

DESCRIPTIVE NOTES

Map-area 46 O/1 is underlain by rocks of the basement complex in its southern one-third and by the Penrhyn Group overlying the complex elsewhere. The complex consists primarily of granodioritic and other granitic igneous rocks (Aggdn) with high infrequently interbedded intrusive rocks. Small lensoidal bodies of amphibolite (Am) some lying near the upper surface of the complex, are common. Along the southern border an irregular belt of paragneiss (An) and amphibolite (Am) with much foliated and massive igneous rock forms a migmatitic terrain of uncertain affinity. Lack of carbonate units suggests that it is not part of the Penrhyn Group.

The basal sequence of orthoquartzite (Ano) with rare feldspathic grit and sillimanite, cordierite and garnet-bearing paragneiss (Ano) of the Penrhyn Group overlies the basement complex in a structurally complex belt that crosses the central part of the area. Rocks near the basement-cover contact usually possess highly sheared and cataclastic textures suggesting some differential movement. Mobilization and resulting interfingering of gneiss, granitic rock and orthoquartzite in a narrow (less than 10 km) zone is common. The basal sequence is followed by a marble unit (Am) that thickens markedly to the north and by paragneiss (An) which is the uppermost unit seen in this region. The basal sequence has been repeatedly brought to the surface by folding in the northern part of the area. There it includes grey-green diopside-actinolite-calcite agglomeritic rock possibly a meta-regolith and garnet-sillimanite schist. The overlying marble unit is thick and generally massive, but in its northernmost outcrops, toward the base of basement gneiss near the north border, apparently passes through a facies change into an interbedded succession of marble, paragneiss and calcium-silicate gneiss. Cross-section B-B' accompanying map-sheet 46 O/2 illustrates these variations.

Commonly leucocratic intrusive rocks of granodioritic, quartz monzonitic and granitic composition bearing biotite and, in places, hornblende are plentiful. In this area they occur only in small bodies. Many are found along the complex southern margin of the Penrhyn Group where they obscure structural and stratigraphic features. Significant amounts of intrusive rock were intruded along the basement-cover contact along the southern border of the large elongate domal structure in the north part of the area. Possibly the presence of water in the sediments facilitated melting near the unconformity or else that plane favoured passage of the plutons.

Early deformation of the basement complex and the Penrhyn Group (D1) produced pervasive foliation nearly parallel to compositional layering and is believed responsible for attenuated recumbent folds and gneissic sheets found along the southern margin of the Penrhyn Group. Shallow dip of most structures and low relief combine to complicate the outcrop pattern and make understanding of the structures difficult. The major fold shown on the cross-section may be an extension of the folded gneissic sheet southwest of Quartzite Lake in map-area 46 P/3. Structures associated with D2 are found only in western parts of the area and are best seen on cross-section B-B' (map-sheet 46 O/2). The small recumbent folds found west of the major river in the northwestern part of the area may be related to either D2 or D3. Relationships between early foliation (D1) and these structures were not observed. Large upright and open folds trending north-northeasterly are assigned to D4 and clearly affect earlier structures. Variations in plunge of these folds to form elongate domal bodies of basement gneiss mantled by the Penrhyn Group are attributed to the effects of D2. Long, north-northeasterly trending faults with apparent left-lateral displacements of about 1/2 km formed during D4.

The third and fourth phases (D3 and D4) produced prominent meso- and megascopic folds that impose an east-northeast structural grain in the Foxe Fold Belt. D3 folds are tight to nearly isoclinal and usually recumbent. Axial plane foliation (S3) is nearly parallel to lineation (S4) and hence to S2 rendering separation of phases D2 and D3 very difficult. D4 folds are coaxial, or nearly so, with D3 but are more open and generally upright or slightly overturned. Mesoscopic D4 folds can often be observed to have deformed earlier structures. North to northeasterly trending broad transverse flexures (D5) alter the plunges of pre-existing folds. A few mesoscopic structures associated with this phase were observed. Steeply dipping fractures and faults, many with northerly and northeasterly trends are evidence of the last phase of deformation (D6). Most fault displacements appear to be left-lateral and east-side-up. Metamorphism is believed to have accompanied all phases of deformation except D3 and D6. It possibly reached its zenith during D3 but mineral recrystallization outlasted much of the penetrative deformation.

Massive and foliated plutonic rocks (Ag), chiefly of hornblende and biotite granodiorite, quartz monzonite and granite intrude the basement complex and the Penrhyn Group. Resembling granitic rocks of the complex, separation of these is based largely on field relationships. Foliated plutonic rocks, except where observed to have intruded the Penrhyn Group, are assigned to the complex (Ag). Where intrusive into the group they (Ag) are considered to be pre- or syntectonic with the main phases of deformation. In some localities intimate mixing with and partial melting of paragneiss was greater than elsewhere (Am). Although mullion and fold axes associated with folds observed to have deformed bedding or early foliation. Lineation (plunging, horizontal), formed by bedding-foliation intersection, mineral growth, rodding and mullion/earliest or only observed. Lineation (plunging, horizontal), formed by bedding-foliation and foliation intersection, mineral growth, rodding and mullion and fold axes associated with folds observed to have deformed bedding or early foliation. Lineation (plunging, horizontal), formed by cleavage-foliation intersection and fold axes associated with great folds believed to have formed late in the tectonic history.

Available results of radiometric analyses indicate formation of the basement complex prior to 2500 Ma with some events occurring as long as 3000 Ma ago (G.K. Hales, personal communication, 1979). Deformation of the basement and the Penrhyn Group may have taken place 2134 Ma ago (Jackson and Taylor, 1972) and again during the Hudsonian Orogeny (c.1820 Ma ago). Post-tectonic plutons (1600 Ma old; Heywood, 1966) were emplaced into the fold belt late in the orogenic history. Following extensive uplift and erosion, diabase dykes (1841) presumed to be part of the Mackenzie dyke swarm of c.1800 Ma age (Fahrig, 1970) cut rocks of the fold belt. These are spatially associated with faults and fractures formed during D6 tectonism. Subsequent uplift and erosion was followed by deposition of Proterozoic carbonate rocks, remnants of which lie north and south of the fold belt.

Geological mapping by: A.V. Okulitch. Cross-sections by: A.V. Okulitch. Descriptive notes by: A.V. Okulitch. Drafting by: A.V. Okulitch. Notice of any revisions or additional geological information known to users of these maps would be gratefully received by the authors.

Fahrig, W.F., 1970. Diabase Dyke Swarms: Geology and Economic Minerals of Canada. Geological Survey of Canada, Economic Geology Report Number 01, pp. 131 - 139. Heywood, W.H., 1967. Geological Notes: Northwestern District of Keewatin and Southern Melville Peninsula, District of Franklin, Northwest Territories (Parts of 46, 47, 56, 57). Geological Survey of Canada, Paper 66-40. Jackson, G.B. and Taylor, F.C., 1972. Correlation of Major Archaean Rock Units in the Northern Canadian Shield. Canadian Journal of Earth Sciences, Volume 9, pp. 1650 - 1660.

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GENERAL GEOLOGY

The Foxe Fold Belt extends in an east-northeast direction from southern Melville Peninsula to central Baffin Island. It is composed of granitoid gneissic rocks of Archean age (2500 Ma and older) overlain by meta-sedimentary rocks of Early Proterozoic age (approximately 2500 to 1700 Ma) of the Penrhyn and Filling Groups. These have undergone polyphase deformation and metamorphism mostly during the Hudsonian Orogeny. Generation and emplacement of plutonic rocks occurred, accompanied and followed deformation. Diabase dykes of presumed Late Proterozoic age cut older rocks.

The Archean rocks form a basement complex predominantly of granitoid gneiss (Aggdn) and foliated granitic rocks (Ag) with relatively minor amphibolite (Am) and paragneiss (An) and other meta-sedimentary rocks (An-Am). The gneissic and plutonic rocks are largely of quartz monzonitic to granodioritic composition; leucocratic and mafic varieties of gneiss are also common but do not constitute a large volume of the complex. Gneissic layering and mineral foliation formed by D1 and D2 are generally massive but not always clearly visible. Plutonic rocks emplaced during at least three episodes of igneous activity can be differentiated locally but cannot easily be mapped regionally because they are compositionally similar to one another and to the gneiss which appear to be both host and parent to the plutonic rocks. The Penrhyn Group have in some places been assigned to the basement complex but their affinity is uncertain given the evident intricacy of stratigraphy and structure. Some lithologic similarity to rocks of the Prince Albert Group in northern Melville Peninsula exists but such correlation is tenuous at best.

The Penrhyn Group consists of paragneiss (An-Am) and marble (Am) with some quartz-biotite schist (An-Amb) and calcic amphibole gneiss (An) and minor quartzite (Anq) and gneiss (An). Biotite and sillimanite schists (Am) and meta-volcanic rocks (Am-Amb-Ar) are also present. Complete understanding of the stratigraphic succession is lacking as most units are discontinuous and lensoidal and the possibility of the existence of facies changes, unconformities and cryptic early structures renders its delineation difficult. A general order to units can be indicated, nonetheless. A thin (50-100 m) basal sequence includes orthoquartzite, rusty sillimanite schist, a suspected meta-regolith and minor amphibolite, marble and quartz-feldspathic grit. This sequence is overlain by a predominantly calcareous unit of marble, calc-silicate gneiss and interbedded paragneiss. The calcareous unit is followed by a thick unit of paragneissic rocks with a thin bed of schistose paragneiss at its base, and a unit of marble, calc-silicate gneiss and biotite quartzite. At the highest observed structural and stratigraphic levels, orthoquartzite, quartz-feldspar, white mica and phlogopite, quartz-biotite schist, amphibole gneiss, a marble group, mineral and tremolite, interbedded with and passing laterally into unit Am. Includes small beds of unit Am. Foliation of unit Ag is nearly parallel to lineation.

The Penrhyn Group appears to lie unconformably on the basement complex. Tectonism has obliterated any angular discordance and unconformable relationships because of the clear lithologic contrast and the common presence of the thin orthoquartzite unit with rare feldspathic grit beds lying upon a variety of rock types in the complex. The uppermost unit of the Penrhyn Group may be separated from the rest of the group by an unconformity. Contrasts in intensity of metamorphism and undeformed structural discordance support such an interpretation but rapid transitions or faulting remain viable alternate explanations.

Metamorphism of the Penrhyn Group produced two lithologic suites. Most of the group is in uppermost amphibolite facies and contains the assemblages garnet-biotite-sillimanite and cordierite-sillimanite-garnet in paragneiss and in marble, diopside-forsterite-calcite as well as scapolite and a white mica group mineral. Rocks of the uppermost unit of the group are in greenschist facies and contain chlorite-muscovite-quartz in pelitic units. Porphyroblasts of a mineral tentatively identified as andalucite are common in meta-sillimanite units and tremolite occurs sporadically in calcareous rocks.

Polyphase structures indicating numerous episodes of deformation of the basement complex and the Penrhyn Group exist throughout the fold belt but unequivocal sequential relationships among them are rare. The earliest deformational phase (D1) is inferred to have affected the basement complex prior to deposition of the Penrhyn Group. Little is known of this phase. The second phase (D2), the earliest, coarse grained, is believed to have formed attenuated isoclinal folds and ubiquitous foliation (D3). In all but a few outcrops S2 is parallel to, and nearly so, with D3 but are more open and generally upright or slightly overturned. The effects of D2 on the Penrhyn Group remain problematical, but may be responsible for some of the observed discontinuity of units described above.

The third and fourth phases (D3 and D4) produced prominent meso- and megascopic folds that impose an east-northeast structural grain in the Foxe Fold Belt. D3 folds are tight to nearly isoclinal and usually recumbent. Axial plane foliation (S3) is nearly parallel to lineation (S4) and hence to S2 rendering separation of phases D2 and D3 very difficult. D4 folds are coaxial, or nearly so, with D3 but are more open and generally upright or slightly overturned. Mesoscopic D4 folds can often be observed to have deformed earlier structures.

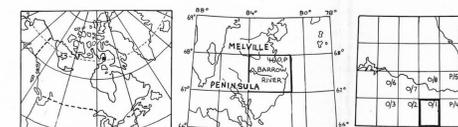
North to northeasterly trending broad transverse flexures (D5) alter the plunges of pre-existing folds. A few mesoscopic structures associated with this phase were observed. Steeply dipping fractures and faults, many with northerly and northeasterly trends are evidence of the last phase of deformation (D6). Most fault displacements appear to be left-lateral and east-side-up. Metamorphism is believed to have accompanied all phases of deformation except D3 and D6. It possibly reached its zenith during D3 but mineral recrystallization outlasted much of the penetrative deformation.

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LEGEND

- LATE (?) PROTEROZOIC HUDSONIAN. Hd: Brown weathering, dark green to black, fine to medium grained pyroxene diabase. Intrusive Contact: Orange and buff weathering, white, tan and grey, massive and foliated, medium to coarse grained, biotite and hornblende granodiorite, quartz monzonite, quartzite and leucocratic equivalents. Microgabbro composed of units An1 and An2 in lit-par-lit, xenolithic and texturally transitional variations. Intrusive Contact: White and light green weathering, light grey, massive, layered amphibatic siliceous rock (acid volcanic rock).

- EARLY PROTEROZOIC ARCHAEN PENRHYN GROUP. Arva: Green to dark green, green actinolitic greenschist (basic to intermediate). Arvb: Green, fine to medium grained, thin to thick bedded, quartz-muscovite-feldspar paragneiss, some with andalucite (?). Arvc: Black, fissile, very fine grained, "sooty" pelite. Arvd: Grey, fine to medium grained, thin to thick bedded, quartz-biotite-feldspar paragneiss and meta-igneous rocks (gabbro and quartz-cordierite-granulite). Arve: Brown, rusty and tan weathering, buff and grey, fine to medium grained, quartz-biotite-feldspar, quartz-feldspar-biotite-garnet-sillimanite and quartz-feldspar-biotite-garnet-paragneiss and minor schist. Arvf: Grey, fine to medium grained, medium to coarse grained thin bedded, calc-silicate gneiss and marble-quartzite with quartz-cordierite and accessory schistose paragneiss at its base, and a unit of marble, calc-silicate gneiss and biotite quartzite.

- POSSIBLE UNCONFORMITY. Arn: Brown, rusty and tan weathering, buff and grey, fine to medium grained, quartz-biotite-feldspar, quartz-feldspar-biotite-garnet-sillimanite and quartz-feldspar-biotite-garnet-paragneiss and minor schist. Arnc: Grey and grey-green, medium to coarse grained thin bedded, calc-silicate gneiss and marble-quartzite with quartz-cordierite and accessory schistose paragneiss at its base, and a unit of marble, calc-silicate gneiss and biotite quartzite. Arnd: White, grey to grey-blue, medium to coarse grained, massive and bedded marble with calcite-diopside-microcline-quartz and minor dolomite. Arne: White to grey-blue, medium to coarse grained, massive and faintly bedded, orthoquartzite with minor feldspar, white mica and phlogopite. Arnf: Quartz-biotite schist, amphibole gneiss, a marble group (meta-regolith ?), biotite-garnet-sillimanite schist and amphibolite. Arng: Grey and grey-green, medium to coarse grained, layered and foliated, biotite and hornblende granodioritic, quartz monzonitic and leucocratic gneiss. Includes rocks of units Ar1 and Ar2.

- ARCHAEN. Ag: Orange and buff weathering, grey and pink, medium to coarse grained gabbro and quartz-cordierite, diorite and syenite. Includes rocks of units Ag1 and Ag2. Aggdn: Orange, grey and tan, medium to coarse grained, layered and foliated, biotite and hornblende granodioritic, quartz monzonitic and leucocratic gneiss. Am: Dark green, fine to medium grained, massive and foliated amphibolite. Amb: Dark green, coarse grained, serpenitized pyroxene-bearing ultramafic rock. An: Quartz-biotite-feldspar paragneiss, some with hornblende/migmatite with unit Am. An1: Rusty weathering, schistose biotite paragneiss. An2: White, medium to coarse grained, massive orthoquartzite.

- PROBABLE UNCONFORMITY. Geological boundary (defined, approximate) boundary of areas extensively drift-covered. Planar Structures: Bedding and compositional layering (horizontal, inclined, vertical). FOLiation, schistosity, gneissic layering, cleavage and axial planes (inclined, vertical); dip unknown, earliest or only observed. FOLiation, cleavage and axial planes (inclined, vertical); associated with folds of later phases observed to have deformed bedding or early foliation. FOLiation, cleavage and axial planes (inclined, vertical); associated with folds observed to have deformed bedding, early foliation and/or pre-existing structures. Cleavage and axial planes (inclined, vertical); associated with gentle folds observed to have deformed earlier structures and believed to have formed late in the tectonic history. Structural form line (on cross-sections). LINEATION (plunging, horizontal); formed by bedding-foliation intersection, mineral growth, rodding and mullion/earliest or only observed. LINEATION (plunging, horizontal); formed by bedding-foliation and foliation intersection, mineral growth, rodding and mullion and fold axes associated with folds observed to have deformed bedding or early foliation. LINEATION (plunging, horizontal); formed by cleavage-foliation intersection and fold axes associated with great folds believed to have formed late in the tectonic history. High angle fault (defined, approximate); arrows indicate apparent relative movement. Low angle fault (defined, approximate); small triangles in direction of dip.

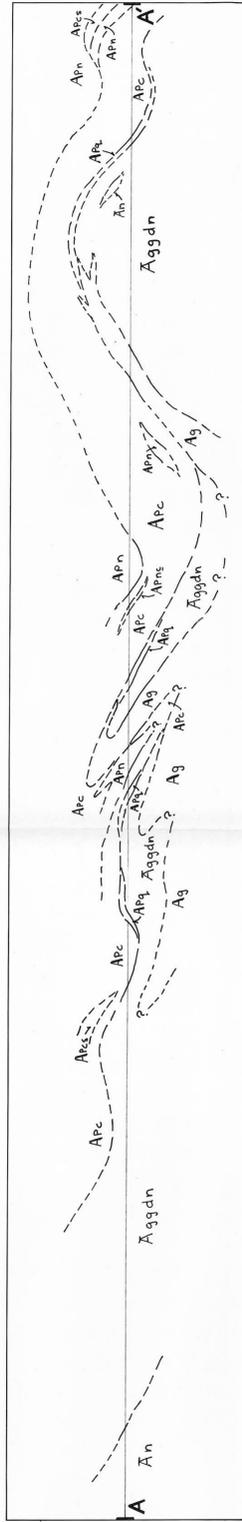
- REFERENCES: A high degree of uncertainty or interpretation in the position or the nature of the symbol used is indicated. Position of ends of cross-section. METAMORPHIC MINERALS: A ANDALUCITE, C CORDIERITE, G GARNET, P PHLOGOPITE, Sc SCAPOLITE, T TREMOLITE, Ac ACTINOLITE, Ch CHLORITE, M MUSCOVITE, S SILLIMANITE, St STAUROLITE. STRUCTURAL NOTE: Description of structures is facilitated by separation into six phases using criteria such as fold style and orientation and sequential relationships among folds. FOLiation and LINEATION. Such separation does not imply presence of discrete events, particularly in the case of phases D2, D3, and D4 which may well be partly or wholly synchronous in some areas. No bounds on the time spans represented by the phases are implied as only the broadest limitations can be placed on the beginnings and ends of orogenic events. Large folds on the map are given numbers corresponding to the phase postulated to be responsible for them. Mesoscopic structures are assigned a position in the tectonic hierarchy based on interpretation of local field relationships within the unit within which they were observed. This position cannot be directly related to deformational phases which formed the large folds. For these preliminary maps, no attempt has been made to integrate all mesoscopic features into a megascopic structural synthesis. CROSS-SECTIONS portray the inferred form of structures and show apparent dips (in the section) of lithologic contacts and foliation. Vertical proportions are not to scale. The horizontal reference line represents an approximate mean elevation along the line of section and is usually within 200 m. of sea level. Structures appearing on cross-sections are highly interpretative. Some features, particularly faults, will often not appear on the map as they were not observed in the field.

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- PRODUCTION BY SURVEY AND MAPPING BRANCH, DEPARTMENT OF ENERGY, MINES AND RESOURCES, OTTAWA, CANADA. ELEVATIONS IN METRES ABOVE MEAN SEA LEVEL CONTOUR INTERVAL: 10 METRES. DISTRICT OF FRANKLIN NORTHWEST TERRITORIES. Scale 1:50,000 Echelle. SYSTEME DE REFERENCE GEODESIQUE NORD-AMERICAIN 1983 PROJECTION TRANSVERSE DE MERIDIEN.

- ETABLI PAR LA DIRECTION DES LEVÉS ET DE LA CARTOGRAPHIE, MINISTÈRE DE L'ÉNERGIE, DES MINES ET DES RESSOURCES, OTTAWA, QUÉBEC. ÉLEVATIONS EN MÈTRES AU-DESSUS DU NIVEAU MOYEN DE LA MER ÉQUIDISTANCE DES COURBES: 10 MÈTRES. DISTRICT DE FRANKLIN TERRITOIRES DU NORD-OUEST. Échelle 1:50 000. SYSTÈME DE RÉFÉRENCE GÉODÉSIQUE NORD-AMÉRICAIN 1983 PROJECTION TRANSVERSE DE MÉRIDIEN.

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LEGEND - LEGENDE. ROADS AND RELATED FEATURES. ROUTES ET OUVRAGES CONNEXES. LANDMARK FEATURES. POINTS DE REPÈRE. BOUNDARIES AND SURVEY CONTROL. FRONTIÈRES ET POINTS DE RÉFÉRENCES. PHOTOGRAPHY. PHOTOGRAPHIE. REVISION. REVISION. ONE THOUSAND METRE UNIVERSAL TRANSVERSE MERCATOR GRID. QUADRILLE DE MILLE MÈTRES UNIVERSAL TRANSVERSE DE MERIDIEN. Conversion scale for elevations. Échelle de conversion des élévations. METERS 0 100 200 300 400 500 600 700 800 900 1000. Feet 0 100 200 300 400 500 600 700 800 900 1000.