

35 H 14

35 H 14



NOTES DESCRIPTIVES / DESCRIPTIVE NOTES. This map is one of a series of sixteen 1:50,000 scale geological maps (Fig. 1) for the eastern Early Proterozoic Cape Smith Belt (Fig. 2). The maps were compiled from the results of field work completed by the Geological Survey of Canada during the summers of 1982 to 1987 (St-Onge et al., 1987, 1988, 1989). The Open File map (Fig. 2) is accessible by scheduled flight from Kuluas, Québec (distance of 200 km) or by chartered aircraft from Iqaluit, Northwest Territories (distance of 350 km). Bedrock exposure in the mapped area is generally excellent, varying from continuous in the Wakeham Bay - Burgoyne Bay and Lac Watts - Lac Cross - Rivière Déception region to sufficient in the vicinity of Lac Bombardier and Lac Vicienza. The geological data presented in this Open File map were gathered during ground-level traverses at a spacing of 2 km or less. Tectono-stratigraphic and structural relationships are presented on the maps of this Open File series directly on the geology mapped during the three summers of field work. In contrast, the position of metamorphic mineral isograds was determined by follow-up petrographic and microprobe work (Bégin et al., 1988). The Open File map provides structural constraints for future mineral exploration projects in the area (St-Onge et al., 1988) and complements the MER (Quebec) for the western portion of the belt (Lamothe, 1986). The ca. 1.9 Ga (R. Parrish, pers. comm., 1986) Cape Smith Belt is a 600-km-long, south-vergent thrust-fold belt (Hynes and Francis, 1982; Lamothe et al., 1984; Hoffman, 1982) which is exposed in a west-plunging oblique section (13 km of structural relief) from low structural levels in the Wakeham Bay area (St-Onge et al., 1986) to high structural levels in the Lac Watts - Lac Cross area (St-Onge et al., 1987). The tectono-stratigraphic record of the Cape Smith Belt documents the evolution of an Early Proterozoic epicontinental rift which ultimately led to the formation of oceanic crust (Hynes and Francis, 1982). The continent-derived sediments of the lower Povungnituk Group (units 2, 3a, 3b and 4) record the opening and infilling of a rift margin basin, which at least in part overlies continental crust. The epicontinental rifting setting for the accumulation of the Povungnituk Group is supported by the similarity of upper Povungnituk Group mafic magmas (unit 3) to modern, within-plate continental basalts with respect to overall major element ratios, ranges in TiO2 content and trace element ratios (Hynes and Francis, 1982; Francis et al., 1983). The Chulotat Group (units 5, 6, 7, 8 and 9) records the formation of a transitional oceanic crust. The geochemistry of Chulotat Group mafic volcanics ranges from komatiitic basalts (unit 6) to low-Ti tholeiites (unit 8) with MORB affinities (Francis and Hynes, 1979; Hynes and Francis, 1982; Francis et al., 1983). The tholeiitic lavas are low in incompatible elements and have trace element characteristics very similar to those shown by modern ocean-floor basalts. Thrust sheets in the most internal (northern) part of the Cape Smith Belt carry the lower Proterozoic sediments (unit 13) and the Watta Group (units 10 and 11) and ultramafic cumulates (unit 9). These units are interpreted to constitute the lowermost and metamorphic remnants of Early Proterozoic oceanic crust, preserved in the thrust-fold belt as the Paratungit (Scott et al., 1981; St-Onge et al., 1988). The ophiolitic suite completes the tectono-stratigraphic record of the northern margin of the Superior craton, which evolved ca. 1.9 Ga ago from an epicontinental rift system to a true oceanic domain. The continental-rift, transitional-crust and ophiolitic suites of the Cape Smith Belt are defined by three generally east-trending, geometrically distinct sets of structures (Lucas and St-Onge, 1987; St-Onge and Lucas, 1988a). The cumulative effect of the D1, D2 and D3 deformation events is to preserve the thin-skinned thrust-fold belt in an east-trending, D1-syn, D2- and D3-antiform double-plunging as the result of D1 northwest-trending cross-folds (Fig. 2). The earliest set of D1 structures recorded in the region includes (Jugaback-style) south-verging thrust faults which root on a basal décollement localized at the Archaean Archean-Archaean Proterozoic cover interface (St-Onge and Lucas, 1986). Transport of the thrust along the basal décollement during thermal relaxation following the early unroofing of the cover resulted in the growth of a ductile basal shear zone (Lucas and St-Onge, 1988a, b). This post-thermal post-orogenic double-plunging range through the previously assembled thrust stack to achieve late D1 crustal thickening (Lucas and St-Onge, 1988a). Incorporation of laterally discontinuous basement slices (unit 1) into the thrust belt is associated with north-south trending, thick-skinned D2 and D3 folding of basement (unit 1) and cover (units 2 to 14) produced a dome and basin fold interference pattern. The distribution of D2 and D3 axial traces shown on the geological completion maps of this Open File emphasizes the importance of mesoscale layering in producing buckle folds and determining their wavelength. Features of this phenomenon include the D2 fold development where large layered gabbro-peridotite sills intrude the sediments of unit 13.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

Hot-side-dome metamorphic mineral isograds in the mapped area document a normal distribution of isotherms in the thickened Early Proterozoic crust following D1 unroofing (St-Onge and Lucas, 1988a). The lower granulite to upper amphibolite facies isograds (unit 1) are associated with the D1 deformation in the southern part of the Cape Smith Belt (Bégin et al., 1983). In contrast, along the northern hinterland margin of the belt, the mineral zones are truncated by the late D1 (retrograde) out-of-sequence thrusts. Extensive overprinting retrograde mineral assemblages are developed in the high-grade units of the late D1 faults.

LEGEND / COLONNE TECTONOSTRATIGRAPHIQUE. Notes: 1. This legend is common to all of the map sheets of this Open File. However, not all map units will appear on each map sheet. / La colonne tectonostratigraphique est commune pour toutes les cartes de dossier publique. 2. The prefix "meta" applies to all lithologies in units 1 to 13 / Le préfixe "méta" s'applique à l'ensemble des lithologies des unités 1 à 13. LITHOLOGIES / Lithologies: Pelite / phylite, Semipelite / siltstone, Micaceous sandstone / grès micacé, Conglomerate / conglomérat, Dolomite / dolomite, Calc-siltstone / calcilite, Sandstone / grès, Dolomitic sandstone / grès dolomitique, Basalt / basalte, Layered gabbro / gabbro stratifié, Sheeted gabbroic dykes / dykes de gabbro en feuilles, Volcanic breccia / brèche volcanique, Pillowed flow / coulée coussinée, Pyroxenite / pyroxénite, Layered gabbro / gabbro stratifié, Layered peridotite / péridotite stratifiée, Dominantly phylitic basalt / basalte à phénocristaux de phylite, Basaltic gabbro sills / basalte à phénocristaux de gabbro, Olivine-phyric basalt / basalte à phénocristaux d'olivine, Rhyolite / rhyolite, Plagiogranite / plagiogranite, Dominantly plagioclasic basalt / basalte à phénocristaux de plagioclase, Dominantly pyroxenitic basalt / basalte à phénocristaux de pyroxène, Dominantly olivine-phyric basalt / basalte à phénocristaux d'olivine, Granitic-gneiss / schiste métamorphe avec granite et gneiss, Granitic-dolomitic schist / schiste métamorphe avec granite et dolomite, Tonalite / tonalite, Amphibolite xenoliths / xenolithes d'amphibolite, Disrupted gneiss / gneiss chaotique, Straight gneiss / gneiss régulier, Rusty, iron-stained / ferrugineux, Gabbro / gabbro, Peridotite / péridotite, Pyroxenite / pyroxénite, Ironstone / miner sandstone and semipelite / sédiments ferrugineux, Layered peridotite / péridotite stratifiée, Sandstone / grès, Conglomerate / grès, Siltstone / argilite, Gneiss / gneiss, Amphibolite / amphibolite, Garnet or clinopyroxene / garnet ou clinopyroxène.

Geology by / géologie par: M.R. St-Onge, S.B. Lucas, D.J. Scott, N.J. Bégin.

Compilation by / compilation par: M.R. St-Onge, S.B. Lucas, D.J. Scott, N.J. Bégin.

Notation bibliographique conseillée: St-Onge, M.R., Lucas, S.B., Scott, D.J., and Bégin, N.J., 1988.

Recommended Citation: St-Onge, M.R., Lucas, S.B., Scott, D.J., and Bégin, N.J., 1988.

Figure 1. National Topographic System reference and index to Geological Survey of Canada maps / référence cartographique et index des cartes de la Commission Géologique du Canada.

Figure 2. Location of Geological Survey of Canada maps in the eastern portion of the Cape Smith Belt, northern Québec / localisation des cartes de la Commission Géologique du Canada dans la partie est de la zone de Cape Smith, Québec.

FEUILLE 3 DE 16 / SHEET 3 OF 16 GÉOLOGIE / GEOLOGY. SECTEUR ORIENTAL DE LA ZONE DE CHEVAUCHEMENT ET DE PLSSEMENT DU CAPE SMITH; PARTIE DES CARTES DE WAKEHAM BAY, CRATÈRE DU NOUVEAU-QUÉBEC ET NUUVILIK LAKES, QUÉBEC SEPTENTRIONAL. EASTERN PORTION OF THE CAPE SMITH THRUST-FOLD BELT; PARTS OF THE WAKEHAM BAY, CRATÈRE DU NOUVEAU-QUÉBEC AND NUUVILIK LAKES MAP AREAS, NORTHERN QUÉBEC. Echelle 1:50 000 - Scale 1:50 000. Projection universelle de Mercator / Universal Transverse Mercator Projection. Droits de la Couronne réservés / Crown Copyrights reserved.

SYMBOLS / SYMBOLES. Boundaries / contacts: Geological boundary (defined, approximate) / contact géologique (fortain, probable), Thrust fault / faille de chevauchement D1, Oblique-slip fault / faille avec décrochement oblique. Structure / Structures: Bedding, tops known (inclined) / stratification, sommet déterminé (incliné), Gneissosity (Archean) / gneissosité (archéenne), D1 schistosity (inclined) / schistosité D1 (incliné), D1 stretching lineation / linéation d'allongement D1, D1 syncline / synclinal D1, D1 anticline / anticlinal D1, D2 schistosity (inclined) / schistosité D2 (incliné), D2 minor-fold hinge (N - northward fold vergence, S - southward fold vergence) / charnière de pli secondaire D2 (N - vergence à l'ouest, S - vergence à l'est), D2 synform / synforme D2, D2 antiform / antiforme D2, D3 schistosity (inclined) / schistosité D3 (incliné), D3 minor-fold hinge (E - eastward fold vergence, W - westward fold vergence) / charnière de pli secondaire D3 (E - vergence à l'est, W - vergence à l'ouest), D3 synform / synforme D3, D3 antiform / antiforme D3, homblande-in / apparition de la homblande, oligoclase-in / apparition de l'oligoclase, actinolite-out / élimination de l'actinolite, garnet or clinopyroxene-in / apparition du garnet ou du clinopyroxène. Geology by / géologie par: M.R. St-Onge, S.B. Lucas, D.J. Scott, N.J. Bégin. Compilation by / compilation par: M.R. St-Onge, S.B. Lucas, D.J. Scott, N.J. Bégin.

Figure 1. National Topographic System reference and index to Geological Survey of Canada maps / référence cartographique et index des cartes de la Commission Géologique du Canada.

Figure 2. Location of Geological Survey of Canada maps in the eastern portion of the Cape Smith Belt, northern Québec / localisation des cartes de la Commission Géologique du Canada dans la partie est de la zone de Cape Smith, Québec.

FEUILLE 3 DE 16 / SHEET 3 OF 16. OPEN FILE / DOSSIER PUBLIC 1730. GEOLOGICAL SURVEY OF CANADA / COMMISSION GÉOLOGIQUE DU CANADA 1988.