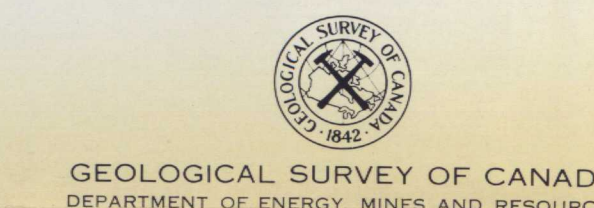
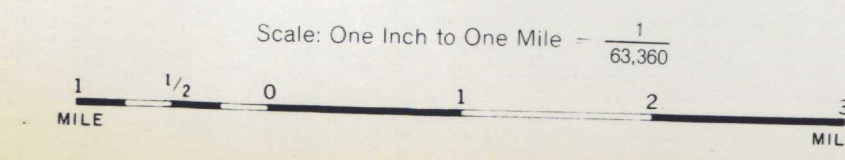


POINT LEAMINGTON

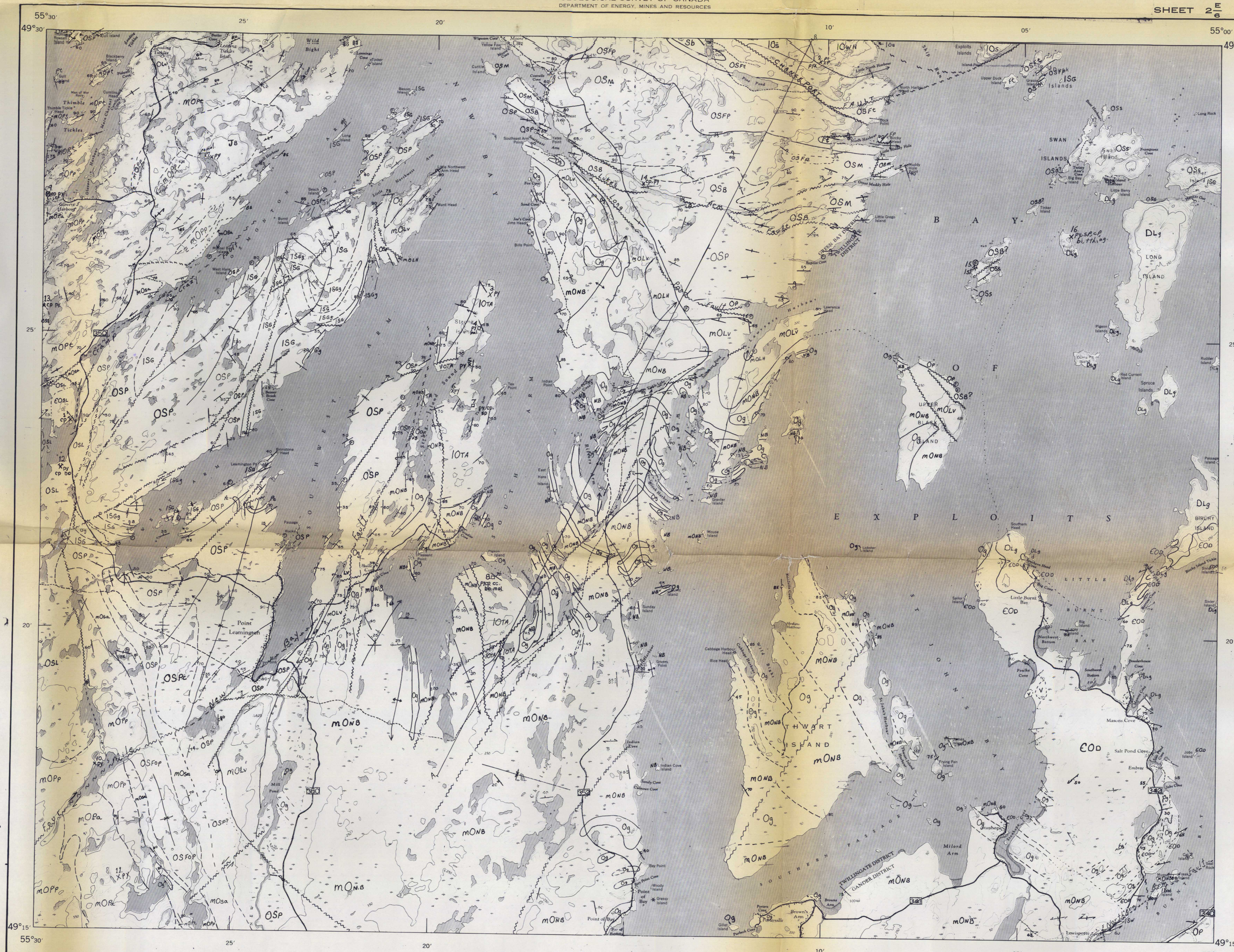
NEWFOUNDLAND



SHEET 2

LEGEND - Point Leamington Map-Area

- JURASSIC
DEVONIAN
SILURIAN
ORDOVICIAN AND SILURIAN
EXPLOITS GROUP
MIDDLE AND LOWER ORDOVICIAN
LOWER AND/OR ORDOVICIAN
CAMBRIAN OR LOWER ORDOVICIAN
WILD BIGHT GROUP
MORÉTONS HARBOUR GROUP
PENNY'S BROOK FORMATION
SHERNY ISLAND TERRANE
MORÉS COVE FORMATION
SANSOM GREYWACKE
SOUTH LAKE OPHIOLITE



Mineral Occurrences, Prospects and Mines

- 1. South Lake #1: Chalcopyrite stringers in sheeted diabase of the South Lake Ophiolite.
2. Little North Harbour: Disseminated and semi-massive pyrite and chalcocopyrite in chloritized altered basalt.
3. Strong Island: Disseminated pyrite in mafic volcanics.
4. Strong Island: Disseminated pyrite in mafic tuff and agglomerate with faint malachite staining.
5. & 6. Strong Island Sound: Disseminated pyrite in mafic volcanics.
7. Tea Arm: Heavily disseminated pyrite with minor chalcocopyrite and sphalerite in a thin acidic pyroclastic unit within mafic volcanics and associated cherts.
8. Saunders Cove: Small sulphide pods and stringers in altered mafic volcanics and in calcite-epidote veins cutting the volcanics.
9. Lock Port Mine: Massive pyritic sulphide body with low copper values and local sphalerite rich zones in acidic volcanic rocks and associated chert.
10. Four Mile Pond: Disseminated and stringers of pyrite in chloritized mafic volcanics.
11. Big Asco Pond Area: Disseminated pyrite throughout acidic volcanic rocks over an extensive area.
12. South Lake #2: Veinlets of pyrite, chalcocopyrite and bornite in hornblende diorite and in quartz veins intruding diorite.
13. Crag Gray North: Disseminated pyrite with minor chalcocopyrite in quartz diorite near contact with tuffs.
14. Cully's Pond: Disseminated and stringers of pyrite and chalcocopyrite in altered basalt.
15. Hummocky Island: Lenses of limestone up to 15 ft wide in chloritized mafic volcanics and grewackes.
16. Pond Island: Blebs, crystals and stringers of pyrite and sphalerite with minor chalcocopyrite, tetrahedrite and bismuthinite in narrow quartz veins in an intrusive rhyolitic phase of the Lusk Bight Group.
17. Budge's Harbour Area: Disseminated pyrite with local malachite staining in the Budge's Harbour gabbro.

- Geological boundary (defined, approximate and assumed)
Bedding, tops known (horizontal, inclined, vertical, overturned)
Bedding, tops unknown (inclined, vertical)
Dyke trend (inclined, vertical)
Schistosity, gneissosity, cleavage, foliation (horizontal, inclined, vertical)
Fault (defined, approximate, assumed)
Anticline (upright, overturned)
Syncline (upright, overturned)
Fossil locality
Mineral occurrence
Mineral prospect, test pit or trench
Shaft or quarry (exploration, abandoned production, production, producing)
Adit or tunnel

CAMBRIAN AND/OR ORDOVICIAN

- 60d DUNNAGE MELANGE: Black to green and black chaotic argillite and pebbly to cobble mudstone in exotic blocks and boulders, greywacke, bedded clastic rocks, agglomerate, gabbro, limestone and granite, v. large volcanic masses in shaly melange.

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Explanatory Notes - Point Leamington Map-Area

The oldest rocks of the map-area are the layered gabbros and diabase dykes of the South Lake Ophiolite. These gabbros and dykes are generally amphibolitized but are otherwise identical to other Cambro-Ordovician ophiolitic rocks from elsewhere in the Central Mobile Belt (Updchay et al., 1971; Strong, 1973; Norman and Strong, 1975). Other phases of the South Lake Igneous Complex, which is in part intrusive but dominantly fault-bounded, may also be ophiolitic. The volcanic rocks of the Morétons Harbour Group, exposed north of the Swillingate area to the east (2E/10). The basal unit of the Group in Budge's Harbour Gabbro which is oceanic tholeiite like that area is the Sleepy Cove Formation which is transitional with the Lusk Bight Group to the west (2E/12, 12H/9). By lithologic and stratigraphic similarity with the Western Arm and Cutwell Groups, the Morétons Harbour Group is Lower Ordovician in age and overlies oceanic mafic rocks of the Tea Arm Volcanics (8 km) in the Morétons Harbour area (Strong and Payne 1973). Its island arc tholeiite chemistry and the abundance of diabase dykes, the Morétons Harbour Group probably represents the core of a volcanic island arc built on oceanic crust.

In the map-area, the lower unit of the Morétons Harbour Group, the Shevny Island Terrane, is characterized by numerous dykes which act as feeders to the pillow lavas of the overlying Western Head Formation. The contact between the two units is transitional over 30 m and the lowermost pillow lavas of the Western Head Formation often have thin lenses of red chert and associated sulphide showings.

South of the Chanceport Fault, Ordovician rocks of pre-Caradoecian age are exposed in three distinct areas: (1) the Wild Bight Group in the north, (2) the lower half of the Exploits Group in the centre, and (3) the Dunnage Melange in the east.

Only the uppermost formation of the Wild Bight Group, the Penny's Brook Formation, is exposed in the map-area. The full sequence of the group is exposed in the Roberts Arm area (2E/5) to the west. The volcanic and volcanogenic sedimentary rocks of the Wild Bight Group overlies oceanic crust, represented in the map-area by the South Lake Ophiolite.

The lowermost unit of the Exploits Group (Helwig, 1969) is the Tea Arm Volcanics. This unit of dominantly mafic volcanics can be correlated with the lower Ordovician of the Wild Bight Group which underlies the Penny's Brook Formation to the west (2E/5). The Tea Arm Volcanics is overlain by dominantly mafic volcanics of the New Bay Formation. Tuff, argillite and chert of the New Bay Formation previously named the Saunders Cove Formation (Helwig, 1969) are presently included in the New Bay Formation since similar lithologies occur throughout the New Bay Formation, and the Saunders Cove Formation is not lithologically distinct and the name is abandoned.

Although the New Bay Formation consists dominantly of graded sandstones and conglomeratic turbidites, tuff, tuffaceous sandstone and argillite are very common. These volcanogenic sedimentary rocks are identical to lithologies of the Penny's Brook Formation to the east. The New Bay Formation then probably represents a more distal sedimentary facies to the volcanic and volcanogenic sedimentary rocks of the Penny's Brook Formation. The New Bay Formation overlies conformably by pillow lavas and flows can be correlated stratigraphically and lithologically with pillow lavas at the top of the Penny's Brook Formation west of Point Leamington.

The New Bay Formation in part underlies and transitionally grades eastward into the chaotic Melange (Williams and Hibbard, 1976). The Lawrence Head Volcanics occur as large blocks in the melange indicating a mid-Ordovician age for the formation of the melange in the area (Eldredge, 1968).

Pre-Caradoecian units south of the Chanceport Fault are believed to be Ordovician in age. The South Lake Igneous Complex, which is dominantly fault-bounded, locally intrudes the Wild Bight Group, although some phases are clearly ophiolitic and older than the Exploits Group. The Wild Bight Group, the lower half of the Exploits Group and the Dunnage Melange are intruded by gabbro sills which pre-date the formation of the Dunnage Melange (Hibbard, 1976).

Lower to Middle Ordovician rocks south of the Chanceport Fault are overlain by carbonaceous and cherty argillites of Caradoecian age. The Penny's Brook Formation of the Wild Bight Group is overlain conformably by the Shoal Arm Formation consisting of black argillite with red and black chert near the base. Graptolites from the Shoal Arm Formation range the entire span of the Caradoecian.

The Lawrence Head Volcanics of the Exploits Group is overlain by the Lawrence Harbour Shale, redefined by Helwig (1969) graptolite assemblage. Younger Caradoecian shales of the Exploits Group, included in the Lawrence Harbour Shale by Heyl (1936), are apparently faulted against the Lawrence Harbour (Helwig 1969) at Lawrence Harbour. These shales are correlative of the younger parts of the Shoal Arm Formation and are designated 'Op' on the map. Caradoecian shales of the Dark Hole Formation apparently overlie the Dunnage Melange to the east (Williams and Hibbard, 1976).

Southwest of Point Leamington, the Shoal Arm Formation is apparently conformably overlain by mafic pillow lavas. These lavas are in an analogous stratigraphic position as mafic pillow lavas of the Frozen Ocean Group which overlie the Shoal Arm Formation in the Lewis Lake Area to the west (2E/5). For this reason these lavas are also included in the Frozen Ocean Group.

With the exception of the above-mentioned pillow lavas, the Caradoecian argillites are everywhere overlain by a thick sequence of greywacke of Upper Ordovician to Silurian age. The greywacke is known as the Point Leamington Greywacke in the New Bay Area (Helwig, 1969) and the Sansom Greywacke in Bay of Exploits. A large lentil of polymictic conglomerate appears within the greywacke between Point Leamington and West Arm. The Point Leamington and Sansom Greywackes are both overlain conformably and gradationally by coarse conglomerate of the Goldson Formation of Lower Silurian age. The Goldson Formation there is no major break in the type of sedimentation - at the Sansom-Goldson boundary.

On the Fortune Harbour Peninsula, the Point Leamington Greywacke is overlain by the Cottrells Cove Group of probable Silurian age (Dean 1973). The base of the Cottrells Cove Group is a complex sedimentary-volcanic melange known as the Boones Point Complex. The Boones Point Complex was interpreted by Helwig (1967) as a tectonic melange in the Luke's Arm Fault zone and redefined the Luke's Arm Fault as the base of the melange. The Boones Point Complex is best regarded as a stratigraphic unit - a complex slump melange that possibly marks the beginning of Cottrells Cove Group volcanism. Blocks and lenses of Goldson-type conglomerate are common and a fault-bounded wedge of volcanic-limestone melange on upper Black Island, tentatively included in the Boones Point Complex, contains Silurian fossils (Bonocot 1969). The Boones Point Complex can be correlated with the upper part of the formation can be correlated with the Crescent Lake Formation of the Roberts Arm Group to the west (2E/5). The Moore's Cove Formation is overlain conformably by dominantly volcanic rocks of the Fortune Harbour Formation. The top of the Fortune Harbour Formation is faulted against the lower Ordovician Morétons Harbour Group along the Chanceport Fault.

Gabbro and diabase sills, possibly contemporaneous with the Fortune Harbour Formation.

South of the Northern Arm Fault, near Lewisporte, Silurian red sandstones of the Migwam Formation of the Botwood Group are exposed on the coast and islands of Burnt Bay in fault contact with the Dunnage Melange.

The Long Pond granitic batholith of Devonian age intrudes Ordovician and Silurian rocks post-tectonically. The youngest rocks in the map-area are Jurassic-Cretaceous lamprophyre dykes and the Budge's Harbour Gabbro which has a K-Ar date of 155.7 m.y. (Strong and Harris, 1974). This alkalalic igneous event heralds the opening of the present Atlantic Ocean (Helwig et al., 1974).

Segments of two major faults, the Chanceport Fault and the Northern Arm Fault, are exposed in the map-area. The Chanceport Fault, continuation of the Lobster Cove Fault to the west (2E/12) and is the major fault of northern Notre Dame. This major fault always separates the lower Ordovician volcanic rocks on the north from north-facing Upper Ordovician-Silurian volcanic rocks to the south. It is presently interpreted as an early Acadian thrust fault with the Morétons Harbour Group. The thrust fault was subsequently steepened by later Acadian folding. The anticline at Little North Harbour is asymmetrical with a steeply overturned south limb. This structure was probably a nappe-like feature related to thrusting.

Acadian folds are dominantly northeast-trending and northeast-plunging, typical of Acadian structures. However, in the southwest part of the map-area, fold axes trend northwest and plunge southeast. This area is on the east limb of a major southeast-plunging anticline. North of the Long Pond Fault on the Fortune Harbour Peninsula, Acadian folds are east-trending and east-plunging, parallel to the Chanceport Fault.

The Northern Arm Fault, in the southeast corner of the area, is a continuation of the Reach Fault to the northeast (2E/7, 2E/10). It post-dates the main phase of Acadian folding and granitic intrusion and is possibly a major transcurrent fault. Other northeast-trending faults such as those at the head of the arms of New Bay are probably related to the same phase of faulting. Silurian red sandstones and subvolcanic volcanics of the Botwood Group are found only southeast of the Reach-Northern Arm Fault system which marks the boundary between the Exploits and Botwood zones in the central Newfoundland Appalachians (Williams et al., 1972).

- Boucot, A.J., 1969. Silurian-Devonian of Northern Appalachians-Newfoundland. In: North Atlantic-geology and continental drift. Rev. (Ed.), Am. Assoc. Petrol. Geol., Mem. 12, pp. 477-483.
Dean, P.L., 1973. The Geology of the Northern half of the Fortune Harbour Peninsula. Report to Geol. Surv. Can., 31 p.
DeZoysa, T.H., 1969. Geology and Base Metal Mineralization of Lockport Area, Notre Dame Bay, Newfoundland. M.Sc. thesis. Memorial Univ. of Nfld., 99 p.
Helwig, J., 1967. Stratigraphy and Structural History of the New Bay Area, North Central Newfoundland. Ph.D. thesis, Columbia Univ., 211 p.
Helwig, J., 1969. Redefinition of Exploits Group, Lower Paleozoic, Northern Newfoundland. In: North Atlantic-geology and continental drift, Kay (Ed.) Am. Assoc. Petrol. Geol., Mem. 12, pp. 408-413.
Helwig, J., Anonon, J., and Bay, D.S., 1974. A Late Jurassic Mafic Pluton in Newfoundland. Can. J. Earth Sci., 11, pp. 1314-1319.
Heyl, G.R., 1936. Geology and mineral resources of the Bay of Exploits Area, Newfoundland Dept. Nat. Resources, Geol. Sect., Bull. 8, 42 p.
Heyl, G.R., 1937. Geology of the New Bay map area, Newfoundland. Unpub. report on File Geol. Surv. Can.
Hibbard, J.P., 1976. The southwestern extension of the Dunnage Melange and its relations to surrounding units, N.S.C. thesis. Memorial Univ. of Nfld.
Kay, H., and Eldredge, N., 1968. Cambrian trilobites in central Newfoundland volcanic belt. Geol. Mag., 105, pp. 372-377.
Noranda Exploration Co. Ltd. unpublished maps and reports.
Norman, R.B., and Strong, D.F., 1975. The geology and geochemistry of ophiolite rocks exposed at King's Bight, Newfoundland. Can. J. Earth Sci., 12, pp. 771-787.
Oversby, B.S., 1967. Geology of Upper Black Island, Bay of Exploits, Newfoundland. M.A. thesis, Columbia Univ., 48 p.
Strong, D.F., 1973. Lusk Bight and Roberts Arm Groups of central Newfoundland: possibly juxtaposed oceanic and island-arc volcanics series. Bull. Geol. Soc. Am., 84, pp. 3917-3928.
Strong, D.F., and Payne, J.C., 1973. Early Paleozoic volcanism and metamorphism of the Morétons Harbour-Wellington area, Newfoundland. Can. J. Earth Sci., 10, pp. 1363-1379.
Strong, D.F., and Harris, A., 1974. The Petrology of Mesozoic alkaline I.I. pp. 208-219.
Updchay, H.D., Innes, J.P., and Neale, E.R.W., 1971. The Betts Cove ophiolite complex, Newfoundland: Appalachian oceanic crust and mantle. Geol. Assoc. Can. Proc. v. 24, pp. 27-34.
Williams, H., 1963. Botwood Map Area. Geol. Surv. Can., Map 60-1663.
Williams, H., Kennedy, M.J., and Neale, E.R.W., 1972. The Appalachian P.A. Fold and R.W.M. Douglas (Eds.) Geol. Assoc. Can. Spec. Paper 11, pp. 181-24.
Williams, H., and Hibbard, J.P., 1976. The Dunnage Melange, Newfoundland. Geol. Surv. Can., Paper 76-1a.
Geology by Heyl (1937), Helwig (1967, 1969), Oversby (1967), DeZoysa (1969), Noranda Exploration Co. Ltd., (1971-1974), unpublished, P.L. Dean (1973), 1974, unpublished, Hibbard (1976), and Williams and Hibbard (1976).
Compiled by P.L. Dean, Memorial University of Newfoundland, 1976.