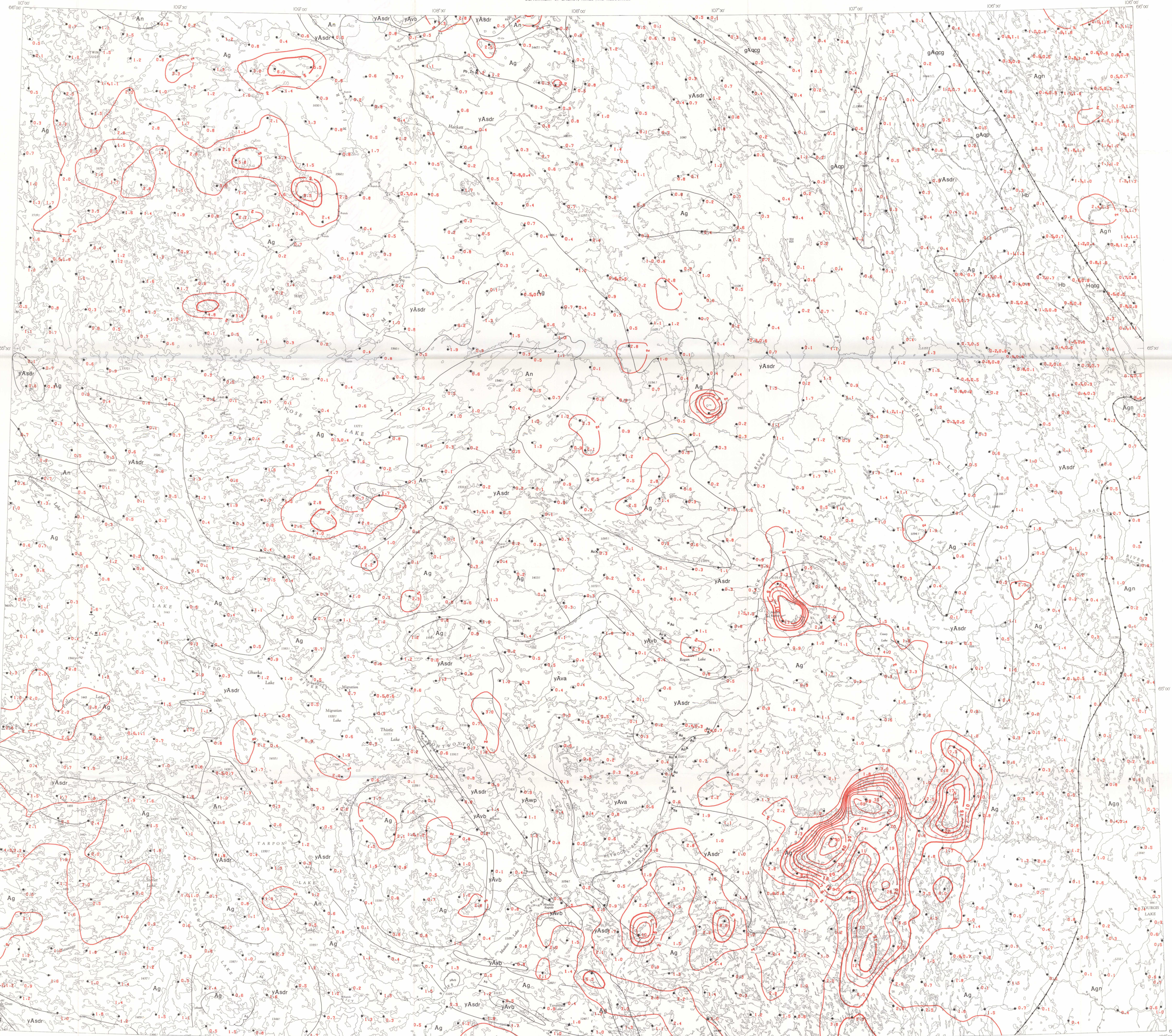


PRELIMINARY SERIES



- LEGEND**
- HADRYMAN**
Hb Gabbro sills, sashes, dykes
- HELIZAN**
Hh ELIZAC FORMATION: quartzite, conglomerate
- APHERIAN**
Aa GOULBURN GROUP
AAsd BURBIDE RIVER FORMATION: pink quartzite, conglomerate, sandstone, shale
AAsp WESTERN RIVER FORMATION: quartzite, greywacke, red siltstone, shale, conglomerate, conglomerate
- AG** Quartz diorite, quartz monzonite, granodiorite, granite
- An** Granitic gneiss, migmatite, mixed gneisses involving Yellowknife rocks
- Ag** Granitoid gneiss, granitoid rocks, minor granites, biotite-hornblende schist and gneiss (within the Churchill Province)
- YELLOWKNIFE SUPERGROUP**
yAsp Greywacke, shale
- yAsdr** Cordierite-andalusite bearing knotted schist and other metamorphic equivalents of yAsp
- yAvb** Acidic lava, tuff, agglomerate
- yAvn** Intermediate to basic lava, tuff, agglomerate, and undifferentiated acidic volcanic rocks
- Boundary between Slave and Churchill geological provinces
Anticline
Syncline
Mineral prospect showing principal element(s)
Lake sample site and metal concentration (sediment sieved to minus 250 mesh)
Lake sample site and metal concentration (sediment sieved to minus 150 mesh)
Geochronological concentration contours as ppm

MINERALS

Copper	Cu	Lead	Pb
Gold	Au	Zinc	Zn
Iron	Fe		

Geology after unpublished map compiled by J. C. McGlynn, 1971

Field work by R. J. Allan, R. M. Cameron, C. C. Durham, R. Benson, R. Collier, R. Cumming, G. Lind, D. Mann, C. Pridie, G. Thomas and R. Worlock

Analyses by J. L. Lynch and J. C. Pelchat

Marginal notes by R. J. Allan and R. M. Cameron

Geochronological contours and metal concentration numbers drawn by computer drum plotter

Geological cartography by the Geological Survey of Canada

Any revisions or additional geological information known to the user would be welcomed by the Geological Survey of Canada

Base-map assembled by the Geological Survey of Canada from maps published at the same scale by the Army Survey Establishment R.C.E. in 1964, 1965

Copies of the topographical maps covering this map-area may be obtained from the Canada Map Office, 615 South Street, Ottawa, Ontario K1A 0S9

Mean magnetic declination 1972, 30°24' East, decreasing 6.0' annually. Readings vary from 27°13' in the SE corner to 34°16' in the NW corner of the map-area

Elevations in feet above mean sea-level

MARGINAL NOTES

Lake Sediment Geochemistry The use of lake sediments as an aid to mineral exploration and geological mapping within the Canadian Shield is based on two principal concepts of their origin. The first is that the detrital portion of a fine-grained lake sediment is a good composite sample of the rocks in the vicinity of the lake. In perhaps a majority of cases, the material forming the sediment has passed through an intermediate stage as a component of locally detrital till or other glacial sediments before being transported to the lake. The second concept is that the fine-grained particles of the sediment are an excellent medium for the sorption of metal ions released during the weathering of nearby soluble ore deposits or similar mineralization. Most of the known sulphide deposits in the Bear-Slave survey area show moderate to high degrees of oxidation.

At the wide reconnaissance sampling interval used, it is unlikely that many samples will be taken from lakes within the limits of the secondary dispersion halo of a single ore deposit. However, country containing such deposits may be defined by the trace element dispersion from the much more extensive non-economic mineralization that is often associated with economic deposits. Similar trace metal patterns may also be derived from mineralization that is not associated with ore deposits or from rock units of unusual chemical composition.

Lake Sediment Sampling The lake sediment samples were collected by post-hole auger from a helicopter. They were taken near the edge of the lake to water 2 to 3 feet deep. They comprise approximately the top 8 inches of sediment, less the surface layer. Of the variety of sediment types that may occur in lakes, the type of sample sought was of clay to silt grade and low in organic material.

Environmental Effects Certain heavy metals, such as uranium, zinc, and copper, may be enriched in sediment samples containing organic matter or iron and manganese oxides. The content of Fe and Mn has been determined for all samples, together with an index of the organic content (organic density). These data will be released separately in the form of a composite map. These data may be examined when assessing anomalous areas on this map. It should be recognized, however, that metal enrichment on these materials is a highly variable phenomenon that depends on a variety of factors, principally the availability of metals in the environment. Also, the Fe and Mn content of a sediment is not a direct measure of the amount of free oxides present. For these reasons the maps for the heavy metals are not presented on a statistically "corrected" basis. Further, since manganese and iron may be associated with these metals in the original rocks or primary mineralization, "correcting" the data for enrichment of mineralization from the resulting map, if it is suspected that free oxides or organic material have created a false anomaly, then for those who do not have access to the original samples, the most satisfactory method of checking is to re-sample the area of the anomaly.

Sample Preparation and Analysis Sediments were dried, then sieved to minus 250 mesh to give a powder suitable for analysis. A few coarse samples were sieved to minus 150 mesh, then bulked before analysis.

For the uranium determination, 100 mg of the sample was digested on a water bath at 95°C for 1.5 hours with 6 ml. of 4N HNO₃ plus two drops concentrated HCl. The sample was diluted to 20 ml with metal-free water, well shaken, and allowed to settle. A 2 ml aliquot of this solution was placed on a platinum dish and evaporated to dryness. The residue was burned for one minute above a Bunsen flame. To this was added 0.2 ml of a 5% Na₂CO₃ solution in the proportions 9:1. The sample was fused for 10 minutes at 650°C, then dissolved into the form of a clear, colorless solution in 10 ml of 2N HCl, then measured by fluorescence. The detection limit of this method is 0.2 ppm U, measured less than this are recorded as 0.1 ppm U.

Uranium in the Rocks and Ore of the Slave Area There is a marked difference in uranium content between the Proterozoic rocks of the Slave Province and the Archean rocks of the Slave Province. This has been demonstrated by the work of Ende and Pelchat (1971) who analyzed their composite samples, including those for which the authors do not report uranium values. The analytical method is similar to that given above, but there is no initial total decomposition of the rocks with hydrofluoric, perchloric and nitric acids. These new data are as follows:

Hardisty Lake (NTS 86A)

Quartz-feldspar porphyry (117 square miles)	7.2 ppm U
Diorite dykes (-)	2.0 ppm U
Undifferentiated gneisses and schists (2 square miles)	2.0 ppm U
Deep level granite to granodiorite (111 square miles)	5.3 ppm U
High level granite, quartz monzonite, monzonite (244 square miles)	6.0 ppm U
Weighted average	6.0 ppm U

Fort Enterprise (NTS 86A)

Volcanics and metavolcanics (104 square miles)	0.7 ppm U
Paragneiss, parashist (209 square miles)	1.0 ppm U
Granitic gneiss, migmatite (1920 square miles)	1.4 ppm U
High level granite, quartz monzonite, monzonite (1750 square miles)	1.7 ppm U
Weighted average	1.5 ppm U

Uranium mineralization is a prominent feature of parts of the Bear Province. Significant uranium mineralization appears to be controlled on both regional and a local scale by major faults of northeast and southeast strike (Ruzhicki, 1971). At the Hardisty Mine at Fort Hardisty, uranium occurs along a prominent northwesterly fault zone. A variety of rocks have been deposited along with the plutonoids including apatite, quartz, hematite, nickel and cobalt arsenides, pyrite, barite, siderite, sphalerite, tetrahedrite, bornite, chalcopyrite, galena, stibiochalcite, native silver, argentic and silver arsenides. The 1400 m.y. age of the mineralization (Ruzhicki, 1971) is younger than the granitic intrusives of the region (1700-1900 m.y.) and is approximately the same age as diorite dykes. Uranium also occurs within extensive stockwork zones within the Bear Province, of which the Hardisty Mine is an example. These contain a much simpler mineral assemblage. Only hematite, pyrite, and chalcopyrite accompany the plutonoids in the Hardisty Mine. There is no known important uranium mineralization within the Slave Province.

It is apparent from the above that the rocks of the Bear Province will contribute a higher uranium background to the lake sediments of this area than those of the Slave Province. Ore-related anomalies may be superimposed on these higher background levels. Because of the generally complex nature of the uranium-bearing ores, these ores should give multielement anomalies. It appears that Cu may be the metal that most consistently accompanies U, but there may also be anomalies for such elements as Pb, Zn, Ag, Ni, Co, As, La and Y.

Uranium in the Surficial Environment Uranium is readily oxidized to the highly soluble uranyl ion (UO₂²⁺). This, in turn, may form soluble complexes with carbonate ions that are stable in the near neutral pH values characteristic of ground and lake waters of the Canadian Shield. In this form the uranium may migrate considerable distances from its source. In lakes there are a number of processes that may cause the uranium to precipitate in the sediments: reduction in the bottom muds to the insoluble U(IV) reaction with carbonates, arsenides and other oxides; sorption onto clay minerals; sorption or coprecipitation with iron and manganese oxides; fixation by organic material.

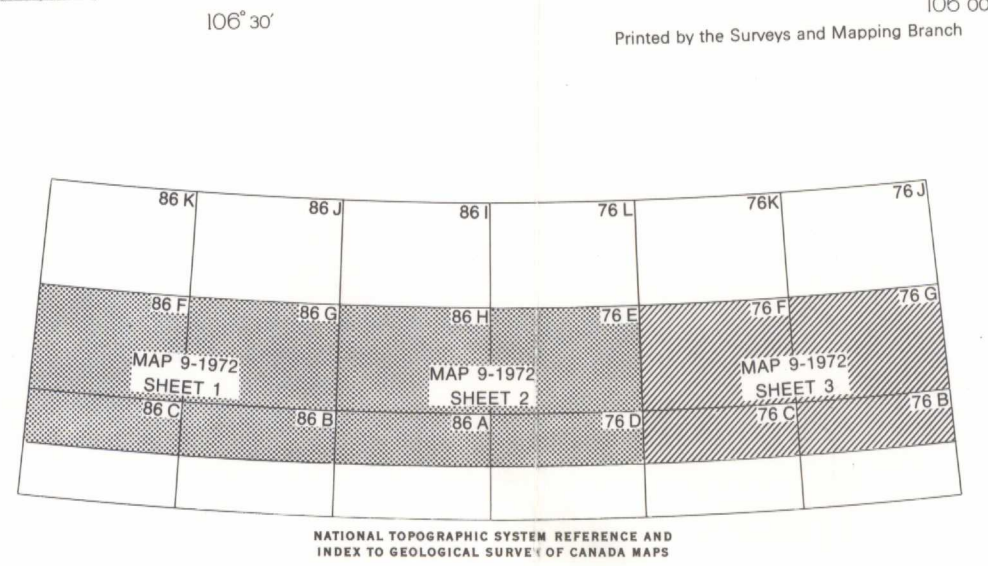
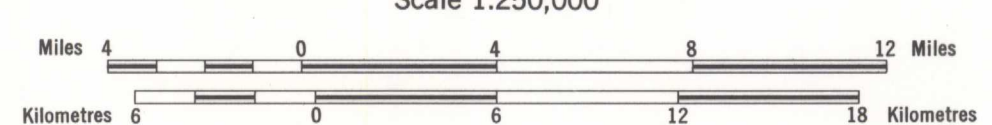
Uranium in Lake Sediments, This Sheet The background levels for uranium within Sheet 3 are lower than that for either of the other maps shown.

There is a large, distinct, anomaly in granitic rocks in the southeastern part of the sheet, near to the boundary between the Slave and Churchill Geological Provinces. This area contains up to 70 ppm U in the lake sediment samples and is also distinguished by relatively high amounts of U, Zn, Co, Cu and Ni. As for most of the anomalies outlined during the Bear-Slave operation, follow-up work on this anomaly should, in the first instance, be aimed at discovering whether the high element values are caused by a distinctive rock type (e.g. a granite of Proterozoic age) or by mineralization. The granitic rocks of this area have been identified as mafic granites and quartz-monzonites (Wright, 1975).

There are a number of other lesser anomalies on the map sheet that are defined by the 2 ppm and the 5 ppm U contours. Most are in granitic rocks. Some of these are associated with higher than average contents of zinc; they should be examined in conjunction with the zinc map.

Geological Survey Paper 79-50, contains a detailed description of the experimental basis for lake sediment sampling in the Shield; the geology and metallogeny of the region; the organization, methods, and costs of the sampling operation; the methods of sample analysis and the references for the articles quoted above.

MAP 9-1972
SHEET 3
URANIUM CONTENT OF LAKE SEDIMENTS
BEAR-SLAVE OPERATION
DISTRICT OF MACKENZIE



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