GENERAL GEOLOGY

The Foxe and Committee Fold Belts extend in an east-northeast direction from southern Melville Peninsula to central Baffin Island. They are composed of granitoid gneissic rocks engulfing metamorphosed sediments and volcanic rocks all of Archaean age, overlain by metasedimentary rocks of Early Proterozoic age of the Penrhyn and Piling Groups. These fold belts suffered polyphase deformation and metamorphism mostly during the Hudsonian Orogeny. Generation and emplacement of plutonic rocks preceded, accompanied and followed deformation. Diabase dykes of presumed Late Proterozoic age cut older rocks.

The Archaean rocks form a basement complex predominantly of granitoid gneiss (Āggdn), layered quartzo-feldspathic gneiss (Āgn) and foliated grantic rocks (Āg$_1$, Āg$_2$ and Āg$_3$), within which are relatively minor amounts of amphibolite (Ām) and metasedimentary and metavolcanic rocks of the Prince Albert Group. The gneissic and plutonic rocks are largely quartz monzonitic to granodioritic in composition; leucocratic and mafic (Āg$_3$) varieties are common but do not constitute a large volume of the complex. Gneissic layering and mineral foliation formed of biotite and hornblende are ubiquitous but not always clearly visible. Metavolcanic (ĀAm$_1$, ĀAb, ĀAva, ĀAub) and metasedimentary (ĀAn, ĀAnm, ĀAif, ĀAq) rocks occuring as discontinuous zones and lenses within the basement complex of the Foxe Fold Belt are correlated with the Prince Albert Group of the northwesterly adjacent Committee Fold Belt on the basis of considerable lithologic similarity. The stratigraphy of that part of the group within the Foxe Fold Belt is unknown but it has been described by Campbell (1974) and Schau (1975a) within the Committee Fold Belt. A lensoid sill of anorthositic gabbro (ĀAb) and layered amphibolite (ĀAm$_1$), as well as possibly related amphibolitic dykes (ĀAm$_2$) may intrude or be genetically related to the group.

Amongst the gneissic rocks of the complex are presumed to be some that form the basement to the Prince Albert Group but unconformable relations, if present are masked by deformation and plutonic activity. Some gneissic units, particularly parts of unit Āgn, may be derived from the Prince Albert Group or some still older metasedimentary succession by migmatitic processes. Granitoid gneiss units of probable plutonic origin (Āg$_1$, Āg$_2$, Āg$_3$) may be older and younger than the Prince Albert Group. In the Committee Fold Belt, porphyritic granite resembling that of unit Āg$_2$ has intruded the group (Schau, 1975a). Elsewhere, age relations are commonly equivocal. More detailed studies of the basement complex have been made by Frisch (1974, 1975) and Schau (1975a, 1975b).

The Penrhyn Group consists of paragneiss (APn, APnc) and marble (APc) with some quartz-mica psammite (APqb, APqm) and calcium-silicate gneiss (APcs). Minor orthoquartzite (APq), amphibolite (APm), pelite (APp) and very minor iron formation (APif) are also present. Complete understanding of the stratigraphic succession is lacking as most units are discontinuous and lensoid and the possibility of the existence of facies changes, unconformities and cryptic early structures renders its delineation difficult. A general order to the units can be indicated nonetheless. A thin (50-100 m) basal sequence includes orthoquartzite, rusty sillimanite schist, a suspected metaregolith and minor amphibolite, dolomitic marble quartzo-feldspathic grit, rusty pyrite-magnetite iron formation and conglomerate with quartz and hematite clasts. This sequence is overlain by a predominantly calcareous unit of marble, calcium-silicate gneiss and some paragneissic interbeds. The calcareous unit is followed by a thick unit of paragneissic rocks and a unit of marble, calcium-silicate gneiss and biotite quartzite. At the highest observed structural and stratigraphic levels is a unit of quartz-biotite and/or muscovite psammite and metagreywacke. This unit is variable in gross lithology and variously interbedded and compositionally gradational with paragneiss (APnqb), calcium-silicate gneiss and minor marble. The top of the Penrhyn Group has not been observed. The relationship between present and original thickness of the group is well disguised by the rival processes of thinning during deformation, repetition by folding and dilation by syntectonic plutonism.

The Penrhyn Group appears to lie unconformably on the basement complex. Tectonism has obliterated any angular discordance and unconformable relationships are inferred because of the clear lithologic contrast and the common presence of the thin orthoquartzite unit with rare feldspathic grit and hematite-clast conglomerate beds lying upon a variety of rock types in the complex.

Metamorphism in the Foxe Fold Belt produced the assemblages garnet-biotite-sillimanite and cordierite-sillimanite-garnet in paragneiss and, in marble, diopside-forsterite-calcite as well as scapolite and a humite group mineral. In pelitic and semi-pelitic rocks, the reactions: muscovite + quartz ⟶ sillimanite + K feldspar
andalucite ⟶ sillimanite
can be delineated in restricted areas. Retrograde, or later low grade metamorphism is probable because of extensive alteration of high grade minerals.

Polyphase structures indicating numerous episodes of deformation of the basement complex, the Prince Albert Group and the Penrhyn Group exist throughout the two fold belts but unequivocal sequential relationships among them are rare. The earliest deformational phase is inferred to have affected the basement complex prior to deposition of the Penrhyn Group. Tectonic trends in the basement complex and the Prince Albert Group within the Foxe Fold Belt are for the most part conformable with those of the overlying Penrhyn Group and pre-Penrhyn structures are not readily distinguishable. Structures in the Committee Fold Belt have been described by Campbell (1973, 1974), Frisch (1974), Reesor, et al. (1975) and Schau (1973, 1974, 1975a, 1975b). Some deformation of the Prince Albert Group may have also preceded deposition of the Penrhyn Group.

A second phase of folding, the earliest observed in the Penrhyn Group, is believed to have formed attenuated isoclinal folds and ubiquitous foliation. In all but a few outcrops this foliation is parallel to bedding. Meagre evidence suggests that the trend of early Penrhyn structures may have been northerly. The effects of this folding episode on the Penrhyn Group remain problematical, but may have resulted in some of the observed discontinuity of units described above.

Later episodes of folding produced prominent meso- and megascopic folds that impose an east-northeast structural grain on the Foxe Fold Belt. Tight to nearly isoclinal recumbent structures are folded by later nearly coaxial, more open, upright to overturned folds. These later folds can often be observed to have deformed earlier structures.

In numerous places gneissic bodies of the basement complex can be seen to lie on and possibly within the Penrhyn Group. Such relationships suggest either the presence of large allochthonous nappes or smaller scale, locally overturned folds and thrust faults. The time of movement of the basement masses is uncertain but as they are folded about northeasterly trending axes, they are presumed to have been emplaced during the early deformation of the Penrhyn Group.

North to northeasterly trending broad transverse flexures alter the plunges of pre-existing folds. Few mesoscopic structures associated with this phase were observed. It may be related to syn- and post-tectonic plutonic intrusion. Steeply dipping fractures and faults, many with northerly and northwesterly trends are evidence of the last major phase of deformation. Most fault displacements appear to be left lateral and east-side-up. Minor evidence of east-west faulting that may have affected Late Proterozoic diabase dykes (Hld) has been observed.

Metamorphism is believed to have accompanied all phases of deformation up to the late northeasterly trending open folding. It possibly reached its zenith during the preceding northeasterly-trending isoclinal phase, but mineral recrystallization outlasted much of the penetrative deformation. Retrogressive metamorphism may have accompanied latest folding episodes or been post-tectonic. Contact metamorphic aureoles are likely present around post-tectonic granitic plutons (Ag).

Massive and foliated plutonic rocks (Āg$_1$, Āg$_2$, Āg$_3$, Ag), chiefly of hornblende and biotite granodiorite, quartz monzonite and granite intrude the basement complex and the Penrhyn Group. Foliated plutonic rocks, except where observed to have intruded the Penrhyn Group are assigned to the complex. Where intrusive into the group (Ag) they are considered to be pre- or syn-tectonic to the main phases of deformation. Common local generation of pegmatite and leucocratic granitic rock (Ag) is believed to be coeval with deformation also. Massive, often cross-cutting plutons (Ag) invaded the Penrhyn Group after cessation of deformation.

Available results of radiometric analyses indicate formation of the basement complex prior to 2500 Ma ago, with some events occurring possibly as long as 2900 Ma ago (R.K. Wanless, personal communication, 1976). Acid volcanic rocks of the Prince Albert Group on the west side of Melville Peninsula have yielded a preliminary date of about 2700 Ma (R.K. Wanless, personal communication, 1977). Deformation of the basement complex and the Penrhyn Group may have taken place 2134 Ma ago (Jackson and Taylor, 1972) and again during the Hudsonian Orogeny (circa 1700 Ma ago). Post-tectonic plutons (1600 Ma old, Heywood, 1967) were emplaced into the fold belt late in the orogenic history. Following extensive uplift and erosion, diabase dykes (Hld) presumed to be part of the Mackenzie dyke swarm of about 1000 Ma age (Fahrig, 1970), cut rocks of the fold belt. These are spacially associated with faults and fractures trending northwest. Subsequent uplift and erosion was followed by deposition of Silurian and Ordovician carbonate rocks (OSc), remnants of which lie north and south of the Committee and Foxe Fold Belts and bordering Foxe Basin.

REFERENCES

Campbell, F.H.A. 1973. Sedimentary Rocks of the Prince Albert Group, District of Keewatin. Geological Survey of Canada, Paper 73-1A, pp. 141-142.

——— 1974. Paragneisses of the Prince Albert Group. Geological Survey of Canada, Paper 74-1A, pp. 159-160.

Fahrig, W.F. 1970. Diabase Dyke Swarms, in Geology and Economic Minerals of Canada, Geological Survey of Canada, Economic Geology Report Number One, pp. 131-134.

Frisch, T. 1974. Geological Studies in the Prince Albert Hills, Western Melville Peninsula, District of Franklin. Geological Survey of Canada, Paper 74-1A, pp. 163-164.

——— 1975. Geological Studies in Western Melville Peninsula, District of Franklin. Geological Survey of Canada, Paper 75-1A, pp. 323-324.

Heywood, W.W. 1967. Geological Notes, Northeastern District of Keewatin and Southern Melville Peninsula, District of Franklin, Northwest Territories (Parts of 46, 47, 56, 57). Geological Survey of Canada, Paper 66-40.

Jackson, G.D. and Taylor, F.C. 1972. Correlation of Major Aphebian Rock Units in the Northern Canadian Shield. Canadian Journal of Earth Sciences, vol. 9, pp. 1650-1669.

Okulitch, A.V., Gordon, T.M., Henderson, J.R., Reesor, J.E., Hutcheon, I.E. and Turay, M. 1977a. 1 50000 Geological Maps. Geological Survey of Canada, Open Files 433-443.

Okulitch, A.V., Gordon, T.M., Henderson, J.R., Reesor, J.E. and Hutcheon, I.E. 1977b. Geology of the Barrow River Map-area, Melville Peninsula, District of Franklin. Geological Survey of Canada, Paper 77-1A, pp. 213-215.

Okulitch, A.V., Gordon, T.M., Henderson, J.R., Hutcheon, I.E. and Turay, M. 1978. Geology of the Barrow River and Hall Lake Map-areas, Melville Peninsula, District of Franklin. Geological Survey of Canada, Paper 78-1A, in press.

Reesor, J.E., LeCheminant, A.N., Henderson, J.R. 1975. Geology of the Penrhyn Group Metamorphic Complex, Melville Peninsula, District of Franklin. Geological Survey of Canada, Paper 75-1A, pp. 349-351.

Schau, M. 1973. Volcanic Rocks of the Prince Albert Group. Geological Survey of Canada, Paper 73-1A, pp. 175-177.

——— 1974. Volcanic Rocks of the Prince Albert Group, District of Keewatin. Geological Survey of Canada, Paper 74-1A, pp.187-188.

——— 1975a. Volcanogenic Rocks of the Prince Albert Group, Melville Peninsula (47A-D), District of Franklin. Geological Survey of Canada, Paper 75-1A, pp. 359-361.

——— 1975b. Gneiss Distinctions in the Hayes River Region, Magnetic and Geochemical Parameters. Geological Survey of Canada, Paper 75-1B, pp. 89-96.

DESCRIPTIVE NOTES

Map-area 46 O/14 is underlain by granitoid gneiss of the basement complex in its northwestern half and the Penrhyn Group containing small bodies of gneiss in the southeastern part. Foliated augen granite (Āg$_2$) and layered gneiss (Āgn) form most of the basement, scattered outcrops of foliated granodiorite (Āg$_1$) are present also. Unit Āg$_1$ is believed to be the oldest, at least in part and appears to have been intruded and partly reconstituted during formation of unit Āg$_2$. Foliation and gneissosity predominantly dip moderately to the northwest.

The Penrhyn Group structurally underlies layered gneiss across the central part of the area and also in the southeast corner, thus being exposed in a complex antiformal structure. The contact with basement gneiss is poorly exposed and commonly obscured by pegmatite (Ag). The basal sequence of orthoquartzite and rusty paragneiss was not observed. Marble (APc), paragneiss (APn) and calcium silicate gneiss (APcs) apparently lie against gneiss in different localities, but marble forms the predominant basal lithology. Proceeding inwards to the core of the major antiform, an interbedded unit of paragneiss and calcium silicate gneiss is encountered, which passes gradationally into paragneiss with marble and rusty paragneiss interbeds. Marble surrounding elongate bodies of gneiss crops out at the core of the antiform and delineates a tight upright fold possibly related to the structure in the southeast corner of map-area 46 O/13.

Leucocratic granite and pegmatite (Ag) intrude the Penrhyn Group and the basement complex and form numerous massive and sometimes foliated bodies, the largest of which lies in the southeast part of the area.

Prominent structures in the area are southwest plunging meso- and megascopic folds and associated linear features such as mullions and fold axes. These structures formed during at least two phases of deformation about nearly parallel axes. Structural trends in the basement complex are for the most part parallel to these folds. Earlier features in the complex are presumed to have been rotated into parallelism with these later structures. Occurence of gneissic bodies within the Penrhyn Group is explained by their exposure in polyphase antiformal fold cores that vary in plunge, and by tectonic interdigitation, either by thrust faulting or isoclinal folding. The stratigraphic succession described above appears to be, on structural grounds, at least partly overturned, and may have been tightly folded upon itself prior to northeast trending upright folding responsible for the major antiform. Basement gneiss may lie above Penrhyn rocks on the limbs of small overturned folds or at the base of a very large nappe, which is exposed in the southeast corner of this area but mainly in map-areas 46 O/9, 10, 15 and 16.

Two diabase dykes (Hld), part of a poorly exposed set also found in map-areas 46 O/13 and 11, cut across the Penrhyn Group in the southern half of the area.

LEGEND

PHANEROZOIC
PALAEOZOIC
ORDOVICIAN AND SILURIAN

OSc	Buff and light grey dolomite

UNCONFORMITY

LATE(?) PROTEROZOIC
HADRYNIAN

Hld	Brown weathering, dark green to black, fine to medium grained pyroxene diabase.

INTRUSIVE CONTACT

EARLY PROTEROZOIC
APHEBIAN AND YOUNGER(?)

Ag	Orange and buff weathering, white, tan and grey, massive and foliated, medium to coarse grained, biotite and hornblende granodiorite, quartz monzonite, granite and leucocratic equivalents. Some porphyritic varieties. Pegmatite commonly containing quartz, feldspar, muscovite and biotite; rarely tourmaline. Pre-and post-tectonic plutons are not differentiated. Includes zenoliths of older units.

INTRUSIVE CONTACT

APHEBIAN
PENRHYN GROUP

APqm	Grey, fine to medium grained, thin to thick bedded, quartz-muscovite-feldspar psammite, some with andalucite(?) porphyroblasts; muscovite schist, meta-greywacke.
APp	Black, fissile, very fine grained, "sooty" pelite.
APqb	Grey, fine to medium grained, thin to thick bedded, quartz-biotite-feldspar psammite and meta-greywacke, some with garnet and rarely cordierite. Gradational contacts with unit APn in some areas.
APn	Brown, rusty and tan weathering, buff and grey, fine to medium grained, quartz-biotite-feldspar, quartz-feldspar-biotite-garnet-sillimanite and quartz-feldspar-biotite-graphite paragneiss and minor schist. Minor cordierite-bearing paragneiss. Includes some interbeds of units APnc, APc and APcs.
APnqb	Units APqb and APn; interbedded, gradational and undifferentiated.
APcs	Grey and grey-green, medium to coarse grained thin bedded, calcium-silicate gneiss and marble-quartzite with quartz-calcite-plagioclase-diopside and accessory scapolite, actinolite, graphite, epidote and biotite. Commonly interbedded with and passing laterally into unit APc.
APc	White, grey and grey-blue, medium to coarse grained, massive and bedded marble with calcite-diopside-microcline-quartz and minor dolomite, scapolite, phlogopite, graphite, a humite group mineral and tremolite. Interbedded with and passing laterally into unit APcs. Includes small beds of unit APn. Pegmatite of unit Ag is ubiquitous.
APnc	Rusty, fine to medium grained, graphitic paragneiss with pyrite and pyrrhotite.
APm	A dark green, fine to medium grained, massive and foliated amphibolite; some biotite-garnet amphibolite.
APq	White to grey-blue, medium to coarse grained, massive and faintly bedded, orthoquartzite with minor feldspar, white mica and phlogopite. Minor quartzo-feldspathic grit, conglomerate with hematite clasts, grey-green diopside-chlorite rock (meta-regolith?). Biotite-garnet-sillimanite schist and amphibolite.
APif	Rusty, massive pyrite, magnetite iron formation.

UNCONFORMITY

ARCHAEAN

Āg$_1$	Foliated, massive and porphyritic granite and granitoid gneiss.
Āg$_2$	Foliated feldspar augen granite; minor granitoid gneiss.
Āg$_3$	Foliated hornblende granodiorite and amphibolitic granitoid gneiss.
Āgn	Layered granodioritic gneiss; minor amphibolite.
Āng	Migmatite of units Āg$_1$, Āg$_2$ and/or Āgn.
Āggdn	Orange, grey and tan, medium to coarse grained, layered and foliated, biotite and hornblende granodioritic, quartz monzonitic and leucocratic gneiss. Includes rock of units Ag, Āg$_1$, Āg$_2$, Āgn, Āng and small bodies of Ām.
Ām	Dark green foliated amphibolite, meta-gabbro and hornblende-plagioclase gneiss.

PRINCE ALBERT GROUP

ĀAm$_2$	Foliated amphibolite dykes.
ĀAm$_1$	Layered amphibolite.
ĀAb	Foliated and massive, dark grey, coarse grained anorthositic gabbro; minor amphibolite.
ĀAub	Foliated, serpentinized ultramafic rock.
ĀAn	Quartz-biotite-feldspar paragneiss, some hornblende-bearing.
ĀAnm	Muscovite-quartz-feldspar paragneiss.
ĀAif	Dark blue-grey layered oxide facies iron formation.
ĀAq	White, medium to coarse grained, massive orthoquartzite; muscovite- and rarely fuchsite-bearing.
ĀAva	Light grey, fine grained layered acid volcanic rocks.

Geological boundary (defined, approximate).

Boundary of areas extensively drift-covered.

Geological boundary gradational, poorly-exposed, imprecisely located or not observed.

PLANAR STRUCTURES

Bedding and compositional layering (horizontal, inclined, vertical)

Foliation, schistosity, gneissic layering and cleavage (horizontal, inclined, vertical, dip unknown); earliest or only observed.

Axial planes (inclined, vertical) associated with folds deforming bedding and compositional layering. Earliest or only observed.

Axial planes (inclined, vertical); associated with folds of later phases observed to have deformed bedding or early foliation.

LINEAR STRUCTURES

Lineation (plunging, horizontal); formed by fold axes, bedding foliation intersection (X), mineral growth or rodding (R), mullion (M), and boudin axes (B); earliest or only observed.

Lineation (plunging, horizontal); formed by bedding-foliation and foliation-foliation intersection (X), mineral growth or rodding (R) and mullion (M) and fold axes associated with folds observed to have deformed bedding or early foliation.

FAULTS

High angle fault (defined, approximate); arrows indicate apparent relative movement.

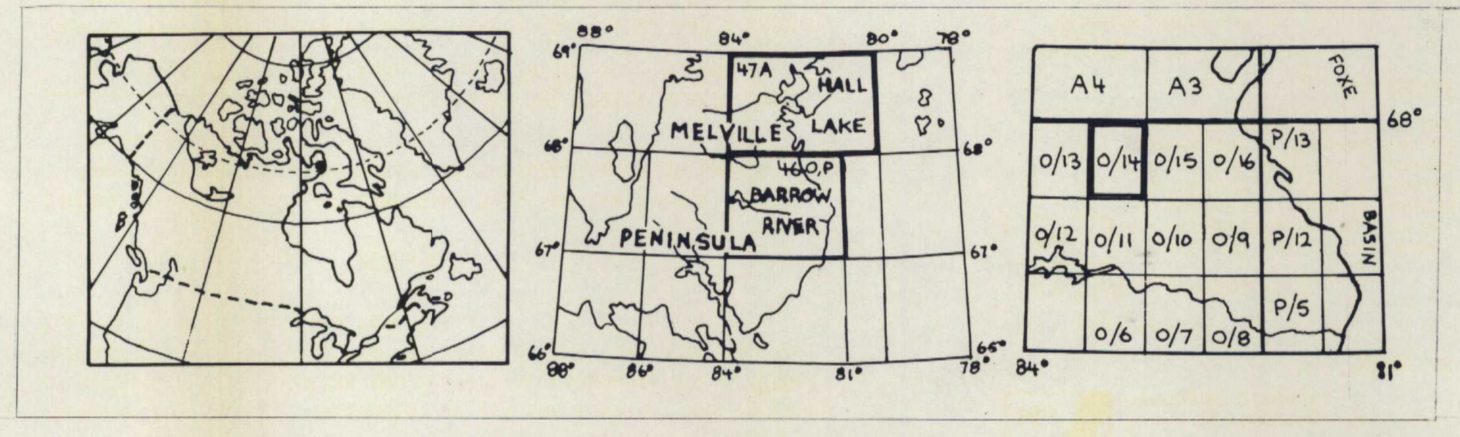
Low angle fault (defined, approximate); teeth in direction of dip.

NOTE ON DATA PRESENTATION

Lithologic and structural data shown is that obtained in the field augmented by only limited interpretation. No attempt has been made to extend lithologic units to construct an integrated stratigraphic and structural model. Lithologic units in the legend are not arranged in temporal order, as this is largely uncertain, except that basal units of the Penrhyn Group are listed below other parts of the group. Description of mesoscopic structures is limited to morphology of common types. Mesoscopic structures are assigned a position (early or late) in the tectonic hierarchy based upon interpretation of local field relationships only. This position cannot be directly related to deformational phases which formed large folds or to mesoscopic structures in nearby areas. No attempt has been made to integrate mesoscopic features into a megascopic structural synthesis. Axial traces of megascopic features are not portrayed as those of early folds are largely interpretative and those of late folds can be deduced from attitudes of planar structures on the map.

Geological mapping by: J. R. Henderson; M. Turay
Geological cartography by: J. R. Henderson
Descriptive notes by: A. V. Okulitch; J. R. Henderson

Notice of any revisions or additional geological information known to users of these maps would be gratefully received by the authors.



OPEN FILE
DOSSIER PUBLIC
FEBRUARY 1978
523
GEOLOGICAL SURVEY
COMMISSION GÉOLOGIQUE
OTTAWA

CANADA

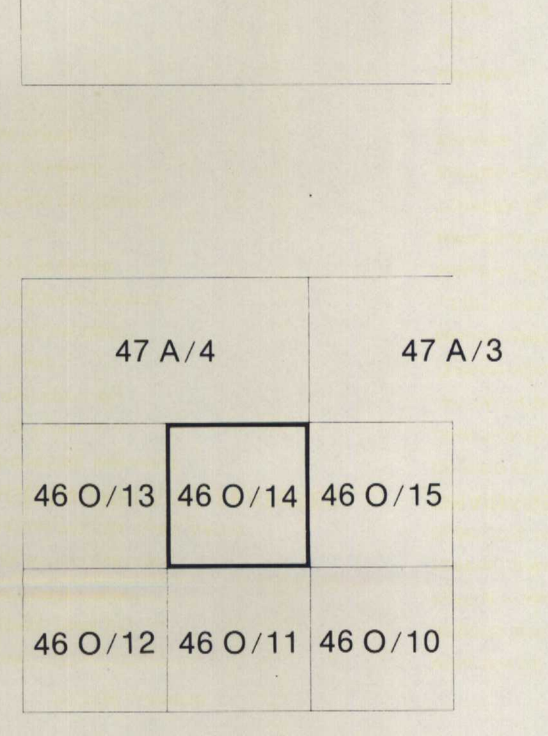
EDITION 1 46 O/14

Military users, refer to this map as: Référence de la carte pour usage militaire:	SERIES A 733 SÉRIE MAP 46 O/14 CARTE EDITION 1 MCE ÉDITION

LEGEND - LÉGENDE

ROADS AND RELATED FEATURES	ROUTES ET OUVRAGES CONNEXES
Hard surface, all weather	Surface pavée, toutes saisons
Loose surface	Gravier
Cart track, winter road or road under construction	Chemin de terre, d'hiver ou route en construction
Trail, cut line, portage	Sentier, percée, portage
Built-up area	Agglomération
Railway, siding, station, stop	Chemin de fer, voie d'évitement, gare, arret
Bridge	Pont
Seaplane base, anchorage	Hydroaéroport, mouillage
LANDMARK FEATURES	**POINTS DE REPÈRE**
House, barn	Maison, grange
Church, school	Église, école
Post office	Bureau de poste
Historical site	Lieu historique
Towers: fire, radio	Tours: feu, radio
Well: oil, gas	Puits: pétrole, gaz
Tank: oil, gasoline, water	Résérvoir: pétrole, essence, eau
Telephone line	Ligne téléphonique
Power transmission line	Ligne de transport d'énergie
Mine	Mine
Cutting, embankment	Déblai, remblai
Gravel pit	Gravière
BOUNDARIES AND SURVEY CONTROL	**FRONTIÈRES ET POINTS DE RÉFÉRENCES**
International, provincial, boundary monument	Internationale, provinciale, borne frontière
County, district	Comté, district
Township, parish-surveyed -unsurveyed	Canton, paroisse - arpentée - non arpentée
Township, DLS - surveyed, unsurveyed	Township, ATC-arpentée, non arpentée
- Section corners	-Coins de section
Municipality	Municipalité
Indian reserve, park, etc	Réserve indienne, parc, etc
Horizontal survey point	Repère planimétrique
Bench mark with elevation	Repère de nivellement avec cote
Spot elevation, precise: land, water	Point coté, précis: sur terre, sur l'eau
DRAINAGE AND RELATED FEATURES	**DRAINAGE ET OUVRAGES CONNEXES**
Stream, shoreline: indefinite	Cours d'eau, rive; imprécise
Direction of flow	Direction du courant
Lake, intermittent lake	Lac, lac intermittent
Inundated land	Terrain inondé
Marsh, swamp (wooded)	Marais, marécage (boisé)
Dry river bed with channels	Lit de cours d'eau tari avec chenaux
Sand: above, in water	Sable: au dessus, dans l'eau
String bog	Marécage en enfilade
Tundra: ponds, polygons	Toundra: étangs, sols polygonaux
Rapids, falls, rapids	Rapides, chutes, rapides
Foreshore flats	Estrans
Rock	Roche
Dam	Barrage
Wharf	Quai
Ditch	Fossé
RELIEF FEATURES	**RELIEF**
Contours	Courbes de niveau
Approximate contours	Courbes de niveau approximatives
Depression contour	Courbe de cuvette
Spot elevation, approximate: land, water	Point coté, approximatif: sur terre, sur l'eau
Esker	Esker
Pingo	Pingo
Sand, sand dunes	Sable, dunes
Palsa bog	Palse
Wooded area	Région boisée
Cleared area	Région déboisée

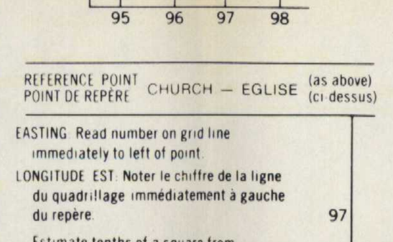
PHOTOGRAPHY PHOTOGRAPHIE
COMPILATION RESTITUTION



REVISION REVISION

GRID ZONE DESIGNATION: DÉSIGNATION DE LA ZONE DU QUADRILLAGE:	100,000 M. SQUARE IDENTIFICATION IDENTIFICATION DU CARRÉ DE 100,000 M.
17W	LF MF 4

EXAMPLE OF METHOD USED TO GIVE A REFERENCE TO NEAREST 100 METRES
EXEMPLE DE LA MÉTHODE EMPLOYÉE POUR FIXER DES REPÈRES À 100 MÈTRES PRÈS



REFERENCE POINT / POINT DE REPÈRE: CHURCH — EGLISE (as above) (ci-dessus)

EASTING: Read number on grid line immediately to left of point.
LONGITUDE EST: Noter le chiffre de la ligne du quadrillage immédiatement à gauche du repère. — 97
Estimate tenths of a square from this line eastward to point.
Estimer le nombre de dixièmes du carré entre cette ligne et le repère en direction est. — 5 — 975
NORTHING: Read number on grid line immediately below point.
LATITUDE NORD: Noter le chiffre de la ligne du quadrillage immédiatement en-dessous du repère. — 98
Estimate tenths of a square from this line northward to point.
Estimer le nombre de dixièmes du carré entre cette ligne et le repère en direction nord. — 4 — 984
GRID REFERENCE SAMPLE / EXEMPLE DU QUADRILLAGE — 975984
Nearest similar grid reference 100,000 metres (about 63 miles)
La prochaine référence similaire est à 100,000 mètres (environ 63 milles)

ONE THOUSAND METRE
UNIVERSAL TRANSVERSE MERCATOR GRID
ZONE 17
QUADRILLAGE DE MILLE MÈTRES
UNIVERSEL TRANSVERSE DE MERCATOR

	47 A/4	47 A/3
46 O/13	46 O/14	46 O/15
46 O/12	46 O/11	46 O/10

The 1974 MAGNETIC BEARING is 43°35′ (775 mils) WEST of GRID NORTH.
ANNUAL CHANGE DECREASING 24.4′
GRID NORTH is 2°05′ (37 mils) WEST of TRUE NORTH for centre of map.

Le REPÈRE MAGNÉTIQUE en 1974 est à 43°35′ (775 mils) OUEST du NORD DU QUADRILLAGE.
VARIATION ANNUELLE DÉCROISSANTE 24.4′
NORD DU QUADRILLAGE est 2°05′ (37 mils) à l'ouest du NORD GÉOGRAPHIQUE au centre de la carte.

THE MAGNETIC COMPASS MAY BE ERRATIC IN THIS AREA
LA BOUSSOLE SERA PEUT-ÊTRE INSTABLE DANS CETTE RÉGION

PRODUCED BY SURVEYS AND MAPPING BRANCH, DEPARTMENT OF ENERGY, MINES AND RESOURCES, OTTAWA, 1974.
COPIES MAY BE OBTAINED FROM THE MAP DISTRIBUTION OFFICE, DEPARTMENT OF ENERGY, MINES AND RESOURCES, OTTAWA.

ELEVATIONS IN METRES ABOVE MEAN SEA LEVEL
CONTOUR INTERVAL .... 10 METRES
NORTH AMERICAN DATUM 1927
TRANSVERSE MERCATOR PROJECTION

DISTRICT OF FRANKLIN
NORTHWEST TERRITORIES
Scale 1:50,000 Échelle
Miles 1 0 1 2 3 Milles
Metres 1000 0 1000 2000 3000 4000 Mètres
Yards 1000 0 1000 2000 3000 4000 Verges

ÉLÉVATIONS EN MÈTRES AU-DESSUS DU NIVEAU MOYEN DE LA MER
ÉQUIDISTANCE DES COURBES .... 10 MÈTRES
SYSTÈME DE RÉFÉRENCE GÉODÉSIQUE NORD-AMÉRICAIN 1927
PROJECTION TRANSVERSE DE MERCATOR

ÉTABLIE PAR LA DIRECTION DES LEVÉS ET DE LA CARTOGRAPHIE, MINISTÈRE DE L'ÉNERGIE, DES MINES ET DES RESSOURCES, OTTAWA, EN 1974.
CES CARTES SONT EN VENTE AU BUREAU DE DISTRIBUTION DES CARTES, MINISTÈRE DE L'ÉNERGIE, DES MINES ET DES RESSOURCES, OTTAWA.

CONVERSION SCALE FOR ELEVATIONS
ÉCHELLE DE CONVERSION DES ÉLÉVATIONS
Metres 30 20 10 0 50 100 150 200 250 300 Mètres
Feet 100 50 0 100 200 300 400 500 600 700 800 900 1000 Pieds

46 O/14
EDITION 1