

- MISSISSIPPIAN
- Mc Greyish red conglomerate and grit
- DEVONIAN
- Dd Greenish grey, medium grained, massive diabase and gabbro; includes some undivided Lower Devonian volcanic and sedimentary rocks
- Ddg JUNIPER GRANITE Ddg1: Grey and pink, medium grained, massive muscovite-biotite and biotite-muscovite granite; includes minor pink apite and pegmatite dykes, and migmatite near contacts Ddg2: Pink and grey, medium- to coarse-grained (porphyritic), massive, cataclastic biotite-muscovite and muscovite-biotite granite; includes minor pink apite dykes Ddg3: Reddish orange, pink and grey, medium- to fine-grained, cataclastic, (mylonitic), massive to foliated biotite and muscovite-biotite (augen) granite
- Dlg LOST LAKE GRANITE Dlg1: Grey and pink, medium grained, massive muscovite-biotite and biotite-muscovite granite; includes minor white pegmatite dykes and in the southwestern part migmatite Dlg2: Pink and grey, fine- to medium-grained, cataclastic, (foliated), (gneissic), biotite and muscovite-biotite (augen) granite (may be in part Ordovician)
- Dng NASHWAAK GRANITE: Light grey, fine- to medium-grained, massive biotite-muscovite granite, minor pink foliated leucogranite and mica schist
- LOWER DEVONIAN
- TOBIQUE GROUP
- WAPESKE FORMATION: Basaltic Unit: IDwb: greenish grey basalt, minor diabase, slate, siltstone and rhyolite. Sedimentary Unit: IDws, grey slate, greywacke, shale, minor basalt, diabase and rhyolite tuff
- COSTIGAN MOUNTAIN FORMATION: Basaltic unit: IDcb: greenish grey basalt flows and tuffs, minor siltstone, shale, greywacke, grit, rhyolite and diabase. Rhyolite Unit: IDcv: pink tuff and red rhyolite and trachyte flows and tuffs, minor basalt and siltstone
- ORDOVICIAN
- Og1 Light pink to grey, fine- to medium-grained, foliated biotite granite cut by minor muscovite-biotite granite dykes and garnetiferous pegmatite veins. Og2: Grey to light pink, fine- to medium-grained, gneissic biotite, feldspar-augen granite, minor migmatite
- CAMBRIAN-ORDOVICIAN
- C-Om Grey and rusty grey, medium- to fine-grained, foliated to gneissic quartz-feldspar-muscovite-biotite (cordierite) (sillimanite) (andalusite) migmatite and paragneiss, commonly containing lenses and clots of biotite and/or metasediment and quartz veins, grey, medium-grained massive biotite granite (Devonian?), minor quartzite, slate, amphibolite and pegmatite dykes

- Rock outcrop (definite, probable)
- Rock rubble indicative of nearby bedrock
- Geological boundary (approximate, assumed, gradational)
- Bedding, tops known (inclined, overturned)
- Bedding, tops unknown (inclined, vertical)
- Cleavage, schistosity, foliation (inclined, vertical)
- Gneissosity (inclined, vertical, dip unknown)
- Shearing (inclined)
- Fault (defined, approximate, assumed)
- Anticline (approximate)
- Glacial striae (direction of flow unknown)
- Mineral prospect

ROCK TYPE			
Basalt	v	Rhyolite	r
Diabase and gabbro	d	Sediments	s
Grit	g	Tuff	t
Pegmatite	p		

MINERAL OCCURRENCES			
Copper	Cu	Molybdenum	Mo

Geology by R. Skinner, 1971. C.J. St. Peter, 1978

Geology compiled by R. Skinner, 1982

Geological cartography by R.Y. Patvin, Geological Survey of Canada

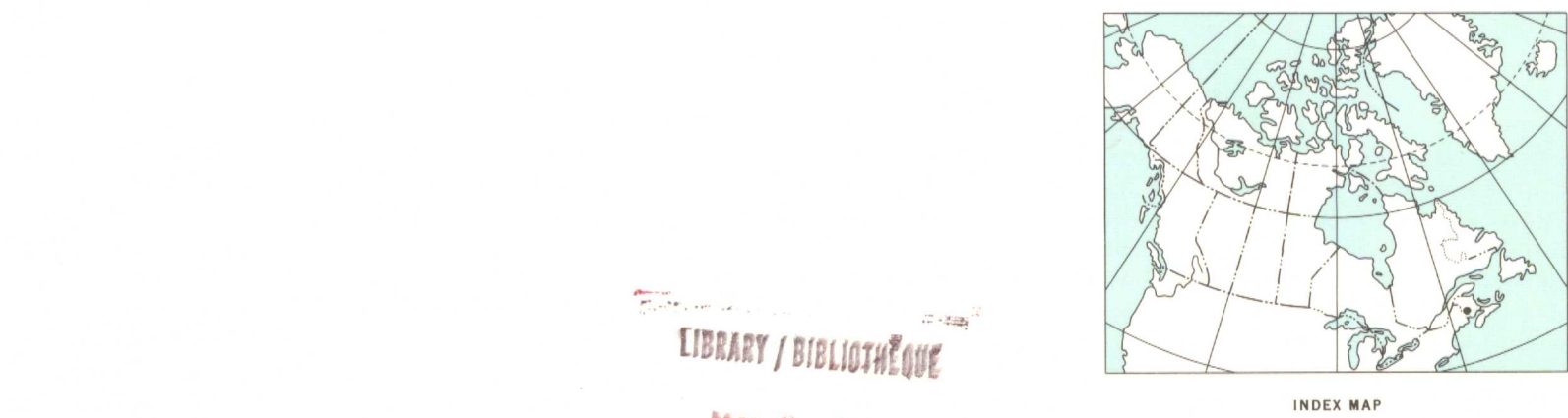
Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada

Base map at the same scale published by the Surveys and Mapping Branch in 1957. Roads were revised by the Geological Survey of Canada for this edition

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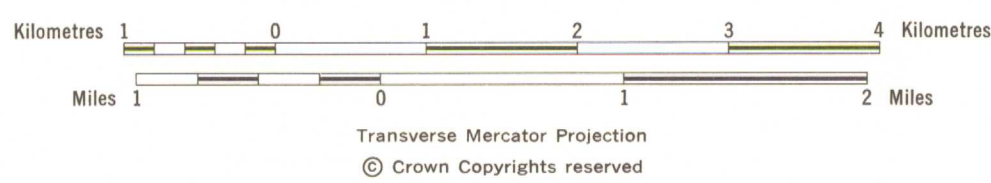


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MAP 1578A  
GEOLOGY  
JUNIPER  
(East Half)  
NEW BRUNSWICK  
Scale 1:50 000



Transverse Mercator Projection  
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21/13	21/14	21/15
1554A	1402A	
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NATIONAL TOPOGRAPHIC SYSTEM REFERENCE AND INDEX TO GEOLOGICAL SURVEY OF CANADA MAPS

MAP 1578A  
JUNIPER  
NEW BRUNSWICK

DESCRIPTIVE NOTES

Juniper (east half) map area lies in the Appalachian Mountain System on the west flank of the New Brunswick Highlands. The map area forms part of the Miramichi Anticlinorium or Massif (Fyffe et al., 1981) - a belt of highly deformed lower Paleozoic volcanic and sedimentary rocks (intruded by Devonian and Ordovician granites) that extends northeasterly across the province from Woodstock to Bathurst. The map area is 50 km northeast of Woodstock, New Brunswick.

The area is underlain mainly by granitic rocks, most of which are of Devonian age and all lie south and east of the Catamaran Fault. The oldest rocks (C-Om) are biotitic migmatites and paragneisses of Cambrian and/or Ordovician age and extend into adjoining areas to the east (Poole, 1963, 1980), northeast (Skinner, 1975) and south (Anderson, 1968). Their age is based on relationships in the Hayesville map area (Poole, 1963; Crouse, 1979) where sillimanite gneiss and pegmatite grade orderly into schist, quartzite and phyllite that pass conformably upward into brachiopod-bearing siltstone, dated Early Ordovician (Neuman, 1967). The migmatites contain lenses and veins of granitic material interlayered with mica-rich layers and clots. The protolith of the micaceous material is probably of Cambro-Ordovician age, whereas the granitic material appears to be identical to the granite in the adjacent granitic bodies and more than likely is of Devonian age.

The oldest granites in the map area are foliated cataclastic "Sugar" (Og1) and feldspar augen (Og2) granites. Poole (1980) dated the "Sugar" granites in the Hayesville map area adjacent to the Juniper map area. Five of the eight whole rock samples defined an isochron age of  $498 \pm 19$  Ma, initial ratio  $0.7034 \pm 0.0029$  and mean square of weighted deviates (MSWD) 1.13. The three other samples fall below this line. An errorchron age of  $484 \pm 33$  Ma, intercept  $0.7039 \pm 0.0052$  and MSWD 4.85 is obtained if all eight samples are included in the regression analysis. These ages fall about early Ordovician.

The Tobique Group rocks are extensions of those present to the north in the Plaster Rock (east half) map area (Skinner, 1981) and have been traced over 110 km from Upper Kintore in the southwest (C. St. Peter, in preparation) to Serpentine River in the north (Irrinki, 1977a,b) and are named after the river along which it outcrops west and northwest of the area. The Early Devonian age, probably Helderbergian and possibly New Scotland (Gedinnian), given these rocks is based on dating by A.J. Boucol, Oregon State University, of brachiopods collected by Skinner in 1970 (1981). Irrinki in 1975 and 1977 (1977c) and St. Peter in 1976 (1978) in the Plaster Rock (east half) map area. Although the Wapke Formation, rich in sediments and mafic volcanics conformably overlies the Costigan Mountain Formation, rich in acid volcanics (in the Plaster Rock map area), fossils collected from each indicate similar ages. It is possible that the Wapke is a lateral facies equivalent of the upper part of the Costigan Mountain Formation.

Most of the Devonian granites in the map area are grey or pink and contain varying proportions of biotite and/or muscovite. The predominantly muscovite-bearing granites are pale grey to leucocratic and are confined largely to the highest parts of Beadle Mountain (Ddg2) and to the Nashwaak body (Dng). The granites adjacent to the Catamaran Fault are cataclastic, commonly foliated, and in places contain feldspar augen, and those along the northwestern border of the Juniper granite body (Ddg3) have a distinctive reddish-orange colour.

The ages of the Devonian granites are based mainly on ages obtained in adjoining areas. The Nashwaak granite (Dng) yielded a K-Ar date of  $386 \pm 10$  Ma from a biotite concentrate obtained from granodiorite in Napadogan map area (Poole, 1958; Anderson and Poole, 1959) 4 km east-southeast of Juniper map area (W.H. Poole in Wanless et al., 1972, p. 71). The age of the Lost Lake granite (Dlg1) is speculative. In the adjoining Hayesville map area, Poole (1963) gave these rocks an Ordovician age based on K-Ar ages of 418 and 443 Ma (in Lowdon et al., 1963, p. 110 and 111) and the fact that they are commonly cataclastic, particularly in the northern part (Dlg2). Later, these dates were re-examined and reserve concentrates of biotite and muscovite were re-analyzed, yielding K-Ar ages of  $394 \pm 18$  and  $363 \pm 16$  Ma respectively (Wanless et al., 1973, p. 73-83). In 1980 Poole carried out a Rb-Sr study of the Lost Lake granite but the results were inconclusive. Of five whole rock samples, four with  $^{87}\text{Rb}/^{87}\text{Sr}$  ratios less than 11 form an errorchron of age  $444 \pm 72$  Ma, initial  $^{87}\text{Rb}/^{87}\text{Sr}$  ratio of 0.7070 and MSWD 11.52, which spans between Late Cambrian and Late Devonian (Poole, 1980).

The only age determinations made within the Juniper map area were on pegmatite cutting migmatite, and on paragneiss (migmatite) from the oldest map unit (C-Om). The ages yielded apparently represent Acadian granitic activity. The pegmatite dated cut migmatite on Highway 107, 1.7 km east of Baggar Brook and gave a K-Ar age of  $372 \pm 18$  Ma from a muscovite concentrate (W.H. Poole in Wanless et al., 1973, p. 80). The paragneiss (migmatite) dated is on a lagging, red 5.6 km east-northeast of the bridge over North Branch Southwest Miramichi River. A biotite concentrate gave a K-Ar age of  $393 \pm 18$  Ma and a muscovite concentrate a K-Ar age of  $408 \pm 18$  Ma (W.H. Poole in Wanless et al., 1965, p. 108 and 109). These ages are considered unreliable by Poole (Wanless et al., 1973, p. 82). The Burnhill granite (Ddg1) and 2 and 10a in Poole, 1963) 4.4 km to the east in the Hayesville map area yielded a whole rock Rb-Sr age of  $352 \pm 9$  Ma from the northern part of the body (C. Brooks, personal communication, 1980, in Fyffe et al., 1981, p. 32). Earlier, Poole obtained a 346 Ma K-Ar age from a biotite concentrate from the eastern part of the body (in Lowdon et al., 1963, p. 112) and a 440 Ma K-Ar age from a muscovite concentrate from a gneiss bordering a quartz vein in the southwestern part of the body (in Leech et al., 1963, p. 100 and 101).

In the Tuadok Lake map area the Redstone granite (Crouse, 1978; unit 4c; Skinner, 1975) 10 km to the north yielded a whole rock Rb-Sr age of  $409 \pm 20$  Ma, essentially the same as the whole rock Rb-Sr age of the Costigan Mountain volcanics (Dcva) (Fyffe and Cormier, 1979, p. 2049) and a smaller granodiorite body 25 km to the east of the Redstone granite yielded a K-Ar age of  $382 \pm 16$  Ma (R. Skinner in Wanless et al., 1972, p. 75).

The Mississippian rocks (Mc) in the map area apparently outcrop in only one place, a low railway cut 0.4 km south of Gould Brook. There the outcrop is poorly exposed and extends for less than 20 m along the west side of the tracks. Not enough strata is visible to estimate the attitude of the highly weathered conglomerate which consists mainly of rounded pebbles and boulders of 45 cm long of red to pink Costigan Mountain rhyolite and granite and green basic volcanics in a brown gritty matrix. The conglomerate is similar to Mississippian conglomerate to the south in the Coldstream map area (Anderson, 1968 and personal communication, 1981).

The Cambro-Ordovician paragneiss and migmatite have pronounced gneissic and schistose structures and the Ordovician granites (Og) commonly have a well developed foliation. The attitude of these structures is highly variable, probably as a result of deformation by both the Taconic and Acadian orogenies. The Tobique Group rocks on the other hand have been folded into complex folds with planar, steeply dipping cleavage. The Devonian granite bodies are so similar petrographically and so close to one another that they no doubt coalesce at depth. The Juniper (Ddg) and Lost Lake (Dlg) bodies have been highly sheared and altered along the Catamaran Fault up to a distance of at least 4 km from the fault in the latter case and are commonly mylonitic adjacent to the fault.

The Catamaran Fault has been traced to the east-northeast of the map area for more than 75 km by Anderson (1972). Movement along the fault produced a right lateral displacement of about 7 km on Middle Ordovician basic volcanics at Catamaran Brook 60 km east-northeast of the map area (Anderson, 1970, 1972; Irrinki, 1979) and 20 km to the west of the map area a probable spilloff of the fault sets Mississippian redbeds against Lower Devonian Basic Formation sedimentary rocks (St. Peter, 1979).

The Cambrian-Ordovician paragneiss and migmatite (C-Om) have been metamorphosed to the amphibolite grade as indicated by the scattered occurrences of cordierite, sillimanite and andalusite, whereas the Tobique Group rocks have only undergone prehnite-pumpellyite facies metamorphism.

No significant metallic mineral occurrences are known within the map area. However, molybdenite and pyrite are present in quartz veins up to 60 m thick cutting migmatite-bearing granite 4 km northeast of Juniper Station. The mineralization can be seen in prospect pits and rubble along the south side of the main logging road about 2.6 km east-northeast of the bridge over the North Branch of the Southwest Miramichi River.

St. Peter (1980) reported disseminated pyrite and malachite in highly fractured Nashwaak granite about 2.5 km west of Diamond Lake in the southeastern part of the map area. He also reported the presence of a uranium anomaly on the north end of Beadle Mountain. A combined geochemistry and scintillometer survey of the area was carried out by J.D. Irving Ltd. as a result of a New Brunswick Department of Natural Resources - Canada Department of Energy, Mines and Resources airborne reconnaissance gamma-ray spectrometric survey of pre-Carboniferous rocks of New Brunswick in 1976 (Woodstock, New Brunswick, Geological Survey of Canada Map 360216). The stream silt samples gave high  $\text{U}_{30}$  values and the scintillometer detected an anomaly three times background (New Brunswick Mines Branch Assessment File No. 472216).

Lead, zinc, copper and fluorine are present in the Rhyolite unit of the Costigan Mountain Formation (IDcv) on Costigan Mountain, 25 km to the north in the Plaster Rock (east half) map area (Skinner, 1981). Probably these rhyolitic rocks in the Juniper map area are also mineralized.

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All K-Ar and Rb-Sr ages quoted from published papers have been calculated or recalculated using the new constants recommended by Steiger and Jager (1977). The time scale used in this report is that of Armstrong (1978), constructed using the new constants.

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