



LEGEND

PROTEROZOIC

- Pt Coarse biotite-hornblende ± orthopyroxene tonalitic orthogneiss, 1450-1300 Ma*
- Pta Medium- to fine-grained tonalitic orthogneiss
- Pm Coarse leucocratic to mesocratic oligopyroxene-synthetic orthogneiss
- Pm Medium- to fine-grained synthetic orthogneiss
- Pm Coarsely layered, very coarse grained phlogopite-hornblende-orthopyroxene-synthetic biotite-hornblende-chloritoid leucocratic mafic gneiss with non-conformable rock inclusions
- Pm Tectonic mafic gneiss with inclusions of Pm
- Poc Undifferentiated sillimanite-kyanite-garnet orthoamphibole cordierite gneiss
- Pa Heterogeneously deformed metarhyolite to metagabbro anorthosite
- Pgr Medium grained, mildly to non-foliated leucocratic, 1000 ± 6 Ma*
- Pgs Regular gneiss to synthetic gneiss
- Pg Irregular granitic gneiss
- Pgb Irregular granitic gneiss
- Ppb Irregular biotite-rich gneiss of beaded aspect
- Ppe Regular gneiss equivalent
- Ppr Proterozoic gneiss
- Pat Granitic orthogneiss, 1065-515 Ma*

*See van Breemen and Hamner (1986)

Structural features: Straight gneiss: an annealed "C" > "L" fabric; Continuous, reclined layering of granitic and amphibolite material resulting from progressive transposition at high strains; Porphyroclastic gneiss: a well layered "L" fabric; "X" > "Z" fabrics perpendicular to foliation and parallel to lineation; "X" > "Z" fabrics show transposition of low angle discordance features which are all preserved in "X" > "Z" sections; Beaded gneiss: a coarse, biotite-rich gneiss of uncertain origin; Beaded aspect is due to clusters of coarse kelp-like porphyroblasts, either evenly distributed or concentrated in isotropic quartziferous pegmatitic layers.

DESCRIPTIVE NOTES

This map is the result of fieldwork within the Ontario part of the northwestern boundary zone of the Central Metasedimentary Belt (CMB) of the Grenville Province (Fig. 1), compiled by the Geological Survey of Canada during the summers of 1984 and 1985. The area is mostly accessible via Highways 121 and 35. Bedrock exposures are excellent along the major roads and fairly good away from the roads.

The CMB boundary zone in Ontario is a southeast-dipping, upper amphibolite facies, ductile shear zone approximately 4 km thick, characterized by a constant south-southwestward extension, rotation and widespread indicators of northward thrusting (Hamner and Gosnell, 1984; Hamner et al., 1985; Hamner, 1988; see also Schwabacher and Mower, 1990; Collin, 1988, 1986; Easton and Van Kesteren, 1984; Easton, 1985, 1982a, b). In the Haliburton region, the CMB boundary zone is composed of a stack of four thrust sheets, labeled and described by Hamner (1986). From east to west, they are: (1) the 1000-1020 Ma (see van Breemen and Hamner, 1986) Part of the first thrust sheet is within the map area (see units Pt and Pta). The age of the tectonic orthogneiss (Pt) comprising the Redstone and Quatt thrust sheets is approximately 1300 Ma (see van Breemen and Hamner, 1986). The lower three zones are marked by quartziferous ductile shear zones (Pgs, Pgb, Ppb, Ppe and Ppr) which are dominantly south-southwestward dipping. The base of the stack is defined by a well-developed zone of porphyroclastic gneiss (Poc) which is approximately 100 m thick. The upper zone is an orthopyroxene-synthetic mafic gneiss (Pm). Thrust sheets and thrust zones alike are cut by a widespread set of 10 m thick extensional ductile shear zones. The shear zones are dominantly north-south trending, make an angle to the bearing which they cut, are internally granoblastic and show no evidence of metamorphic retrogression. They appear to represent major extensional unroofing events. A 500 m thick carbonate (Pb) thrust zone occurs within the footwall, 500 m below the basal porphyroclastic gneiss (Poc). It is apparently late and has little effect on the thrust configuration.

The metamorphic conditions of stack assembly are determined from the complex mineralogy of the sillimanite-bearing rock (Poc) which comprises a mineralogically annealed orthopyroxene and a prograde assemblage. The former consists of sillimanite, garnet, cordierite, biotite, zircon, quartz, albite and enstatite in the biotite zone. The latter, biotite zone, is composed of orthopyroxene, hornblende, garnet, cordierite, biotite, zircon, quartz, albite and enstatite. The assemblage suggests minimum conditions of 650°C at 5 kbar (Baker, 1980). Rare retrograde orthopyroxene within the Redstone thrust sheet (Fig. 2) suggests that the tectonic gneiss (Pt) had experienced granulite facies metamorphism prior to stack assembly.

The alternation of tonalitic (Pt) and mafic (Pm) compositions within the stack, coupled with the prograde character of both tonalitic (Pt) and mafic (Pm) materials within the thrust sheets prior to assembly of the present stack suggests that the ductile thrust zones may represent out-of-sequence thrusts re-stacking an older thrust configuration (Hamner, 1988).

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Geology by N.G. Collin 1980, R.H. Thöni 1984, and S. Hamner 1984, 1985
Compiled by S. Hamner 1984-1988
Geological cartography by the Geological Survey of Canada
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MAP 1688A
GEOLOGY
WESTERN PART OF THE CENTRAL METASEDIMENTARY BELT BOUNDARY ZONE, GRENVILLE PROVINCE, ONTARIO

Scale 1:50 000 - Échelle 1/50 000

Universal Transverse Mercator Projection / Projection transversale universelle de Mercator
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Magnetic declination 1989, 11°18' West, increasing 4.0" annually
Elevations in feet above mean sea level

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