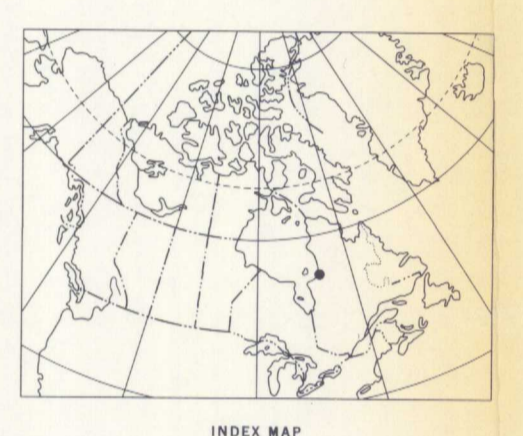




**Geological map of the southern margin of the Richmond Gulf Graben,
Lac Guillaume-Delisle, New Québec**

Scale 1:50,000 Échelle
Kilometres 1 2 3 4
Miles 0 1 2

Geology by F.W. Chandler, André Ciesielski and Robert St. Michel, 1981
Geological Survey of Canada
Base map assembled from parts of 33N/15, 33N/16, 34C/1, and 34C/2



LEGEND

ABHEBIAN NASTAPOKA GROUP	7 Basalt, 25-40 m
	6 Quartz arenite overlying stromatolitic carbonate, 170 m
Angular unconformity	
RICHMOND GULF GROUP	5 S ₁ mafic sill, 25 m cutting 2, and S ₂ mafic dikes cutting 1 and 2
	4 QINGALUK FORMATION, 200 m
	3 PERSILLON FORMATION, basalt, 0-40 m
	2 PACHI FORMATION, mainly pink arkose, 250 m
Nonconformity	
ARCHAIC	
	1 Granitic rocks, pink, homogeneous, foliated or massive

Outcrop, area of outcrop..... X

Geological boundary, defined, approximate, assumed.....

Bedding: horizontal, inclined.....

Syncline: with plunge.....

Fault: down throw marked, defined, approximate, assumed.....

Lead-Zinc deposit; location approximate..... Pb-Zn X

DESCRIPTION OF GEOLOGY

This file essentially completes the geological map of the Richmond Gulf Graben started with Geological Survey of Canada, open file 600 (Chandler 1979). The map covers the southern border zone of the graben which is marked by a set of broadly trending, often listric normal faults, downthrown to the north. On the up-thrown side of the faults the granitic basement (1) is, in many places, raised as rounded cliffs one to two hundred metres high. On the downthrown side of the faults the sedimentary, volcanic and included mafic intrusive rocks of the Richmond Gulf Group (2-5) dip towards the faults except where upturned to form synclines by drag at the fault surfaces. In the northern part of the map area many similar faults transect the Richmond Gulf Group, but the overlying Nastapoka Group is virtually unfaulked. Generally the Richmond Gulf and Nastapoka Group dip gently more or less westward and the Nastapoka Group overlies the more steeply dipping Richmond Gulf Group with gently angular unconformity. Where the lower group has been removed by erosion on the south margin of the graben the upper group lies directly on the granite basement. The above information, as well as absence of paleocurrent and facies changes in the Richmond Gulf Group attributable to the faulting, places graben formation between deposition of the two groups.

Although abundant faulting and insufficient outcrop of the Richmond Gulf Group prevent detailed interpretation of its structure, the low lying eastern part of the map area is seen to be underlain by the Pachi Formation (2) and the included mafic sill (5₁), higher ground to the west exposes the stratigraphically overlying volcanic Persillon Formation (3) and rebeds of the Qingaluk Formation (4a). Further west, pink arkose of the Pachi Formation (2) is preserved in lower parts of steep cliffs capped by the Nastapoka Group (6, 7). A similar succession is present in the southwest near the mouth of the Little Whale River. Microspites, apart from limited contact effects of the volcanics and sills, are not visible in the field.

The basement of the area consists almost entirely of relatively homogeneous granodiorite (1) with varied intensity of deformation and foliation striking at about 290°. Scattered metre-sized mafic and ultramafic inclusions could be fragments of pre-kinematic dikes. The contact with the nonconformably overlying Richmond Gulf Group (2-5) is, in many places, faulted. Where it is undisturbed there is green and white alteration of feldspars and no sign of an oxidized weathering zone. Hematite stain in some chert in the granite may be connected with local diagenetic oxidation of the overlying sandstone (2). In the extreme southwest of the area, where the granite is overlain by carbonates of the Nastapoka Group, it is locally pyritized by reducing sulphurous fluids descending from the carbonates.

The Pachi Formation (2), overlying the basement granite in the graben, consists of crossbedded arkose, usually pink but with minor amounts white and red. The arkose varies from fine sand to pebbly grit. A red sandstone with 5 cm quartz pebbles occurs locally at the base. Small amounts of red and light green mud-cracked mudstone are present at all levels in the formation.

The Persillon Formation (3), a sequence of terrestrial basalt flows, overlies the Pachi Formation with sharp contact. No fragments of the lower unit were seen in the basalt. The basalt is very rarely absent, and where it is red sandstone (unit 4a) contains basalt pebbles directly overlies unit 2 with no apparent sedimentary hiatus. At the base of the basalt there are pipe vesicles, and within it columnar jointing basalt pebbles directly overlies unit 2 with no apparent sedimentary hiatus. At the base of the basalt there are pipe vesicles, and within it columnar jointing basalt pebbles directly overlies unit 2 with no apparent sedimentary hiatus. Pea-sized vesicles and amygdalae, some green, others of Jasper, are also present. The top of the unit has been reddened by weathering.

The Qingaluk Formation (4) commences with one to several metres of cobble-boulder conglomerate (4a) derived from the conformably underlying basalt. The sandy to muddy matrix of the conglomerate is red, and locally boulders show advanced spheroidal weathering. The conglomerate is overlain by the remainder of unit 4a, a fine to medium grained crossbedded arkose interbedded with mud-cracked red mudstone and siltstone, the colour being derived from the underlying basalt. This unit is 80 m thick to the north (Chandler 1979). The bulk of the Qingaluk Formation consists of crossbedded pink arkose and pebbly grit. Pebbles, usually 2-3 cm across are of granite and quartz. Beneath unit 4a the arkose is bleached, rusty-weathering and sporadically pyritic. A local volcanic unit (4c) similar to unit 3 in appearance occurs in the coastal cuesta high in unit 4. It is lenticular in vertical section, up to 21 m thick and was traced 5 km along strike.

A diabase sill (5₁) at least 25 m thick occurs, except in the extreme southwest of the graben, within unit 2. Its stratigraphic position varies from within two metres of the basement to a hundred metres up into the formation. It can be separated in the field at a distance from unit 3 by the well developed columnar jointing that it displays in cliffs. Though smaller stromatolitic sills were noted to the north of the mapped area none were seen in the map area. Mafic dikes are extremely sparse in the area. One 15 cm dike, striking southwest cuts unit 2 and it is on strike with one observed 6.5 km to the northeast that cuts the basement.

Unit 5, present in the east-facing cliffs of the coastal cuestas is composed of about 150 m of peritidal stromatolitic carbonate overlain by 20-25 m of marine quartz arenite. In a several metre-thick rusty zone 2/3 of the way up the carbonate lies the Ruby Lake and Nancy Island deposits (see map) containing 525,000 and 96,000 tonnes respectively of 2' combined lead-zinc (Franklin and Thorpe in press).

The cuestas is capped by a basalt unit (7) that continues westward to the coast. At least four flows, 10 to 20 m thick occur. Rare interflow layers 1-2 m thick consist of breccia and fine-grained sediment. The flows are relatively homogeneous and columnar jointing is very common. Flow tops are brecciated and near them red banded agates and vesicles are concentrated. Zones several metres thick contain plagioclase crystals up to 10 cm in size. Pillows are rare.

References

Chandler, F.W.
1979: Geology of the Abhebian Precambrian rocks, Lac Guillaume-Delisle, Québec: Geological Survey of Canada, Open File 600. See bedrock geological maps, scale 1:50,000.

Franklin, J.M. and Thorpe, R.I.
In Press: Comparative metallurgy of the Superior, Slave and Churchill Provinces.

19: Precambrian Ore Deposits, Geological Association of Canada, Special Volume.