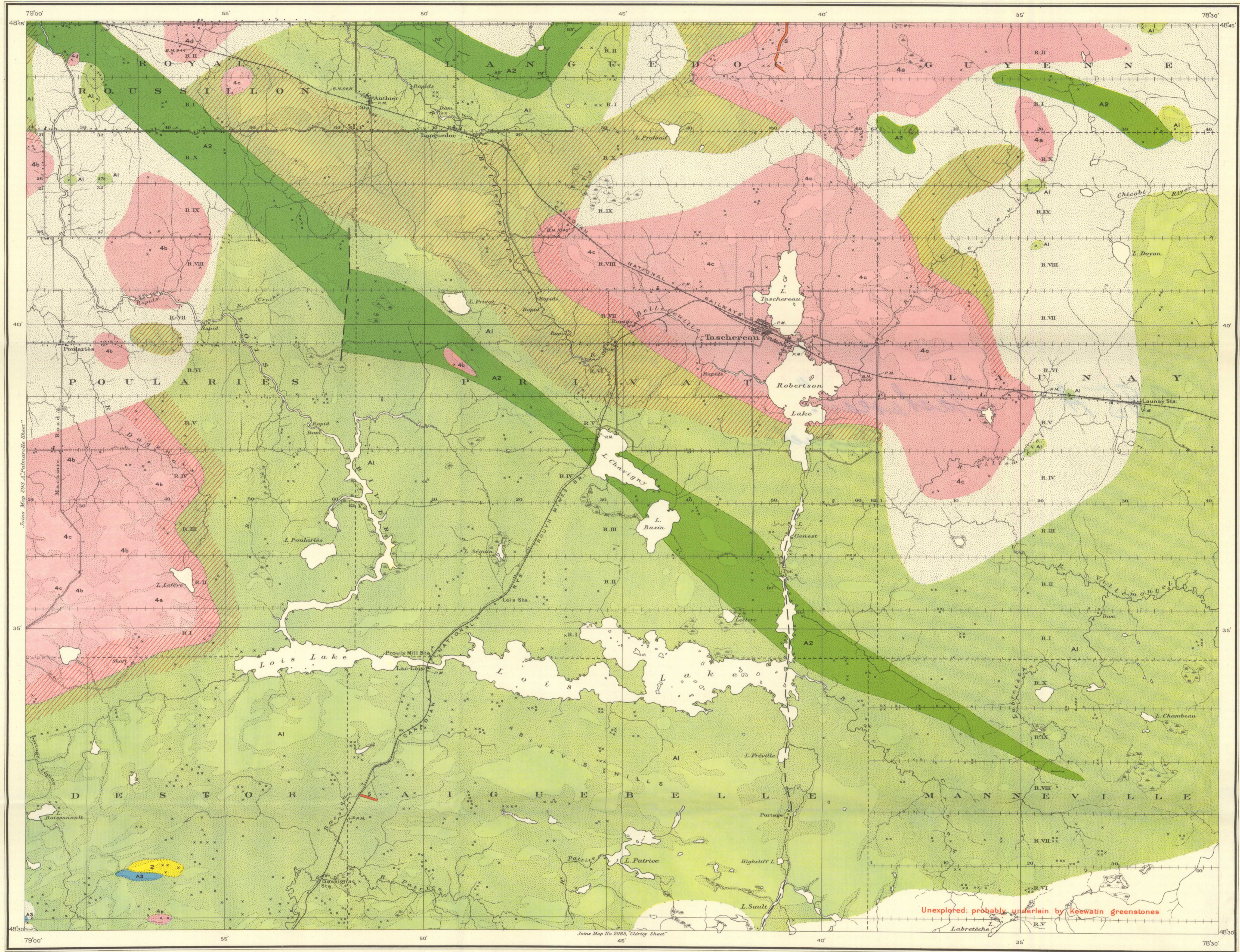


Issued 1934



LEGEND

RECENT AND PLEISTOCENE

- Recent alluvium and glacial deposits. Small rock outcrops.

PRE-HURONIAN (?) INTRUSIVES

- 5 Quartz gabbro dykes (later gabbro)

PRE-HURONIAN INTRUSIVES (relative ages not definitely known)

- 4 Hornblende syenite, 4a hornblende-biotite granite, 4b porphyritic granite, 4c granodiorite and micro basic rocks, 4d serpenitised peridotite possibly unrelated to the granite rocks, 4e

PRECAMBRIAN

- 2 Quartz diorite (older gabbro)

TIMISKAMING SERIES

- A3 Deformed conglomerate

KEEWATIN SERIES

- A2 Chiefly altered tuff
- A1 Chiefly altered andesite, some rhyolite, dacite and basalt, minor amounts of interbedded tuff and breccia, small bodies of basic intrusives

Symbols

- Geological boundary
- Bedding (dip uncertain, inclined)
- Schistosity (dip uncertain, inclined)
- Fault (assumed)
- Contact zone

Note: In Launay and Manneville townships more bedrock outcrops may be present than are indicated. Those marked only represent the outcrops seen along traverse lines one-half mile or more apart. Elsewhere, it is believed, almost all the outcrops are indicated.

Geology by B. S. W. Buffam, 1925, 1926; revised by A. H. Lang, 1933.

PHYSICAL FEATURES

The map-area lies within the "clay belt" and except the southern part is gently rolling, has an average elevation of 1000 feet above sea-level, is floored with clay and is suitable for agriculture. The southern part of the area, in Destor and Aiguebelle townships is traversed by the Abitibi hills rising 500 to 700 feet above the clay lowlands. Deposits of sand, gravel and boulders of glacial origin occur within and bordering the Abitibi hills and form low hills and sand plains on the clay lowlands.

GEOLOGY

In Launay and Manneville townships more bedrock outcrops probably occur than are indicated on the map but elsewhere it is believed that nearly all outcrops have been located.

The oldest rocks are steeply folded KEEWATIN greenstones (A1) accompanied by minor amounts of altered tuffs and breccias forming small, lenticular bodies interbedded with the flow rocks. The lavas are mainly andesites and in many places original textures and structures such as pillows are preserved but, as a rule, the original minerals have been largely destroyed and the rocks are now chiefly composed of chlorite, or grade into chlorite or hornblende schists, and locally have been sheared. Minor amounts of rocks resembling altered rhyolite and dacite, and also basalt are present in the southern part of the area, small bodies of quartz porphyry occur locally. Most of them are considered to be rhyolite flows but some may be intrusives. What are believed to be the youngest Kewatin rocks (A2) are altered, fine-grained tuffs and minor quantities of highly altered rocks resembling phyllites. These sediments are believed to lie in steeply folded synclines and are conformable with the adjoining lavas. The contacts of the Kewatin volcanics and sediments with the larger intrusive bodies are in some places sharp but as a rule broadly developed CONTACT ZONES intervene. Within these zones many inclusions of Kewatin rocks lie in the intrusives and many tongues and other irregularly shaped bodies of the intrusives invade the Kewatin, or the Kewatin rocks are so recrystallized and, apparently, impregnated with matter given off by the intrusives that the intrusive rocks seem to grade into the volcanics or sediments.

The sediments classed as TIMISKAMING (A3) are deformed conglomerates. They presumably occur in a closely folded syncline and probably were once continuous with the larger developments of similar strata present farther west in Duparquet township.

The Kewatin strata are intruded by small and large bodies of different types of rocks all of which are believed to be post-Timiskaming but whose ages relative to one another are, in most cases, unknown. What is believed to be the oldest intrusive is represented by one small body of quartz diorite (2), of so-called older gabbro. The poorly exposed body of hornblende-biotite granite (4a) in Languedoc and Guyenne townships is cut by the porphyritic granite (4c) lying south of it. The hornblende-biotite granite (4b) of the western area of intrusives is also cut by porphyritic granite (4c) and towards the southern margin developments of more basic rocks (4d) occur. Granodiorite and more basic facies (4e) occur in the intrusive areas east of Macamic. Small irregular masses of basic rocks also characterize the contact zones. In the south one body of white-weathering, partly serpenitised peridotite (4e) occupies an isolated position within the Kewatin. This rock resembles types found within the main batholithic areas but may be unrelated.

Small basalt or lamprophyre dykes are present. A few larger dykes of quartz gabbro (5) occur but are not as common as in other districts to the south. No olivine diabase dykes have been found. The quartz gabbro dykes are the youngest of the igneous bodies. Some may be pre-Huronian and others of Huronian or later age.

The structure of the area is imperfectly known. The Kewatin strata lie in large, compressed folds and the sediments, in most cases at least, lie in synclines. An anticlinal axis is believed to exist north of Lois lake but there may be more than one fold between the syncline marked by sediments in Destor township and the syncline similarly indicated to the north in Privat township. The tuffs in Languedoc and Guyenne townships are poorly exposed and their structure can, for the most part, only be inferred. The available evidence suggests a large cross or drag fold.

A prominent depression extends southward from Robertson lake and has been regarded as marking a major fault zone but there is little or no evidence of displacement of the strata. If a fault exists either it is dying out to the north or the displacement is mainly along a vertical direction. The disposition of the outcrops of sediments near the Poularrieux-Privat township boundary suggests the presence of a fault but any physiographic expression of such a fault is lacking.

MINERAL DEPOSITS

There are no producing mines in the map-area. Underground development work was done in 1930 in the contact zone east of the Macamic road on a quartz vein or shear zone carrying chalcocopyrite. Quartz veins, lenses and stringers and also shear zones carrying some pyrite or pyrite and chalcocopyrite have been found in other contact zones and in the Kewatin at some distance from granitic bodies. A wide quartz vein carrying a small quantity of pyrite and molybdenite, and reported gold values occurs in the granite body about Robertson lake.

RELATED PUBLICATIONS

SUMMARY REPORT, PART C, 1925; Destor Area, Abitibi County, Quebec; B. S. W. Buffam.

SUMMARY REPORT, PART C, 1928; Desmeloizes Area, Abitibi District, Quebec; by J. B. Mawdsley.

MEMOIR 166, Geology and Ore Deposits of Rouyn-Harricana Region, Quebec; by H. C. Cooke, W. F. James, and J. B. Mawdsley, 1931.

MAP 271A: Rouyn-Harricana Area, Abitibi and Temiscamingue Counties, Quebec; scale, 1 inch to 4 miles, 1931.

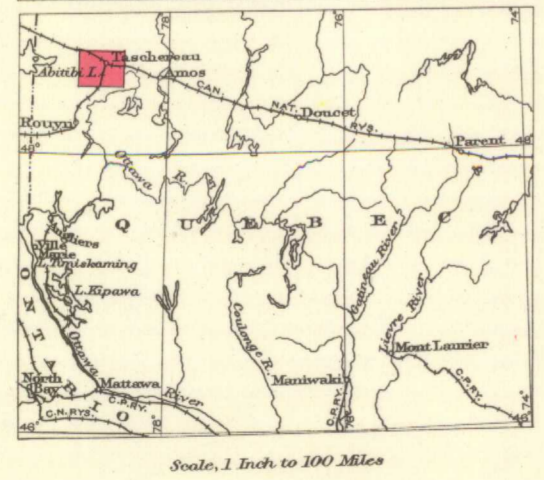
SUMMARY REPORT, PART D, 1932; Palmariolle and Taschereau Map Areas, Quebec; by A. H. Lang.

MAP 281A: Duparquet Sheet, Abitibi and Temiscamingue Counties, Quebec; scale, 1 inch to 1 mile, 1933.

MAP 284A: Desmeloizes Sheet, Abitibi County, Quebec; scale, 1 inch to 1 mile, 1933.

MAP 293A: Palmariolle Sheet, Abitibi County, Quebec; scale, 1 inch to 1 mile, 1934.

MAP 298A: Macamic Sheet, Abitibi County, Quebec; scale, 1 inch to 1 mile (in course of preparation).



MAP 285 A
TASCHEREAU SHEET
ABITIBI COUNTY
QUEBEC

Scale, 63360 or 1 inch to 1 Mile
Miles
Kilometres

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Legend

Road and buildings	Bush road, trail or portage
Road not well traveled	Railway
Road along township boundary	Township boundary
Church	Permanent reference mark
School	Beach mark
Post office	Marsh

Surveys by S. C. McLean, 1926; R. C. McDonald, 1927, and the Department of Lands and Forests, Quebec. Compilation of aerial photographs supplied by the Topographical Survey, Department of the Interior. Beach marks by the Geodetic Survey of Canada.

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