



Scale 1:100 000. Elevation in feet above mean sea level. Contour interval 100 feet. Universal Transverse Mercator Projection. Topographic base: 1:50 000 scale, 1969 (updated 1973) by Surveys and Mapping Branch, Department of Energy, Mines and Resources, Canada. Includes 6 sheets: 105M/1 (Moose Lake), 105M/2 (Clarke Hills), 105M/3 (Sideslip Lake), 105M/4 (Woodburn Lake), and parts of 105M/5 (Ethel Lake) and 105M/6 (Francis Lake).

Note: Symbols with a ? suffix indicate areas not visited but extrapolated to contain this rock unit. Symbols with a subscript 'prime' (e.g., PC'H') lack paleontologic control and have no diagnostic lithologic features. Unit assignment is based upon local structural coherence only.

LEGEND

- QUATERNARY: Unconsolidated alluvial, colluvial, glaciofluvial and moraine gravel, sand, clay and organic deposits.
CRETACEOUS: McArthur pluton: White medium-grained, biotite, hornblende-phyric granite. Buff to grey dykes, sills and small plugs of apite and granite.
TRIASSIC or older: Hornblende-clinopyroxene meta-diorite to meta-gabbro.
DEVONO-MISSISSIPPIAN EARN GROUP: Undifferentiated black siltstone, mudstone and graphitic phyllite with lenses of chert pebble conglomerate (circle pattern), chert breccia with baritic matrix and dark grey, field limestone (hatch pattern).
PALEOZOIC (age range uncertain): 'NOGOLD ASSEMBLAGE': Beige weathering, medium bedded, quartz grit sandstone with interbedded khaki and dark grey mudstone-phyllite; P'm, Maroon and apple green slate, mudstone contains single thin beds (hatch pattern); P'hc, Chloritic gritty schist and foliated breccia; P'hd, Light grey weathering, thick bedded quartzite.
SILURIAN to LOWER DEVONIAN ROAD RIVER GROUP: Undifferentiated black siltstone, mudstone and chert; SDR1, Thick bedded black chert; SDR2, Interbedded black mudstone, dark brown and black quartz siltstone, dark grey shale, with minor lenses of dark grey limestone (hatch pattern); SDR3, Black quartz-feldspar arenite (of igneous provenance).
LATE CAMBRIAN to EARLY ORDOVICIAN(?): Olive and brown weathering siltstone with black laminae, brown sandstone with thin interbeds of black, bioturbated chert; single occurrence of olive grey, clastic limestone; COp, purple and green-white nodular limy siltstone-schist.
PRECAMBRIAN-LOWER CAMBRIAN HYLAND GROUP: may include unexposed CO at top. Undivided, except where lithologic separation as noted. PChm, maroon and green mudstone and siltstone, locally with white quartz sandstone interbeds; PChs, Brown weathering medium to coarse-grained quartz mica schist and metasandstone (psammite) with thin, discontinuous white limestone (hatch pattern).

SYMBOLS

- Geological contact (defined, approximate, assumed)
Fault (defined, approximate, assumed)
Fault (in cross-section: fault block moving toward, away from observer)
Thrust fault (defined, approximate, assumed)
Axial surface trace of folds
Folds (syncline, anticline, overturned anticline, overturned syncline; arrow on trace indicates direction of plunge)
Bedding (tops known, tops unknown, overturned)
Prominent foliation (strike and dip) and lineation (trend and plunge)
Limestone beds (age corresponds to enclosing unit)
Fossil and micro-fossil locality (see legend - Age Constraints)
Mineral occurrence (Yukon MINFILE reference number)
Line of cross-section
Form-line to outline surface of resistant layer, indicating structural trend
Thermal alteration halo around McArthur pluton

NOTES

The area shown on this map includes Late Proterozoic through Mississippian strata which potentially host sedimentary-exhalative zinc-lead mineralization like that of the Anvil and Macmillan Pass districts. The Paleozoic units described here have not been distinguished on previous maps. Furthermore, a middle Paleozoic maroon argillite (the Nogold Assemblage), similar in appearance to the maroon argillite of the Hyland Group (Late Proterozoic to Middle Cambrian) has recently been discovered. Although the extent of this new unit is unknown its existence may eventually result in a reinterpretation of the paleogeography and structure of this tectonic belt.
Mayo map area (105M) lies within the Selwyn Basin at the northern edge of the Selwyn Fold Belt, which was deformed in Middle Jurassic and Early Cretaceous time (Roots, 1991). All rocks have been tectonically displaced from their place of origin. The Robert Service Thrust which underlies most of this area at shallow depth (less than 5 km) comes to surface 30 km north (Roots and Murphy, 1992b). The thrust sheet of northward-displaced strata is truncated by Tintina Fault. Rocks in the extreme southwest area of this map are part of the Nisutlin allochthonous assemblage of Paleozoic or Mesozoic age (Tempelman-Kluit, 1979, 1984).
In the map area in place rock exposure is sparse and steep bush with talus hinders foot travel below 1200 m (4000') elevation. The distribution of pro- and postglacial deposits is modified from Hughes (1982).

'Nogold Assemblage' (P'n provisional) is a dominantly siliciclastic succession of probable mid-Paleozoic age. In the headwaters of Nogold Creek a composite section includes: at the base more than 400 m of beige weathering fine-grained meta-sandstone including chlorite schist (up to 40 m) at the top, overlain by purple argillite (100 m, locally thickened) with isotatic, centimetre-thick beds of black limestone containing rare late Silurian to Devonian organic remains (F'), capped by clean, uniformly fine-grained quartz sandstone. Bedding is rarely observed and top-indices are unknown.
The contact of the Nogold Assemblage with older and younger rocks has not been identified. The brown meta-sandstone appears to conformably overlie black chert of Road River Group northeast of Sideslip Lake (63°10'N 135°25'W). Meta-sandstone of the Nogold Assemblage resembles that of the Hyland Group and the contact is obscure in a succession near the headwaters of Crooked Creek (63°12'N 135°50'W). The Nogold Assemblage is presently thought to be a facies-equivalent of Earn Group, although the two units are nowhere in contact. However the Nogold Assemblage may underlie a larger area than shown because its contact with PChs has not been determined. On this map the Nogold Assemblage is restricted to laterally continuous lithology in the area of (F'), and adjacent areas considered to have been offset by younger faults.

McArthur pluton belongs to the Tombstone suite (Woodsworth et al., 1991) on the basis of age and mineralogy. Unlike other intrusions of the suite it has an irregular shape and is oriented parallel to the Tintina Trench. It contains nodular fragments of hornfelsed SDR-DMc strata. The thermal metamorphic aureole, 500 to 1500 m wide, contains andalusite, staurolite and chloritoid 'spotted' slates. The granite near Woodburn Lake truncated by Tintina Fault is poorly exposed and reported by Bostock (1947) to be coarse-grained, non-porphyritic, and more altered than other intrusions of Mayo map-area.

STRUCTURE

A gradual transition north of Nogold Creek headwaters and across the Clarke Hills separates open, upright folds in the south from penetratively foliated, isoclinal and recumbent folds in the north. This transition may reflect exposure of a deeper structural level in the north.
The micaceous and friable Hyland Group meta-sandstone commonly displays a mineral stretching lineation (typically plunging northwest) and locally a fine crenulation. Overlying shale units are pervasively cleaved, and steeply dipping cleavage is axial planar to right, nearly isoclinal folds that verge both north and southward on northwest- and southeast-plunging axes. These folds reflect the northward-ward (and possibly coincidental) motion of the Robert Service Thrust in Late Jurassic or Early Cretaceous time (Roots and Murphy, 1992b).

Siliciclastic rocks of the Nogold Assemblage are characterized by a flaser fabric in which are locally preserved isoclinal, rootless folds. Although thick layering resembles bedding in north-facing cliffs, the strata has been internally thickened by layer-parallel thrusts and folds. Four Rams Bluff (63°12'N 135°35'W) is a south-facing exposure of the vertically complexly (ramped west-verging chevron folds at all scales) that belies this poorly exposed and enigmatic unit.

Truncation of structural trends (such as north of Sideslip Lake and south of North Crooked Creek) indicate faults buried beneath unconsolidated valley sediments. The northwest trend of many straight faults suggests that they may have released strain associated with Late Cretaceous and Tertiary movement of Tintina Fault.

MINERAL OCCURRENCES

- 39. Sideslip W, Cu, Zn - skarn unknown
40. Great Horn Ag, Pb - vein
41. Ram Cu-sulphide unknown
42. Hot Spring Cu-sulphide unknown
43. Lost Wernicke Copper Cu, W, Mo - skarn
45. Able (Dope)
51. Friesen

Numbered with Yukon MINFILE (105M) reference numbers (updated 1992). Occurrences #39, 40, 42 and 51 occur within the contact metamorphic aureole of the McArthur pluton. Chalcopyrite and other sulphides are relatively common in the aureole, particularly near discontinuous limestone pods. Lost Wernicke Copper (#43) may be a legend or erroneously located from stories of a large, low grade deposit in the Dawson Range. The Road River-Earn Group strata at the northern edge of Selwyn Basin hosts numerous stratobanded zinc-lead occurrences. The nearest that have been extensively explored are at Dromedary Mountain (105L M/52, 54), 25 km southeast of Clarke Peak. Rams Bluff (#42) was not visited but there are probably numerous showings of this type among the low hills and steep ravines on the flanks of the Tintina Trench.

McArthur hot springs (in 1993 protected as wildlife refuge and under selection by the Selkirk Band) is of great natural beauty. An area at least 100 m x 50 m contains numerous cold water springs and at least three springs of >40° C (estimated 3-10 l/min, 1992). The hot springs appear to be re-circulated groundwater driven by residual heat of the adjacent McArthur pluton, because no faults are apparent and the Tintina Fault is 10 km southwest. Although iron oxide coatings are present limonite deposits appear minimal in the 40 m x 60 m clearing around the hot springs. At other margins of the McArthur pluton, such as southwest of Grey Hunter Creek fenestrate talus may indicate abandoned springs. Limonite-cemented breccia that pre-dates Holocene glaciation on the south side of Clarke Peak (Able/Dope, #45) was described by Bostock (1947).

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INDEX TO MAP AREAS table with columns: LARSEN CREEK (116A), NASH CREEK, NADALEEN RIVER (105G), Green, 1972; Green, 1972; Blussion, 1974; McQUESTEN (115F), MAYO (105M), LANSING (105N), Bostock, 1964; Roots and Murphy, Blussion, 1974; CARMACKS (115I), GLENLYON (105L), TAY RIVER (105K), 1992a; Tempelman-Kluit, 1984; Campbell, 1967; Gordey and Irwin, 1987.

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GEOLOGICAL MAP OF SOUTHERN MAYO MAP AREA (105M/1,2,3,4 and parts of 105M/5 and M/6) YUKON
by Charles F. Roots
Canada/Yukon Mineral Development Agreement / Geoscience Office and Geological Survey of Canada (Contribution #33494)

