

MONASHEE MOUNTAINS

QUATERNARY

Qa Fill, alluvium

PROTEROZOIC

UPPER PROTEROZOIC

WINDERMERE SUPERGROUP

MICA CREEK SUCCESSION (Pg - Ppa)

Ppa Mica Creek succession: pelite, psammite, micaceous psammite, amphibolite, pelitic schist, and gneiss; local matrix-supported pebbles to boulder conglomerate with clasts of quartzite, marble, calcisilicate, and granitoid rocks; minor calcareous psammite, marble, and micaceous quartzite; minor pyroxene; amphibolite; calcisilicate; calcisilicate

Pp Lower calcareous unit: marble, calcareous psammite, calcisilicate, pelite, minor semipelite and amphibolite (stratigraphic diagram only)

Pp Lower pelite unit: pelitic schist, semipelite, minor local amphibolite and marble

Pp Lower gneiss unit: pelitic schist, conglomerate, minor psammite, gneiss and amphibolite

MALTON COMPLEX

PROTEROZOIC

UPPER PROTEROZOIC

WINDERMERE SUPERGROUP

COVER EQUIVALENT TO MALTON GNEISS: pelite, semipelite, minor gneiss, local quartzite and/or conglomerate at base (similar to lower pelite unit of Mica Creek succession and lower Mica Group)

Ppc Basal conglomerate unit: Matrix- to clast-supported pebbles to cobble conglomerate with brownish-biotope calcareous matrix. Clasts of underlying quartzite, amphibolite, and granitoid units. Abundant background quartz veins appear as clasts. Ppc unconformably overlies quartzite or gneiss and forms the base of the Windermere Supergroup (stratigraphic diagram only)

UPPER PROTEROZOIC?

Pp Quartzite unit: fine-grained foliated to mylonitic muscovitic quartzite, minor quartz-pebble conglomerate, orthoquartzite

LOWER PROTEROZOIC

Pmg3 Malton Gneiss UNIT 3: foliated granitic augen orthogneiss, probably 1870 Ma, minor amphibolite unit (Pmg3 of northern Malton gneiss; see Note 6)

Pmg2 Malton Gneiss unit 2: foliated biotite-hornblende granodioritic to dioritic augen orthogneiss; U-Pb zircon ages are 1987 Ma and 2060-2100 Ma (Note 7)

Pmg1 Malton Gneiss unit 1: layered quartzose paragneiss, mafic amphibolite gneiss, and amphibolite, locally intruded by granitic gneiss of Pmg3, and augen gneiss of Pmg2

CARIBOO MOUNTAINS

TERTIARY

INTRUSIVE ROCKS

Tg Muscovite, biotite granite, quartz monzonite (see Note 9)

MESOZOIC

INTRUSIVE ROCKS

Jkpgg Jurassic pegmatite dykes and dykes

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PROTEROZOIC

UPPER PROTEROZOIC

WINDERMERE SUPERGROUP

KAZA GROUP

Ku Massive conglomeratic sandstone (unit); sandstone; minor grey green to grey pelite; intervening grey-green and lesser black pelite map units (stratigraphic diagram only)

Kp Old Fort Plain Formation Equivalent: pink quartzite; calcareous cross-bedded psammite; fine-grained pelite; calcareous psammite; rounded pelite, and buff sandy marble; black graphic phyllite (stratigraphic diagram only)

Kpm Middle Kaza Group

Kp Massive conglomeratic sandstone (unit); sandstone and conglomeratic sandstone; minor grey green to grey pelite; intervening grey-green and lesser black pelite map units; minor sulphide-rich graphic pelite near the base

Kps Lower Kaza Group

Kp Carbonate unit: Division Carbonate of Peil, (1984), cycles of semipelite and gneiss, graphic pelite, impure buff-colored sandy marble

Kp2 Black graphic phyllite, dark grey to black silty pelite; minor slate

Kp1 Semipelite, biotite schist; minor gneiss, calcareous amphibolite

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NOTES

1. Geology of the Malton Range by Campbell (1968), Morrison (1982), Doucet et al. (1983), Dugal et al. (1985), and updated mapping by S. Simony. Geology of the Cariboo Mountains by Peil (1984), Dechenne (1986), Currie (1988), and updated mapping by P.S. Simony.

2. Bedding parallel to sub-parallel S. Isolation of the Rocky Mountains is equivalent to S. of the Monashee and Cariboo mountains; S. of the Rockies is equivalent to S. of the Monashee and Cariboo mountains (see Simony et al. 1980; Morrison, 1982; McConough and Simony, 1982; McConough, 1989).

3. The gneisses of the Malton Complex are typically separated from their cover by thermally annealed pre-metamorphic shear zones that indicate the Windermere cover rocks are detached from the basement (Morrison, 1982; McConough and Simony, 1982; McConough and Simony, 1989; McConough et al. 1982; McConough and Peil, 1991). Although the contacts are generally and thermally annealed, the Early Proterozoic gneisses of the Malton Complex are unconformably overlain by a pre-Windermere quartzite, which in turn is unconformably overlain by the basal conglomerate of the Windermere Supergroup (Murphy et al. 1991; McConough et al. 1991, 1992). Inward facing stratigraphy within exposures of cover rocks that are enclosed by basement gneisses suggest that the cover is deformed into large folds. These are represented as D₁ structures (possible westerly vergent, see D₁ fold closures have been identified in the Malton Complex). Above the Malton structures, possible westerly vergent, and related by north-south vergent D₂ structures (Simony et al. 1980). The slope ranges in the largest of these D₂ ranges, having a dip length of about 50 km (Rasside and Simony, 1982; Dugal et al. 1985; Sevgiry and Simony, 1989). The Malton Complex was probably active during formation of D₁ structures. Fabric analysis indicates that it was active during D₂, indicating north-westward translation of the lower rocks of the cover over the carbon basement of the Malton Complex (Morrison, 1982; Sevgiry et al. 1989). Note the continuation of the aureole over and silicified augenites with the Malton décollement.

4. The timing of metamorphism and syn-metamorphic north-south vergent D₂ deformation in the Valmont region is constrained by U-Pb zircon and zircon dates of Currie (1988), Sevgiry et al. (1988), Scarmel (1991), and Currie et al. (in prep.) to between 135-100 Ma (Early Cretaceous).

5. Orogen parallel (OP) D₂ stretching lineations are ubiquitous in the gneisses of the Malton Complex, and where not annealed consistently represent distinct tectonic events. In the Rockies, such structures (D₂ structures) are part of a distinct oblique-slip thrust system that carried basement gneisses of the Malton Complex into the Rocky Mountain belt (see McConough and Simony, 1982; McConough and Simony, 1989). OP stretching lineations also occur in conglomerates and in the Thunder River thrust zone in the Cariboo Mountains (McConough et al. 1982), and in the Windermere cover rocks in the Monashees (McConough et al. 1991).

6. The North Thompson-Alberda normal fault is a brittle fault with flow data indicative of west-side-down displacement and lowest Creek map sheet), but as they are not reported for this zone. However, if the segment of the fault south of Chemina Creek is essentially slip-free, then the Alberda River section of the fault must have a substantial dextral strike-slip component of displacement. Mesoscopic scale fault faults in the northern Malton Range are parallel to the strike of the Alberda segment of the NTFAN are dextral oblique slip normal faults (McConough et al. 1992).

7. U-Pb zircon and zircon dates for orthogneisses of the Malton Gneiss Complex are reported by McConough and Peil (1991). K-Ar dates for muscovite and hornblende in the Cariboo Mountains are reported by Currie (1988), Currie et al. (in prep.), and Hart and Rodick (1990, 1992). The dates for the Lempreire map sheet (83D/6) are:

A: zircon U-Pb 2060-2100 Ma (biotite-hornblende granodioritic augen orthogneiss, Malton gneiss unit 2; sample MGM 01, collected by MW).

B: zircon U-Pb 1870 Ma (biotite-hornblende granodioritic augen orthogneiss, Malton gneiss unit 3; sample MGS 01, collected by MW).

8. Dioritic to granodioritic layered augen gneisses of unit Pmg2 occur as map units in the southern Malton gneiss (Lampson and lowest Creek map sheet), but as they are not reported for this zone. However, if the segment of the fault south of Chemina Creek is essentially slip-free, then the Alberda River section of the fault must have a substantial dextral strike-slip component of displacement. Mesoscopic scale fault faults in the northern Malton Range are parallel to the strike of the Alberda segment of the NTFAN are dextral oblique slip normal faults (McConough et al. 1992).

9. The Malton Pluton is a weakly deformed to undeformed body of granite and quartz monzonite that appears to truncate earlier mylonites of the Cariboo. K-Ar cooling ages for biotite and muscovite crystallized at 92-84 Ma and 54-52 Ma, respectively (Wallace et al. 1986; CSQ24-13, U-Pb zircon reference 723530N, 865450M, Argus Home map-sheet, 83D/6). A new U-Pb zircon date is 551 Ma (L.C. Strak and R.R. Parrish, unpublished data, 1991).

10. The stratigraphy of the Mica Creek succession is discussed in detail by Simony et al. (1980), who referred to the package as part of the Horizontal Creek Group. The term Horizontal Creek Group is used here to refer to the Malton Gneiss, Monashee Mountains, British Columbia. In current research, Part 6, Geological Survey of Canada, Paper 88-1E, pp. 95-100. Doucet, P.J., Ghent, E.D., and Simony, P.S. 1986. Metamorphism in the Monashee Mountains of Blue Hill, British Columbia. In Current research, Part 6, Geological Survey of Canada, Paper 86-1A, pp. 69-71.

11. Structure sections 22 and 23 show a fan in the vergence direction of D₂ structures mapped by Peil (1984). The Slide Mountain synform forms the axis of the fan. It is proposed here that the change in D₂ vergence may be related to a westerly wedge (westward) to the west vergent structure, in the hanging wall of the Cariboo Creek thrust. The upper displacement of such a wedge would be a blind west-vergent thrust that is out by the Murtle pluton, before merging with the east-vergent Thunder River thrust on the west side of the pluton.

12. Neotectonic data: Campbell, R.B. 1968. Cariboo River (83D), British Columbia. Geological Survey of Canada, Map 15-187, 1:250,000 scale (1 sheet).

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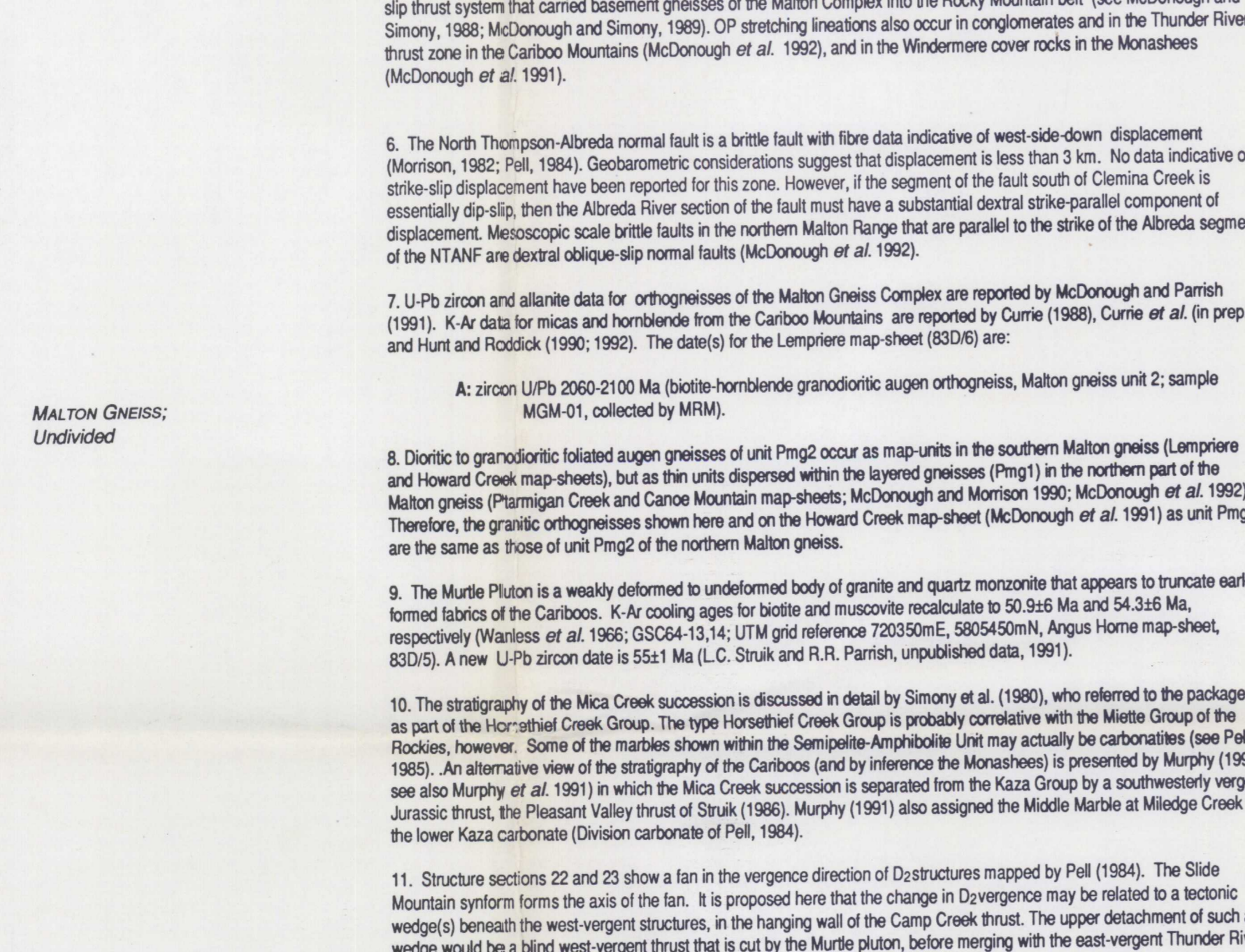
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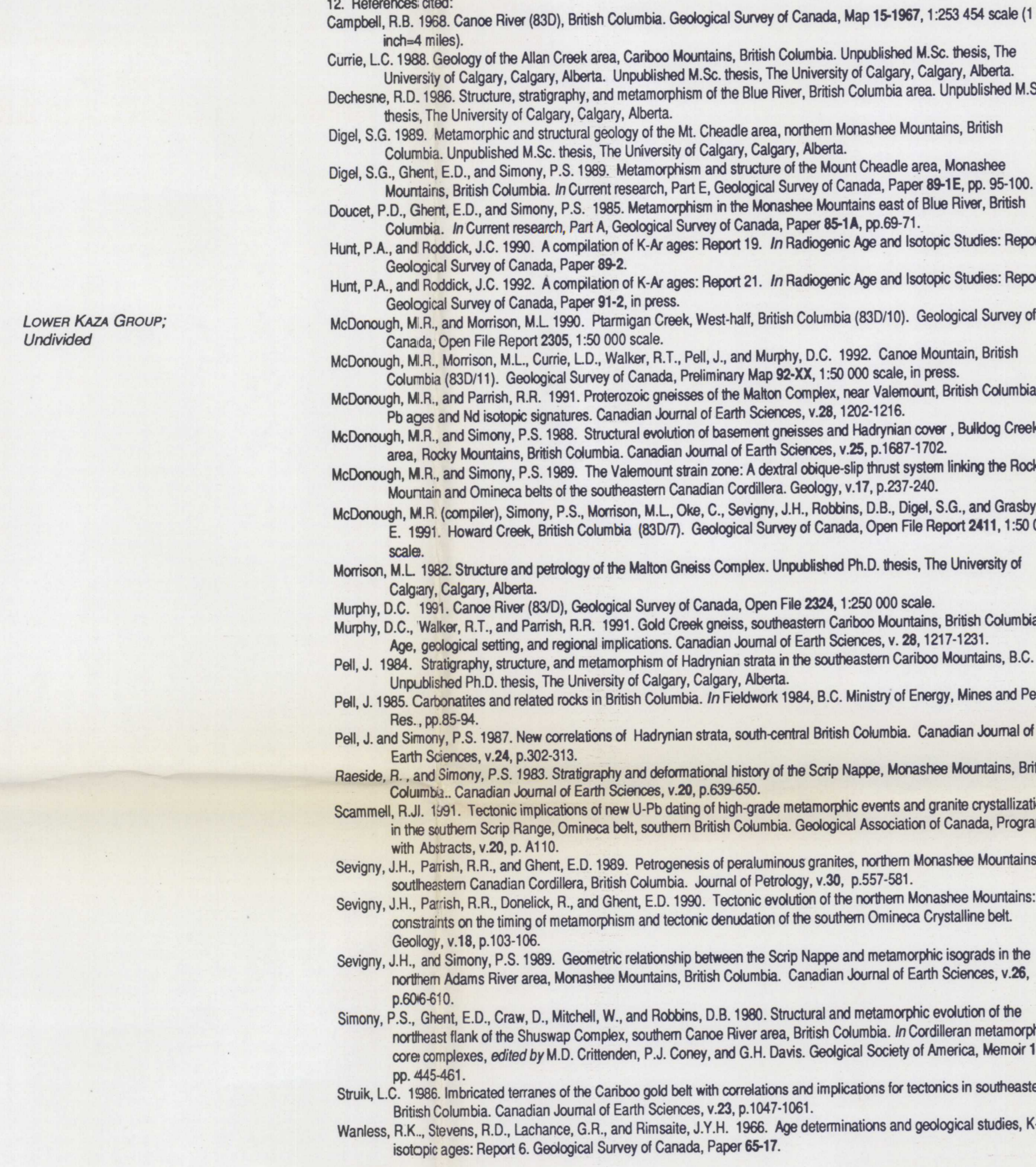
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22 Lempreire, B.C. 83D/6

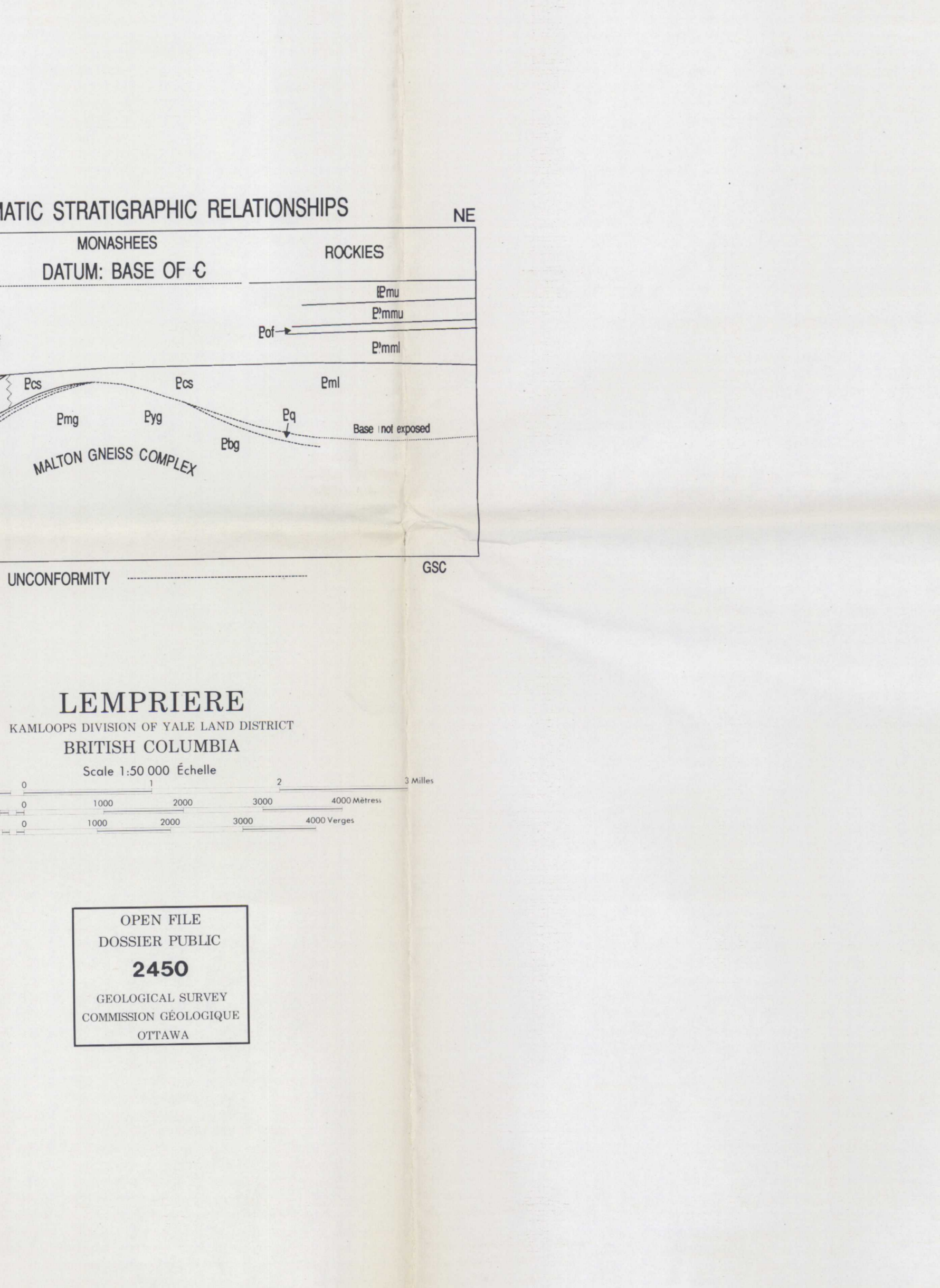


23 Lempreire, B.C. 83D/6



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