

DESCRIPTIVE NOTES

The Timmins map area of about 186 000 km² in Ontario and Quebec south of James Bay, includes parts of the Precambrian Superior and Southern provinces of the Canadian Shield, and the Phanerozoic Moose River Basin. The Canadian Shield, the peneplained Abitibi Upland, has low relief and extensive cover of Pleistocene till and lacustrine deposits. Fluvial Phanerozoic strata of the Moose River Basin form the southern part of the Hudson Bay Lowland, an area of bog and swamp poorly drained by river systems incised in thick Pleistocene glaciogenic deposits.

The Superior Province Subprovince consists mainly of tuffaceous volcanic rocks of the Quetico, Opatica, Nemiscau and Pontiac metamorphic subprovinces, and the Kapuskasing Structural Zone, and in the Partridge River area, Ontario, and along the Nottaway River, Quebec are probably part of the largely buried eastern extension of the volcano-plutonic Wabigoon Subprovince. The Quetico, Opatica and Nemiscau subprovinces consist mainly of migmatitic paragneiss, tonalitic and granitic orthogneiss, minor mafic gneiss of volcanic origin and abundant biotite granite and peraluminous, palineogenic intrusions. Medium grade metasediments and metavolcanics in southern Quebec Subprovince have relatively well-preserved primary structures such as bedding and pillows. These subprovinces, their major lithological units and the internal geology generally trend east-west. Eastern Wawa Subprovince has several remnant metamorphic belts (e.g. Michipicoten belt) but consists mainly of tonalitic gneiss with variable amounts of mafic and metasedimentary xenoliths of probable supracrustal origin. Gneiss and greenstone are intruded by foliated to foliated quartz diorite to granite and syenite plutons with ages of 2677-2623 Ma. Polyphase deformation and diagenetic domal structures with major culminations at approximately 30 km intervals. Metamorphic grade increases eastward from lower greenschist-amphibolite facies in the Michipicoten, greenstone to upper amphibolite facies in the Kapuskasing Structural Zone.

Kapuskasing Structural Zone, characterized by northeast-trending, north-west-dipping granitic gneiss, a large layered gabbro-anorthositic complex, and positive gravity and magnetic anomalies, trends the east-west trends of adjacent subprovinces. It is bounded on the east by the western gneiss belt and on the west by the LePage and Saganash Lake faults. The zone represents the basal part of an oblique crustal section about 120 km wide, from low grade supracrustal rocks of the Michipicoten belt north of Hearst and at depths of 5-10 km (2-3 kb), through amphibolite gneiss of eastern Wawa Subprovince formed at 15-20 km (4-6 kb), to granulite of the zone, subjected to pressures of 7-9 kb at depths of 30 km. This section was exposed by major displacement on a west-dipping thrust, of which the Ivanhoe and Bad River faults are remnants, followed by normal movements on subparallel faults, the LePage and Saganash Lake structures, that broke the thrust sheet into several blocks. Locally allochthonous sedimentary and volcanic rocks of the Kapuskasing Zone are 2765-2611 Ma; metamorphic zircon systems generally become younger eastward, from 2690 to 2611 Ma. Seismic refraction experiments indicate that Moho depth increases eastward from about 40 km beneath eastern Wawa Subprovince to over 44 km beneath the southern Kapuskasing Zone and decreases to about 35 km beneath western Abitibi Subprovince.

Abitibi Subprovince is dominated by the Abitibi greenstone belt which comprises several major volcanic megacycles each consisting of a lower komatiitic-tholeiitic sequence, a middle tholeiitic sequence and upper diverse tholeiitic, calc-alkalic mafic to felsic sequences; locally a fourth, alkali-shoshonitic division is present. The lower and middle sequences form extensive submarine lava plains, whereas the upper mixed sequences form central volcanic complexes of subaqueous and subaerial pyroclastics and flows. Turbiditic, volcanogenic wacke and conglomerate units are interstratified with and form aprons about the volcanic accumulations, notably the central complexes. Locally allochthonous sedimentary and volcanic sequences unconformably overlie older rocks. Sequences of this latter type, notably the Timiskaming Group, occur along major faults such as the Kirkland Lake-Cadillac Fault and may have been deposited in rift or pull-apart basins. Abitibi volcanic rocks have U-Pb zircon ages of 2730-2697 Ma; the age of Timiskaming Group volcanics is possibly about 2695 Ma.

Plutonic rocks include early, pre- to syn-kinematic, in part synkinematic, and foliated quartz diorite, tonalite and granodiorite suites and younger, syn- to post-kinematic, foliated to massive granodiorite, granite and syenite suites. Ages of Abitibi plutonic rocks are 2718-2676 Ma; many late- to post-kinematic plutons have ages of about 2680 Ma, and constrain the last major deformation and metamorphism in this part of the Superior Province to 2700-2680 Ma.

Early ductile folding and faulting formed major east-west trending anticlines and synclines, complex interference folds and weakly deformed domains of various sizes, each with distinctive lithological and structural characteristics, and separated by narrow, linear high-strain zones and faults. Later deformation was increasingly brittle, culminating in large-scale transcurrent faulting, deposition of Timiskaming-type sequences, shearing, alteration and formation of gold-bearing quartz-carbonate vein systems.

Abitibi greenstone belt is primarily in low greenschist facies. Several large areas of subgreenschist (prehnite-pumpellyite) facies rocks are present but upper greenschist and amphibolite or hornblende-hornfels facies are present only in relatively narrow aureoles about some intrusions. Gravity data over part of Abitibi Subprovince indicate an average upper crustal density of 2.73 g/cm³ corresponding to average tonalitic composition. Greenstone belts have been interpreted to extend down to about 5-8 km, and tabular granitoid batholiths are 3-6 km thick.

Northern Pontiac Subprovince consists mainly of tuffaceous wacke with minor units of conglomerate and komatiitic-tholeiitic volcanics, metamorphosed to low and medium grade, and cut by granitic and syenitic granulate facies gneiss. Metavolcanics and plutonic rocks of ages of about 2713, 2920 and 3000 Ma. The grade of Barrovian-type metamorphism increases rapidly and uniformly southward from low greenschist to middle and upper amphibolite facies whereas apparent geotherms of paragneiss and orthogneiss plus granitic, pegmatitic and syenitic dykes, sheets and plutons are present. Polyphase deformation formed east-trending, south-verging isoclinal and recumbent folds and thrust faults.

In the northern extension of the Cobalt Embayment of the Southern Province, glaciogenic rocks of the Acheban Gowganda Formation, Huron Supergroup, unconformably overlie Precambrian rocks and occupy several narrow flood-like, north-trending basins that may have been controlled by pre-existing faults. Superior Province is intruded by Acheban and Helikian alkalic rock-carbonatite complexes, notably along the Kapuskasing Zone, and by several dyke swarms, including the Archean Malachew and Hearst, the Acheban Priestess, and the Helikian Sudbury and Abitibi swarms. Cretaceous(?) kimberlitic dykes intrude the northern Kapuskasing Zone.

Strata of the Moose River Basin are Ordovician to Cretaceous. The basin is bounded on the south by Precambrian rocks of the Fraserville Arch, an east-northeast-trending tectonic element that underwent vertical movements in the Paleozoic and Mesozoic which truncated strata by erosion and faulting.

Oldest strata are Late Ordovician, Gamachião(?) laminated dolostone with interbeds of feldspathic sandstone, shale and minor gypsum lenses of the Red Head Rapids Formation. The Fraserville Arch shed terrigenous sediments into the basin producing channel fill deposits in tidal flat and supratidal carbonates.

Silurian, mid- to late Llandovery rocks discontinuously succeed Ordovician strata and locally overlap the latter in the Precambrian. They consist of thin to massive bedded, fossiliferous limestone of the Severn River and succeeded by Silurian, Ludlow, Ludlow and Gedinnian redbed shale, siltstone, sandstone, dolostone and minor gypsum of the Kenogami River Formation. Kenogami River Formation is fault contact with Precambrian rocks of the basin, but unconformably overlies the latter in the southern part of the basin and lies on the Precambrian. The Severn River, Ewan River and Kenogami River formations are a fault contact with Precambrian rocks of the Fraserville Arch.

Uplift of the arch in the Silurian, Wenlock and Devonian, Gedinnian, Siegenian and Emsian red and grey conglomerate, sandstone, siltstone and shale of the SEXTANT Formation which was deposited unconformably on Precambrian, Ordovician and Silurian rocks are intertongued with Emsian marine cherty limestone of the Stopping River Formation in the lower part of the basin. Reefal limestone of the Kwatobahegan Formation occurred during Moose River deposition. Associated faulting triggered extensive salt dissolution, collapse and development of megabreccias at the same stratigraphic horizon as the Mackinac Breccia of northern Michigan. Limestone of the Murray Island Formation lies unconformably on the Moose River. Gweltaun rocks consist of fossiliferous limestone and shale of the Williams Island Formation, followed by Frasnian and Famennian black petroliferous shale and minor limestone of the Long Rapids Formation.

Throughout the Siegenian Stage, seaways occupied similar channels onto the Hudson Platform from the central and eastern of the arch, over the Fraserville Arch, the latter connection being largely continuous until the end of the Devonian. Upper Devonian black shales of the basin are similar to equivalent rocks in the Great Lakes region and both were presumably derived from the Appalachian Orogen.

Renewed uplift of the Fraserville Arch in the Jurassic and the Early Cretaceous provided a source for sand and clay of the Missisquoi beds and Mattagami Formation deposited along the southern margin of the Moose River basin. Lamprophyre dykes and sills (Kf ages of 125 Ma) that intruded Siegenian to Eifelian strata along the Abitibi River and presumably eugeine in the basin were associated with Cretaceous movements of the arch.

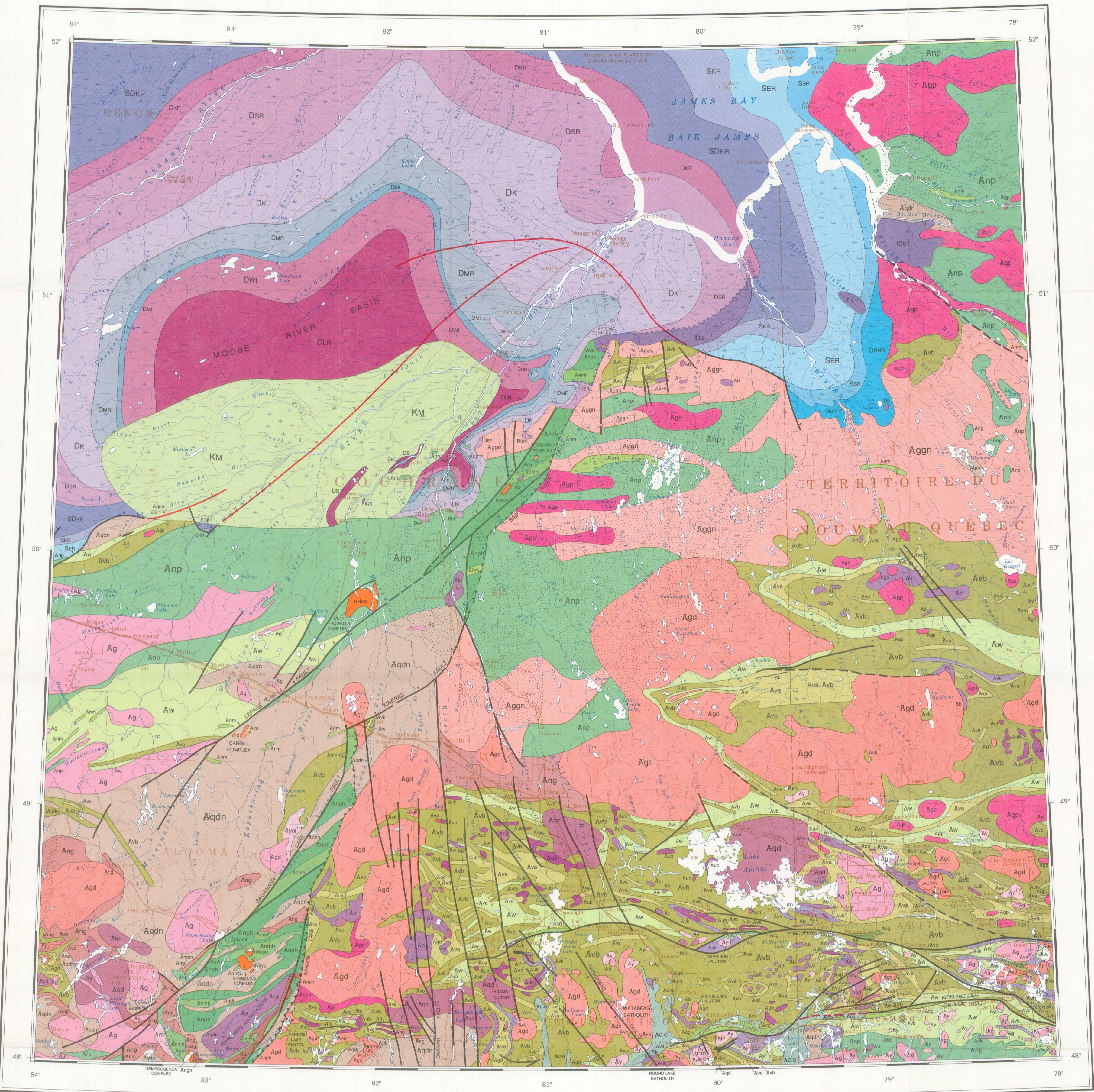
The Timmins map area is an important mining region in Canada with major production of gold and base metals, chiefly from the Timmins, Kirkland Lake and Noranda areas of the Abitibi greenstone belt. The Moose River Basin has potential for oil, gas, lignite, lead and zinc, gypsum, limestone, clay and sand.

Selected references

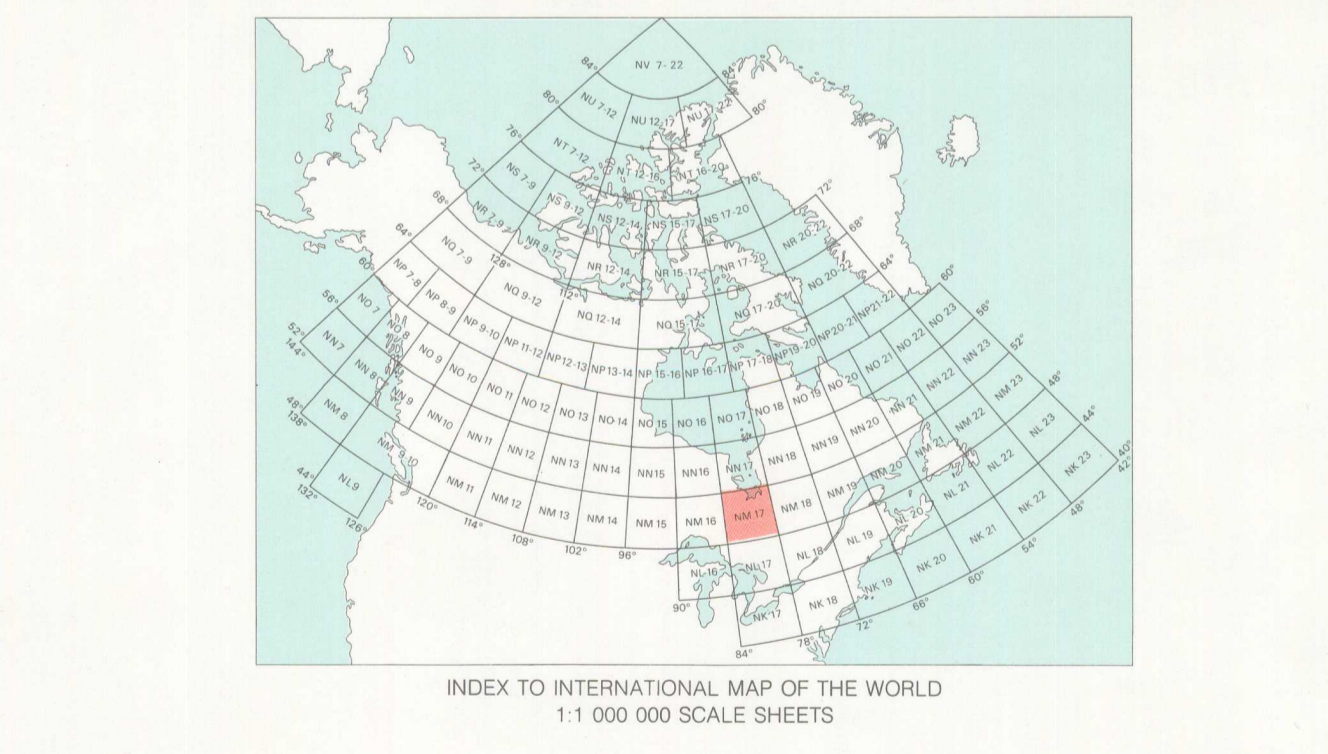
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GEOLOGY TIMMINS ONTARIO-QUÉBEC GEOLOGICAL ATLAS, MAP NM-17-G SHEET 1 OF 5 GENERAL CO-ORDINATOR: A.V. OKULITCH Scale 1:1 000 000 - Échelle 1/1 000 000 Lambert Conformal Conic Projection Standard Parallels 48°40'N and 51°20'N © Crown copyrights reserved



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