



**GEOLOGICAL SURVEY OF CANADA
COMMISSION GEOLOGIQUE DU CANADA**

Open File 2336

**JURASSIC LITHOLOGIES AND SOME FACTORS
AFFECTING SEDIMENT DISTRIBUTION IN
NORTHWESTERN ALBERTA**

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EMR Contract No. 23294-0-06-00/01-XSG

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Resources Canada

March, 1991

JURASSIC LITHOLOGIES AND SOME FACTORS AFFECTING SEDIMENT
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Abstract

The Lower Jurassic strata of northwestern Alberta consist of basal, highly radioactive marlstones (Pliensbachian?) overlain by green or brown (Toarcian to Aalenian) Poker Chip Shales. Both were deposited on a passive cratonic margin, although a brief erosional or non-depositional period may have intervened between the two.

A regional unconformity (probably erosional) separates these sediments from the overlying Late Jurassic or Upper Fernie shale/sandstone package. These marine clastics (mainly shale) were deposited in and adjacent to a subsiding foreland basin which developed in response to the onset and continued activity of the Columbian orogeny to the west.

A highly to moderately glauconitic transgressive unit at the base of the Upper Fernie package (Early Oxfordian to Late Callovian 'Green beds' equivalent) is present throughout the study area, and in the region coincident with the ancient Peace River Arch, is expressed by relatively thick quartz sands. To the west, thick marine Nikanassin sand and shale deposition extended into Late Jurassic and possibly Early Cretaceous times in response to rapid subsidence of the foreland basin.

A Lower Cretaceous (Hauterivian) erosional event subsequently bevelled increasingly older Jurassic sediments eastward.

Re-initiation of Cordilleran uplift in Early Cretaceous (Barremian) times produced a fresh sediment supply in the west and continued subsidence of the Alberta foredeep trough. A major trunk fluvial system developed in the trough (Spirit River Channel), and received sediment from the Cordillera to the west, and from a broad highland to the east. This highland, including the western Fox Creek Escarpment, was a relief feature which exposed mainly Jurassic sediments to erosion in the Early Cretaceous. Continued subsidence and further accumulation of Gething fluvial and delta plain deposits eventually buried these highlands in late Aptian time.

Introduction

The purpose of this paper is to reiterate and provide more detail to lithologic descriptions discussed in a previous paper co-authored by Poulton, Tittlemore and Dolby (1990). The intra- and interrelationships of Jurassic units will be summarized, and the effects of Lower Cretaceous erosion on the current subsurface distribution of Jurassic sediments will be discussed.

The area of study is shown in Figure 1, and shows the location of the stratigraphic cross sections utilized by Poulton et al (1990), and in this report. These sections have been slightly modified by the author subsequent to detailed sample examination.

The Jurassic section in northern Alberta is a complex, eastward tapering wedge of clastics and minor carbonates that represent the dynamic pre-orogenic events and onset of the Columbian Orogeny in the early Late Jurassic. The complexity arises from the numerous unconformities that occur within the section, and by its substantial structural modification in pre-Lower Cretaceous and Lower Cretaceous times. The dynamics are expressed by relatively quiescent periods that are interrupted by these widespread erosional events. Figure 2 is a table of formations reflecting current thoughts on Jurassic stratigraphy.

From the standpoint of an oil and gas explorationist, there are three major considerations that lend support to the examination a generally non-producing collection of deposits. Firstly, the development of a Nordegg carbonate bank in west central Alberta probably influenced the accumulation of major Cretaceous and Jurassic oil and gas deposits in this area. Secondly, the high organic content of Jurassic shales and marls were important source rock contributors for many Cretaceous and Triassic clastic plays found throughout Alberta and in British Columbia [Jackson (1984), Riediger (1990), Masters (1984)]. Thirdly, an understanding of Jurassic lithologies and their distributions has a bearing on the delineation of Lower Cretaceous clastics.

The figures in the Appendix 2 of this report contain lithologic descriptions which incorporate core and chip sample examinations. Cores were sampled by the author and dated palynologically by G. Dolby (Poulton et al, 1990). Although they are essential for understanding the stratigraphic relationships, they are less than representative of the total Jurassic section north of Township 55. Most cores include only thin Upper Fernie representatives below Cretaceous sands or above

Rock Creek sands, Poker Chip Shale below Rock Creek sands or above Nordegg carbonates, or Nordegg clastics above the Triassic formations of interest.

To get a complete lithologic picture, these cores were thus supplemented by chip sample descriptions. Because there are no cores in the wells of Section C-C' (Fig. 6), chip samples were examined in selected wells, although these descriptions are not included in this report.

It should be noted that in Appendix 2, the Triassic has quotation marks ("Triassic") in the lithologic descriptions. Since the Triassic is a period, and the column heading indicates formations or members, these quotations are meant simply to mark this exception.

Five cores were described and sampled south of Township 60 (secondary study area of Figure 1), but these wells are not described in detail here because several comprehensive studies have been published regarding the Jurassic in west central Alberta. [Losert (1986,1990), Marion (1984), Rosenthal (1988)]. Summaries of core data are included in Appendix 1.

This report basically divides the Jurassic into 5 units. The lowest four are members of the Fernie Formation. The basal, oldest member is the Lower Fernie shelf limestone/clastic unit (Plienbachian?), or Nordegg Member. As discussed in Poulton et al (1990), descriptions of the Nordegg Member in Alberta encompass more than one rock type and age. While at its type area at Nordegg, Alberta it is a black, cherty limestone (Early Sinemurian or older) (Spivak, 1949; Frebald, 1957), in west central Alberta it is a carbonate bank (undated - Sinemurian?)

(Poulton et al, 1990) and is a series of dolomites and sandstones (undated) on the eastern Jurassic subcrop edge of the same area (Bovell, 1979). In the study area of this report, the marlstones have been interpreted as being a younger stable shelf facies north of the carbonate bank of west central Alberta, and will be referred to as Lower Fernie shelf facies or limestone/clastic unit, or simply Nordegg Member.

This unit lies unconformably on Late Carboniferous carbonates in the south, and on Triassic carbonates and clastics in the study area. The Poker Chip Shale rests unconformably on the Nordegg, and while overlain by Rock Creek sands south of the study area, it is blanketed by the Upper Fernie shale/sandstone unit in northwestern Alberta. In some areas, it may be entirely missing due to erosion. The Nikanassin Formation, a sandstone package, is in part, syndepositional with Upper Fernie shales and sandstones, but has not been examined by the author, and will only be touched on briefly in this report.

Acknowledgments

This paper draws on information contained in a previous paper by Poulton et al (1990). The sections (Figs. 4 to 8) are taken directly from this paper, with slight modifications. Some of these wells, in turn, were taken from more detailed sections correlated by Jiri Losert (Alberta Department of Energy) and Shaun O'Connell (Alberta Geological Survey). All of the information regarding Lower Cretaceous geology was gleaned from studies presented by Paul Jackson; David Smith, Carl Zorn, and Robert Sneider; and Robert Gies in A.A.P.G. Memoir 38 on the Elsworth deep basin (see references). The geology of Figure 3 is copied, with slight modifications, from Paul Jackson's study of Lower Cretaceous paleogeography in the aforementioned memoir. However, the interpretations and data

provided in this study are entirely the responsibility of the author. The Jurassic portion of the correlation chart (Fig. 2) is after Poulton et al (1990), and the Cretaceous part is modified from Smith et al (1984) and Rosenthal (1988). Cores were examined at the core storage facility of the Energy Resources Conservation Board, and samples were provided by the sample repository at the Institute of Sedimentary and Petroleum Geology. Financial support for this study, which is a contribution to the Peace River Arch and Western Canada Atlas projects of the Geological Survey of Canada and the Alberta Geological Survey, was provided by S. Bell and R. McQueen (Institute of Sedimentary and Petroleum Geology). A special thanks to Terry Poulton (I.S.P.G.) for providing opportunities, guidance, and valuable information and insights.

Lithologic Descriptions of Units

NORDEGG MEMBER

Described in the Lexicon of Canadian Stratigraphy - Western Canada (1990) and by Spivak (1949) and Frebold (1957) as "a black, cherty limestone". Springer et al. (1964) add that "in the subsurface it extends eastward to a line from Pincher Creek through Calgary to the Gilby area southwest of Edmonton and consists of speckled dark brown shale of bituminous appearance, with thin beds of shaly limestone."

In the study area, the Nordegg just south of Township 54 is a calcareous multi-coloured chert varying to a cherty banded limestone. These beds are 45 to 60 m. thick, thinning northward and eastward, and are generally massively-bedded in the middle and lower portions. The upper two to three meters (noted in cores from wells 10-12-57-19W5M and 14-8-56-19W5M) are brecciated to varying degrees,

and abundant collapse structures, stylolitic sutures and shale intraclasts were observed. These features would indicate a period of emergence, and the contact with the overlying Poker Chip Shale is therefore probably erosional. As discussed in Poulton et al (1990), the Nordegg carbonate bank is undated, but regarded as Sinemurian in age. Extensive sampling of the Nordegg in the subsurface have produced no conclusive dates due to the poor quality of diagnostic fossils. They are either too thermally altered, too far-ranging or too few in number for practical dating purposes.

Between Townships 59 and 60, this chert/limestone sequence changes northward to interbedded dark marlstones and highly calcareous dark shales with abundant calcareous white specks and fish remains. Thin, argillaceous limestone stringers are not uncommon. Figs. 4 and 5 demonstrate this lithologic change from the carbonate bank facies to northward thinning shelf shales and marls. As discussed in Poulton et al (1990), these sediments are perhaps younger (Pliensbachian) than the Sinemurian(?) carbonate bank of central Alberta.

The gamma ray log signature of the Nordegg north of Township 60 shows a high radioactive count, and can be similarly seen in two well-known marker beds in Alberta; the Base of Fish Scales, with its high concentration of fish remains, and the Second White Specks, which contains abundant calcareous white specks.

Average thickness of this carbonate/clastic facies is 30 to 40 meters, and is maintained northward to Twp. 85, then thins to the erosional zero-edge at Township 91.

East to west relationships seem to be less dramatic north of Township 60, and do not show the lithologic transitions apparent in the northwest/southeast direction on strike with the Alberta Trough. The thickness increases from 10 to 15 meters at the Alberta/B.C. border, to 30 meters just east of the Sixth Meridian, then tapers down to the zero edge eastward beyond that.

POKER CHIP SHALE MEMBER

Described in the Lexicon of Canadian Stratigraphy (1990) as typically black, platy, fissile shale. In north central Alberta, it can also be dark green, dark brown, and greyish-green, and is variously silty and/or pyritic. Average thickness is 10 to 25 meters (thinning southward), and thin, clean sandstone stringers are not uncommon. In well 14-8-56-19W5M, the Poker Chip Shale contains a thin limestone bed. The Poker Chip Shale unconformably overlies the Nordegg in the subsurface of north central Alberta.

North of Township 60, these shales are generally green or brown, and are variably bentonitic, silty, arenaceous and pyritic, and often contain siderite stringers. Thicknesses range from one or two meters to over 60 meters in well 15-34-77-1W6M (Fig. 5). This unusually thick Poker Chip accumulation also yielded the first documented Aalenian age found in Western Canada, whereas Toarcian dates are given for Poker Chip Shales in the rest of the province. The nature of the contact with the Lower Fernie shelf facies north of Twp. 60 is less clear than that further south.

ROCK CREEK MEMBER

Described in the Lexicon (1990) as quartzose siltstones and sandstones with calcite cement. They can be heavily bioturbated. Marion (1984) reports that in west central Alberta, they are basal coquinoid sandstones overlain by burrowed, silty sandstone with clean flaser and cross-bedded, and planar-laminated to wave rippled sandstones at the top.

The Rock Creek sand is not found north of Township 55, and its northern erosional edge in north central Alberta can be seen in Figure 4. The core in well 10-14-54-13W5M was dated by C. Singh for Losert (1986), and confirms the presence of Niton 'B' and Rock Creek as far north as Twp. 54. By log correlation alone, wells 14-8-56-19W5M (Fig. 5) and 7-5-58-18W5M (Fig. 4) could easily be assigned to the Rock Creek Member. However, cores were sampled and described by the author, and on the basis of resultant dates and lithologies, these sands have been assigned to the Poker Chip Shale and Green Beds, respectively (see Appendix 1 for sample depths and dates). This also points out the importance of core examination for the delineation of sometimes subtle but often significant lateral variations in Jurassic sands.

Lackie (1958) described Rock Creek sands in northwestern Alberta and northeastern British Columbia. However, re-examination of these sands has led to the conclusion that they can easily be placed in the Upper Fernie shale/sandstone package (Poulton et al., 1990). Poulton (1984) noted that no Middle Jurassic fossils have been documented north of 54°N latitude (Twp. 58). However, Aalenian sediments have been found in three wells sampled in this study (wells 14-8-56-19W5M, 4-32-69-19W5M and 15-34-77-1W5M). Based on lithologies and correlations, these sediments (predominantly shales) have been designated as Poker Chip Shale Members.

UPPER FERNIE SHALE/SANDSTONE UNIT

The upper parts of the Fernie Formation are described in the Lexicon (1990) as "glaucopitic sands, concretionary bands and brown weathering siltstones and sandstones". They include the Green Beds ("dark to bright green glauconitic sandstone or siltstone with calcareous concretions weathering yellowish-brown"). Upper Fernie shales in northwestern Alberta are not described in the Lexicon, but in the south they are included in the poorly defined Ribbon Creek Member as dark grey, silty or clayey shales. The Passage Beds in the south are described as dark grey shales with yellow-brown weathering concretions.

In the south end of the study area, the cored sands in the Upper Fernie package were dated as Late Callovian to Early Oxfordian, and are therefore probably Green Beds equivalents (see well 6-27-60-19W5M; Appendix 1 and 2). They are represented by thin (up to 5 meters thick), highly to moderately glauconitic and often argillaceous sandstones. Small-scale, low angle ripple cross bedding was observed in these sands. Shale interbeds are dark grey, greenish-grey and black, and are also often highly glauconitic. Siderite stringers are common, as are arenaceous and/or argillaceous siltstone stringers. These sands are equivalent to the "Niton B" sands of north central Alberta, which are prolific hydrocarbon producers (Losert, 1990).

Northward, from Twps. 60 to 74, basal Green Beds equivalents persist, but are much thinner and less discernable on logs. The top of this argillaceous sand unit is marked by a deflection on the gamma ray log which reflects a highly glauconitic black shale and siderite interval with minor sideritic quartz sand, and probably represents the maximum transgressive episode in Green Beds

deposition (Poulton et al, 1990). Interbedded shales are grey, and green or brown with a greyish cast, and are variably glauconitic, arenaceous and silty.

North of Twp. 74, very well-developed, thick (10 to 50 meters) quartz sands appear at the same interval. Lackie (1958) named them Rock Creek, but the glauconitic and sideritic nature of these sands, which often have thin bands of black, glauconitic shales, argues for their affinity to the Green Beds equivalents further south.

Deposition of interbedded sand, shale and siltstone continued into at least Late Kimmeridgian times (as in well 6-35-64-21W5M). The shales grade from grey to green, and are silty and arenaceous. The grey, quartzose sands are occasionally glauconitic and cemented with clay. Siltstones are most often very arenaceous and/or argillaceous. The thickness of the Upper Fernie package ranges from 10 to 75 meters.

Summary of Jurassic Unit Relationships

In the study area, the Lower Fernie shelf facies is probably Pliensbachian in age (Poulton et al, 1990). Being perhaps younger than the Sinemurian(?) carbonate bank in central Alberta, it may have been deposited on the stable shelf while an erosional event occurred over the carbonate bank. While the contact of the carbonate bank with the overlying Poker Chip Shale is therefore erosional in central Alberta, the contact between the Pliensbachian(?) shelf marls and the Toarcian to Aalenian Poker Chip Shales of northwestern Alberta is less clear. It may well be conformable, but for purposes of consistency with outcrop data to the west (Poulton et al, 1990), and because there is no evidence to the

contrary, the contact is shown in this report as an unconformity.

A probable unconformity exists between the Toarcian/Aalenian Poker Chip Shales and the youngest (Late Callovian or Early Oxfordian) beds of the Upper Fernie shale/sandstone unit. Whether this is a period of non-deposition or erosion is debatable, but Poulton (1984) and Davies and Poulton (1986) suggested that uplift in northwestern Alberta and northeastern British Columbia may have taken place prior to Upper Jurassic sedimentation. Erosion is further indicated by substantial thickness variations and local absence (well 10-9-74-20W5M) of the underlying Poker Chip Shale. Whether erosional or non-depositional, the contact on geophysical logs is expressed subtly, and without sufficient core control, the uncertainty of the exact contact position is indicated on the sections and lithologic descriptions by a dashed unconformity line.

The varying and relative thickness of the basal Green Beds sand north of Twp. 74 may well be related to the movement of fault blocks over the Peace River Arch. It is possible that local Arch related tectonism occurred with the onset of the Columbian orogeny to the west and the associated development of the Alberta Trough foredeep east of it. The collapse or downwarping phase of the Arch took place in the Late Paleozoic. Cant (1988) found evidence of further collapse in Lower Cretaceous sediments, demonstrated principally by positive thickness anomalies. The possibility cannot be ruled out that similar developments may have occurred in the early Late Jurassic while orogenic activity was initiated westward, and while the Green Beds (interpreted to be a transgressive basal sand unit- Poulton, 1984) were being deposited in the newly developing foredeep.

Uninterrupted deposition of shales, siltstones, and minor sandstones continued from the Late Callovian into at least Late Kimmeridgian times in the east part of the study area, and in the west part, may have carried on into the Early Cretaceous as thick Nikanassin sands and shales accumulated in the rapidly subsiding foreland basin.

Lower Cretaceous Erosional Influence on Jurassic Sediment Distribution

The present day distribution of Jurassic strata in the subsurface was influenced considerably by Lower Cretaceous erosional events. The basic information regarding the geologic setting of the Lower Cretaceous in this report is distilled from comprehensive studies by Jackson; Smith, Zorn and Sneider; and Gies (1984 - see references).

The sub-Cretaceous unconformity represents an erosional event developed during Hauterivian (Early Cretaceous) time. It was the result of a brief cessation of Cordilleran orogenic activity in the west which reduced sediment supply and minimized tectonic movements in the Alberta basin (Smith et al, 1984). Consequently, pre-Cretaceous strata were truncated to produce a subcrop of Jurassic sediments in the west, Devonian carbonates roughly east of the Fifth Meridian (the 'Axial High' ancient highland marked 'Y' in Figure 3 was mainly composed of these sediments), and a thin subcrop of Mississippian shales and carbonates in between the two.

A strong re-initiation of Cordilleran uplift in Barremian time resulted in the deposition of eastward prograding Cadomin alluvial fan conglomerates. These, in turn, contributed coarse clastic sediment to a developing northwest/southeast

trending fluvial braid plain (Spirit River Channel) along the west edge of the basin. The eastern edge of these re-worked Cadomin fluvial sediments is marked by the Fox Creek Escarpment. This was the western edge of a broad positive feature which divided, and in part, supplied sediment to the developing Spirit River Channel system in western Alberta and the Edmonton Channel system of central Alberta. The Spirit River Channel system became a broad valley incised into Jurassic sands and shales (Jackson, 1984).

Figure 3 is a lithofacies map showing the major Lower Mannville fluvial systems and highland areas. Most of the highlands (depicted on the map as white or hatched areas) were exposed Upper Jurassic and uppermost Poker Chip Shale sediments in southern, central and western Alberta, and were underlain by lowermost Poker Chip Shales and the Nordegg Member. The hatched areas were extreme relief features, probably consisting of relatively thick Upper and Middle Jurassic, and thin Poker Chip Shales. These would seem to have been the last areas that were inundated during maximum flooding of the southeasterly transgressing Boreal Sea in Late Gething times, and are therefore overlain by relatively thin accumulations of Lower Mannville sediments. Locally, these high areas are unconformably overlain by post-Aptian deposits, thus indicating a period of exposure extending through Lower Mannville and early Middle Mannville deposition.

The Alberta Trough or Basin east of the Cordilleran uplift continued to subside during Aptian time, resulting in further accumulation of fluvial, and eventually, of delta plain sediments of the Upper Gething formation in northern Alberta. Rivers flowed eastward from the Cordilleran uplands into the Spirit River braid plain, then were re-distributed northward along this trunk river system to the Boreal Sea. Smith et al (1984) note that fluvial Gething sandstones of the Spirit River Channel are mainly restricted to the Lower Gething, and Jackson

(1984) adds that they were deposited eastward to the position of the Fox Creek Escarpment over the Cadomin Formation. In Late Gething time, the Fox Creek Escarpment and highlands were buried due to a significant change in drainage patterns which expanded river drainage flow eastward over the ancient highland. The highland thus became a low-lying, swampy drainage basin, with river flow meandering randomly across this plain from west to east (Smith et al, 1984). Jackson (1984) submits that this area subsequently became a marine basin with the transgression of the Boreal Sea. However, he states that the outline of marine areas is quite speculative because "marine upper Gething or Ostracod equivalents cannot be distinguished from younger marine Bluesky or Wabiskaw Formations". As mentioned previously, some areas of the 'extreme relief' highlands (X, Y and Z - Fig. 3) were still subaerially exposed.

The study area of this report demonstrates some of these paleogeographic features; the Fox Creek Escarpment and highlands, the Spirit River Channel, the Peace River Channel, and the eastern approach to the western Cordilleran uplands (see Figure 3 for locations of these features with superimposed stratigraphic cross section locations).

Section A - A' (Figure 4) cuts diagonally (southwest to northeast) across the highlands of Barremian time, from the western Fox Creek Escarpment edge to the northern edge of the highlands and into the Peace River Channel.

It is interesting to note that the 'Pembina High' (indicated by an 'X' on Figure 3) roughly coincides with the position of the Sinemurian(?) carbonate bank of Poulton et al (1990). It reflects an area of thin to non-deposition of Lower Cretaceous strata, but also represents a relatively thick accumulation of Middle

and Upper Jurassic sands discussed by Losert (1986, 1990), Marion (1984) and Rosenthal (1988) among others (see also well 10-14-54-13W5M: Figure 4).

Hypothetically, the carbonate bank may have acted as a structural repository for shoaling Jurassic sands which were relatively resistant to erosion (compared to 'Ferne Shales') during the Early Cretaceous. Subsequent burial by Gething/Ellerslie continental deposits completed a scenario for hydrocarbon accumulation including: structural relief, proximity to source beds, stratigraphic/structural trapping mechanisms, and unconformities which may have acted as conduits for the migration of oil and/or gas.

Moving northward from the Pembina High on the same section (Figure 4), the apparent erosional truncation of the Poker Chip Shale between wells 4-32-69-19W5M and 10-9-74-20W5M argues for an erosional event pre-dating Upper Fernie deposition. Section D - D' (Figure 8) shows a similar relationship. Figure 4 also indicates that the northern edge of the ancient highlands sloped gently into the Peace River Channel.

Section B - B' (Fig. 5), on the other hand, demonstrates relief of up to 25 meters (80 ft.) as you come off the Fox Creek Escarpment (well 11-12-69-22W5M) into the Spirit River Channel (well 7-33-73-1W6M), and then back onto the escarpment (well 15-34-77-1W6M). As suggested by Smith et al. (1984), relatively thin (less than 60 m.) Upper Gething sediments overlie a relatively thickly preserved Jurassic sequence over the highlands area. This study shows that this Jurassic sequence consists of Upper Fernie shales and sandstones, and in some cases, uppermost Poker Chip Shales. Vertically thick Gething deposits, on the other hand, infill areas where the Jurassic is thinly preserved (lowermost Poker Chip Shale and Lower Fernie shelf facies only - see Figs. 4 and 5).

Section C - C' (Fig. 6) shows Cadomin alluvial fan and braid deposits overlying, in the south end, a wedge of basal Nikanassin marine sands and shales probably derived from the craton (Stott, 1984). They are Late Jurassic sediments infilling the rapidly subsiding Alberta foreland basin which were subsequently truncated by pre-Cadomin (Hauterivian) erosion. This bevelling, as indicated in Section C - C', removed successively older Jurassic units northward, and in the northern highland area (Figure 3 - white), Lower Fernie marlstones were probably exposed to erosion. It is interesting to note that the relatively thick sands of the basal Upper Fernie unit (Green Beds) in this section coincide with the location of the Peace River Arch (Cant, 1988) between Twps. 75 and 85 near the B.C. border.

The east-west cross sections (Figs. 7 and 8) both demonstrate the thick Upper Jurassic section in the west resulting from rapid subsidence and infilling of the foreland basin. The resulting wedge tapers eastward due to depositional thinning, and later, to truncation of successively older units by sub-Cretaceous (Hauterivian) erosion.

Section A - A' (Fig. 7) shows the expression of the ancient northern highlands (well 10-1-82-6W6M), where Mannville strata are relatively thin. Mannville deposits increase vertically as you move into the Peace River Channel to the east and the Spirit River Channel to the west. Once again, thick quartzose sands at the base of the Upper Fernie are coincident with the location of the Peace River Arch.

Section D - D' (Figure 8) illustrates the abutment of Cadomin sands against the Fox Creek Escarpment (well 10-9-74-20W5M), the thick accumulation of Upper Jurassic sediments in the foreland basin, and eastward tapering due to depositional thinning and Hauterivian truncation.

Conclusions

This report is an attempt to characterize Jurassic sediments which, for the most part, have been considered non-descript 'Ferne shales' situated between more economically interesting Cretaceous and Triassic or Paleozoic deposits.

As can be seen, the Jurassic sequences in northwestern Alberta are unique in their own right, and not simple northward extensions of their 'counterparts' further south.

That they represent a fundamental tectonic change from a passive epicratonic margin to an orogenic belt and foredeep make them worth study on an academic level. From an economic standpoint, Jurassic reservoirs in Alberta, Saskatchewan and Manitoba "are hosts for about 4 per cent of western Canada's total initial reserves of oil, and 1.4 per cent of its marketable natural gas" (Poulton, 1989). In addition, they served as watersheds which supplied sediments to major Lower Mannville fluvial deposits. As mentioned earlier, they also made significant hydrocarbon contributions to Paleozoic and Lower Cretaceous oil and gas reservoirs.

This report, then, contains previously unavailable data which contributes to a better understanding of Jurassic strata. Some of the ideas put forth here are speculative, but highlight the need for further study.

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FIGURES:

- Fig. 1 - Location of study area and index map of subsurface sections.
- Fig. 2 - Correlation chart
- Fig. 3 - Lower Mannville fluvial systems
- Fig. 4 - North-south stratigraphic section A-A'
- Fig. 5 - North-south stratigraphic section B-B'
- Fig. 6 - North-south stratigraphic section C-C'
- Fig. 7 - East-west stratigraphic section A-A''
- Fig. 8 - East-west stratigraphic section D-D'

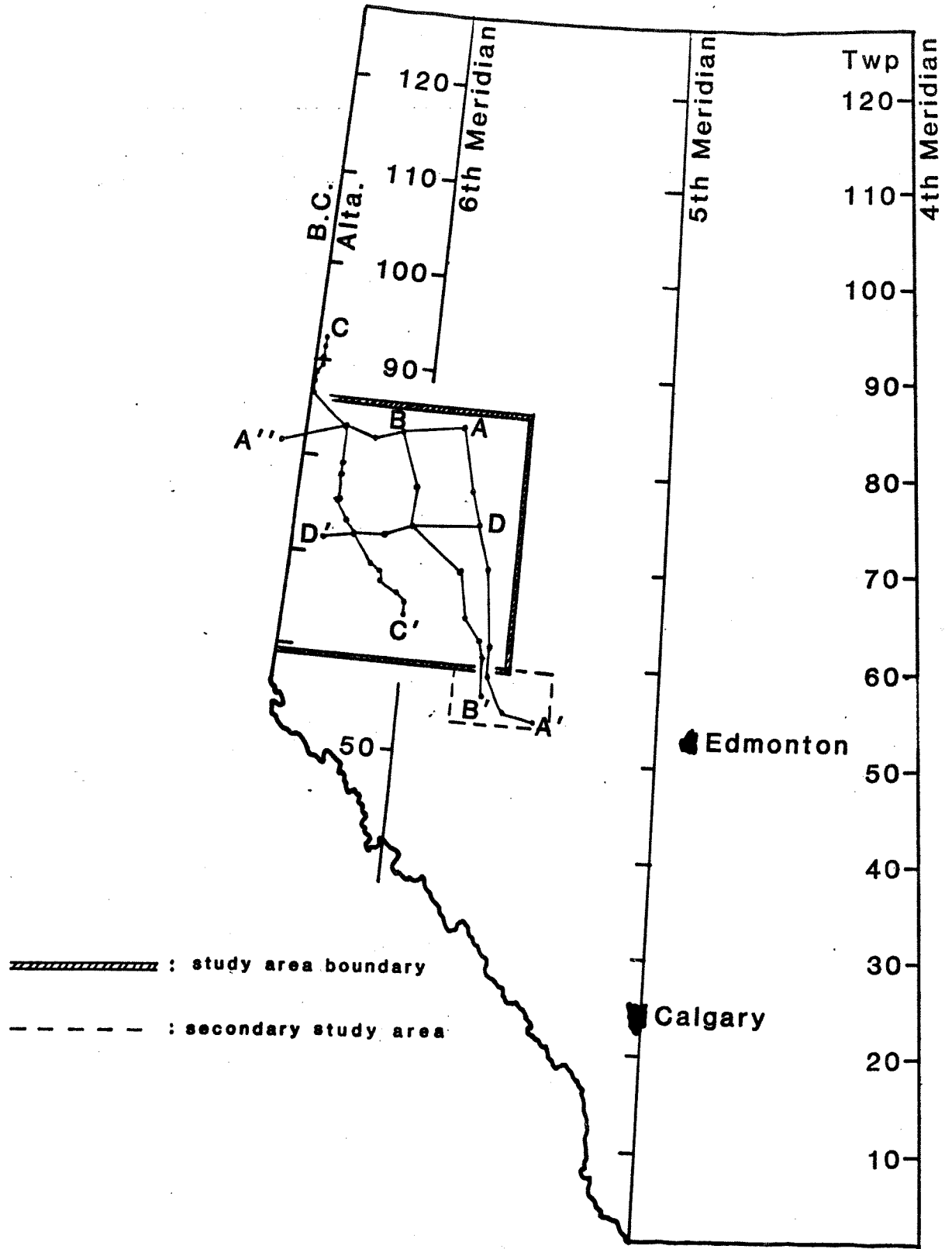


Figure 1 : study area and cross section index map

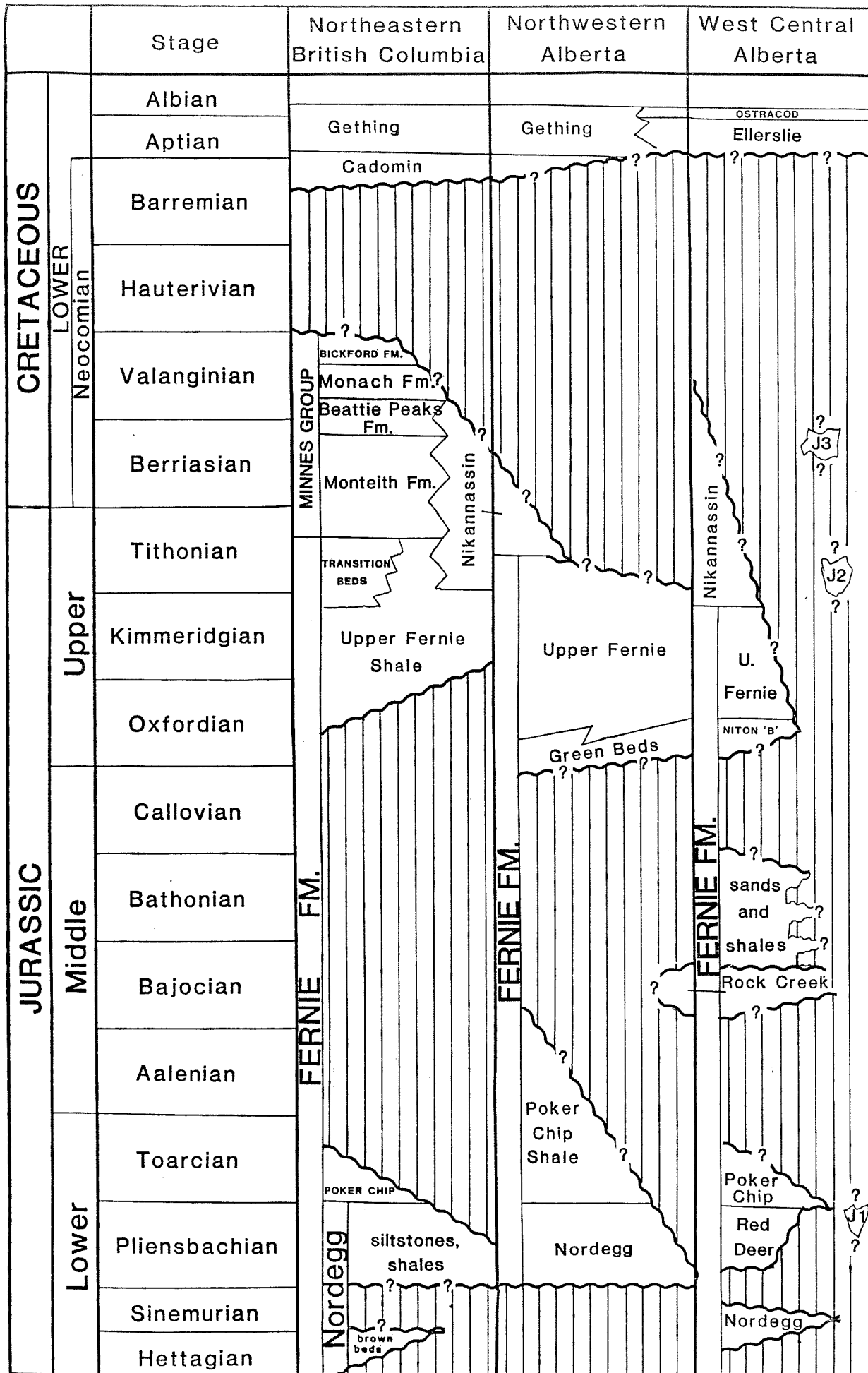


Figure 2 : Correlation chart

FIGURE #3: Lower Mannville Fluvial Systems

- 1 - Spirit River Channel
- 2 - Peace River Channel
- 3 - Edmonton/Bellshill Lake Channel
- F - Fox Creek Escarpment/highlands
- V - Cadomin alluvial fans
- X - Pembina High
- Y - Axial High
- Z - Southern highlands

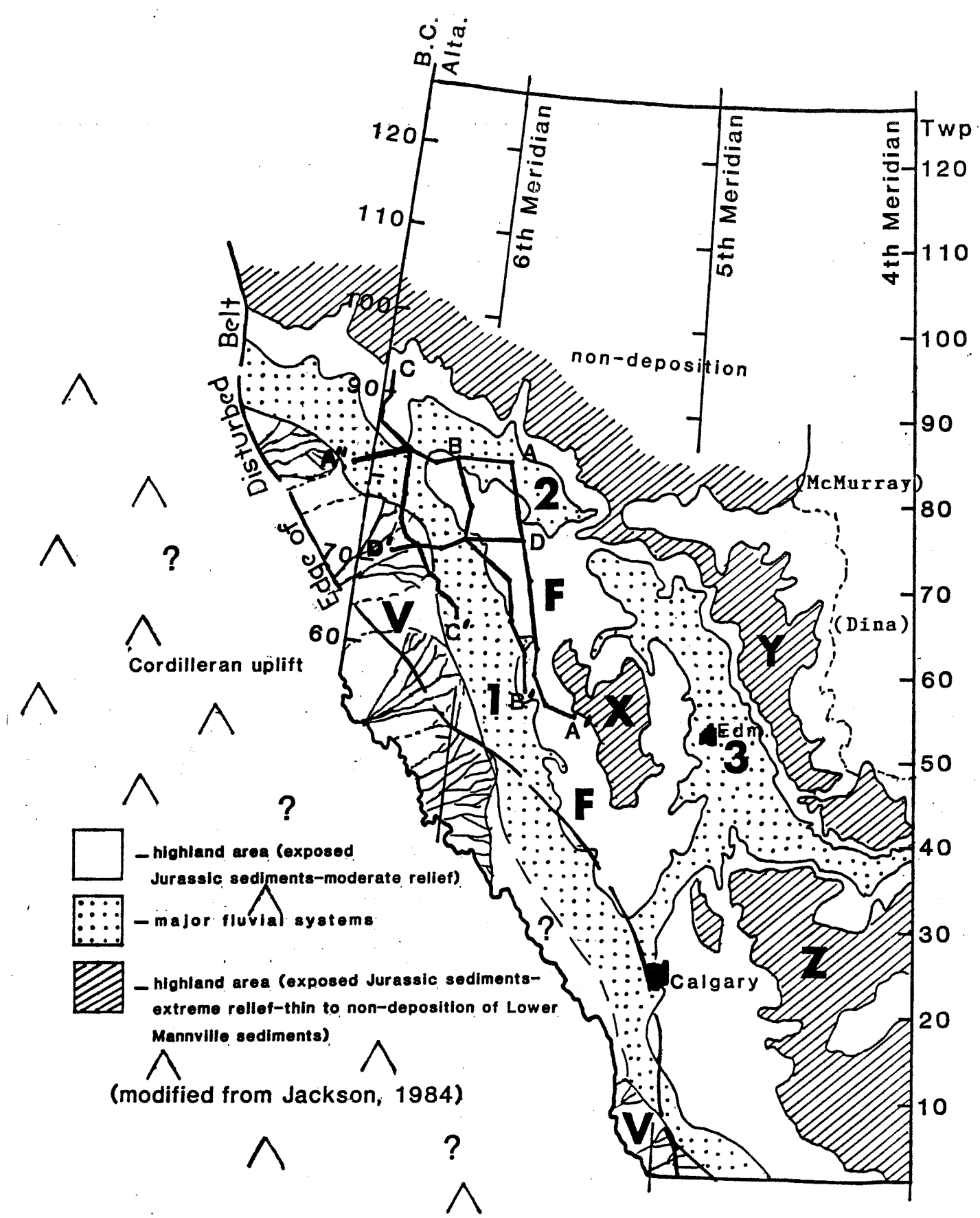


Figure 3 : Lower Mannville Fluvial Systems

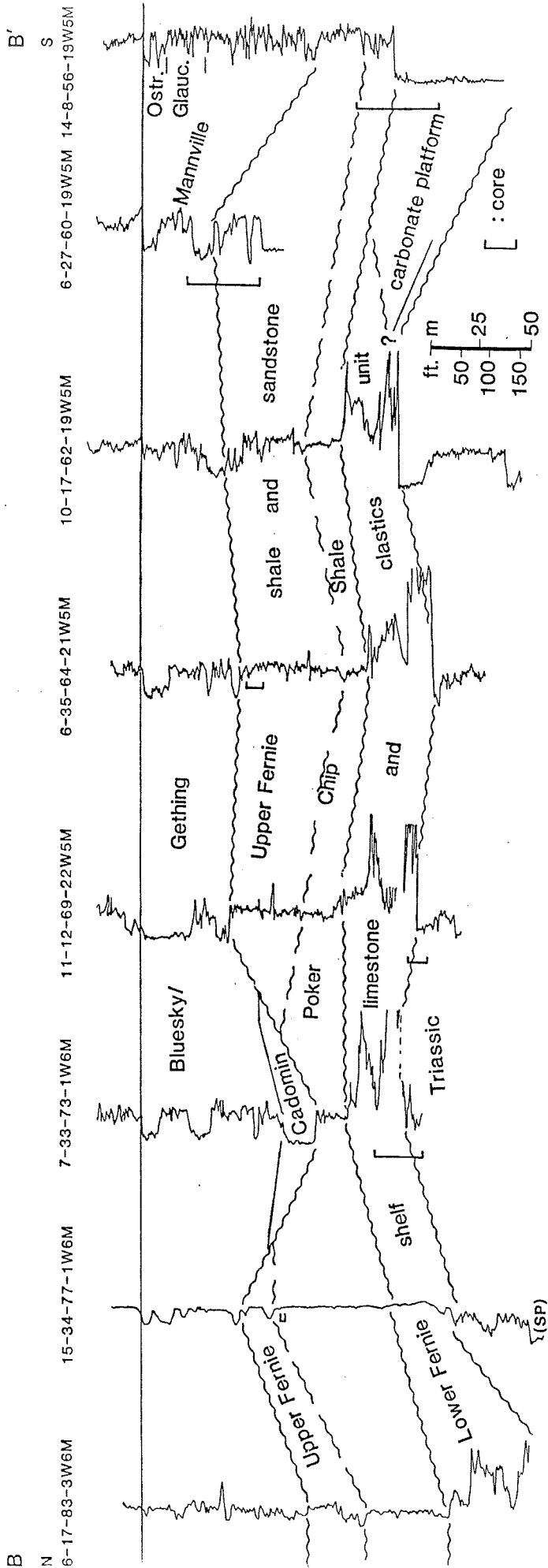


Figure 5 : Stratigraphic section B-B'

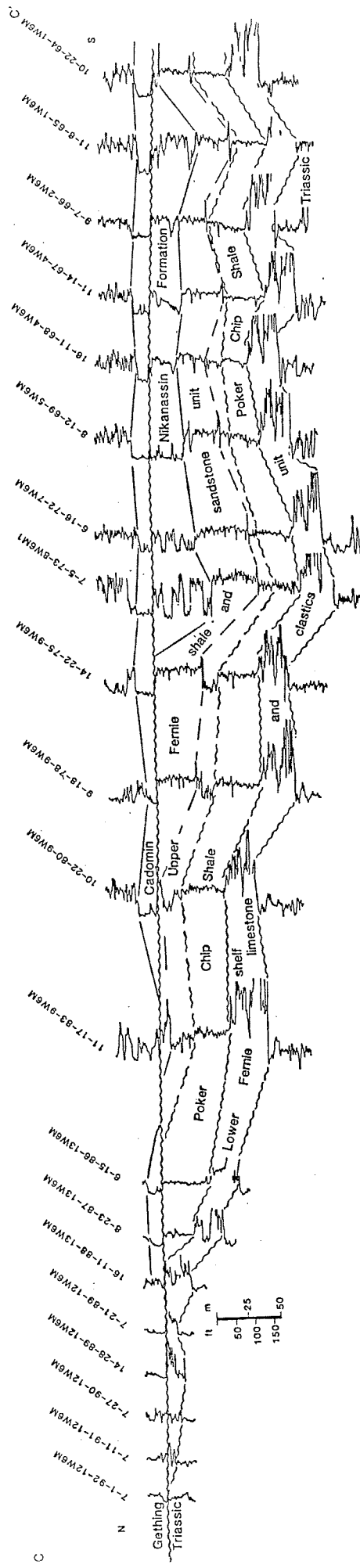


Figure 6 : Stratigraphic section C-C'

A" A

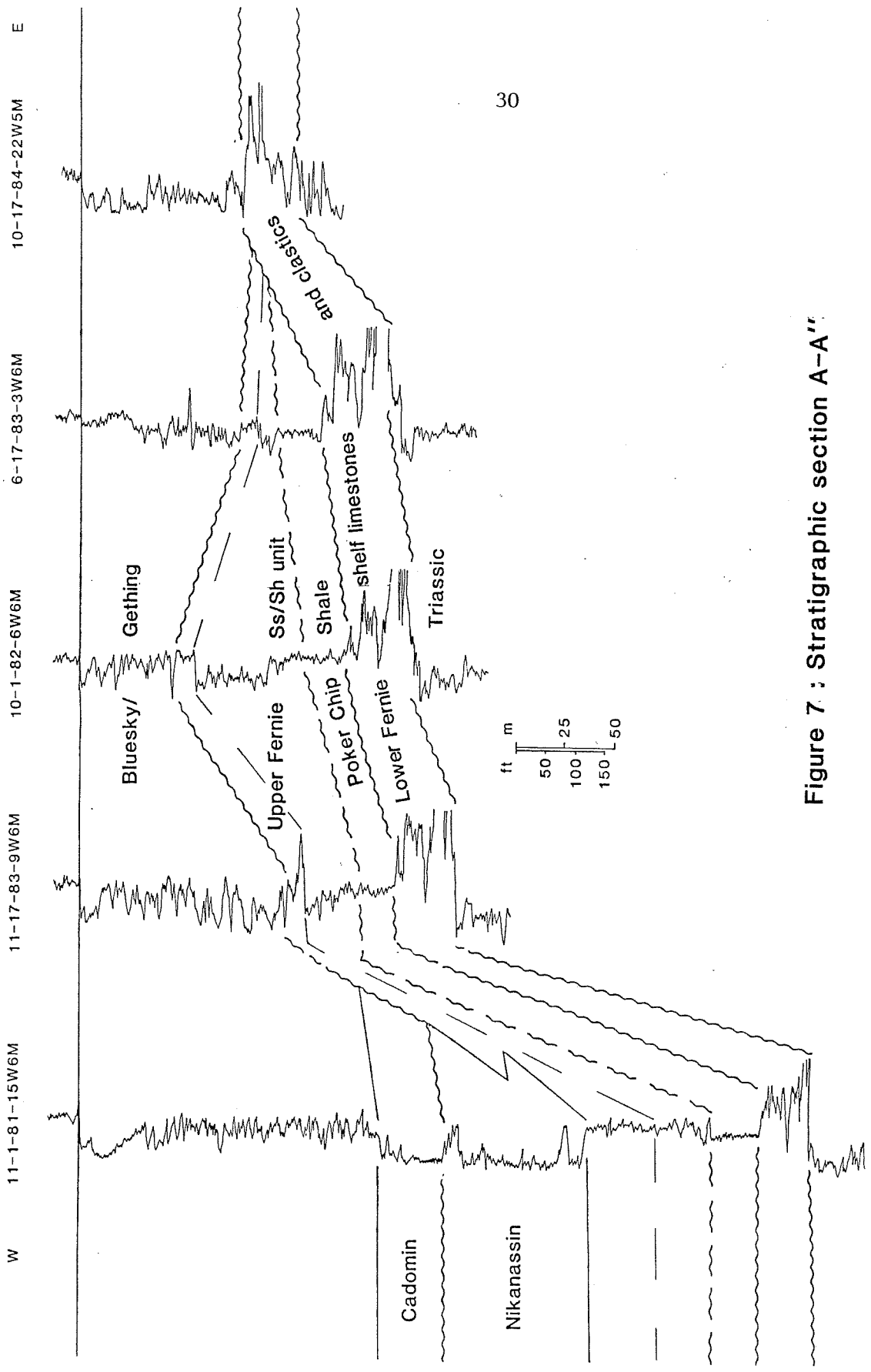


Figure 7 : Stratigraphic section A-A"

D

E

10-9-74-20W5M

7-33-73-1W6M

6-9-72-4W6M

6-16-72-7W6M

6-25-71-10W6M

W

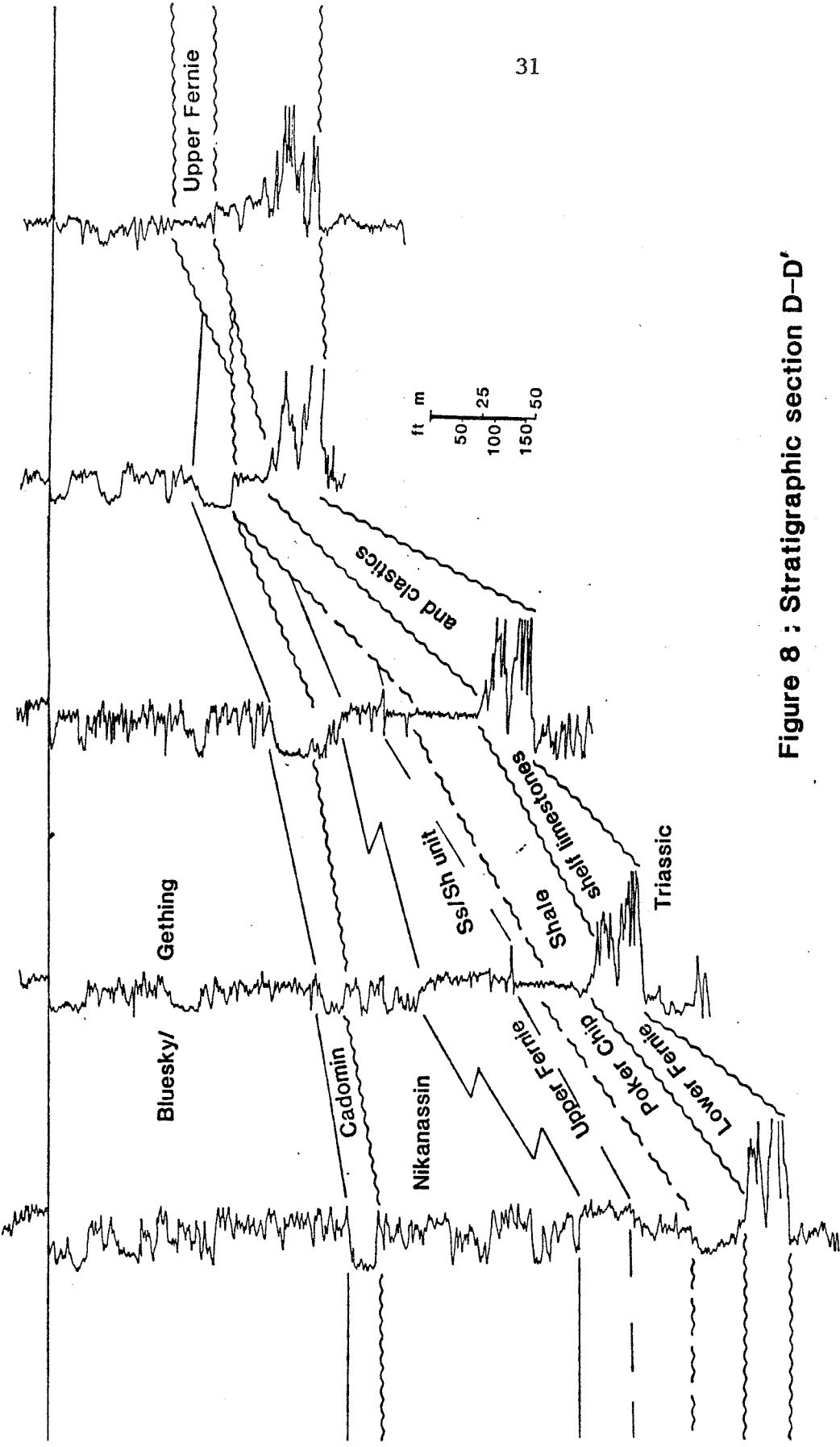


Figure 8 : Stratigraphic section D-D'

APPENDIX 1: CORES AND PALYNOLOGICAL SAMPLES EXAMINED

APPENDIX 1: CORES AND PALYNOLOGICAL SAMPLES EXAMINED

Location: 14-6-56-13 W5M

Interval: 1906.5 - 1920.1 m.

cut: 13.6 ; rec. 13.6 m.

Formation: Green Beds

sample 1: GSC loc. no. C-171160

depth: 1907.5 m.

age: Middle Oxfordian

2: GSC loc. no. C-171161

depth: 1910.9 m.

age: Middle Oxfordian

3: GSC loc no. C-171162

depth: 1916.3 m.

age: Middle Oxfordian

Location: 14-8-56-19 W5M

Interval: Core #1: 2829.2 - 2847.0 m.

cut: 17.8; rec. 17.8 m.

#2: 2847.6 - 2865.6 m.

cut: 18.0; rec. 4.0 m.

#3: 2865.8 - 2866.6 m.

cut: 0.8; rec. 0.7 m.

#4: 2866.6 - 2872.4 m.

cut: 5.8; rec. 3.8 m.

Formation: Poker Chip Shale to Nordegg

Sample 1: GSC loc. no. C-171166

depth: 2829.2 m.

age: Jurassic undifferentiated

2: GSC loc. no. C-171167

depth: 2830.8 m.

age: Jurassic undifferentiated

3: GSC loc. no. C-171168

depth: 2831.8 m.

age: Jurassic undifferentiated

4: GSC loc. no. C-171169

depth: 2834.8 m.

age: Aalenian to Toarcian

Sample 5: GSC loc. no C-171170

depth: 2836.3 m.

age: Aalenian to Toarcian

6: GSC loc. no. C-171171

depth: 2839.1 m.

age: Aalenian to Toarcian

7: GSC loc. no. C-171172

depth: 2841.3 m.

age: Aalenian to Toarcian

8: GSC loc. no. C-171173

depth: 2843.8 m.

age: Aalenian to Toarcian

9: GSC loc. no. C-171174

depth: 2846.8 m.

age: Aalenian to Toarcian

10: GSC loc. no. C-171175

depth: 2847.8 m.

age: Indeterminable

Location: 12-26-57-15 W5M

Interval: 2351.5 - 2362.7 m.

cut: 11.6; rec. 11.2 m.

Formation: Ellerslie (Cretaceous)

Sample 1: GSC loc. no. C-171164

depth: 2354.35 m.

age: Cretaceous

2: GSC loc. no. C-171165

depth: 2359.25 m.

age: Cretaceous

Location: 10-12-57-19 W5M

Interval: Core #3: 2780.0 - 2787.0 m.
cut: 7.0; rec. 7.0 m.
#4: 2787.0 - 2788.2 m.
cut: 1.2; rec. 1.2 m.

Formation: Poker Chip (Red Deer?) to
Nordegg

Sample 1: GSC loc. no. C-171177
depth: 2780.0 m.
age: Early Toarcian to Late
Pliensbachian

2: GSC loc. no. C-171178
depth: 2781.0 m.
age: Early Toarcian to Late
Pliensbachian

Location: 7-5-58-18 W5M

Interval: Core #1: 2399.0 - 2417.0 m.
cut: 18.0; rec. 17.4 m.
#2: 2417.0 - 2435.0 m.
cut: 18.0; rec. 6.0 m.

Formation: Cretaceous to Green Beds

Sample 1: GSC loc. no C-171181
depth: 2406.0 m.
age: Early Oxfordian to Latest
Callovian

2: GSC loc. no. C-171184
depth: 2414.5 m.
age: as above

3: GSC loc. no. C-171185
depth: 2416.0 m.
age: as above

4: GSC loc. no. C-171186
depth: 2417.0 m.
age: as above

5: GSC loc. no. C-171187
depth: 2419.0 m.
age: as above

Sample 6: GSC loc. no. C-171188
depth: 2421.0 m.
age: Early Oxfordian to Latest
Callovian

7: GSC loc. no. C-171189
depth: 2421.4 m.
age: as above

Location: 6-27-60-19 W5M

Interval: Core #2: 2204.0 - 2222.25 m.
cut: 18.25; rec. 18.25
#3: 2222.5 - 2240.5 m.
cut: 18.0; rec. 18.0 m.

Formation: Gething to Upper Fernie/Green
Beds

Sample 1: GSC loc. no. C-171190
depth: 2204.9 m.
age: Indeterminable

2: GSC loc. no. C-171192
depth: 2217.35 m.
age: Jurassic undifferentiated

3: GSC loc. no. C-171193
depth: 2219.0 m.
age: as above

4: GSC loc. no. C-171195
depth: 2224.25 m.
age: as above

5: GSC loc. no. C-171196
depth: 2225.25 m.
age: as above

6: GSC loc. no. C-171197
depth: 2226.5 m.
age: as above

7: GSC loc. no. C-171198
depth: 2229.5 m.
age: Early Oxfordian to Latest
Callovian

Sample 8: GSC loc. no. C-171199
 depth: 2232.0 m.
 age: Early Oxfordian to Latest
 Oxfordian

9: GSC loc. no. C-171200
 depth: 2232.5 m.
 age: as above

10: GSC loc. no. C-187101
 depth: 2235.75 m.
 age: as above

11: GSC loc. no. C-187102
 depth: 2237.0 m.
 age: as above

12: GSC loc. no. C-187103
 depth: 2238.0 m.
 age: as above

13: GSC loc. no. C-187104
 depth: 2240.0 m.
 age: as above

Location: 10-32-61-18 W5M

Interval: 6515 - 6540 ft.
 cut: 25'; rec. 25'

Formation: Gething to Green Beds

Sample 1: GSC loc. no. C-187105
 depth: 6527.5'
 age: Early Oxfordian to Latest
 Callovian

2: GSC loc. no. C-187106
 depth: 6532'
 age: as above

3: GSC loc. no. C-187107
 depth: 6535'
 age: as above

4: GSC loc. no. C-187108
 depth: 6536'
 age: as above

Sample 5: GSC loc. no. C-187109
 depth: 6537'
 age: as above

Location: 6-35-64-21 W5M

Interval: Core #2: 5914 - 5941 ft.
 cut: 27'; rec. 27'
 #3: 6200 - 6261 ft.
 cut: 61'; rec. 61'

Formation: Upper Fernie to Upper Fernie
 Nordegg to Triassic

Sample 1: GSC loc. no. C-187110
 depth: 5914'
 age: Late Kimmeridgian

2: GSC loc. no. C-187111
 depth: 5922'
 age: as above

3: GSC loc. no. C-187112
 depth: 5928'
 age: as above

4: GSC loc. no. C-187113
 depth: 5934'
 age: as above

5: GSC loc. no. C-187114
 depth: 5940'
 age: as above

6: GSC loc. no. C-187116
 depth: 6202'
 age: Indeterminable

7: GSC loc. no. C-187119
 depth: 6220'
 age: as above

8: GSC loc. no. C-187121
 depth: 6224'
 age: as above

Location: 4-32-69-19 W5M

Interval: Core #1: 4509 - 4567 ft.
 cut: 58'; rec. 55'
 #2: 4734 - 4769 ft.
 cut: 35'; rec. 35'

Formation: Gething to Poker Chip Shale
 Nordegg to Triassic

Sample 1: GSC loc. no. C-187122
 depth: 4512'

age: Cretaceous

2: GSC loc. no. C-187123
 depth: 4531'

age: as above

3: GSC loc. no. C-187124
 depth: 4545'

age: Aalenian to Toarcian

4: GSC loc. no. C-187125
 depth: 4552'

age: as above

5: GSC loc. no. C-187126
 depth: 4559'

age: as above

6: GSC loc. no. C-187127
 depth: 4564'

age: as above

7: GSC loc. no. C-187128
 depth: 4734'

age: Indeterminable

8: GSC loc. no. 187129
 depth: 4736'

age: as above

9: GSC loc. no. C-187130
 depth: 4738'

age: as above

Sample 10: GSC loc. no. C-187131
 depth: 4740'

age: Indeterminable

Location: 11-12-69-22 W5M

Interval: 1498.0 - 1508.0 m.
 cut: 10; rec. 9.5 m.

Formation: Nordegg to Triassic

Sample 1: GSC loc. no. C-187132
 depth: 1498.0 m.

age: Indeterminable

2: GSC loc. no. C-187133
 depth: 1499.5m

age: as above

3: GSC loc. no. C-187135
 depth: 1502.75 m.

age: as above

Location: 7-33-73-1 W6M

Interval: 4693 - 4753 ft.
 cut: 60'; rec. 60'

Formation: Nordegg to Triassic

Sample 1: GSC loc. no. C-187149
 depth: 4693'

age: Indeterminable

2: GSC loc. no. C-187150
 depth: 4698'

age: as above

Location: 10-9-74-20 W5M

Interval: 3510 - 3533 ft.
 cut: 23'; rec. 23'

Formation: Nordegg

Sample 1: GSC loc. no. C-187136
 depth: 3510'

age: Indeterminable

Sample 2: GSC loc. no. C-187137
 depth: 3519'
 age: Indeterminable

3: GSC loc. no. C-187138
 depth: 3523'
 age: as above

4: GSC loc. no. C-187139
 depth: 3528'
 age: as above

5: GSC loc. no. C-187140
 depth: 3532'
 age: as above

Location: 12-33-77-21 W5M

Interval: 2964 - 3042 ft.
 cut: 78'; rec. 51'

Formation: Nordegg to Triassic

Sample 1: GSC loc. no. C-187141
 depth: 2965'
 age: Indeterminable

2: GSC loc. no. C-187143
 depth: 2975'
 age: as above

3: GSC loc. no. C-187146
 depth: 3008'
 age: as above

Location: 15-34-77-1 W6M

Interval: 3370 - 3377 ft.
 cut: 7'; rec. 7'

Formation: Poker Chip Shale

Sample 1: GSC loc. no. C-187151
 depth: 3370'
 age: Aalenian

Sample 2: GSC loc. no. C-187152
 depth: 3375'
 age: as above

3: GSC loc. no. C-187153
 depth: 3376'
 age: as above

Location: 10-17-84-22 W5M

Interval: Core #2: 2206 - 2256 ft.
 cut: 50'; rec. 50'

#3: 2305 - 2362 ft.
 cut: 57'; rec. 48'

Formation: Gething

Gething to Nordegg

Sample 1: GSC loc. no. C-187154
 depth: 2207.5'
 age: Cretaceous

2: GSC loc. no. C-187155
 depth: 2218'
 age: as above

3: GSC loc. no. C-187156
 depth: 2226'
 age: as above

4: GSC loc. no. C-187158
 depth: 2241'
 age: as above

5: GSC loc. no. C-187159
 depth: 2253'
 age: as above




























6: GSC loc. no. C-187160
 depth: 2256'
 age: as above

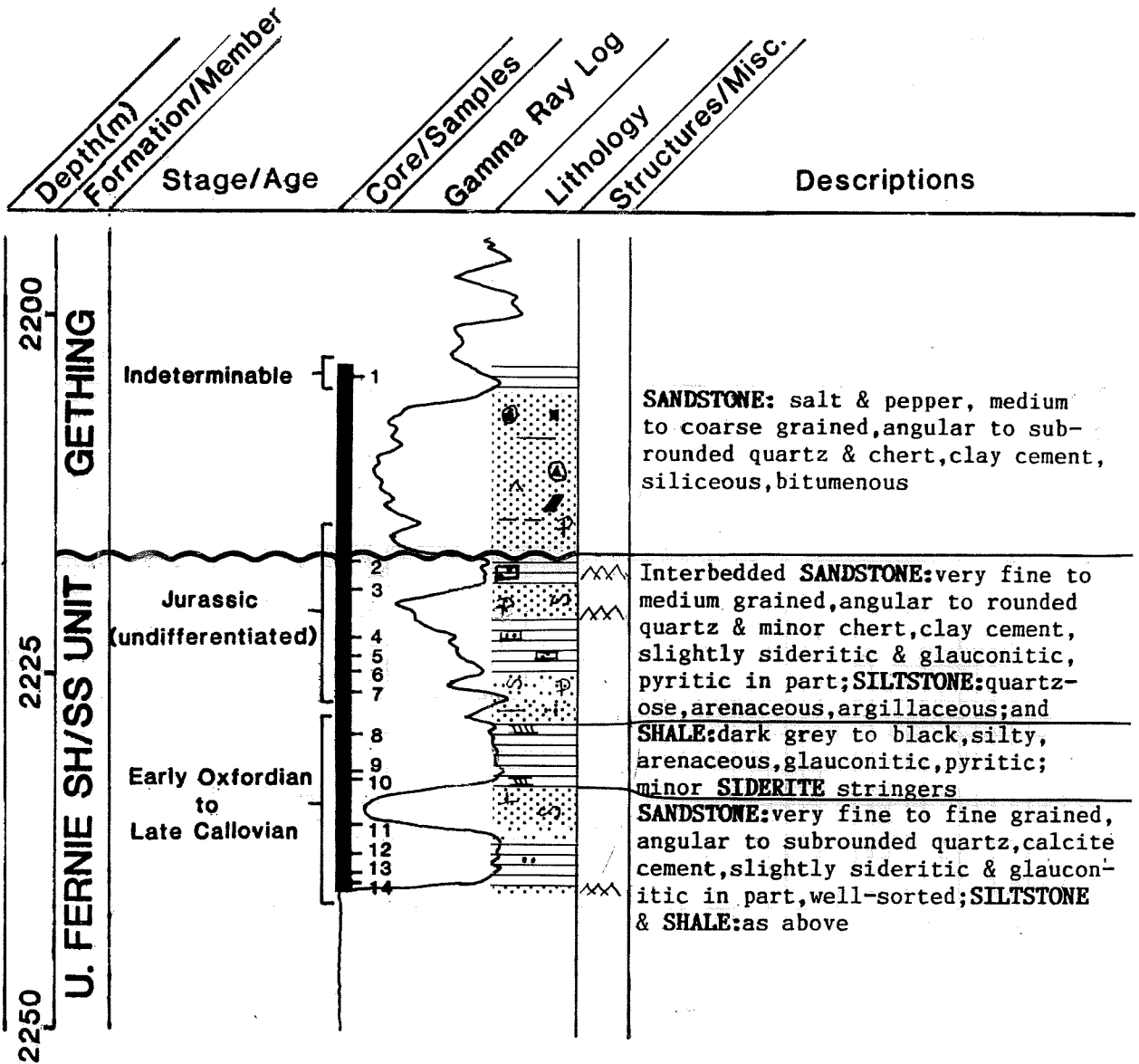
7: GSC loc. no. C-187161
 depth: 2314'
 age: as above

- Sample 8: GSC loc. no. C-187162
depth: 2319'
age: Cretaceous
- 9: GSC loc. no. C-187163
depth: 2322'
age: as above
- 10: GSC loc. no. C-187164
depth: 2326'
age: as above
- 11: GSC loc. no. C-187165
depth: 2333'
age: as above
- 12: GSC loc. no. C-187166
depth: 2340.5'
age: as above
- 13: GSC loc. no. C-187167
depth: 2343'
age: Indeterminable
- 14: GSC loc. no. C-187168
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age: as above
- 15: GSC loc. no. C-187169
depth: 2352'
age: as above

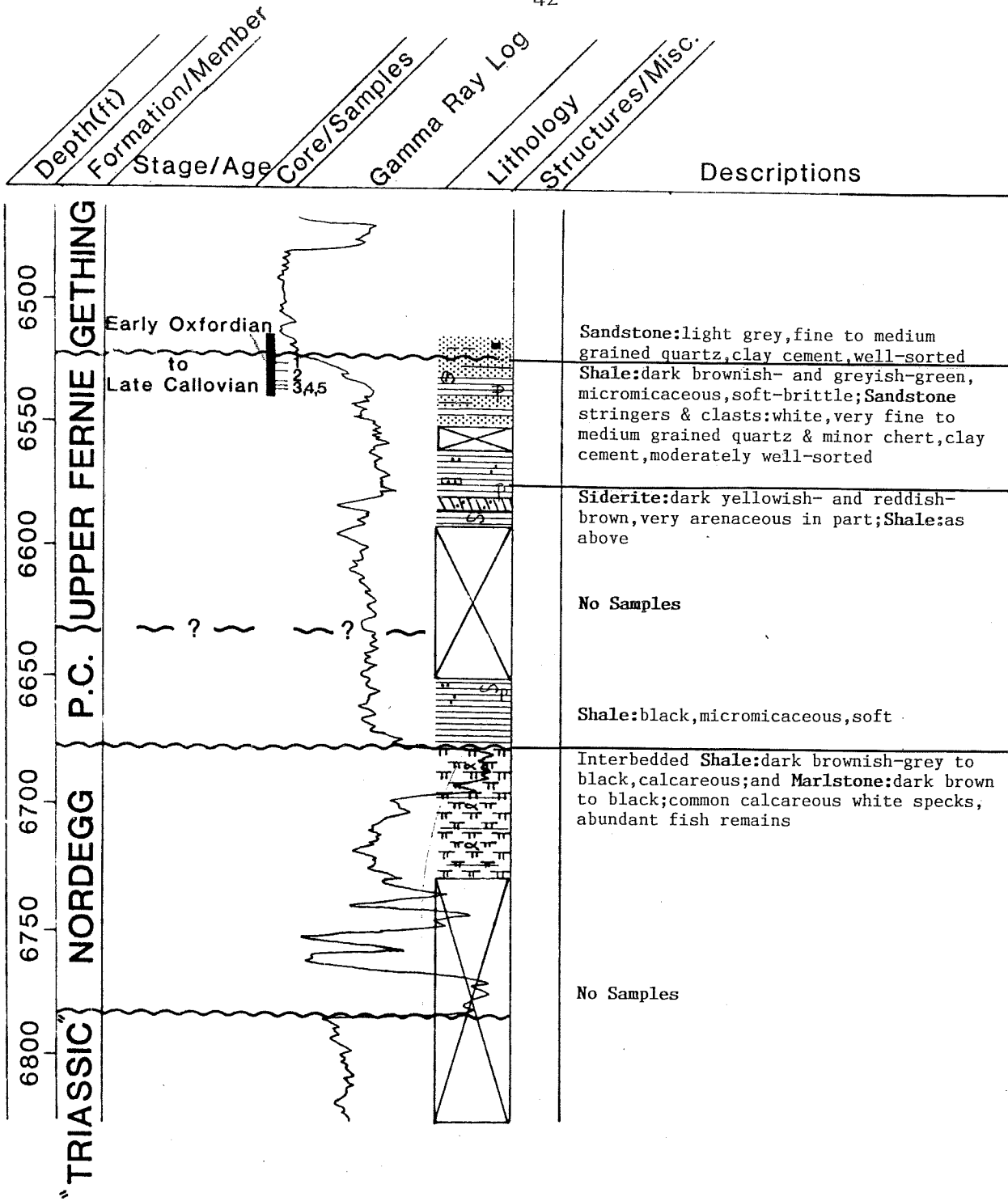
APPENDIX #2: DETAILED LITHOLOGIC DESCRIPTIONS

LEGEND

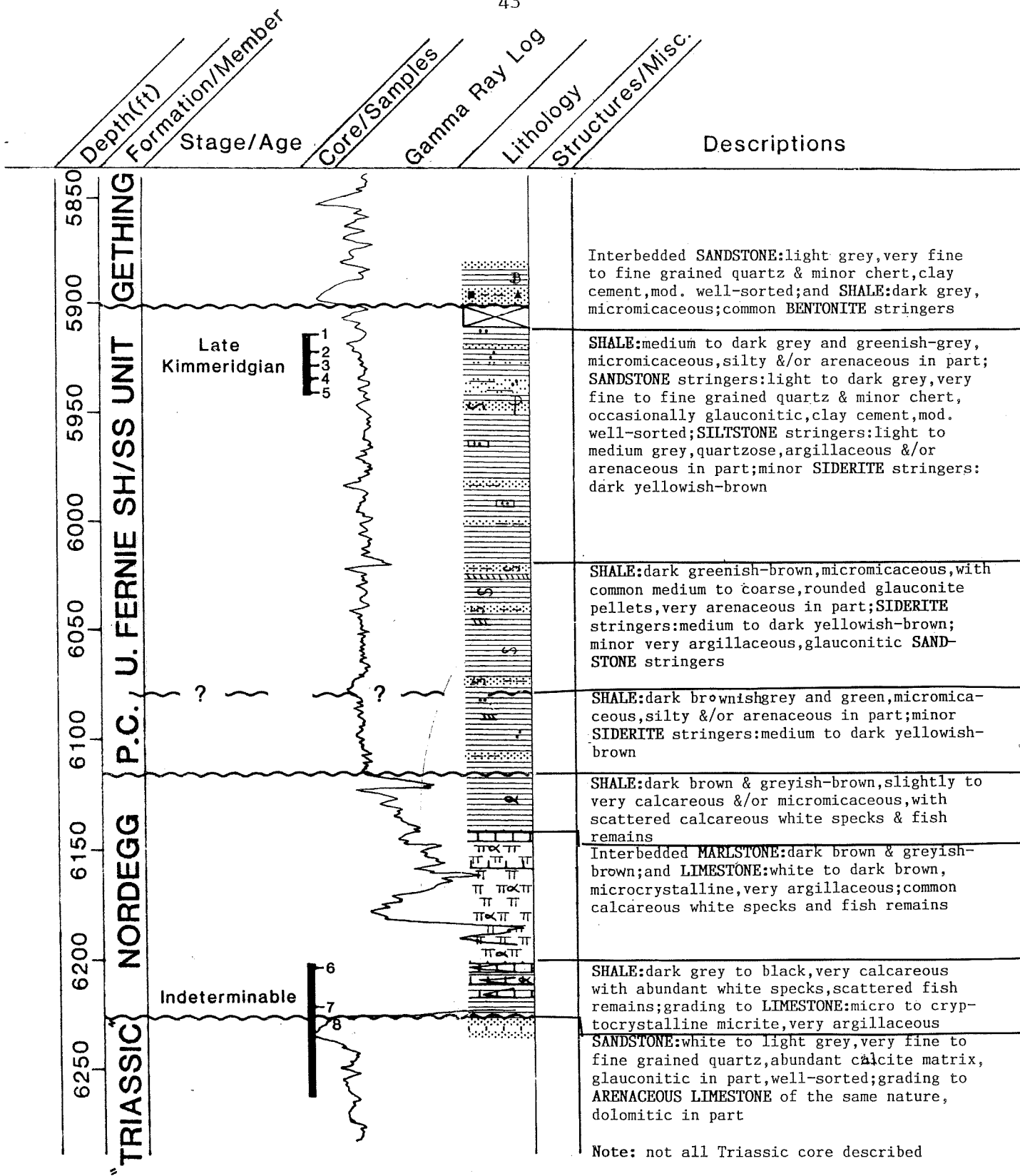
	- Sandstone		- Ripple cross-bedding
	- Shale		- Bioturbated
	- Siltstone		- Deformed bedding planes
	- Limestone		- Clast
	- Marlstone		- Bentonite
	- Bedded Chert (light, dark)		- Ammonite
	- Clay (stone)		- No sample
	- Siderite (bedded, itic)		- Plant remains
	- Argillaceous		- Coal/carbonaceous
	- Silty/stringers		- Glaucconitic
	- Arenaceous/stringers		- Siliceous
	- Chert grains/pebbles		- Bitumenous
	- Pyritic		- Fish remains
	- Calcareous		



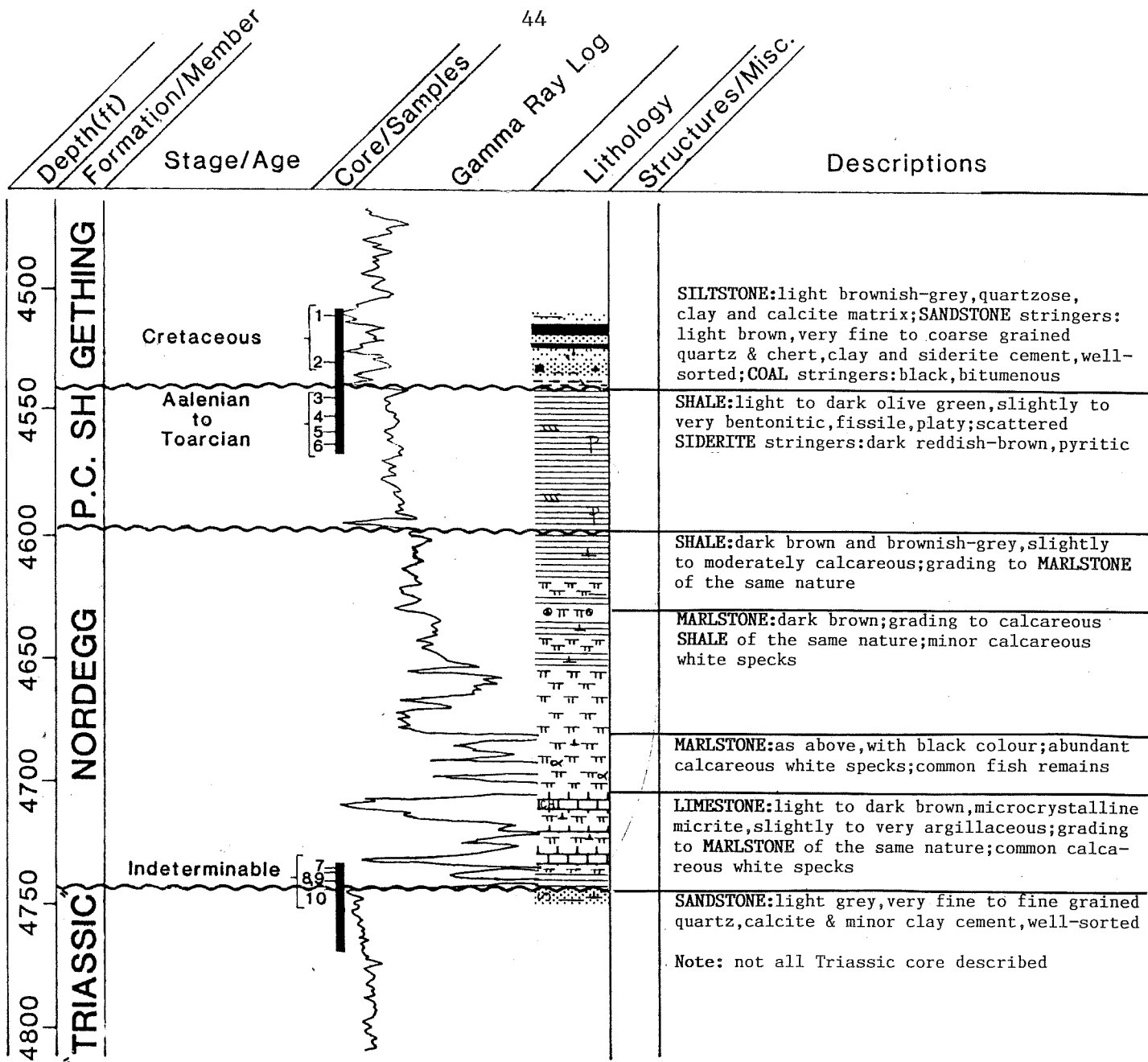
6 - 27 - 60 - 19W5M



10 - 32 - 61 - 18W5M



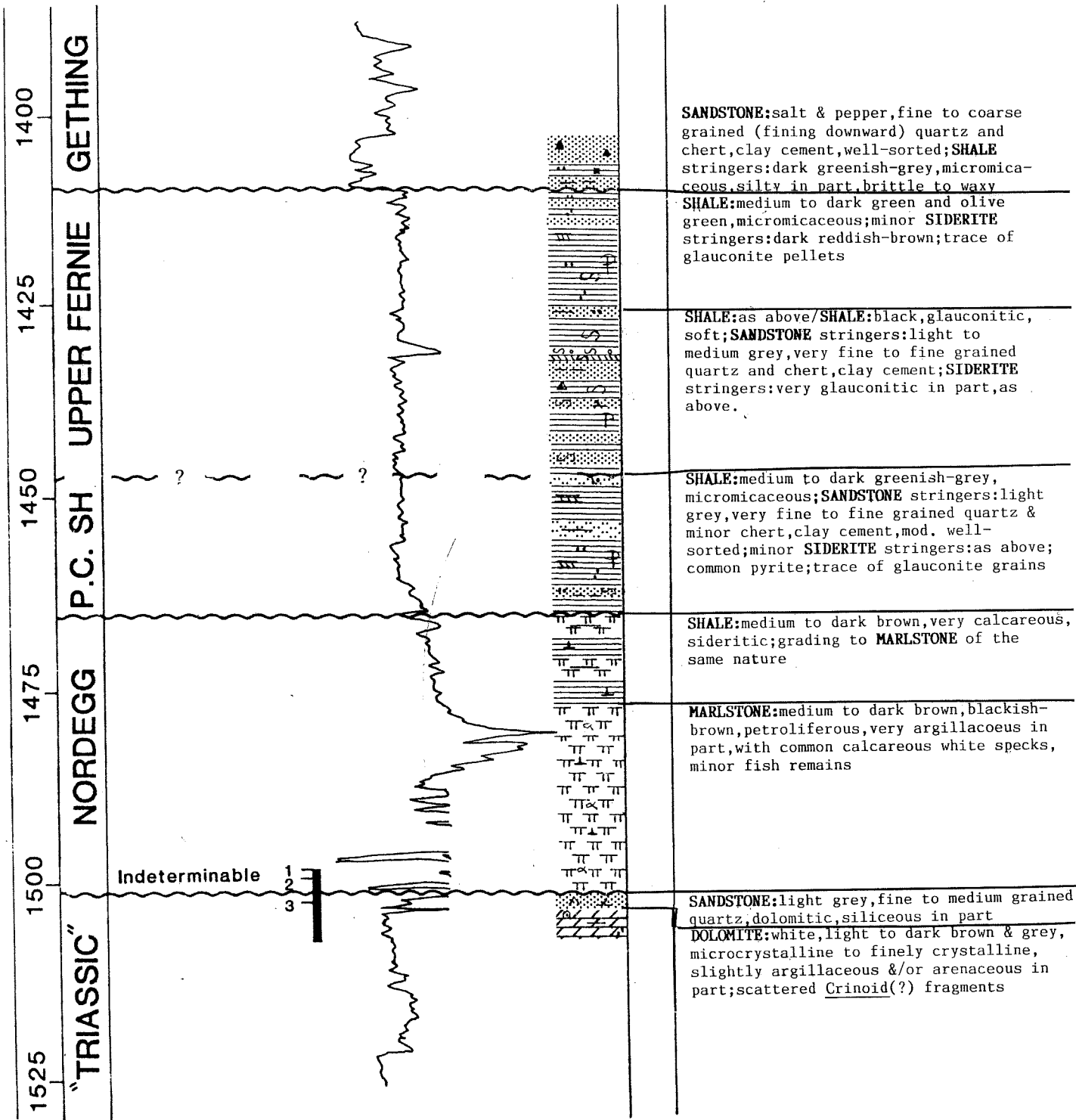
Note: not all Triassic core described

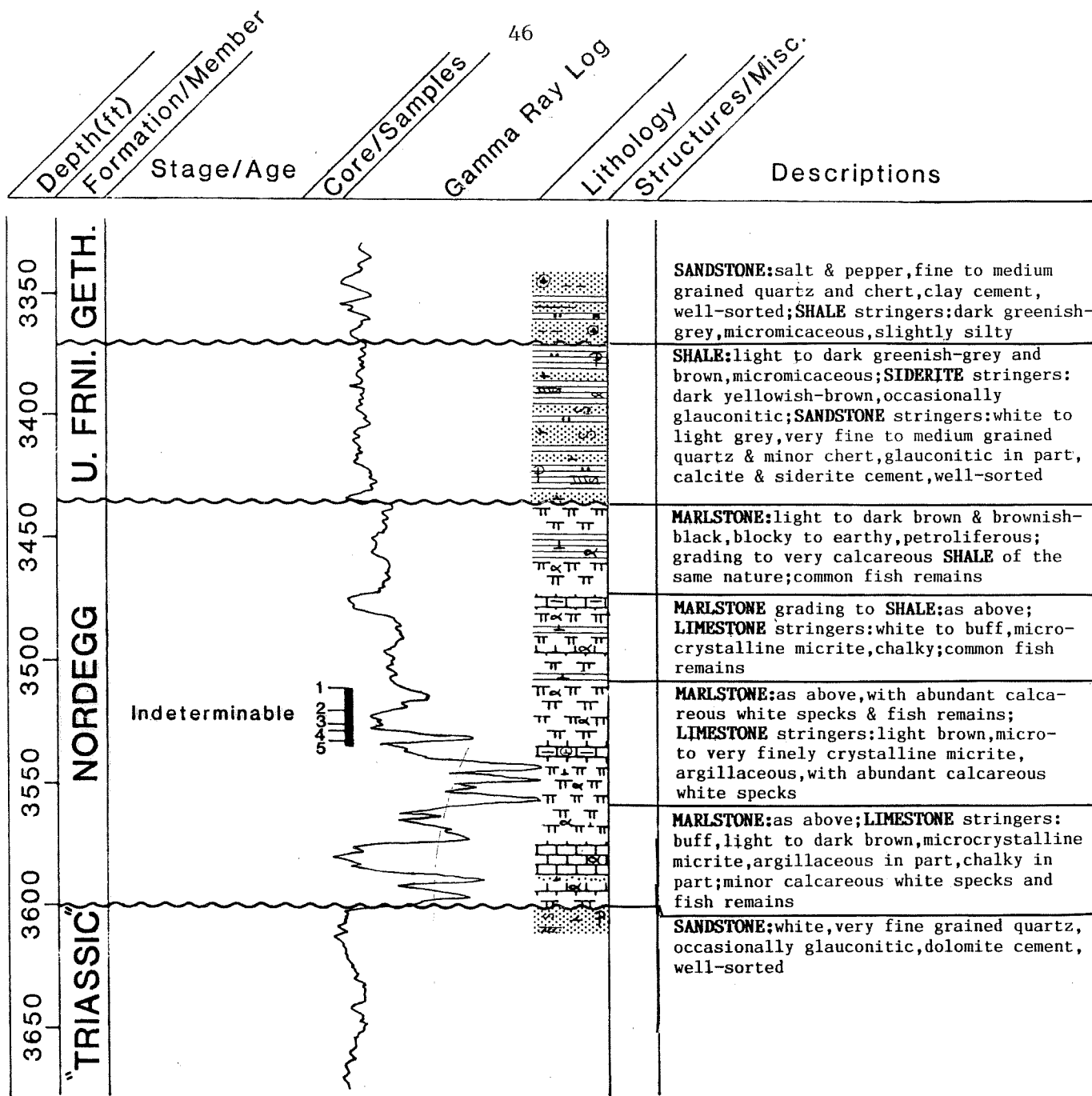


4 - 32 - 69 - 19W5M

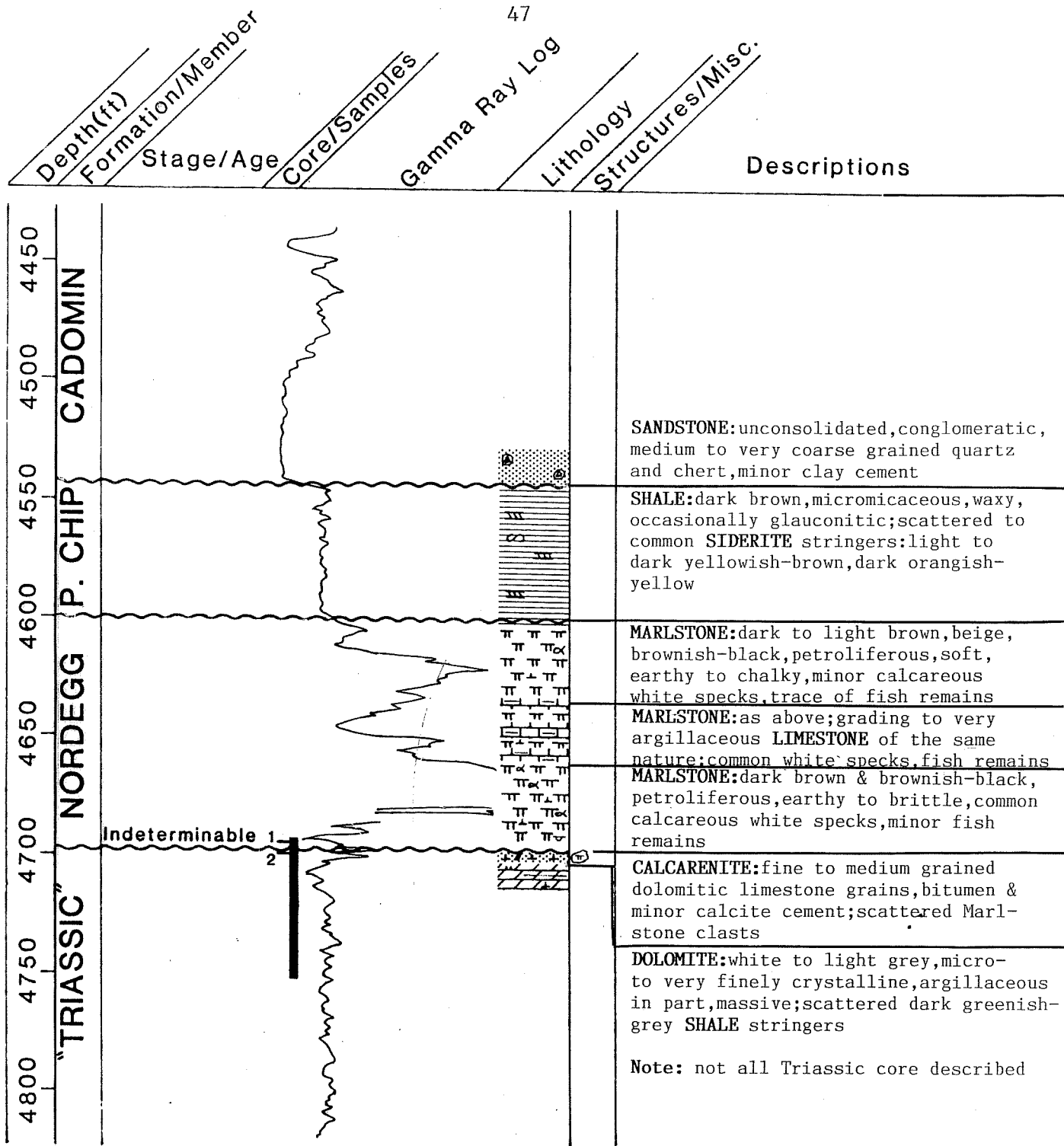
Depth(m)	Formation/Member	Stage/Age	Core/Samples	Gamma Ray Log	Lithology	Structures/Misc.	Descriptions
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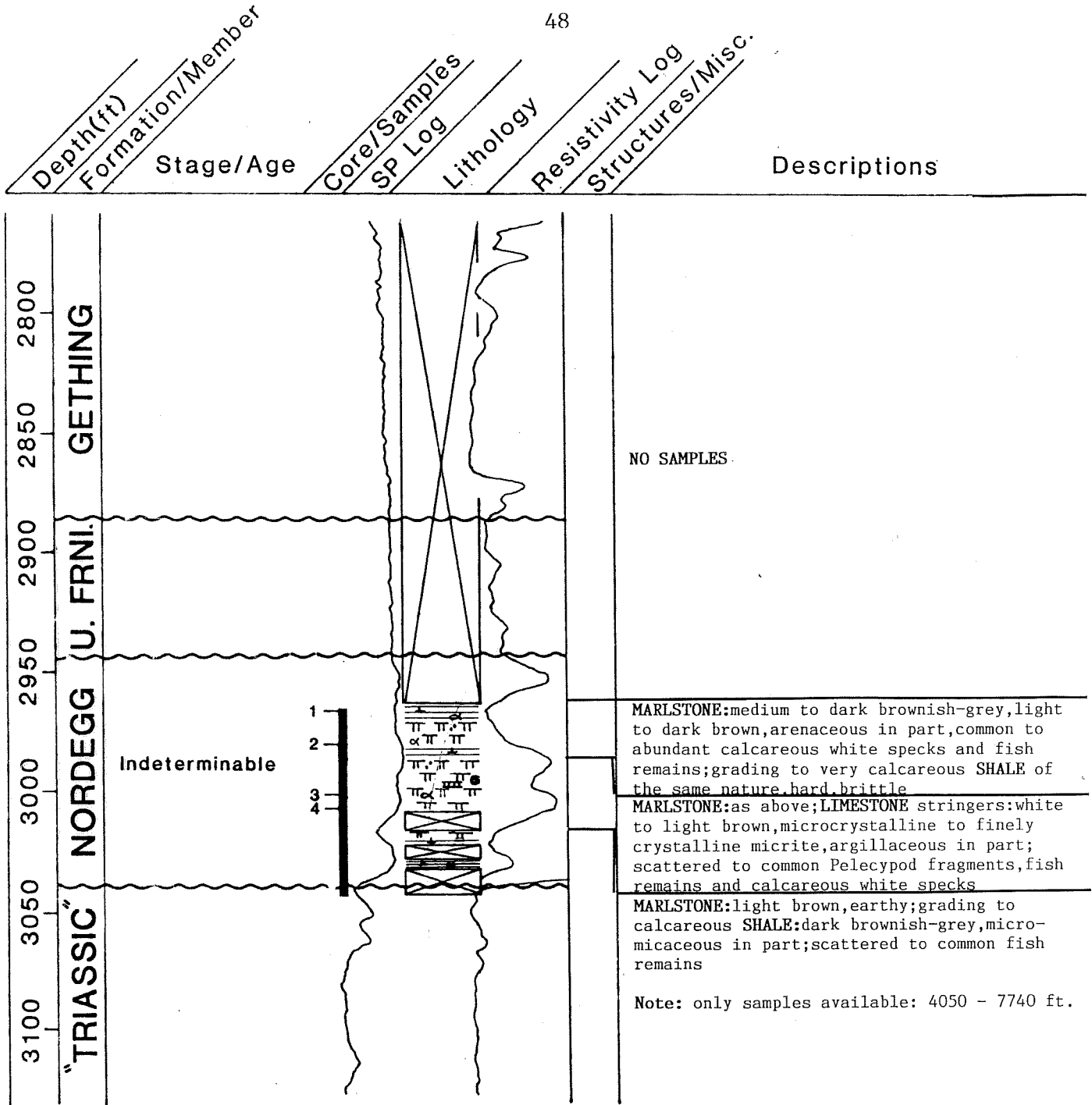
45





10 - 9 - 74 - 20W5M





12 - 33 - 77 - 21W5M

