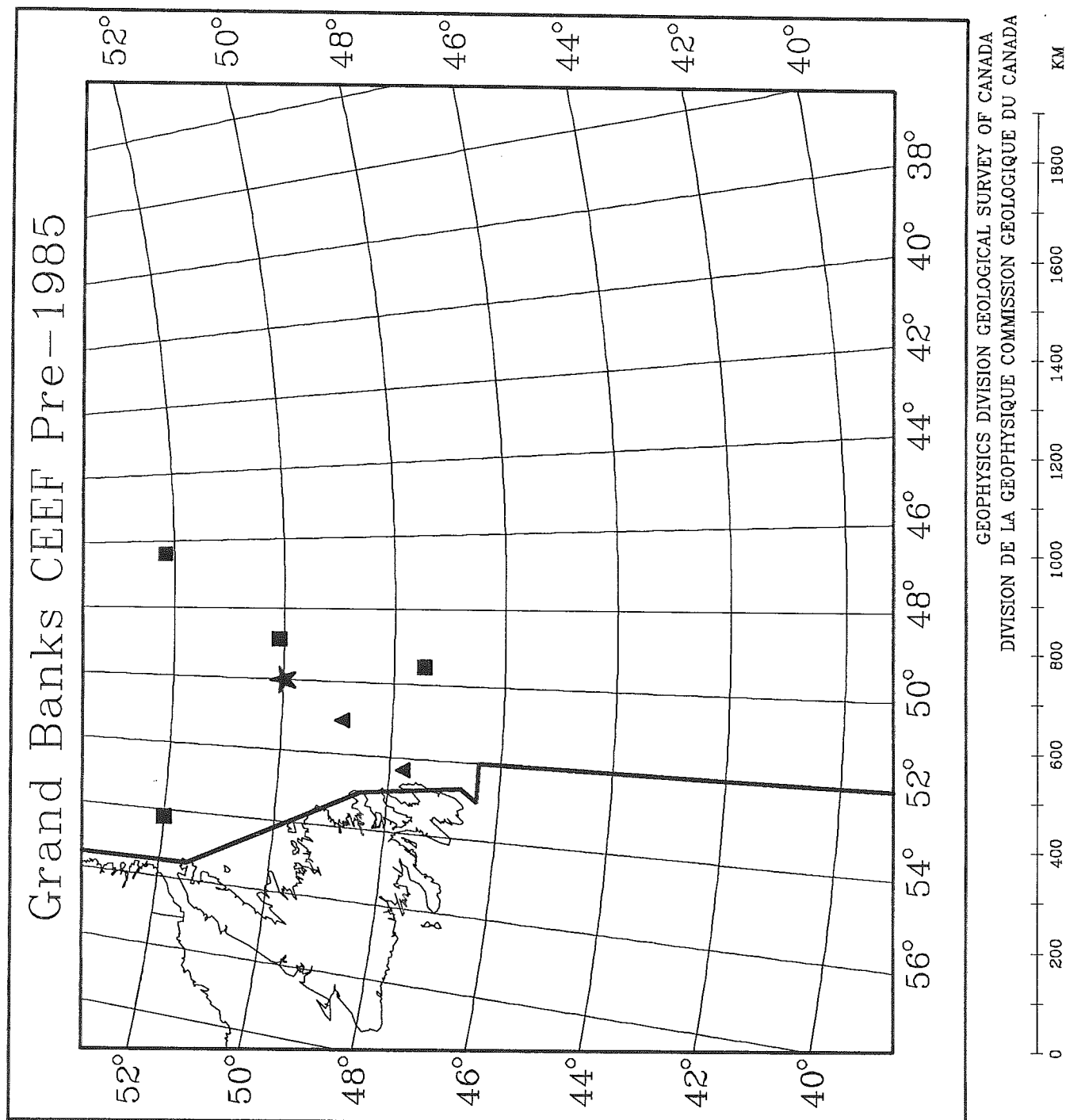


Figure 1. Map of the Eastern Margin showing the geographic scope of the two published reports and the region of the present study.

Figure 2. Seismicity of the Grand Banks as known in 1985, taken from the then-current Canadian Earthquake Epicentre File. Only earthquakes within the study area (east of the heavy line) are shown. Symbols in this and subsequent figures are:— star, magnitude (M) ≥ 5.0 ; square, $4.0 \leq M \leq 4.9$; triangle, $3.0 \leq M \leq 3.9$; and cross, $M \leq 3.0$.

DEFINITIONS

$M < 3$	×
$3.0 \leq M \leq 3.9$	▲
$4.0 \leq M \leq 4.9$	■
$M \geq 5.0$	★



Relocated Grand Banks Earthquakes

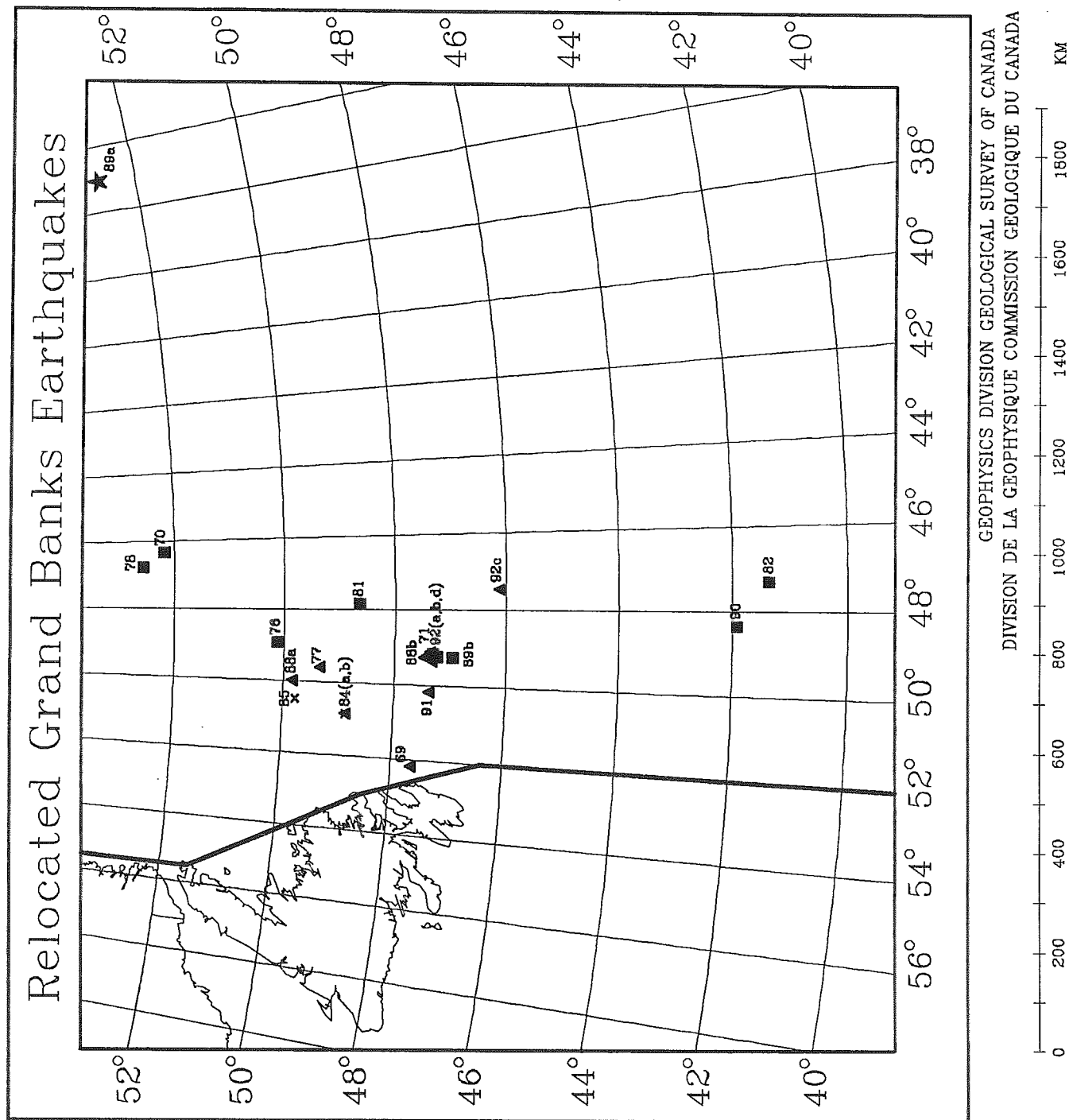


Figure 4. Relocation vectors for the studied earthquakes. The heavy lines extend back from the revised epicentre (also marked by a symbol) to the pre-revision epicentre. For most earthquakes the revisions are small, less than the symbol size, but two events have moved north, off the map area (see text).

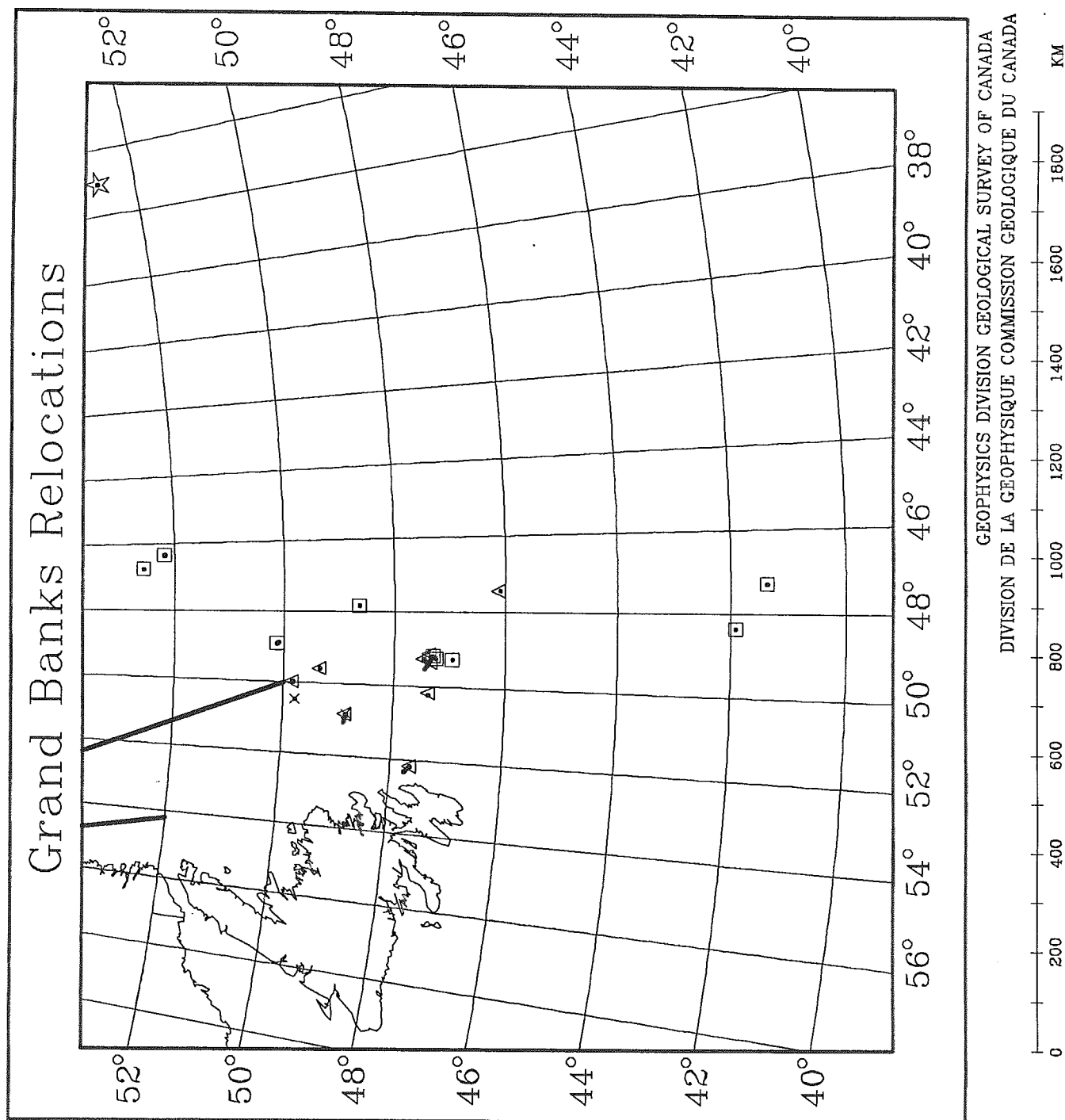


Figure 5. Computed uncertainty for the revised epicentres, as represented by precision diamond (the axes represent the standard errors in latitude and longitude). For almost all the earthquakes, the computed precision is less than the space occupied on the map by the earthquake symbol (see also Fig. 6 and 7), though the plotted epicentres may be inaccurate by an amount at least as large as the symbol size.

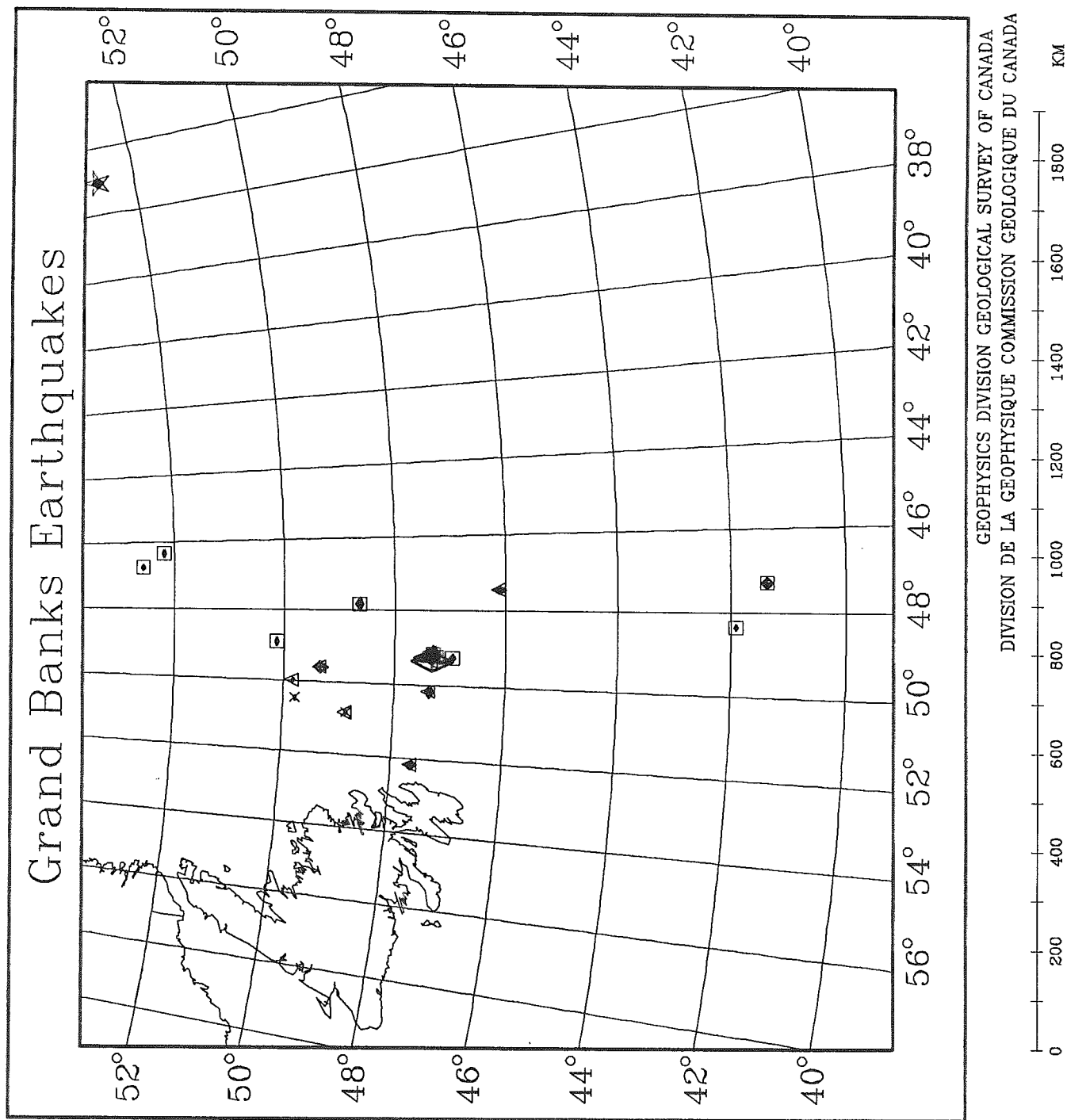


Figure 6. Detail of Figure 5, showing the seismicity of the Grand Banks proper, together with the precision diamonds.

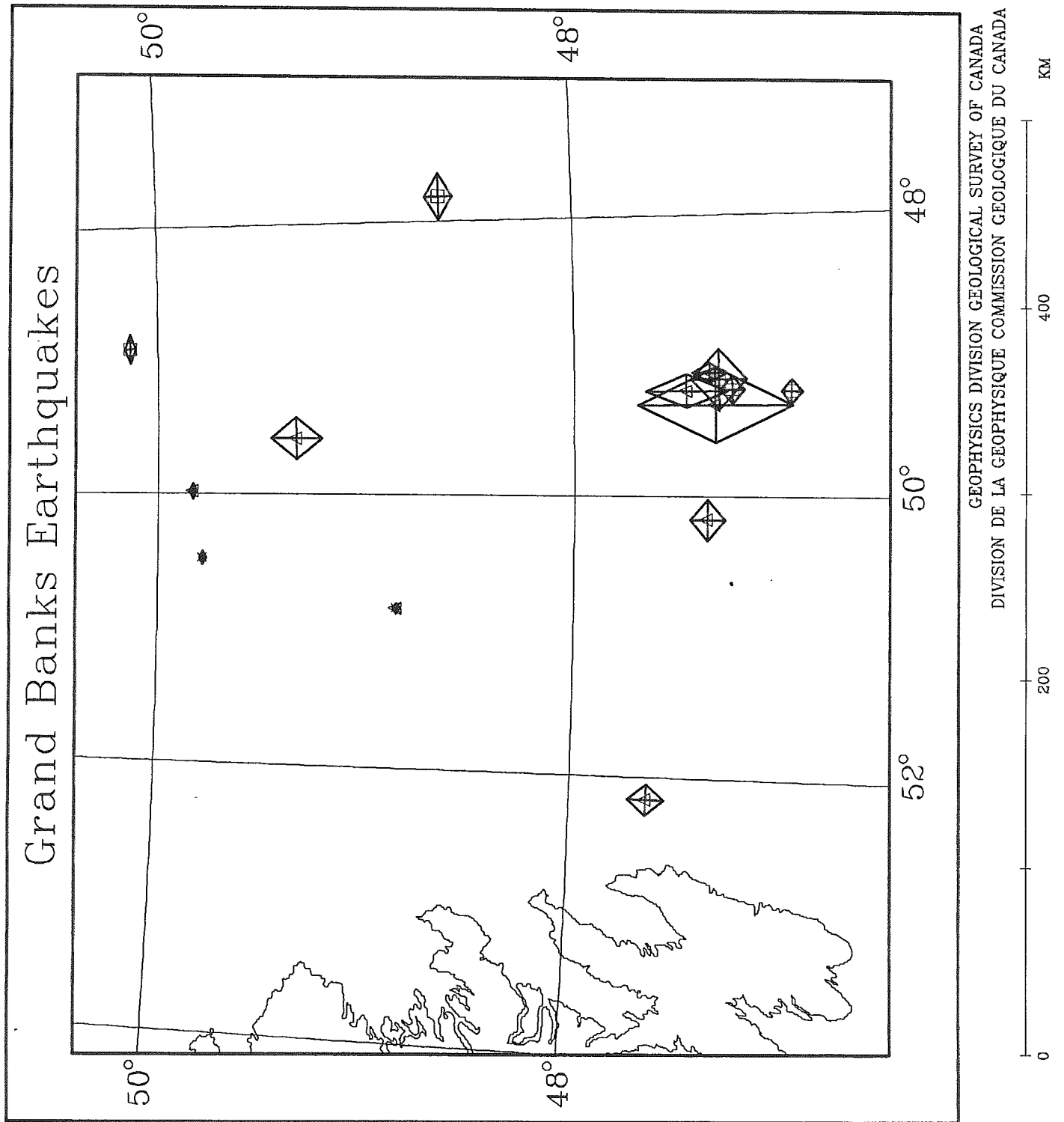


Figure 7. Detail of Figure 5, showing the epicentres in the cluster of earthquakes near Hibernia, together with their precision diamonds.

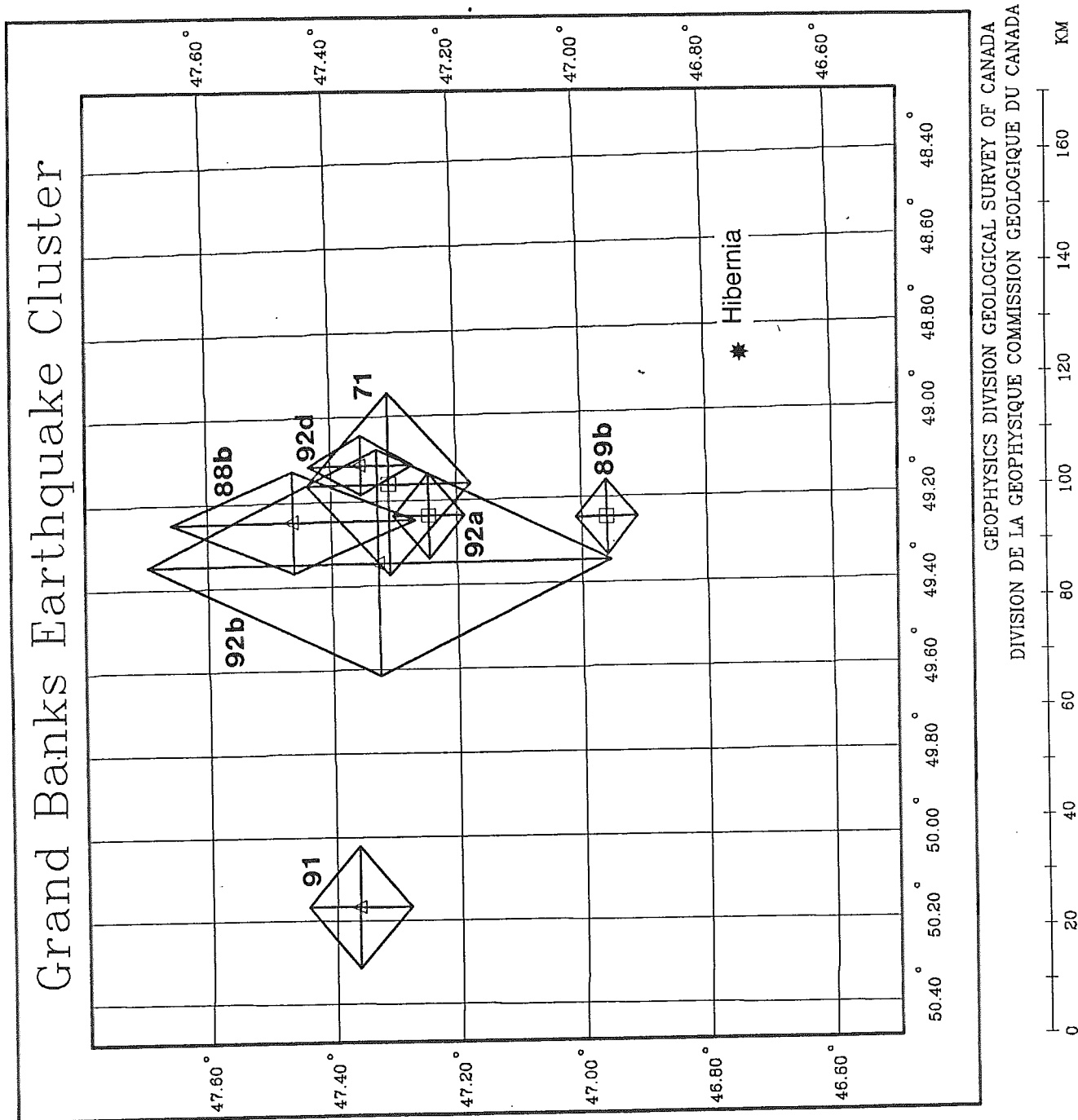
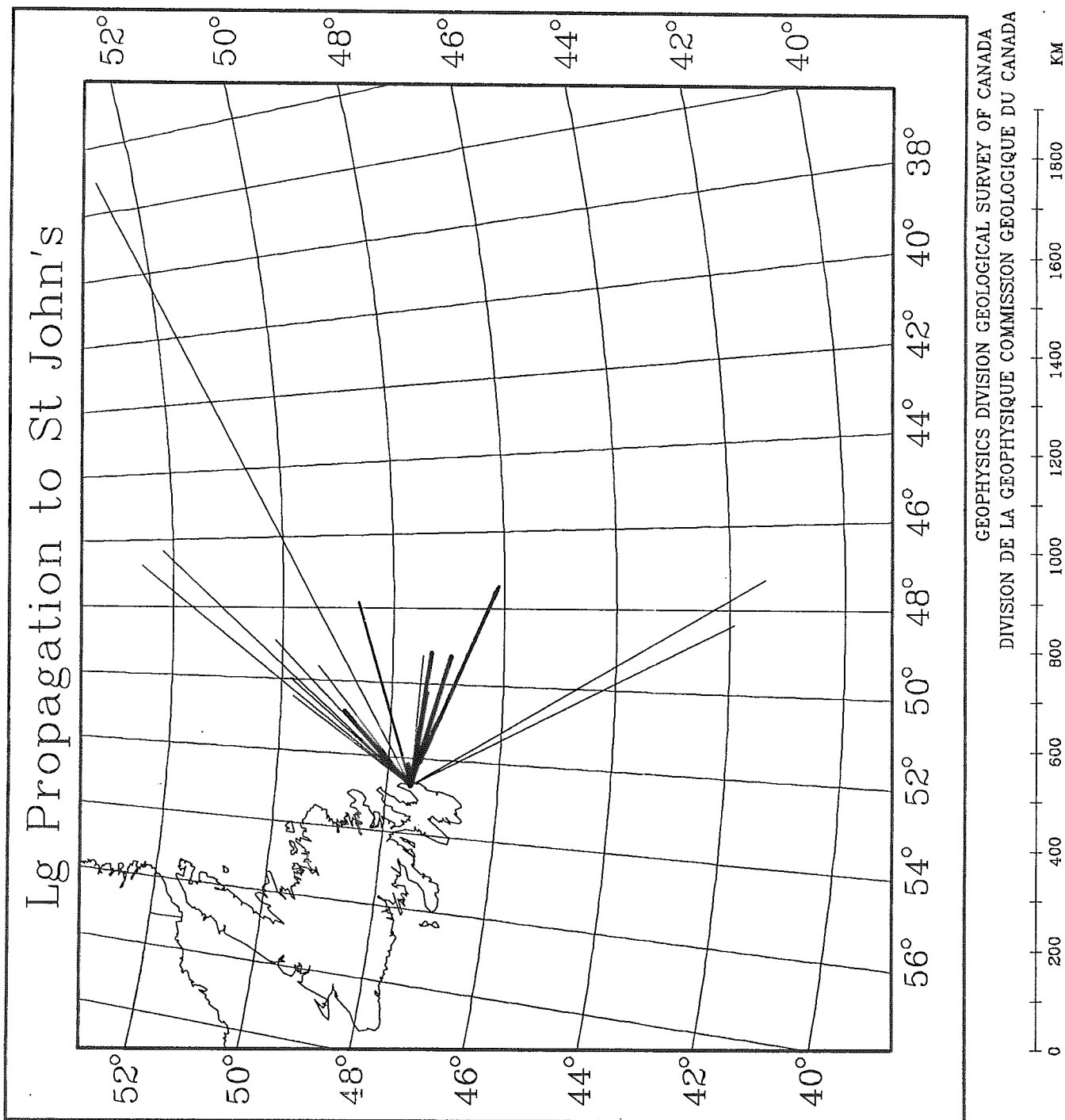


Figure 8. Propagation of Lg-wave energy to St. John's from offshore earthquakes. Thick lines represent paths with efficient propagation, single moderate-thickness line represents partial attenuation of Lg, and thin lines represent paths that completely attenuate Lg.



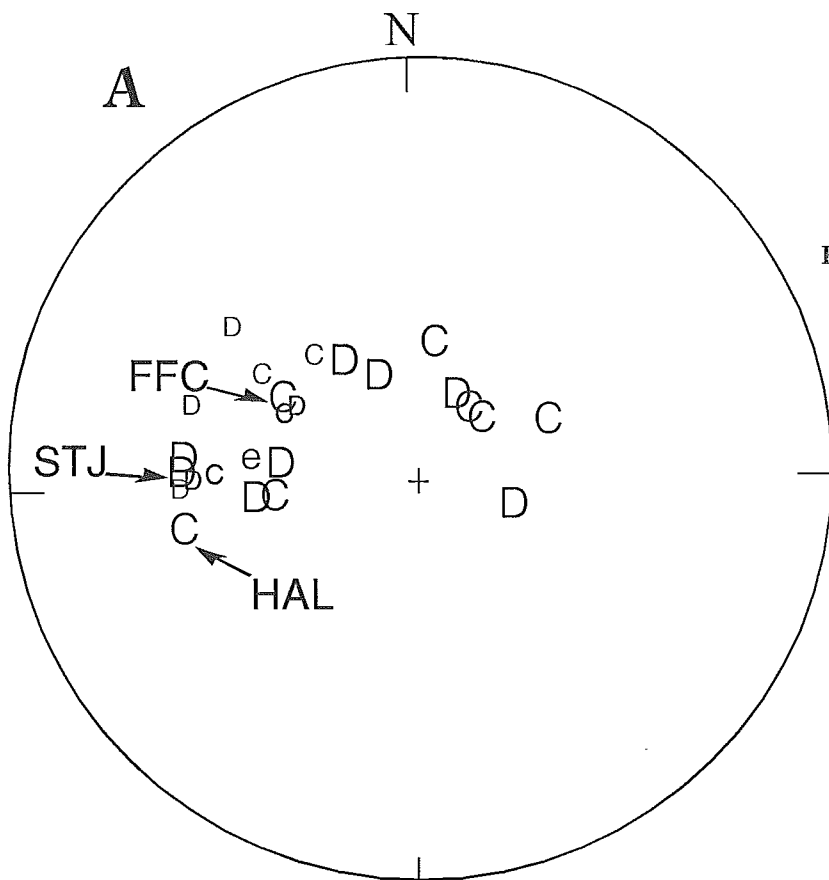


Figure 9. Focal mechanism information for event #71. A) Data (see Table 5). C and D represent full-weight data, c and d represent half-weight data. B) Possible nodal planes, representing strike-slip faulting on a N-S or an E-W plane. P, B, and T represent the axes of the solution.

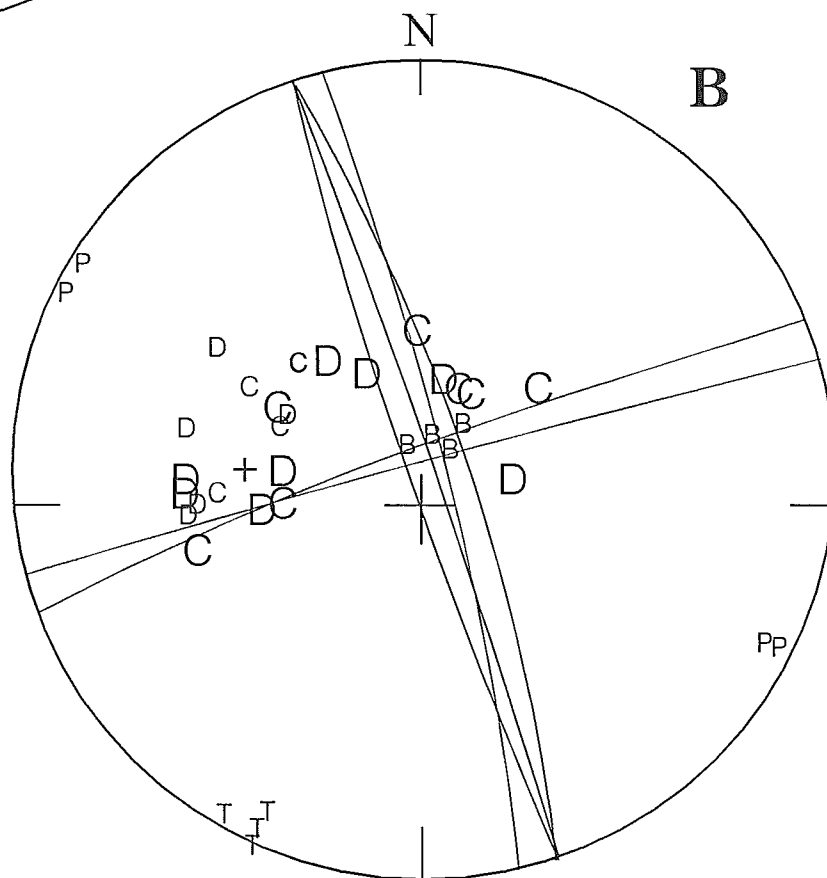
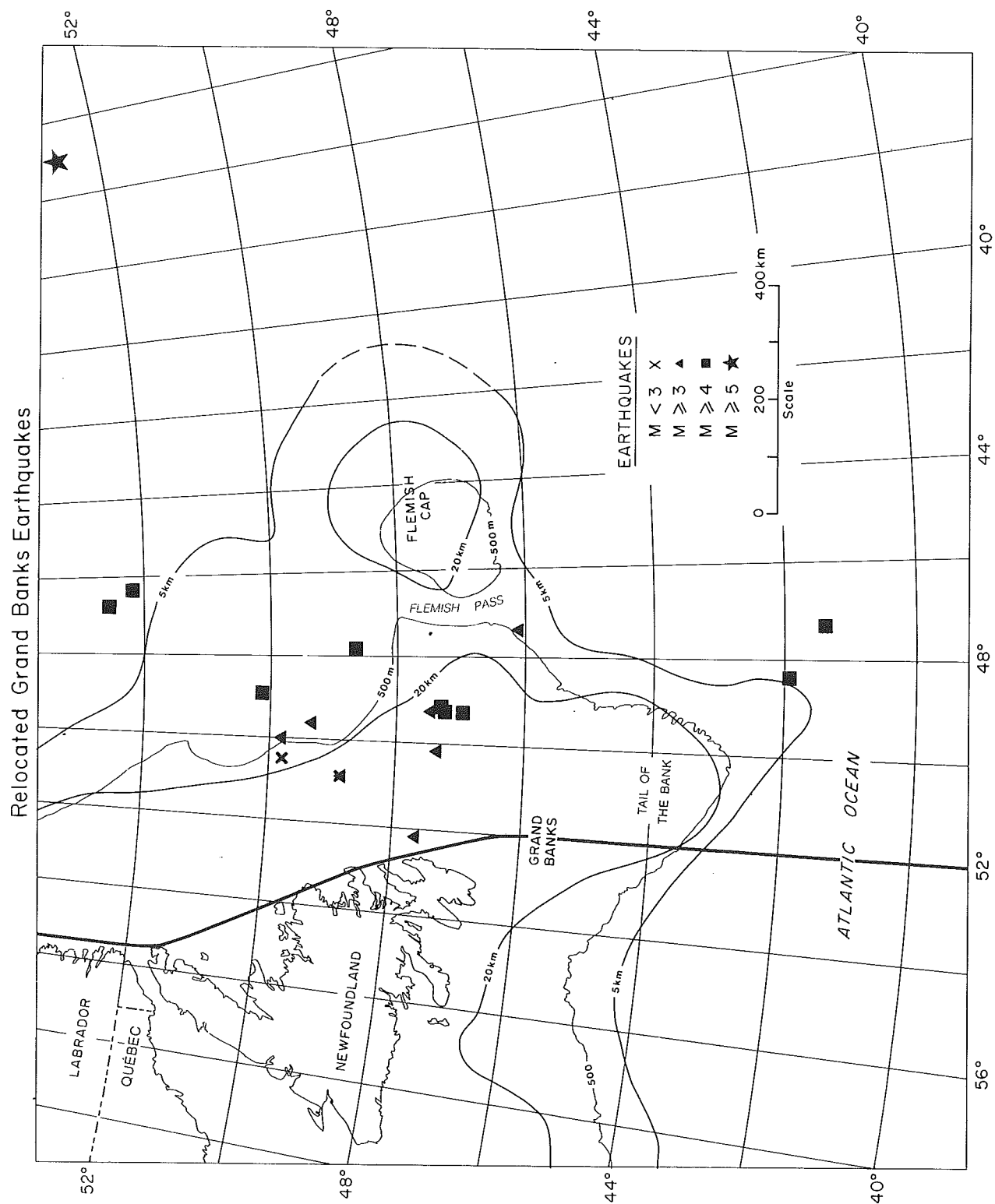


Figure 10. Relocated offshore earthquakes, crustal thickness (taken from Shih et al., 1988), and 500-m isobath. Compare with Figure 3 to determine the identity of the events. As in previous maps, earthquakes west of the heavy line are not plotted.



APPENDIX

Event files in this Appendix are arranged chronologically. We have tried to provide a full documentation of our data and results, so for each earthquake there is a "PIK" listing of the earthquake phase data used to compute the epicentre (a description of the format is given on the final pages), together with comments on the data and epicentre.

22

```
+70.000- 75.00001MS=5.1 0628550 26071922 00.0000.000 0.0 2 3 00.00N218.00 0 1ML=0.0 00 0L3.60
$50.0 - 50.0 ML=5.3 063108 CEEF
BAFFIN ISLAND REGION
$ ORIGINALLY PLACED BY ISS AT 50N 50W
$ EPICENTRE PEGGED AT LOCATION FOUND ROUGHLY BY WAHLSTROM
$ SEE GSC OPEN FILE 3043 BY ADAMS AND WAHLSTROM
$ DOUBTFUL THAT ANY BETTER CAN BE DONE WITHOUT THE ORIGINAL RECORDS
$ SAS OPERATING BUT RECORD NOT KEPT
$ HAL READINGS BY WAHLSTROM FORCES EPICENTRE TO NORTH
$ CHI, WAS, ESK READINGS FROM ISS, REPRESENT RAYLEIGH WAVE MAXIMUM
$ SEE TEXT OF GSC OPEN FILE 3043 FOR DISCUSSION
$ OTT P TAKEN WITH ASSUMED 1 MIN READING OR REPORTING BLUNDER FITS HAL S IF PHASE IS FIRST S
$ LOCATION FROM HAL P AND SN, OTT SN AND (P-1MIN) IS:
$68.837- 73.76301MS=5.1 0629094 26071922 00.3742.100 0.0 3 4 02.52 218.00 0 1ML=0.0 00 0 3.62
$ GSC DOES NOT CURRENTLY HAVE A TELESEISMIC LOCATION PROGRAM TO ALLOW
$ FITTING THE DISTANT PHASES
$
OTT 2207260631P X3514 3837 X4006
OTT S 2744KM 00 6012$181 35 00-1148$00-9123$ 0000000 00ML00MN
HAL 2207260631P 3433 3905 X3905
HAL S 2899KM 00 622 161 34 00-1646$00-9999$ 0000000 00ML00MN
CHI 2207260631P X4215
CHI S 3215KM 200 33 00-9307$ 0000000 00ML00MN
WAS 2207260631P X4400
WAS S 3468KM 183 33 00-5841$ 0000000 00ML00MN
ESK 2207260631P X4400
ESK E 3748KM 078 32 00-9999$ 0000000 00ML00MN
Z
```

62

```
+61.027- 58.2250 ML=4.8 0131042 03081962 00.0000.000 0.0 2 3 10.00 218.00 0 1MN=0.0 00 0 3.65
$+52 -54.2 ML=4.8 CEEF AFTER SMITH
$ LABRADOR RIDGE (PROBABLY)
$ NEEDS GREENLAND STATIONS AND RES, MBC, ALE, PNT.
$
$ SMITH'S COMMENTS
$ SCH CLEAR PN AND SN; NO LG - OCEANIC EVENT
$ HAL (SN) TRACE "VERY SMALL AND MAY NOT BE ASSOCIATED".
$ WESTERN EPICENTRE NOT POSSIBLE BECAUSE OF LACK OF RECORDS ON SOUTHERN STATIONS.
$
$ WAHLSTROM COMMENTS:
$ HAL: RE-READ AT LAMONT, LARGE T.C. USED TO REPORT PHASE IN BULLETIN
$ CONFIRMED WITH TELESEISM, CHECK AGAIN FOR S.
$ SCH: ARR. TIMES AND AMP/PER FROM SMITH CONFIRMED (ROUGHLY) BY MICROFILM,
$ RECORD NOT FOUND? (CHECK)
$ SFA: MICROFILM NOT FOUND, RECORD NOT FOUND, INTERMITTENT OPERATION
$ SHF: NO SIGNAL FROM MICROFILM, RECORD NOT FOUND? (CHECK)
$ OTT: STATION DISCONNECTED.
$ MNT: POSSIBLE S PHASE AT 3904 ON EW, NOT ON NS. 44 SEC TOO LATE FOR SN
$
$ ADAMS COMMENTS
$ NO TRACE OF SECOND (EARLIER OR LATER) PHASE ON HAL
$ HAL PHASE RE-INTERPRETED AS PN, PLACES EPICENTRE ON LABRADOR RIDGE.
$ CURRENT EXPERIENCE SUGGESTS SMITH EPICENTRE JUST N OF NEWFOUNDLAND
$ SHOULD HAVE GENERATED LG ON SCH
```

\$ RIDGE LOCATION IS CONSISTANT WITH OCEANIC PATH FOR SCH.
 \$ RIDGE IS LOCUS OF MANY EARTHQUAKES; SMITH'S LOCATION IS HIGHLY UNUSUAL.
 \$ ABSENCE OF SN ON HAL IS STILL PUZZLING.
 \$ COMPARE WITH ML=5.0 EVENT ON 621202.
 \$ COMPARE WITH ML=5.1 EVENT ON 621026. THIS EVENT ALSO LACKS HAL SN READING
 \$ THOUGH AS IT WAS LARGER HAD PN ON SHF AND RES TO CONFIRM NORTHERN EPICENTRE.
 \$
 \$

SCH	6208030132P	32545		3416		30 175 150	8 1
SCH	SW 0857KM 10	000 220 49		10 000		0001795 48ML42MN	
HAL	6208030132P	3456					1
HAL	S 1860KM 10	000 193 50				0000000 00ML00MN	

Z

69

+47.632- 52.156FLMN=3.4 2153207 05081969 00.0860.110 0.2 5 9 40.94 218.00 0 1ML=0.0 00 0 3.62
 \$47.74 - 52.31 MN=3.3 215324 CEEF
 \$ OFF ST JOHN'S
 \$ ST JOHN'S NEWSPAPERS DO NOT REPORT IT AS FELT, DESPITE CLOSENESS
 \$ COMMENT FROM COMPUTER CARDS: FM IS CNW - EVENT FROM ALMOST DUE EAST
 \$ HAL: LG USED FOR MAGNITUDE, MAY BE ATTENUATED
 \$

STJ	6908052153P		A53288 C		A53333		1
STJ	W 0044KM	261-68	03 036		08 -059	0000000 00ML00MN	
HAL	6908052153P	B5523		B5652	X5746 040 66 6		1
HAL	W 0944KM 03	142 254 49		01 089	00 446 0000143 40ML32MN		
SIC	6908052153P	A55421			B58275 050 106 10		1
SIC	W 1105KM 18	088 290 49			03 154 0000119 42ML32MN		
SCH	6908052153P	A56032			B5922 050 72 11		1
SCH	NW 1294KM 32	-117 314 49			20 365 0000192 46ML36MN		
SFA	6908052153P	B56190					1
SFA	W 1409KM 01	071 275 49			0000000 00ML00MN		
GWC	6908052153P						1
GWC	NW 1959KM	305 47			XB6224 090 35 4		

Z

70

+52.189- 46.33301ML=4.3 1744440 31101970 00.0310.110 0.3 2 4 20.29 218.00 0 1MN=0.0 00 0 3.62
 \$ 52.17 46.37 ML=4.3 174445 CEEF
 \$ GRAND BANKS
 \$ ORIGINAL COMMENTS: NO RECORDS FOR SIC; NOT AT HAL, FBC, SFA, GWC
 \$ WAHLSTROM COMMENTS: NOT ON HAL,OTT,GWC,FBC,SFA,SUD
 \$ SIC CLOCK WRONG
 \$ NO SIGNAL ON HAL EXCLUDES THE OTHER POSSIBLE SOLUTION: 43.3N 59.3W
 \$ ML ASSUMED - EARTHQUAKES CLOSE BY DO NOT HAVE LG
 \$

STJ	7010311744P	46135		47205		30 34 9	8 1
STJ	SW 0689KM 15	-036 224 49		05 021		0000554 41ML36MN	
SCH	7010311744P	4739		4948		40 94 9	8 1
SCH	W 1384KM 15	036 290 49		05 -021		0000150 45ML35MN	

Z

71

+47.308- 49.160FLMN=4.8 0617095 15081971 00.1310.215 0.2 10 17 62.50 218.00 0 1ML=0.0 00 0L3.62
 \$ 47.46- 49.53 MB=4.8 061715 CEEF (FROM ISC)
 \$ 47.46- 49.53 MB=4.8 0617152 ISC
 \$ GRAND BANKS
 \$ READING BY WAHLSTROM IN 1985 ADDED BY ADAMS OCT 1988
 \$ POLARITIES CHECKED BY ADAMS OCT 1988
 \$ WAHLSTROM COMMENTS: STJ ONLY GIVES DIST=255KM, BACKAZ=084DEG, H=18KM ASSUMED
 \$ I.E. 47.8N 49.4W OT=061713
 \$ STJ: FM IS D-E. EVENT FROM SLIGHTLY NORTH OF EAST
 \$ WELL-RECORDED ON LPS
 \$ HAL: FM IS C. LG SLIGHTLY LARGER THAN SN
 \$ SCH READ BY WAHLSTROM AS C; BY ADAMS AS -
 \$ SIC READ BY WAHLSTROM AS C; BY ADAMS AS +
 \$ OTT READ BY WAHLSTROM AS ; BY ADAMS AS +
 \$
 \$
 \$ CHECK ISC
 \$ CHECK REST OF CSN
 \$ WAHLSTROM HAS TRIED A FM
 \$

STJ	7108150617P	A17498 D	X17508	B18172	X18192		1
STJ	W 0271KM 05	147 278 49	00 -259	00 042 00 -544	0000000 00ML00MN		

HAL	7108150617P	A19371 C	A21278	100	18	155	3	1
HAL	W	1156KM 02	087 260 49	17	277	0005411	60ML49MN	
SIC	7108150617P	A19560 C				00000000	00ML00MN	
SIC	W	1329KM 04	-137 290 49			XA23336	100 35 166	
UNB	7108150617P	A19533 -				00 1232\$	0002980 59ML48MN	
UNB	W	1346KM 81	-604 270 49			A22344	80 62 405	
SCH	7108150617P	A20167 C				02 -098	0005130 63ML51MN	
SCH	NW	1486KM 00	015 311 49			A23069	X25001 90 27 45	
SFA	7108150617P	A20330 D				01 -054	00 1826\$ 0001164 57ML45MN	
SFA	W	1637KM 08	-193 277 49			B24410	110 17 28	
OTT	7108150617P	A21213				06 650	0000941 58ML46MN	
OTT	W	2047KM 01	-074 274 47			B24578	100 26 115	
GWC	7108150617P	A21376 +				01 -194	0002779 63ML51MN	
GWC	NW	2165KM 21	305 305 45			B25114	90 39 18	3
FBC	7108150617P	B21389				09 827	0000322 54ML42MN	
FBC	NW	2181KM 01	269 334 43				110 20 49	
SUD	7108150617P	A22011 +					0001399 61ML49MN	
SUD	W	2411KM 12	229 280 39					
		Z						

76

+50.133- 48.910F1ML=4.1 1923304 28081976 00.0260.108 0.3 5 8 30.54 218.00 0 1MN=0.0 00 0 3.62
\$50.10 - 48.85 ML=4.0 192330 CEEF
\$ GRAND BANKS
\$ MUNF IS MEMORIAL SEISMOGRAPH IN STJ VAULT
\$ WAHLSTROM COMMENTS:
\$ ISC: AS OTT SOLUTION, I.E., 50.10 N 48.85 W OT=192330 ML=4.0 "POOR SOLUTION"
\$ MIQ, SUD, FRB, OTT, QCQ, LGQ: NOT RECORDED, CBK, HAL: RECORD NOT FOUND
\$ PBQ: VISIBLE 1932 BUT NO CLEAR ONSET
\$ UNB, SIC: PROB DISTURBANCES AT APPR 2940
\$
MUNF 7608281925P 40 73 156 8 1
MUNF SW 0400KM 226 49 0003357 42ML39MN
STJ 7608281925P A24258 D A25048 40 32 20 8 1
STJ SW 0400KM 34 086 226 49 04 -028 0000982 37ML34MN
SCH 7608281925P A26168 A28209 30 100 8 8 1
SCH NW 1318KM 01 -017 300 49 06 037 0000168 43ML35MN
MNQ 7608281925P A26281 B28427 50 5 3 1
MNQ W 1412KM 04 -031 279 49 14 219 0000000 00ML00MN
POC 7608281925P A26483 0000000 00ML00MN 1
POC W 1580KM 12 -052 267 49 0000000 00ML00MN 1
CHQ 7608281925P B27006 0000000 00ML00MN
CHQ W 1688KM 05 -137 266 49
Z

77

+49.337- 49.58601ML=3.4 0805261 25031977 00.1230.155 0.0 2 4 10.12 218.00 0 1MN=0.0 00 0 3.62
\$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
\$ GRAND BANKS
\$ ALTERNATIVE LOCATION (LESS LIKELY BECAUSE HAL SHOWS NO TRACE) FOLLOWS:
\$44.88 - 52.59 01ML=3.4 080527. 25031977 1 0.06 0.54 0.0 2 5 1 0.5 218MUN 300
\$ CBK PEN AND INK RECORD
\$ MUNF NO LG
\$
MUNF 7703250806P A0609 XB06175 A06405 030 108 85 8 1
MUNF SW 0305KM 00 -001 231 49 00 213 00 001 0001648 34ML34MN
CBK 7703250806P C0649 C0745 1
CBK W 0614KM 30 227 269 49 10 -130 0000000 00ML00MN
HAL 7703250806P 1
HAL W 1185KM 249 49 0000000 00ML00MN
Z

78

+52.573- 46.752F1ML=4.1 1435140 03011978 00.0280.073 0.1 3 6 30.36 218.00 0 1MN=0.0 00 0 3.62
\$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
\$ GRAND BANKS
\$ ADAMS RE-READ RECORDS OCT 1988
\$ MUNF, STJ NO LG
\$
MUNF 7801031436P XA36458 - X37542 030 108 30 8 1
MUNF SW 0702KM 00 038 220 49 00 120 0000582 41ML36MN
STJ 7801031436P A36455 - A37530 040 29 6 8 1
STJ SW 0702KM 01 008 220 49 00 000 0000325 40ML33MN
CBK 7801031436P A3710 -2.0 A38350 1
CBK SW 0889KM 08 -029 247 49 01 010 0000000 00ML00MN

SCH	7801031436P	B38066	A40094	050 120 7	8 1
SCH	W 1343KM 50	295 289 49	01 -008	0000073 42ML32MN	
FRB	7801031436P				1
FRB	NW 1773KM	323 49		0000000 00ML00MN	
	Z				

81

+48.643- 47.820F1ML=4.2 1306126 26091981 00.0670.170 0.2 5 10 61.16 218.00 0 1MN=0.0 00 0 3.62
\$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
\$ GRAND BANKS
\$ READ BY WAHLSTROM AND ADAMS
\$MUNF: TC UNKNOWN, APPEARS TO BE 0.0 AS TIMES AGREE
\$MUNF: LG=0.5SN, NO AFTERSHOCKS
\$CBK: T.C. UNKNOWN, T.C. -2 SEC APPLIED
\$ LG WEAK OR ABSENT
\$HAL: V QUIET TRACE
\$SIC: QUIET, NO PHASES READABLE
\$UNB: BOTH READING SIMILAR TO NOISE
\$LPQ,LMQ: HIGH FREQUENCY ON LOW FREQUENCY NOISE
\$FRB: NOT RECORDED
\$

STJ	8109261307P	A07068	A07155	A07431	50 36 65	8 1
STJ	W 0385KM 24	153 254 49	05 073	09 -095	0002269 41ML37MN	
MUNF	8109261307P				30 108 250	8 1
MUNF	W 0385KM	254 49			0004848 42ML41MN	
CBK	8109261307P	X07523		X09031	50 89	3 1
CBK	W 0746KM 00	291 276 49		00 208	0000000 00ML00MN	
HAL	8109261307P	X09046		B10566	50 83 6	8 1
HAL	W 1285KM 00	957 256 49		01 106	0000091 43ML32MN	
UNB	8109261307P	XB09168				1
UNB	W 1452KM 00	139 265 49			0000000 00ML00MN	
SCH	8109261307P	A09180		A11369	40 95 6	8 1
SCH	NW 1474KM 00	-016 305 49		10 100	0000099 44ML34MN	
MNQ	8109261307P	A09224		B11437	40 220 8	8 1
MNQ	W 1526KM 43	-204 286 49		06 -315	0000057 42ML31MN	
LPQ	8109261307P	XA09410				1
LPQ	W 1658KM 00	050 273 49			0000000 00ML00MN	
LMQ	8109261307P	A09427		B12200	30 190 3	8 1
LMQ	W 1676KM 00	-006 274 49		01 119	0000033 39ML30MN	
	Z					

82

+41.367- 47.282F1ML=4.0 0407239 22071982 00.1040.124 0.1 6 8 41.05 218.00 0 1MN=0.0 00 0 3.62
\$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
\$ SOUTHEAST OF TAIL OF THE BANK, GRAND BANKS
\$ FIRST EARTHQUAKE KNOWN IN THIS AREA
\$ GOOD ON MUNF, STJ AND CBK, KLN
\$ HAL: VERY WEAK
\$ CBK TC SEVERAL HOURS AND ABOUT 26 SECONDS
\$ SCH, UNB: PRESENT, BUT UNCLEAR
\$ SIC, GSQ, LMN, GGN, LMQ: RECORDS NOT CHECKED
\$

STJ	8207220409P	X09114	A10267	50 36 4	8 1
STJ	NW 0815KM 00	239 330 49	00 -009	0000140 39ML31MN	
MUNF	8207220409P	A09087	X10277	20 173 19	8 1
MUNF	NW 0815KM 01	-031 330 49	00 090	0000345 39ML35MN	
CBK	8207220409P	X1021	X12123	50 11	3 1
CBK	NW 1186KM 00	022 319 49	00 -006	0000000 00ML00MN	
HAL	8207220409P	XB10284	B12265		1
HAL	W 1377KM 00	1084\$291 49	00 011	0000000 00ML00MN	
KLN	8207220409P	A10491	A13210	20 628 18	8 1
KLN	NW 1641KM 05	-070 298 49	28 -164	0000090 41ML34MN	
EBN	8207220409P	A11082	A13573		1
EBN	NW 1795KM 00	-018 299 52	32 177	0000000 00ML00MN	
MNQ	8207220409P	A11264	XC14327	50 161 2	8 1
MNQ	NW 1944KM 14	115 309 50	00 557	0000016 39ML28MN	
	Z				

84a

+48.852- 50.83701MN=3.6 1925475 26051984 00.0190.037 0.3 8 18 60.48 218.00 0 1ML=4.1 30 0 3.62
\$48.89 - 51.06 MN=3.5 192540 CEEF
\$ GRAND BANKS
\$ KLN,SIC NOISY.
\$ LG>SN FOR THIS OFFSHORE EVENT
\$ AFTERSHOCK ONLY ON MUNF AT 2204:50, SAME S-P, MAGNITUDE = 2.6

84b

85

88a

46

KLN	8801090715P	XC1827		C2022		8
KLN	W	1258KM	00 305 261 49	00 -002	0000000	00ML00MN
SCH	8801090715P	B1824		B2024	30 137 4	8
SCH	NW	1270KM	27 -141 302 49	04 -056	0000061	39ML30MN
MNQ	8801090715P			B2040	20 628 3	8
MNQ	W	1342KM	281 49	00 003	0000015	31ML25MN
		Z				

88b

+47.463- 49.24601ML=3.1 0016451 09081988 00.1950.120 0.0 5 8 20.76 218.00 0 1MN=3.2 20 0 3.62
\$47.463- 49.293 ML=3.1 0016455 CEEF
\$ ON THE GRAND BANKS, EAST OF ST JOHN'S
\$ STJ HAS GOOD SHARP ONSETS
\$ THIS EARTHQUAKE WOULD NOT HAVE BEEN MISSED ON MUNF! (& PERHAPS NOT ON STJ)
\$ MUNF LG NOT EVIDENT
\$ NOTHING ON HAL, SIC, LMQ, SLQ, GGN
\$ SCH MEAREST WIGGLE AT 2218
\$ MNQ SN MAY BE LOST NEAR PAPER EDGE
\$ CBK HAS LARGER AMPLITUDE PHASE 20 TO 40 SEC AFTER SN
\$ KLN NO SN; GBN SN AMP TOO SMALL FOR MAGNITUDE
\$ POOR AZIMUTHAL DISTRIBUTION OF STATIONS
\$ FITS ALMOST AS WELL AT 48.96N 49.98W (THIS MISFITS THE VERY POOR MNQ PHASE)
\$ STJ P ONSET AND CODA IS STRONGER ON EW COMPONENT THAN ON NS, SUPPORTING
\$ THE CHOSEN EPICENTRE OVER THE ALTERNATIVE EPICENTRE TO THE NW OF STJ
\$

STJ	8808090018P	A17235		A17509	40 30 21	8
STJ	W	0263KM	15 059 274 49	03 029	0001100	31ML32MN
MUNF	8808090018P	XA17225		XA17512	30 91 70	8
MUNF	W	0263KM	00 -041 274 49	00 058	0001611	32ML33MN
CBK	8808090018P	B18105		B19132		
CBK	W	0668KM	09 -188 287 49	37 -373	0000000	00ML00MN
GBN	8808090018P	A18487 C		B20195		
GBN	W	0970KM	06 -037 261 49	05 -144	0000000	00ML00MN
KLN	8808090018P	B1928				
KLN	W	1299KM	04 -126 273 49		0000000	00ML00MN
MNQ	8808090018P	C1951				
MNQ	W	1466KM	00 137 291 49		0000000	00ML00MN
		Z				

89a

+52.730- 35.200F1MS=5.4 2253371 14051989 00.0390.030 0.0 00 00 250.00H610.00 0 1MB=4.9 60 0L0.00
052.671- 35.037F1MB=5.0 2253387 14051989 00.1010.155 0.0 13 19 00.00H510.00 0 1ML=5.9 60 0L0.00
052.618- 34.687F1ML=5.9 2253369 14051989 00.0820.126 0.2 14 20 60.96 218.00
MID-ATLANTIC RIDGE RIDE MEDIO-ATLANTIQUE

MAG (NEIS) 5.0 MB (50 OBS)

\$ ADOPTING THE NEIS ORIGIN TIME AND EPICENTRE LEADS TO A 5-6 S ERROR

\$ ON ALL CANADIAN PHASES. THIS IS LIKELY NOT DUE TO AN ERROR IN

\$ THE TIME SERIES FILE TRIGGER ERROR, SINCE ALSO ON STD STATIONS

\$ RATHER IT IS LIKELY DUE TO INAPPROPRIATE VELOCITY MODEL

SCH	8905142257P	A57565 -	0.00	B0120	060 86 117	0 8
SCH	W	2092KM	00 -072 289 45	00 468	0001425	60ML48MN
LMN	8905142257P	B581489	-0.29	C615482	053 100 61	0 8
LMN	W	2278KM	00 -167 262 43	00 -034	0000723	57ML45MN
GSQ	8905142257P	B581647	-0.22	C615538	077 100 239	0 8
GSQ	W	2282KM	00 -042 272 41	00 -056	0001950	62ML50MN
FRB	8905142257P	A58163 +	0.00		000 0 0	0 0
FRB	NW	2287KM	00 -083 316 41		0000000	00ML00MN
MNQ	8905142257P	B582045	-0.10	C620766	080 100 173	0 8
MNQ	W	2330KM	00 -112 278 41	00 158	0001359	61ML48MN
KLN	8905142257P	B582036	-0.29	B620663	053 100 74	0 8
KLN	W	2330KM	00 -139 266 41	00 037	0000877	58ML46MN
HTQ	8905142257P	B582390	-0.07	C620999	060 100 50	0 8
HTQ	W	2357KM	00 -013 274 41	00 -170	0000524	56ML44MN
EBN	8905142257P	B583069	-0.22		000 0 0	0 0
EBN	W	2426KM	00 -003 270 39		0000000	00ML00MN
LPQ	8905142257P	B584521	-0.22		000 0 0	0 0
LPQ	W	2551KM	00 296 271 37		0000000	00ML00MN
DPQ	8905142257P	C590130	-0.06		000 0 0	0 0
DPQ	W	2772KM	00 006 271 35		0000000	00ML00MN
TRQ	8905142257P	B591363	-0.09		000 0 0	0 0
TRQ	W	2917KM	00 029 272 34		0000000	00ML00MN
GRQ	8905142257P	B591813	-0.09		000 0 0	0 0
GRQ	W	2985KM	00 -067 273 33		0000000	00ML00MN
WBO	8905142257P	XC593216	-0.06		000 0 0	0 0
WBO	W	3033KM	00 950 270 33		0000000	00ML00MN

EEO 8905142257P B593575 -0.06
EEO W 3200KM 00 -035 276 33
Z

000 0 0 0 0
0000000 00ML00MN

89b

+46.960- 49.247F1ML=4.2 0515398 03121989 00.0500.088 0.4 6 9 40.71 218.00 0 1MN=0.0 00 0L3.62
OFFSHORE NEWFOUNDLAND AU LARGE DE TERRE-NEUVE

MAG (GSC) 3.9 MB (1 OBS)
\$ MB 3.9 FROM YELLOWKNIFE ARRAY
\$ MNQ DOWN, TOO WEAK FOR OTHERS
\$ LOCATION FROM YELLOWKNIFE ARRAY: 49.7 51.4 O.T. 05:15:58
\$ SCH LG>=SN AMP; SN PRECEDED BY NOISE
\$ STJ UNABLE TO DECIDE IF LG>SN
\$ FRB NOT VISIBLE
\$ CBK NOT OPERATING
\$ ADAMS ADDED STJ AND REVISED SCH
\$ ADAMS DID NOT USE KLN, GGN, EBN BECAUSE TOO MANY STATIONS ON THIS AZIMUTH
\$

STJ	8912030515P	B16197		XB16215	C1647		
STJ	W	0273KM	19	092 286 49	00 -239	00 -035	0000000 00ML00MN
GBN	8912030515P	B17435	-		B19145	040 216 20	8
GBN	W	0962KM	09	062 264 49	04 043	0000145 40ML32MN	
HAL	8912030515P	B1804					
HAL	W	1144KM	22	-100 262 49		0000000 00ML00MN	
LMN	8912030515P	B181142	-0.29		XC2005	000 0 0 0 0	
LMN	W	1201KM	20	-096 270 49	00 -034	0000000 00ML00MN	
KLN	8912030515P	XB182287	-0.29		XC2031	050 204 4 0 8	
KLN	W	1303KM	00	-192 276 49	00 402	0000025 37ML27MN	
GGN	8912030515P	XB183188	-0.29		XC2043	050 102 7 0 8	
GGN	W	1374KM	00	-147 268 49	00 109	0000086 43ML32MN	
EBN	8912030515P	XB1838	0.00			000 0 0 0 0	
EBN	W	1437KM	00	-285 279 49		0000000 00ML00MN	
SCH	8912030515P	B18500	-		B21095	050 102 11 8	
SCH	NW	1507KM	09	065 312 49	06 -054	0000136 46ML35MN	
YKA	8912030515P	B2304	0.00			000 0 0 0 0	
YKA	NW	4329KM	00	-010 318 31		0000000 00ML00MN	

Z

90

+41.918- 48.34101ML=4.7 2140041 24041990 00.0270.062 0.1 14 23 130.67 218.00 0 1MN=0.0 00 0L3.62
042.000- 48.48001MB=4.3 2140030 24041990 00.1800.072 0.1 00 00 81.31 618.00 0 1MN=0.0 00 0L3.62
SOUTHERNMOST PART OF THE GRAND BANKS LA PARTIE LA PLUS AU SUD DU GRAND BANC
DE TERRE NEUVE

MAG (GSC) 3.9 MB (1 OBS)
MAG (ISC) 4.3 MB (8 OBS)
\$ DETECTED ON THE YELLOWKNIFE ARRAY
\$ MAGNITUDE MB=3.9 AT 1 HZ ON YK
\$ STJ: NO LG; POLARITY POSSIBLE +
\$ SCH: STRONG P, WEAK SN; PN READING IS VERY SHARP
\$ GBN: EMERGENT BEGINNING, POLARITY NOT READABLE
\$ CBK: PRECURSOR TO SN MAKES ONSET UNCERTAIN
\$

STJ	9004242140P	A41379		A42465	020 173 145	8
STJ	NW	0718KM	27	049 333 49	00 005	0002633 46ML43MN
CBK	9004242140P	B4221		XB4405	040 132 46	8
CBK	NW	1082KM	05	-082 319 49	00 106	0000547 48ML39MN
GBN	9004242140P	B42295 E		B4412	020 563 200	8
GBN	NW	1130KM	25	190 295 49	28 -203	0001116 48ML42MN
HAL	9004242140P	B42455		XC44405	030 207 60	8
HAL	W	1273KM	01	047 289 49	00 -394	0000607 49ML40MN
LMN	9004242140P	A425987D	-0.29	B451097	013 100 13	0 8
LMN	NW	1391KM	01	008 294 49	07 100	0000628 46ML41MN
GGN	9004242140P	C432033	-0.29	C453877	017 100 13	0 8
GGN	W	1533KM	05	328 290 49	01 -131	0000480 47ML41MN
UNB	9004242140P	C43175		X4532	050 74 12	8
UNB	NW	1533KM	00	072 293 49	00 -783	0000204 48ML37MN
KLN	9004242140P	B431925	-0.29	C454049	0131000 88	0 8
KLN	NW	1534KM	28	202 297 49	00 010	0000425 46ML40MN
GSQ	9004242140P	B433283	-0.22	C460650	040 100 14	0 8
GSQ	NW	1657KM	03	069 304 49	00 004	0000220 48ML38MN
EBN	9004242140P	B433497	-0.22	C461354	0201000 77	0 8
EBN	NW	1689KM	07	-101 298 49	00 038	0000242 46ML38MN
LPQ	9004242140P	XB435397	-0.22		000 0 0 0	
LPQ	NW	1816KM	00	289 297 52		0000000 00ML00MN
DAQ	9004242140P	B440216	-0.06	C470492	0271000 83	0 8
DAQ	NW	1921KM	04	-077 298 50	02 241	0000193 48ML38MN

SCH	9004242140P	A44076 C	C4713	060 94 12	8
SCH	NW 1971KM 74	-082 323 47	00 -003	0000134	49ML37MN
DPQ	9004242140P	C441495 -0.06	X472281	0331000	46 0 8
DPQ	NW 2013KM 02	194 294 47	00 089	0000088	45ML35MN
YKB9	9004242140P	XB480459C			
YKB9	NW 4803KM 00	117 322 30		0000000	00ML00MN
YKA	9004242140P	X4804 0.00		000 0 0	0 0
YKA	NW 4804KM 00	053 322 30		0000000	00ML00MN
YKR3	9004242140P	XB480555C			
YKR3	NW 4816KM 00	118 322 30		0000000	00ML00MN
YKR2	9004242140P	B480575C			
YKR2	NW 4819KM 10	119 322 30		0000000	00ML00MN

Z

91

+47.364- 50.167F1MN=3.2 1123013 23071991 00.0820.146 0.2 6 10 30.79 218.00 0 1ML=3.8 20 0L3.62
 GRAND BANKS
 \$ EAST OF ST JOHN'S
 \$ LG ON STJ, CBK, SMQ, SCH, NOT ON GBN OR LMN
 \$ NOT ON LMQ, CIQ, DAQ
 \$ MNQ DEAD
 \$ NO ECTN TRIGGER (?)
 \$ STJ: ONLY FIRST MOTION VISIBLE; NO AFTERSHOCKS EVIDENT
 \$ CBK: P PRECEDED BY ?NOISE AT 24122
 \$ CBK: NO CLEAR S PHASES
 \$ HAL: NOISY RECORD, READING IS PROBABLY JUST NOISE
 \$ SMQ: NOISY RECORD
 \$ SCH: LG SEEMS OK
 \$ KUQ: ?PHASE AT 26455 IS 10 SEC TOO LATE FOR P
 \$
 \$ SOLUTION BY ADAMS 910814
 \$ POOR STATION DISTRIBUTION; THIS LOCATION FITS PN'S AT STJ, CBK, GBN, AND SCH
 \$ ACCEPTABLY, ALSO SN'S AT GBN AND SCH (BUT NOT CBK S PHASES)
 \$

STJ	9107231123P	A23308 D			
STJ	W 0195KM 00	-005 278 49		0000000	00ML00MN
CBK	9107231123P	B2420	C2515 XC2535	030 193 27	
CBK	W 0606KM 07	-100 289 49	10 -486 00-1382\$	0000293	37ML32MN
GBN	9107231123P	B24575	B26205	020 563 25	8
GBN	W 0899KM 04	077 260 49	20 -171	0000140	36ML31MN
HAL	9107231123P		X2715		
HAL	W 1082KM	259 49	00 1383\$	0000000	00ML00MN
LMN	9107231123P	B2525	B27145	040 300 13	8
LMN	W 1133KM 01	-028 267 49	43 248	0000068	39ML30MN
SMQ	9107231123P	XB25325	C2734 X2832	040 300 13	
SMQ	W 1254KM 00	-752 291 49	06 -373 00-1579\$	0000068	40ML31MN
SCH	9107231123P	B2602	B2814 XB29155	060 94 10	
SCH	NW 1425KM 09	112 312 49	00 -013 00-1950\$	0000111	45ML34MN

Z

92a

+47.237- 49.226F1MN=4.0 0607280 13011992 00.0560.097 0.2 16 24 140.98 218.00 0 1ML=4.5 20 0L3.62
 SOUTHERN GRAND BANKS
 OFFSHORE NEWFOUNDLAND AU LARGE DE TERRE-NEUVE
 \$ EXACT EPICENTRE RATHER UNCERTAIN BECAUSE OF EXTREMELY POOR STATION
 \$ DISTRIBUTION
 \$ ISC 47.2N 49.2W OT=060727.2; added 3 stations to GSC's 15;
 \$ "EXACT EPICENTRE RATHER UNCERTAIN BECAUSE OF EXTREMELY POOR STATION DISTRIBUTION"
 \$ ADAMS SOLUTION
 \$ VLG OF 3.62 USED FOR THIS SOLUTION
 \$ LG GREATER THAN SN ON MOST ECTN EXCEPT LMN AND DAQ WHERE AMPLITUDE SN = LG
 \$ LG ON PWM AND FCC, NOT IN TRIGGER WINDOW
 \$ SCH, FRB CLOSED
 \$ CBK ASSYMETRIC TRACE; MAX DOWNSWING USED
 \$ HAL LONG CODA, PROBABLY REPRESENTS ATTENUATED LG
 \$ GBN NO LG
 \$ KUQ, IGL, FCC AMP FROM LG
 \$ FCC EARLIER PHASE AT 2141, STONGEST ON SPZ MAY BE BE TELESEISMIC S OR
 \$ UNRELATED TELESEISM
 \$ CLOSE TO LARGE EARTHQUAKE IN 1971
 \$ RELATIVE ARRIVALS ON STJ AND HAL (ONLY COMMON STATIONS) SUGGEST
 \$ SIMILAR LOCATION TO 1971 EARTHQUAKE
 \$

STJN	9201130610P	A08074 D	X0833		
STJN	W 0268KM 22	063 279 49	00 -197	0000000	00ML00MN
CBK	9201130610P	A08575	B1033	040 135 170	

92b

```

$ CLOSE TO CLUSTER OF EARTHQUAKES IN 1971, 1988, 1989, AND 1992, BUT SLIGHTLY
$ FARTHER TO THE WEST
$ HAL TOO NOISY
$ NOT ON LMN,ICQ,MNQ,KUQ,DAQ
$ STJN GOOD SIGNAL BUT POOR LG, HALF AMP OF SN
$ CBK AMPLITUDE OF SN WAS READ
$ CBK PROMINENT PHASE FITS AS LG WITH AMP 1.1 MM Z-P AT 030; TWICE AMP OF SN
$ GBN POOR PHASES, NOISY RECORD.
$ GBN POSSIBLE OTHER ARRIVALS AT 6028 AND 6228
$ ARE NOT CONSISTANT WITH CBK AND STJN
$
$

```

50

GBN 9207061658P C6017 C6150 030 317 02 8
GBN W 0960KM 01 -192 262 49 00 010 0000013 29ML22MN
Z

92c

+46.118- 47.438F1MN=3.9 0420227 17071992 00.1120.087 0.1 9 17 20.77 218.00 0 1ML=4.2 30 0L3.62
EASTERN MARGIN OF GRAND BANKS
\$ FIRST EARTHQUAKE KNOWN FROM HERE; APPROX 200 KM SE OF NEAREST CLUSTER
\$ OF EARTHQUAKES
\$ STJN VERY CLEAR, GOOD PHASES; LG SEEMS OK
\$ CBK GOOD LG
\$ LG FROM THIS SITE TO NEWFOUNDLAND STATIONS IS UNEXPECTED
\$ ALSO THE LG HAS QUITE LONG PERIODS
\$ GBN VERY WEAK LG; LG AT ALL ECTN SITES APPEARS ATTENUATED
\$
\$ ECTN TRIGGER CONTAINS ONLY PN FOR STATIONS BEYOND LMN
\$ KURILE AFTERSHOCK AT THIS TIME INTERFERS: MNQ, MOQ, DAQ READINGS WERE OF THIS EVENT
\$ JAQ DEAD
\$ KUQ LG ARRIVAL NOT READABLE, BUT AMPLITUDE READ AT TIME CORRESPONDING TO LG
\$
STJN 9207170421P A21223 - A22052 C22185 040 62 127
STJN NW 0437KM 31 060 294 49 00 002 08 -497 0003218 43ML40MN
CBK 9207170421P A22122 D B2332 C2418 070 43 36
CBK NW 0852KM 01 -013 295 49 12 -152 00 -006 0000751 47ML38MN
GBN 9207170421P B22407 + C2424
GBN W 1097KM 12 -151 271 49 01 -165 0000000 00ML00MN
HAL 9207170421P C23045 B25055 050 116 8 8
HAL W 1274KM 00 066 268 49 24 210 0000087 42ML32MN
LMN 9207170421P B231104 -1.52 C251825 0571000 51 8 25243
LMN W 1344KM 44 -287 275 49 01 -157 0000056 41ML30MN
SMQ 9207170421P B232985 -0.10
SMQ NW 1500KM 14 -163 295 49 0000000 00ML00MN
GSQ 9207170423P B233080 -0.10
GSQ W 1511KM 21 -197 289 49 0000000 00ML00MN
ICQ 9207170423P XB233268 -0.10 XC2604
ICQ W 1529KM 00 -229 292 49 00 636 0000000 00ML00MN
CNQ 9207170423P XB233972 -0.10
CNQ W 1584KM 00 -203 290 49 0000000 00ML00MN
SLQ 9207170421P C23505 C2624 X2810 100 68 4 3
SLQ W 1649KM 00 099 284 49 00 092 00 1181\$ 0000037 42ML30MN
MNQ 9207170421P -0.10 C26225 100 96 8 3
MNQ NW 1651KM 295 49 01 -125 0000052 44ML32MN
DPQ 9207170421P X2723 XC2938 050 140 2 3
DPQ W 1943KM 281 50 00 -258 00 1863\$ 0000018 40ML28MN
KUQ 9207170421P 090 44 1 3
KUQ NW 1948KM 321 47 0000016 40ML28MN
Z

92d

+47.353- 49.114F1MN=3.4 1131520 10081992 00.0830.070 0.0 5 9 10.39 218.00 0 1ML=3.6 30 0L3.62
GRAND BANKS
\$ IN CLUSTER OF EARTHQUAKES (4 PREVIOUS) NW OF HIBERNIA
\$ VERY POOR AZIMUTHAL COVERAGE FOR THESE EARTHQUAKES
\$ NOT ON DAQ, ICQ
\$ STJ SN>LG
\$ CBK SG=3*SN
\$ GBN AMP READ FROM SN, THERE IS WEAK LG FOLLOWING THE SN
\$ CBK LG MIGHT BE 4 SEC EARLIER
\$
STJ 9208101132P A32312 B33012 XC3308 040 73 165 8
STJ W 0274KM 00 000 276 49 44 129 00 006 0003550 37ML37MN
CBK 9208101132P B33211 B34260 XC3458 040 135 32
CBK W 0682KM 01 017 288 49 12 -068 00 -244 0000372 40ML34MN
GBN 9208101132P B33572 B3529 X3627 030 370 12 8
GBN W 0978KM 01 021 262 49 09 -059 00 483 0000068 36ML29MN
HAL 9208101132P
HAL W 1160KM 260 49 0000000 00ML00MN
LMN 9208101132P B34255 C3618 030 428 5 8
LMN W 1213KM 01 -017 268 49 04 -162 0000024 34ML26MN
MNQ 9208101132P C3455 XC3933
MNQ W 1479KM 17 -324 291 49 00 5223\$ 0000000 00ML00MN
Z

94a

+52.127- 48.523F1ML=3.2 1052279 08011994 00.0380.065 0.2 6 10 40.56 218.00 0 1MN=0.0 00 0L0.00
 NORTHEAST NEWFOUNDLAND SLOPE
 \$ LIES WITHIN LINEAR TREND OF EPICENTRES ALONG EXTENSION OF CARTWRIGHT FRACTURE ZONE
 \$ SEEN FIRST ON DRLN, PICKED UP WITH DIFFICULTY ON LMN
 \$ ANALOG ECTN READ - TRACES ARE MERE WRIGGLES
 \$ NOT ON HAL, CBK
 \$ KUQ IS WRIGGLE
 \$

STJN 9401081055P	B53455	B5444	040 62 9	8	
STJN SW 0591KM 07 -030 213 49		08 033	0000228	35ML31MN	
DRLN 9401081054P	B540047 +0.00	B550865	0152240	46 0000 8	559.62
DRLN SW 0710KM 02 016 247 49		08 -033	0000086	30ML28MN	
GBN 9401081055P		C56564	030 317 2	8	
GBN SW 1210KM	237 49	05 105	0000013	31ML23MN	
ICQ 9401081055P		C5726			
ICQ W 1350KM	265 49	05 104	0000000	00ML00MN	
LMN 9401081055P	C552005 +0.00	C572900	0452560	11 0000 8	5736.35
LMN SW 1378KM 12 -163 246 49		18 -195	0000006	31ML21MN	
MNQ 9401081055P	C5529	C57415			
MNQ W 1419KM 23 222 271 49		13 165	0000000	00ML00MN	
KUQ 9401081055P		X5738			
KUQ NW 1428KM	306 49	00 -362	0000000	00ML00MN	
FRB 9401081055P					
FRB NW 1743KM	326 49		0000000	00ML00MN	

Z

94b

+47.827- 52.67301MN=3.1 1813492 11081994 00.0270.064 0.2 6 11 30.50 2 0.00 0 1ML=3.3 10 0L3.62
 ST JOHN'S NFLD ST JOHN'S T.-N.
 \$ LOCATION PERHAPS 15-30 KM N OF ST JOHN'S, BUT N-S RESOLUTION
 \$ IS POOR BECAUSE OF STATION DISTRIBUTION
 \$ BECAUSE OF TIME OF DAY AND PROXIMITY TO CITY A BLAST WAS SUSPECTED
 \$ ENQUIRIES BY PAUL BARNES (MEMORIAL UNIV., PAUL@CONVEX.ESD.MUN.CA)
 \$ TO A LOCAL EXPLOSIVES SUPPLIER INDICATED:
 \$ 1. APPROXIMATELY 20 KM WEST OF ST. JOHN'S (LONG POND), THERE IS A
 \$ PYROPHYLITE MINE THAT USES CHARGES ON THE ORDER OF 25 KG.
 \$ 2. ON 24 OCTOBER, APPROX 12-15 KM WEST OF ST. JOHN'S, 300X25KG WERE
 \$ DETONATED OVER A 2 SECOND INTERVAL TO BLAST AWAY A ROCK FORMATION TO
 \$ GENERATE ROADBED GRAVEL. THE MAXIMUM CHARGE SIZE IN ANY ONE HOLE WAS
 \$ 250KG. THIS WAS THE LARGEST BLAST EXPLODED. THE DATE MAY BE IN ERROR.
 \$ THE EVENT WAS NOT LOCATED ON THE STJN RECORD.
 \$ 3. ABOUT 5-8 KM N OF THAT SITE, CONSTRUCTION OF A BYPASS ROAD IS BEING
 \$ UNDERTAKEN. ACCORDING TO THE EXPLOSIVES SUPPLIER, THEY ARE DETONATING
 \$ CHARGES ON THE ORDER OF 50 KG.
 \$ 4. THERE IS ALSO BLASTING FOR ROAD CONSTRUCTION OCCURRING ALONG THE TCH
 \$ APPROXIMATELY 50 KM WEST OF ST. JOHN'S.
 \$ 5. UNAWARE OF ANY OTHER CONSTRUCTION.
 \$
 \$ ADAMS CHECKED ALL THE STJN RECORDS HERE FOR AUG 10 TO EARLY OCT AND
 \$ IDENTIFIED 10 EVENTS, 3 WITH A S-P OF 5.5 SEC AND AROUND 2030-2230 UT
 \$ (PROBABLY SINGLE SOURCE, POSSIBLY YOUR TCH BLASTING?), AND THE REST
 \$ WITH S-P AROUND 2 SEC AT 1200-2045 UT (DAYLIGHT HOURS, LIKELY MULTIPLE
 \$ SOURCES WITHIN 20 KM). NOT A SINGLE ONE WOULD BE CLASSIFIED AS A LARGE
 \$ BLAST, AND THE LARGEST EVENT IS AT LEAST 1 MAGNITUDE BELOW THE AUG 11TH
 \$ EVENT. IF THE AUG 11 EVENT WAS A BLAST, WE SEEM FORCED TO CONCLUDE THAT
 \$ IT WOULD HAVE NEEDED TO INVOLVE C. 10* THE 300X25 KG = 87 TONNES, AND
 \$ THEREFORE WOULD HAVE BEEN REMEMBERED!
 \$ ---> CONCLUSION CLASSIFY AS AN EARTHQUAKE
 \$
 \$ MISSED DURING REGULAR DAILY PROCESSING OF CSN
 \$ STJN CODA LASTS 2.5 MIN
 \$ STJN: SUSPECT S-P IS ABOUT 2.8 SEC, BUT UNREADABLE
 \$ HAL NOT EVIDENT ON NOISY RECORD
 \$ SCH NOISY RECORD
 \$ LMN DEAD
 \$ MNQ FROM ANALOG
 \$ DONE BY ADAMS
 \$ LAST COMPARABLE EVENT WAS IN 1969:
 \$ +47.632- 52.156F1MN=3.4 2153207 05081969 00.0860.110 0.2 5 9 40.94 218.00
 \$ THIS WAS FROM ALMOST DUE EAST OF STJ ACCORDING TO 3 COMPONENT MOTIONS
 \$

STJN 9408111814P		A13542 C			
STJN S 0029KM	192-90	06 024	0000000	00ML00MN	
DRLN 9408111814P	A144492C -1.00	A152635 B153947	0252303	672 0000 0	1544.58
DRLN NW 0391KM 32 -055 296 49		11 033 11 129	0000733	33ML33MN	

94c

95

53

[SAMDEV.LOC]PICKF
C. WONG MAY 1

PIK FILE

>>>>>> NEW FORMAT VERSION AUGUST 1987
>>>>>> 5 CHARACTER STATION ID

The PIK file is the input file to and also the output file (o version newer) from the CANSESS MULTILAYER epicenter location prog (LOC). SAM PIK (or PK4) command generates a PIK file automatically the event. These PIK files can be modified/created by the EPK progr or by the DEC text editor EDT.

It contains four types of records:

1. ESR - earthquake solution record.
2. ECR - earthquake comment record.
3. ODR - observed data record.
4. CDR - calculated data record.

The ESR (if it exists) has to be the first record in the file. It is an output record from the previous LOC. Some of its fields can b to modify the LOC parameters for the next run. The ECR records come the ESR, they have to be before the first ODR. There is one ODR for station with picked information. This is an input record, LOC progr not modified any field on this record. The CDR contains only the ca results for this station. If it exists, it is always right after it corresponding ODR. You can have as many ECR, ODR or sets of ODR & C you want. However, the current EPK program can handle total of 100 at most, and the LOC program is dimensioned only 100 for all the ph picked. If you have any deficuilties with the above limitations, ple me know. The detail layout of these records starts on the next pag

EARTHQUAKE SOLUTION RECORD (ESR)

(solution record is defined as having "+" or "-" on col.1 and "M" ON

CLOUNS	ENTRY	FORMAT	DEFINITION
1-1	+	A1	PRIME SOLUTION BY EPB
	-		PRIME SOLUTION BY OTHER AGENCY
	0		SUPPLEMENTARY SOLUTION
2-7	45.233	F6.3	NORTH LATITUDE, DEGREES
8-15	-123.300F8.3		LONGITUDE, DEGREES
16-16		A1	HYPOCENTRE QUALITY INDICATOR.
	O		POOR QUALITY SOLUTION
	F		GOOD QUALITY SOLUTION
17-17		I1(A1)	OBSERVED DATA FORMAT INDICATOR,
	BLANK		PRE-1979 DATA FORMAT USED
	1		1979 DATA FORMAT USED.
18-19		A2	PRIME MAGNITUDE TYPE
	ML		RICHTER
	MLE		EBEL
	MN		NUTTLI (DEFAULT)
	MLG		H. & K.
	MB		BODY-WAVE
	MS		SURFACE WAVE
	MC		CODA LENGTH
20-20	BLANK		
21-23	3.1	F3.1	AVERAGE MAGNITUDE VALUE (MN)
24-24	BLANK		
25-26	18	I2	ORIGIN TIME HOUR, U.T.
27-28	23	I2	ORIGIN TIME MINUTE
29-31	323	I3	ORIGIN TIME SECOND*10 (OR F3.1 IN SECOND)
32-32	BLANK		
33-34	12	I2	DAY
35-36	03	I2	MONTH
37-40	1979	I4	YEAR
41-41	BLANK		
42-42	2	I1	STANDARD DEVIATION ORIGIN TIME, SECONDS

43-47	0.122	F5.3	STD IN LATITUDE, DEGREES
48-52	0.333	F5.3	STD IN LOGITUDE, DEGREES
53-53	BLANK		
54-56	0.3	F3.1	STD IN MAGNITUDE FOR EPB
	XXX	A3	AGENCY CODE FOR EXTERNAL MAG, DEPENDS ON COL.
57-59	34	I3	NUMBER OF STAIONS USED FOR HYPOCENTER
60-62	14	I3	NUMBER OF PHASES USED FOR THIS HOPOCENTER.
63-65	14	I3	NUMBER OF AMPLITUDE USED FOR MAGNITUDE.
66-69	0.33	F4.2	RMS OF HYPOCENTER SOLUTION, SECONDS.
70-70		A1	SUOLUTION TYPE INDICATOR
	BLANK		FIXED DEPTH.
	Z		FREE DEPTH
	X		NO ACTION FOR THE WHOLE FILE
	N		ASSIGNED HYPOCENTER AND TIME
	H		ASSIGNED HYPOCENTER, BUT CALCULATED ORIGIN TIME.
71-71		I1	AGENCY CODE
	1		USGS
	2		EPB
	3		PGC
	4		SEA UNIVERSITY OF WASHINGTON
	5		NEIS NATIONAL EARTHQUAKE INFORMATION CENTER
	6		ISC INTERNATIONAL SEISMOLOGICAL CENTER
	7		LDGO LAMONT-DOHERTY GEOLOGICAL OBSERVATORY
	8		WES WESTON GEOPHYSICAL OBSERVATORY
	9		UAGI UNIV. OF ALASKA, GEOPHYSICAL INSTITUTE
72-76	18.33	F5.2	FOCAL DEPTH, KM IF AND ONLY IF COL.=Z, FREE DEPTH SOLUTION
77-80	3.0	F4.1	STD IN DEPTH, KM (OR I4 FORMAT IN 100-METERS)
81-82	03	I2	MODEL NUMBER
83-85	ML=		AVERAGE ML
86-88	1.3	F3.1	RICHTER MAGNITUDE
89-91	008	I3	NUMBER OF STATIONS USED TO CALCULATE AVERAGE M
92-92	1	I1	MULTILAYER HYPO SIMULATION FLAG, 0-OFF, 1-ON.
93-93		A1	
	F		FELT ! FL001
	N," "		NOT FELT ! FL001
94-95	10	I2	NB. OF ASSOCIATED EVENTS ! FL001
96-96		A1	
	E," "		EARTHQUAKE " ! FL001
	B		BLAST

	R		ROCKBURST
	P		POSSIBLE ROCKBURST
97-100	3.56	F4.2	S VELOCITY USED BY SINGLE LAYER MODEL ! FL00
<			FORMAT(A1,F6.3,F8.3,2A1,A2,1X,F3.1,1X,I2.2,I2.2,I3.3,1X,2I2.2,
<	&		'19',I2.2,1X,I1,2F5.3,1X,A3, ! FOR AG
<	&		'19',I2.2,1X,I1,2F5.3,1X,F3.1, ! FOR ST
<	&		3I3.3,F4.2,A1,I1,F5.2,I4,T81,I2,'MC=',F3.1,I3,I1
<	&		A1,I2,A1,F4.2)

EARTHQUAKE COMMENT CARDS (ECR)

COLUMNS	ENTRY	FORMAT	DEFINITION
1-40	40A1		EARTHQUAKE DESCRIPTION IN ENGLISH
41-80	40A1		EARTHQUAKE DESCRIPTION IN FRANCH

OBSERVED DATA RECORD (ODR)

COLUMNS	ENTRY	FORMAT	DEFINITION
1-5	OTT	A5	STATION CODE ! FLO01
6-7	79	I2	YEAR ! FLO01
8-9	12	I2	MONTH ! FLO01
10-11	23	I2	DAY ! FLO01
12-13	12	I2	HOUR, U. T. ! FLO01
14-15	14	I2	MINUTE OF 1ST RECORDED P PHASE, NOT NECESSARY @
16-16		A1	INSTRUMENT CODE ! FLO01
	P		SHORT PERIOD INSTRUMENT READ
	L		LONG PERIOD INSTRUMENT READ, AMP. & 1ST MOTION
17-17	BLANK		! FLO
18-18		A1	PN WEIGHT ! FLO
	" "		USED IN CALCULATION
	X		NOT USED IN CALCULATION
19-19		A1	PN QUALITY DESIGNATOR ! FLO
	A		SHARP CLEAR BEGINNING (+- 0.25 SEC.)
	B, " "		GOOD BEGINNING (+- 1.0 SEC.)
	C		WEAK POOR BEGINNING (+- 4.0 SEC. OR MORE)
	X		PHASE NOT USED IN SOLUTION, LARGE RESIDUAL.
	0		PHASE NOT READ
20-21	14	I2	MUNITE OF PN ARRIVAL ! FLO
22-25	2341	F4.2	SECOND OF PN ARRIVAL ! FLO
26-28	CNM ??	3A1	FIRST MOTION OF PN ARRIVAL ! FLO
29-33	+0.03	F5.0??	TIME CORRECTION ! FLO
34-34		A1	PG WEIGHT ! FLO
	" "		USED IN CALCULATION
	X		NOT USED IN CALCULATION
35-35	A, B. .	A1	PG QUALITY DESIGNATOR, SEE 16 ! FLO
36-37	14	I2	MINUTE OF PG ARRIVAL ! FLO
38-41	264	F4.2	SECOND OF PG ARRIVAL ! FLO
42-44	DSE	3A1	FIST MOTION OF PG ARRIVAL ! FLO
45-45		A1	SN WEIGHT ! FLO

" " USED IN CALCULATION
 X NOT USED IN CALCULATION

46-46	A,B...	A1	SN QUALITY DESIGNATOR, SEE 16	! FL0
47-48	14	I2	MINUTE OF SN ARRIVAL	! FL0
49-52	52	F4.2	SECOND OF SN ARRIVAL	! FL0
53-53	" "	A1	LG WEIGHT USED IN CALCULATION NOT USED IN CALCULATION	! FL0
54-54	A,B...	A1	LG QUALITY DESIGNATOR, SEE 16	! FL0
55-56	14	I2	MINUTE OF LG ARRIVAL	! FL0
57-60	589	F4.2	SECOND OF LG ARRIVAL	! FL0
61-61	BLANK			! FL0
62-64	031	F3.2	PERIOD OF MAX. TRACE AMPLITUDE, SECOND	! FL0
65-68	150	F4.0	MAGNIFICATION OF INSTRUMENT AT GIVEN PERIOD, I	
69-72	125	F4.1	TRACE AMPLITUDE(ONE-HALF MAX. PEAK-TO-PEAK) I	
73-73	BLANK			! FL0
74-77		I4	DURATION IN SECONDS.	! FL0
78-79	BLANK			! FL0
80-80		I1	MAGNITUDE CODE	
	BLANK		AMPLITUDE SUITABLE FOR NUTTLI OR RICHTER SCALE	
	1		AMPLITUDE SUITABLE FOR RICHTER ONLY, CORDILLER	
	2		AMPLITUDE SUITABLE FOR EBEL	
	3		AMPLITUDE UNRELIABLE, NOT USED FOR MAGNITUDE	
	4		AMPLITUDE SUITABLE FOR HUEN & KISCO	
	5		AMPLITUDE SUITABLE FOR MS SCALE ONLY	
	8		SN AMPLITUDE READ, USE RICHTER SCALE ONLY BEYOND 600 KM IF REQUIR	
81-83	BLANK			! FL0
84-85	15	I2	MINUTE OF THE MAX. AMPLITUDE	! FL0
86-89	155	F4.2	SECONDS OF THE MAX. AMPLITUDE	! FL0

< FORMAT(A5,5I2,A1,1X,2A1,I2,F4.2,A3,F5.0,2A1,I2,F4.2,A3,2A1
 < & I2,F4.2,2A1,I2,F4.2,1X,F3.2,F4.0,F4.1,1X,I4,2X,I
 < & 3X,I2,F4,2)

CALCULATED DATA RECORD (CDR)

COLUMNS	ENTRY	FORMAT	DEFINITION	
1-5	OTT	A3	STATION CODE	! FL001
6-6	BLANK			
7-8	NW	A2	QUADRANT OF STATION	
9-9	BLANK			
10-13	1305	I4	EPICENTRAL DISTANCE, KM	
14-15	KM	A2	RECORD FLAG	
16-16	BLANK			
17-18	28	F2.1	PN WEIGHT USED FOR CALCULATIONS	! FL001
19-19	BLANK			
20-23	0107	F4.2	PN RESIDUAL, SECOND	! FL001
24-24	BLANK,#	A1	LARGE RESIDUAL FLAG	! FL001
25-27	235	I3	AZIMUTH TO STATION, DEGREES	
28-30	049	I3	EMERGENT ANGLE PN POSITIVE PG NEGATIVE	
31-34	BLANKS			
35-36	14	F2.1	PG WEIGHT	! FL001
37-37	BLANK			
38-41	-091	F4.2	PG RESIDUAL, SECOND	! FL001
42-42	BLANK,#	A1	LARGE RESIDUAL FLAG	! FL001
43-45	BLANKS			
46-47	07	F2.1	SN WEIGHT	! FL001
48-48	BLANK			
49-52	0024	F4.2	SN RESIDUAL, SECOND	! FL001
53-53	BLANK,#	A1	LARGE RESIDUAL FLAG	! FL001

54-55 07 F2.1 SG WEIGHT ! FL001

56-56 BLANK

57-60 -434 F4.2 SG RESIDUAL, SECOND ! FL001

61-61 BLANK,# A1 LARGE RESIDUAL FLAG ! FL001

62-63 BLANKS

64-70 0001356 I7 GROUND VELOCITY, NM/SEC

71-71 BLANK

72-73 35 F2.1 RICHTER OR SURFACE WAVE MAGNITUDE

74-75 ML,MS A2 MAGNITUDE DESIGNATOR

76-77 34 F2.1 NUTTLI MAGNITUDE

78-79 MN A2 MAGNITUDE DESIGNATOR

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<  FORMAT(A5,1X,A2,1X,I4.4,'KM',1X,F2.1,1X,F4.2,A1,2I3.3, >
<  &          4X,F2.1,1X,F4.2,A1,3X,F2.1,1X,F4.2,A1,F2.1,1X,
<  &          F4.2,A1,2X,I7.7,1X,2(F2.1,A2)) >
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