



CHÉTICAMP RIVER
NOVA SCOTIA

Scale 1:50 000 Echelle

1041
GEOLOGICAL SURVEY
COMMISSION GÉOLOGIQUE
OTTAWA

CONVERSION SCALE FOR ELEVATIONS
Echelle de conversion des altitudes

FOOT METERS
0 100 200 300 400 500
0 100 200 300 400 500

LEGEND

STRATIGRAPHIC UNITS AND THEIR ALTERNATIVE EQUIVALENTS

CAMBRIAN

DEVONIAN - EARLY CARBONIFEROUS

DEVONIAN

LATE PROTEROZOIC

LATE PROTEROZOIC

SYMBOLS

Strike and dip of bedding, top known, vertical, overturned

Strike and dip of foliation, first, second, vertical

Strike and dip of unconformity, first, second, vertical

Strike and dip of joints, inclined, vertical

Trend and plunge of minor folds, first, second

Trend and plunge of lineation, down dip, vertical

Geological contact, known, estimated

Steep fault, known, estimated (sense of offset shown where known)

Break fault, known, estimated (sense of offset shown where known)

Outcrop locality

Fossil locality

Mafic dyke

RECONNAISSANCE GEOLOGICAL MAP OF THE SOUTHERN CAPE BRETON HIGHLANDS -
Marginal notes to accompany Map Sheets 11K/10, 11K/11, 11K/12

The Cape Breton Highlands form a steep-sided, fault-bounded plateau underlain by Proterozoic to Carboniferous crystalline rocks and surrounded by low-lying Carboniferous sedimentary rocks. These three maps sheets show the geology of the crystalline rocks of the southwestern Highlands between the southern boundary of Cape Breton Highlands National Park and the Cabot Trail in the vicinity of Middle River. The area is heavily forested with outcrop restricted to stream valleys and logging roads. Detailed mapping of streams and roads was undertaken between 1979 and 1981 in the area bounded by Fielding Road, the Cabot Trail, and the North Boreas River (parts of 11K/7 and 11K/21); a more detailed map of this area is available on a scale of 1:25 000 (Geological Museum - Middle River area). The rest of the map (parts of 11K/7, 11K/10) represents a survey of the Nova Scotia Forest Industries road network carried out in 1981-82. No detailed mapping has been done by this author north of Fielding Road. Overlapping areas mapped in detail by Currie (in press), Blamont (1982), Barr and Macdonald (1982), Doucet (1983), Fawcett et al. (in press), and Crow (in press) are indicated on the map.

The oldest rocks in the area are late Proterozoic metamorphic rocks including undeformed gneisses (E_g), George River Group metasediments (E_g), and low to medium grade metamorphic and metasedimentary rocks (E_g). The gneisses are mainly semi-pelitic to pelitic in composition, with small areas of amphibolite (E_g) and tonalitic orthogneiss (E_g) recognized locally. They are compositionally banded and well foliated and contain several generations of pegmatite bodies. The main north-south trending foliation of the gneisses, defined by muscovite and biotite, is generally a second generation retrograde overprint of the original anisotropy. Horizons of feldspathic quartzite within the gneisses near First Fork Brook and marble north of Middle River suggest that the high grade rocks of the central Highlands are in part equivalent to metamorphoses of the George River Group which occur in the eastern margin of the mapped area. Lithologic, orthostructural and elastic characteristics of the George River Group occur in the map area, semi-pelitic compositions are more common and amphibolite and calc-silicates zones of uncertain origin have also been recognized; the correlation of these metasediments with the George River Group should be regarded as tentative. Low to medium grade metamorphic and metasedimentary rocks underlie large parts of the western and southern Highlands. Rock types include crystal-litic tuffs, felsic pyroclastic rocks of various types, minor rhyolitic and mafic flows, fine grained pelitic sediments, and minor pegmatites and conglomerates. The distribution of these rocks is complex in detail and presumably partly reflects an original complex interfingering of various units. However, the present outcrop pattern of these units must be regarded as tectonic - in both the Chéticamp River and Middle River areas they have been affected by the same extensional or tectonic folding followed by upright folding and shortening along shear zones (Crow, in press; Doucet, 1983). The rocks have been affected by greenschist to lower amphibolite facies metamorphism with grade increasing toward the central Highlands. The age of the volcanic and sedimentary rocks with respect to the gneisses is uncertain. There is no evidence of a stratigraphic unconformity in the study area, and in the Chéticamp River and Middle River areas the transition from gneisses and amphibolites to gneisses appears gradational, with telescoping along high strain zones (Crow, in press; Doucet, 1983). The high grade rocks are intruded by the Late Proterozoic Boreas Lakes tonalite complex which shows faulted contacts against the lower grade rocks. The low grade rocks are intruded by the Cambrian Chéticamp pluton which is brecciated but not foliated and are thus probably also Proterozoic. These rocks are tentatively interpreted as a volcanic and sedimentary complex coeval with the quartzites to the west on the basis of lithologies transitional between silicic pyroclastics and quartzites that occur near Rocky Brook and Sugarloaf Mountain. This implies that the entire metamorphic complex formed from a single proterozoic package that underwent low to high grade metamorphism in late Proterozoic times.

The metamorphic rocks are intruded by dioritic to granitic plutons of various ages. The Boreas (Bell) Lakes complex (E_g) is a variably foliated, medium grained tonalite to diorite complex that has been dated at 752 ± 26 Ma by Rb-Sr whole rock methods (Jamieson, in press). (Note that the National Topographic Survey map used in 1979-82, the Lake "Boreas Lakes" and reports using this name have been published (Jamieson and Crow, 1982; Jamieson, in press.) This pluton intrudes metamorphic rocks and is cut by mafic dykes of unknown age. It is faulted against metamorphic rocks on its western margin, and is cut by major shear zones in the east. Its eastern limit has not been determined. Other dioritic and tonalitic rocks to the north (E_g) have not been mapped in detail. They are generally similar to the Boreas Lakes complex except for a somewhat more mafic composition. In the vicinity of North River they intrude gneisses and quartzites of the Chéticamp River Group. These plutons are part of a belt of diorites that extends from Ingonish through the eastern Highlands (Jamieson and Crow, 1983); plutons of various ages are present but few have been mapped in detail (Jamieson et al., in press).

In the central part of the map area variably foliated, fine, medium and coarse grained gneisses (E_g) are common. Their age is uncertain but they must post-date the gneisses and pegmatite undeformed Devonian gneisses. A medium grained, foliated, coarse grained Devonian gneiss, a medium grained, foliated, coarse grained Devonian gneiss (E_g) is present in the northern part of the mapped area (F_g "Four granites" resembles a foliated version of the Cambrian Chéticamp pluton and may be of similar age). Foliated gneisses of this type occur east of the Margaree Valley near John Peers Mountain (E_g). Coarse grained foliated gneisses with prominent F_g-like schists (E_g) are common in the Taylor's Barren - Egypt Highland area (Doucet, 1983).

Lower Paleozoic granitic rocks have been identified in the western Highlands (Cambrian Chéticamp pluton (E_g); Barr and Macdonald, 1982) and the western Highlands (Silurian-Devonian Cape Breton pluton (E_g); Barr et al., 1981; not shown on this map) but these have not been mapped in detail by this author. Granitic rocks of uncertain age have been mapped in the southern Highlands. These include the Robin Brook Intrusion (E_g) which may be a highly altered phase of a Devonian granite to the north, and the Caplan's Mill Brook diorite-microgranite complex (E_g). The Caplan's Mill Brook complex in places appears to be subordinate to Middle Devonian rhyolites (Jamieson and Doucet, 1983) but it is not reformed and altered by shear zones that most distinct relationships are well known. The rocks are generally fine grained and well developed intrusion breccias and local hybrid textures have been observed, particularly near the North Boreas River.

Devonian to Carboniferous granitic rocks are prominent in the central Highlands. The North River monzonite (E_g) is a medium grained biotite-amphibole monzonite to granodiorite with local albite and pegmatitic phases. It is weakly foliated to undeformed and has a distinctive black and white appearance with rare masses of K-feldspar which in the east develop into megacrysts. It intrudes gneisses and foliated granites in the vicinity of the Highland Road - Fielding Road section, and is cut by the northern extension of the Farnon Brook shear zone. It has been dated as 401 ± 19 Ma by Rb-Sr whole rock methods (Jamieson, in press). The Farnon Brook gneiss (E_g) is a fine grained monzonite associated with minor albite extending from the Chéticamp River south to the Campbell's Lake. Its age is uncertain but it bears some similarity to the Leonard's Mill Brook complex, and thus may be as old as Devonian, and has also been reported to intrude the Margaree gneiss east of the mapped area (Currie, in press) and thus may be as young as Carboniferous. The Margaree gneiss (E_g) is the youngest pluton in the study area. It has a distinctive rapakivi texture, and is associated with rhyolite porphyry dykes on its margins. It is offset by a major fault zone and has been dated at 350 ± 1 Ma by Rb-Sr (Jamieson, in press).

Upper Devonian to Lower Carboniferous volcanic and sedimentary rocks occur along the southern and western margins of the Highlands. In the south, rhyolites, felsic pyroclastics, siltstones, conglomerates, and minor intermediate flows occur at MacMillan Mountain (E_g). These have been dated at 384 ± 10 Ma (Jamieson, in press) and appear to be unconformably overlain by surf-weathering erosion and conglomerates of the Horton Group. The Farnon Brook Formation (E_g) in the western part of the study area includes basaltic, rhyolitic, and associated ash flow tuffs and metasediments interbedded with red siltstone and sandstone. Upper Devonian plant fossils of the genus *Archegonites* have been found on siltstones interbedded with volcanic breccias north of Farnon Brook (Crow et al., in press). The volcanic rocks are faulted against and conglomerates of the Horton Group that resemble siltstones within the volcanic sequence. Lower Carboniferous sediments of the Horton and Windsor Groups (E_g) bound the Highlands on all sides but have not been investigated in detail. A single basal flow occurs within the Horton Group on the North Boreas River (E_g). In general, contacts between the Carboniferous sediments and older crystalline rocks are obscured by steep faults.

Ductile shear zones are of major importance in the southern Highlands, particularly in the Chéticamp River and Middle River areas where they juxtapose low grade and high grade metamorphic rocks. Major shear zones cut intrusive and volcanic rocks along Muskrat Brook and Farnon Brook, shear shear zones give an Rb-Sr date of 394 ± 28 Ma which must be regarded as a maximum (Jamieson, in press). The sense of offset is sinistral but down-slip lineation suggests a component of thrusting. North-southward trending shear zones in the central part of the study area are of uncertain age but may represent an early stage of movement along the Agay - Northeast Margaree River fault zone. Steep faults occur throughout the study area; only the major ones are shown on the map. Along the upper reaches of the Northeast Margaree River major zones of oblate lateral extension of the Agay - Northeast Margaree River fault zone. These faults cut the Margaree gneiss with a central sense of offset; most of these faults also show evidence of vertical displacements. Thrust faults occur in the southern part of the area near Leonard's Mill Brook. They are generally not important, but may represent the latest phases of movement on the major ductile shear zones noted above.

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