

LEGEND

11	Gabbro, diabasic gabbro; basalt dykes and sills; 11a, assemblage of several narrow dykes	Drag-fold
10	ATHABASCA GROUP (7-10) Conglomerate, round to subangular fragments, includes some arkose and siltstone fragments, arkosic matrix; 10a, includes some arkose; 10b, includes some siltstone; 10c, brecciated	Fault (defined, assumed)
9	Arkose, orange-red, medium-grained; 9a, includes siltstone beds; 9b, includes some conglomerate; 9c, brecciated	Anticline (position approximate)
8	Siltstone, chocolate-red, fine-grained; 8a, includes some conglomerate; 8b, includes some arkose; 8c, includes some green shale; 8d, brecciated	Syncline (position approximate)
7	Basal conglomerate, angular to subangular fragments, abundant fragments of 1, 3, 5 and 9; 7a, abundant fragments of granitic rocks (5 and 6); 7b, abundant fragments of quartzite rocks (1); 7c, abundant fragments of chlorite-bearing rock or chlorite-rich matrix; 7d, includes beds of siltstone; 7e, brecciated	Glacial striae
6	TAZIN GROUP (1-6) Metasomatic granite, red, massive, low mafic content; 6a, brecciated; 6b, mylonitic, finely gneissic, low mafic content; 6c, coarsely to roughly bedded and massive, biotite-bearing; 6d, includes white granitized quartzite rocks; 6e, hybrid appearance, high mafic content	Rock trench and stripped area (radioactive)
5	Granitic, banded gneiss, generally reddish, high mafic content; 5a, includes some granitic rock	Radioactive occurrence
4	Interbedded mixture of grey quartzite and quartz-feldspar-chlorite-sericite schist and gneiss, augen structure, minor granitic rocks; 4a, locally biotite-bearing	Mineral occurrence
3	Amphibolite, hornblende-feldspar gneiss and schist, in part chlorite-bearing; 3a, hornblende schist and gneiss; 3b, chlorite schist; 3c, includes some quartzite beds; 3d, includes some granitic masses; 3e, granitized amphibolite	
2	Argillite, grey to black, locally silty, in part tuffaceous; 2a, altered to chlorite-epidote rock, green, massive to schistose; 2b, includes granitic material, granitized	
1	Quartzitic rocks, generally white and massive, low mafic content, locally feldspathic; 1a, includes much granitic material; 1b, brecciated; 1c, thinly bedded and interbedded with chlorite-sericite schist; 1d, coarse, rusty white, granite-like, generally roughly bedded and massive, biotite-bearing; 1e, similar to 1d, in part schistose; 1f, similar to 1d, includes a few beds of massive, dense, white quartzite; 1g, quartz-feldspar banded gneiss, quartzite-like, minor areas of granitic rocks; 1h, similar to 1d, includes much hornblende schist; 1i, somewhat schistose and interbedded with abundant chlorite-sericite schist and granitic rocks	

MINERAL SYMBOLS

Garnetgt Pyroxenepy
Pitchblende...U Uraninite.....m

Geology by L. P. Tremblay, 1955, 1956

ARCHAIC (?)

Drift-covered area
Area of gravel and sand
Rock outcrop, area of outcrop
Geological boundary (defined, assumed)
Bedding (inclined, vertical, dip unknown)
Schistosity (inclined, vertical, dip unknown)
Lamination (dipping)
Minor fold (as indicated)

MAP 25-1957
URANIUM CITY
SHEET 6
SASKATCHEWAN

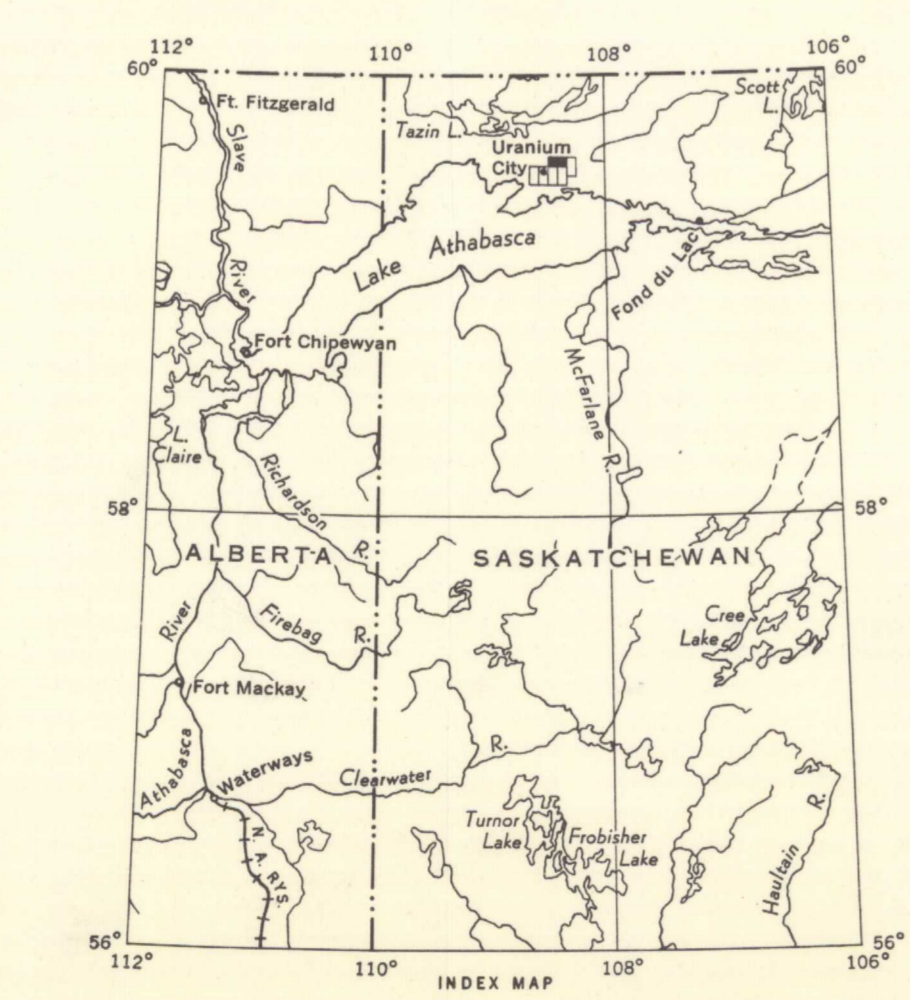
Scale: One Inch to 800 Feet = $\frac{1}{9600}$
800 1600 2400
Feet

Approximate magnetic declination, 25° 25' East

Cartography by the Geological Cartography Unit, 1958

Air photographs covering this area may be obtained through the National Air Photographic Library, Topographical Survey, Ottawa, Ontario

In response to public demand for earlier publication, Preliminary Series maps are now issued in this simplified form, thereby effecting a substantial saving in time. There is no loss of information, but the maps will be clearer to read if all or some of the map-units are hand-coloured.



The centre of the area is about 4 miles northeast of Uranium City and at the time of mapping was best reached by aircraft. A road built in the autumn of 1956 reaches the south end of Fredette Lake, making access easy to the west half of the area. Good and short portage either from Donaldson Lake near the east boundary of the area or from Ace Lake south of its east half lead to Mickey Lake.

The area is rugged. It is dissected into a few broad north-easterly trending ridges separated by deep valleys. The main ridges are further dissected into several narrow minor ridges. The relief is around 150 feet above the lakes in the north half of the area but is from 100 to 250 feet near the southern boundary. All drainage is to the south.

The relative ages of the various rock types of the Tazin group (1-6) as shown in the legend are based on an anticlinal structure east of Donaldson Lake and outside the map-area. The stratigraphic succession west from Donaldson Lake inside the area is from the earliest to the latest rocks, although near the lake some formations are repeated by folding.

The quartzite rocks (1) pass westerly from a white and rusty brown weathering, coarse-grained, granite-like rock (1d) near Donaldson Lake to a narrow zone of finely bedded, buff coloured, dense massive quartzite (1c), and then to a zone of coarse-grained red granite (1a). This granite zone includes numerous remnants and a few somewhat larger masses of ungranitized massive or brecciated, buff and white, glassy quartzite, all with gradational contacts. The stratigraphic position of the granite rocks west of the Black Bay fault relative to those east of the fault is unknown, but in general the rocks appear similar.

Thus, near the fault, they are white and glassy, and massive or locally coarsely brecciated whereas farther west they are somewhat schistose (1h) and are thinly interbedded with chlorite-bearing quartzite and chlorite schist beds, even locally with a few beds of black massive argillite.

The argillite (2), together with its derived metamorphic equivalents (2a), occurs mainly in the east half of the area. It overlies the quartzite (1) and locally is interbedded with it as narrow beds and lenses. Where the argillite is relatively fresh it is grey to black and thinly bedded, but where slightly more metamorphosed it is a green, massive to schistose, chlorite-epidote-bearing rock (2a). Field relations suggest that this chlorite-epidote rock becomes a hornblende schist (3) near some of higher metamorphism.

The amphibolite (3) is a medium- to coarse-grained, dark green rock composed mainly of hornblende and feldspar. It may be massive but is generally gneissic or finely banded. Angular hornblende crystals, all oriented nearly parallel in a common plane, normally define the gneissic structure. The banding, however, is probably related bedded structure as it consists of an alternation of light coloured, feldspar-quartzite-rich bands and dark coloured hornblende-bearing bands. Hornblende schist near the southwestern end of Mickey Lake seems to pass to the south across the lake into the chlorite-epidote rock (2a) described above. The amphibolite and the hornblende schist overlie the quartzite and are interbedded with it in the same way as the argillite and the chlorite-epidote rock.

In the area around Pluton Lake, in the northwestern corner of the map-area, grey, fine-grained chlorite- and/or biotite-bearing quartzites (4) are thinly to coarsely interbedded with quartz-feldspar-chlorite-sericite schist and gneiss. This interbedded mixture forms a belt a few hundred feet wide. The schist and the gneiss have characteristic yellowish green mica streaks and lines, and generally exhibit an augen structure. The augen are made up of red feldspar or of a mixture of white quartz and red feldspar, and locally are lenses or small masses of granitic material.

The granitic banded gneiss (5) is generally red with an obscure, granitic appearance. It is composed mainly of quartz, feldspar, and chlorite but biotite or hornblende may be present locally. The mafic minerals are concentrated in bands of dark and light coloured bands which are composed mainly of quartz and feldspar. This banded gneiss occurs as irregular masses throughout the other rocks and is believed to indicate variations in the original composition of the rock.

The metasomatic granite (6) is a coarse-grained, massive, red rock, composed mainly of quartz and feldspar. Chlorite is the common mafic mineral and accounts for less than 5 per cent of the rock. Locally the granite may be gneissic, and there the chlorite content is somewhat higher and accounts for the gneissosity of the rock. Remnants of all the other Tazin rocks are found throughout the granite and show gradational contacts with it.

The rocks of the Tazin group are overlain unconformably by those of the Athabasca group (7-10). The unconformity is a marked

feature but is not everywhere easily determined as the rocks underneath are locally brecciated. Its trace is somewhat irregular and wavy. The Athabasca formations normally follow one another in stratigraphic succession westerly from the unconformity. This is suggested by the dip of the plane of unconformity and by the dips of beds. The apparent thickness of this succession is at least 800 feet.

The basal conglomerate (7) is a hard reddish rock made up of angular to subangular fragments of quartzite rocks, coarse red granite, and gneiss, all unsorted as to size and kind and generally closely packed in an arkosic matrix. Near the unconformity the fragments are somewhat sorted as to kind as they are mainly of local origin, and the matrix is in places dark coloured and chlorite-rich. Near the overlying siltstone band (8), however, the fragments are unsorted and the matrix resembles siltstone.

The siltstone (8) is a chocolate-red, fine-grained rock, generally thin bedded, but locally massive. It lies immediately above the basal conglomerate and forms a lens 800 feet wide that subsides to the northeast to enclose a lens of conglomerate (10).

The arkose (9) is orange-red to reddish brown. It grades from a fine-grained rock near the siltstone band to a coarse-grained arkose away from it. Locally, however, the grain size is highly variable laterally. The arkose is generally massive and coarsely bedded. It is composed of grains of feldspar and quartz with some interstitial iron oxide. Cross bedding was noted very rarely and where observed was of micro type.

The conglomerate (10) forms lenses interstratified with the arkose (9) and siltstone (8). It is characterized by rounded to subangular fragments fairly well sorted as to size, being generally less than 3 inches in diameter, and resting in an abundant arkosic matrix. Some of the fragments are derived from the arkose and siltstone bands lying below and interstratified with the conglomerate.

The gabbro and basalt (11) is dark green, black, and reddish brown. It is a massive, coarse- to fine-grained rock that occurs mainly as dykes and sills generally less than 15 feet wide. The dykes and sills are sparsely distributed over a wide area east of Fredette Lake but west of the lake their main occurrence is in a swarm less than half a mile across. They occur mainly in areas underlying the Tazin rocks but a few cut the basal Athabasca conglomerate and the arkose directly above it.

The rocks of the Tazin group have been much folded and faulted. East of Fredette Lake the folds are rather open whereas west of the lake they are close, tight, and probably overturned to the northwest. All fold axes are north-south, near their mapped position is only approximate. East of Fredette Lake, folds in the hornblende schist are close and disharmonious in relation to the chlorite-epidote rock (2a) in the underlying quartzite rocks.

The Athabasca formations north of the A B C fault form the east limb of a broad basin plunging gently northwesterly. The axis of this basin is in Fredette Lake about 1,500 feet east of the Black Bay fault, which cuts out most of the west limb. South of the A B C fault, the Athabasca formations are on the west limb of another broad basin, the Martin Lake basin, near its northern apex. This also is truncated on the west by the Black Bay fault and on the north by the A B C fault.

Faults were seen in rocks of both the Tazin and Athabasca groups. They strike mainly northwesterly to easterly, but some strike north-south. Their dip is generally south but held to be generally northwesterly. Some indicate large apparent offsets. The A B C fault, which may be the extension of the St. Louis fault, and the Camdecc fault, which is probably the Tom fault, join near the south end of Fredette Lake and the combined fault seems to converge northwesterly into the Black Bay fault.

Uranium is the only metal of known economic interest in the area and occurs as a pitchblende along the Black Bay fault and in the Martin Lake basin. The pitchblende forms lenses generally less than 2 inches wide along these fractures. The sheer and strike structures strike northwesterly and north-easterly to easterly and dip gently to steeply south. It is believed that these fractures correspond to a system of joints in the area.

Donaldson Lake in the southeastern corner of the map-area at a rate that varies from time to time but is generally less than 15 tons a day. Baska Uranium has an adit on the west shore of Virgin Lake near its south end. This adit was driven to test the ground some 600 feet north of the northern boundary of the area. Another small orebody was apparently outlined near Pluton Lake. Estimates ground a short distance north of Hab Lake beyond the northeastern corner of the map-area.

5.1.10 Uranium City, Sask.
A.G.S.O.
25-1957.