



- LEGEND**
- PENNSYLVANIAN**
CLIFTON FORMATION: grey sandstone,
12 grey-green siltstone; minor olive-green
and red shale.
- BATHURST FORMATION:** red conglomerate
and grit; shales; red sandstone,
siltstone, shale
- DEVONIAN (?)**
10 Pink, coarse- to medium-grained biotite
granite; gneiss; granite porphyry;
10a, gneissic granite to quartz diorite
- 9 Diabasic gabbro, diorite; sills and dykes;
9a, serpentinite
- SILURIAN**
MIDDLE SILURIAN
CHATEAU BAY GROUP (6-8)
8 Limy slate, subgreywacke, limestone;
minor basalt
- 7 Volcanic rocks: minor interbedded
sedimentary rocks; 7a, mainly basalt,
some trachyte; 7b, mainly trachyte
- 6 Sedimentary rocks 6a, green and red
conglomerate; 6b, limy shale,
slate, greywacke
- ORDOVICIAN**
MIDDLE ORDOVICIAN
TETAGOUCHE GROUP (1-5)
5 Grey, black, green, and red slate,
greywacke, conglomerate; granoblastic
schist; minor limestone, basic
volcanic rocks
- 4 Basic volcanic rocks; interbedded slate,
granoblastic schist, iron-formation; 4a, basic
tuff, in places occurs with unit 1
- 3 Quartz-feldspar porphyry; minor
layers of rhyolite, chlorite quartzite
and chlorite
- 2 Acid to intermediate lava and tuff, derived
quartz-sericite schist; locally
abundant agglomerate and breccia with
elongated fragments; minor sedimentary
and porphyritic layers
- 1 Metasedimentary rocks: feldspathic
quartzite, quartz-chlorite and
quartz-sericite schist; interbedded slate
- Contact metamorphic aureoles around
bodies of 10, biotite hornfels,
quartz-biotite gneiss.
- Drift-covered area.
- Geological boundary**
(defined, approximate, assumed)
- Limit of geological mapping
- Bedding (horizontal, inclined,
vertical, overturned)
- Bedding (upper side of bed faces as
indicated, direction of dip unknown)
- Schistosity, cleavage (inclined,
vertical, dip unknown)
- Fault (assumed)
- Anticline axis (assumed)
- Synclinal axis (assumed)
- Glacial striae (direction of ice
movement known, unknown)
- Esker
- Possibly locality
- Open pit
- Mineral prospect or occurrence
- Shaft
- MINERAL SYMBOLS**
- Antimony Sb Pb
Arsenic As Mn
Copper Cu Mo
Gold Au Silver Ag
Iron Fe Zinc Zn
- Geology by E. W. Shaw (1935), R. Skinner
(1949-54), J. D. McAlary (1956), C. Brown
(1954), C. H. Stoddard (1955-56), Charles
H. Smith (1955-56), F. D. Anderson (1956),
R. E. Dawson (1956), Geological Survey of
Canada; A. L. McAllister, J. Smith
(1955-56), New Brunswick Department of
Lands and Mines, Geological compilation
by Charles H. Smith.
- Main highway
- Other roads
- Power line
- Parish boundary
- County boundary
- Indian Reserve boundary
- Intermittent stream
- Approximate magnetic declination 24° 06' West
- Cartography by the Geological
Survey of Canada
Photographic Library, Topographical
Survey, Ottawa, Ontario

DESCRIPTIVE NOTES

The map-area lies in the northern part of a belt of deformed early Palaeozoic sedimentary and volcanic rocks, cut by granite batholiths, that extends southwesterly from Bathurst to the Maine border.

The rocks of the district are divisible into three structural units.

1. Pre-Silurian sedimentary and volcanic beds that have undergone intense igneous metamorphism form the core of the area.

2. Less deformed Silurian and Devonian strata that overlie these unconformably on the north.

3. Flat-lying Carboniferous rocks that lie unconformably on 1 and 2.

The pre-Silurian rocks are subdivided on a lithological basis, on a structural basis, and are interlayered with quartz-feldspar porphyry (1) south and north of Little Bald Mountain. Some areas mapped as unit 1, such as the Heath Steele ore-bearing horizons, may be metamorphosed equivalents of unit 5.

The Silurian volcanic rocks (1) cover a wide area in the south and south-west parts of the map-area. They underlie the siliceous volcanic rocks (2) along the Northumberland River south of South Little River Lake. They also underlie, and are interlayered with quartz-feldspar porphyry (1) south and north of Little Bald Mountain. These rocks are generally schistose. In this section they are seen to be composed of ragged feldspar phenocrysts with or without embayed, brecciated quartz phenocrysts, in a fine, schistose ground-mass of quartz, feldspar, chlorite, and sericite. In some places it is difficult to determine whether the schistose acid rocks are sheared flows or tufts. In some places the phenocrysts in the above rocks are particularly large and abundant. Later-feldspar porphyry (1) has assumed great importance in the district because of its spatial relation to ore deposits. Porphyritic rocks behave locally as normal stratigraphic units, but a regional scale they occur in various stratigraphic positions, either above, below, or interbedded with, the main body of volcanic rocks (1) and for this reason they cannot be directly correlated throughout the district. More than one rock type is called 'porphyry' locally. The type that is represented by most of unit 1 is a close association of unit 2, in a microcrystalline or interstitial texture, with quartz and/or feldspar phenocrysts up to 1/2 inch long in a matrix of quartz, feldspar, chlorite, and sericite. A fourth rock type, possibly a rare variety of porphyry, is called 'quartzite' locally. It is a coarse-grained rock with a well-developed cleavage and contains large phenocrysts of quartz, feldspar, and chlorite. As the size and abundance of phenocrysts decrease in grades into the quartzite, the rock becomes more and more schistose. The schistose rocks are considered to be contemporaneous with them.

A second type of porphyry forms dikes that cut across the Ordovician and Silurian rocks, e.g. at Nigadoo Mines, Limited. It is younger and less completely altered than the Silurian porphyry, and is probably related to the granite intrusions. A third rock type that resembles porphyry in places is actually an altered, or metamorphosed, quartzite. It is a fine-grained rock, in the belt of porphyry containing the Heath Steele and Brunswick No. 12 (12) are deposited. In outcrop they resemble quartzite, but differ from a true porphyry in having a fine-grained rather than microcrystalline groundmass. They are associated with quartz-feldspar granites and quartzite. A fourth rock type, possibly a rare variety of porphyry, is called 'quartzite' locally. It is a coarse-grained rock with a well-developed cleavage and contains large phenocrysts of quartz, feldspar, and chlorite. As the size and abundance of phenocrysts decrease in grades into the quartzite, the rock becomes more and more schistose. The schistose rocks are considered to be contemporaneous with them.

The Silurian volcanic rocks (1) are interlayered with the acid volcanic rocks (2). They are brown weathering, dark green, schistose, fine-grained to medium-grained rocks with rare meso- or macroscopic structures, such as sills and dykes. They consist entirely of secondary mafic minerals, calcite, and magnetite. Layers of red and granitic slate are probably numerous, but these rocks rarely outcrop in interstream areas. Layers of jasper and magnetite-bearing chert occur with the volcanic rocks and cause high magnetic anomalies. The upper sedimentary unit of the Tetagouche group (5) consists chiefly of argillaceous shales. The lower unit (6) is a fine-grained, argillaceous, crumpled, and in places contains secondary quartz veins and disseminated pyrite. Electromagnetic surveys have shown that this argillaceous unit is common. Basic volcanic rocks (4) are interlayered with members of this unit, but only the larger are shown on the map. Lithologically similar rocks also occur interlayered or folded with the acid volcanic complex (2, 3).

The sedimentary rocks of the Chateau Bay group are predominantly fossiliferous, greenish grey limy slates and greywackes. They are divided into two members (7) and (8) separated by a disconformity, although the latter is largely interbedded with the volcanic rocks. The lower unit contains a distinctive fossiliferous member (7a). The volcanic rocks (7a) are dark red, orange or greenish grey, fine-grained basalt (7a) to dark red, orange or grey trachyte (7b). The basalt (7a) is considered to be of Middle Silurian age, and is the basis of fossils collected north of the area. The only diagnostic fossils collected within the area are trilobites and brachiopods from the basal beds of the basalt. The basal beds of the trachyte (7b) are considered to be of Middle Silurian age. As these fossils occur in the upper sedimentary rocks, they are not assigned to the Middle Silurian, although infolded Devonian beds may be present.

The basic sills and dykes (9) are grey or greenish grey, fine-grained, and carry nickel-bearing sulphides. A distinctive petrologic type is found near the Tracy Lumber Company road northeast of Goodwin Lake and also west of Popple Depot. It is banded and contains olivine. It resembles the type found near the Tracy Lumber Company road northeast of Goodwin Lake and carries nickel-bearing sulphides. Smaller sills and dykes (9a) are also present. The basic intrusions cut both the pre-Silurian and Silurian rocks. They are the most abundant near the area of the Nickel Brook, and because of this association, they may thus be basic dykes of two ages. Some relationships between the basic intrusions and some related to Silurian-Devonian volcanism, the absence of good field evidence they can be difficult to distinguish from more altered flows. The principal criteria are the coarse-grained, schistose texture, and massive appearance.

The granitic rocks (10) occur as two distinct types. Gneissic varieties (10a) found near South Little River Lake and Seville River were emplaced during the Middle Silurian, and are composed of clonoproterose, altered glaucoschists, secondary ferromagnesian minerals, and magnetite. A distinctive petrologic type is found near the Tracy Lumber Company road northeast of Goodwin Lake and also west of Popple Depot. It is banded and contains olivine. It resembles the type found near the Tracy Lumber Company road northeast of Goodwin Lake and carries nickel-bearing sulphides. Smaller sills and dykes (9a) are also present. The basic intrusions cut both the pre-Silurian and Silurian rocks. They are the most abundant near the area of the Nickel Brook, and because of this association, they may thus be basic dykes of two ages. Some relationships between the basic intrusions and some related to Silurian-Devonian volcanism, the absence of good field evidence they can be difficult to distinguish from more altered flows. The principal criteria are the coarse-grained, schistose texture, and massive appearance.

Many of the biotite granite bodies are surrounded by a distinctive contact metamorphic aureole (11) characterized by the development of biotite in sedimentary and volcanic host rocks. Contact metamorphic effects may extend as far as a mile or more from the nearest granite outcrop. Near the granite the sedimentary rocks are altered to a fine-grained, rusty weathering, purplish biotite hornfels carrying cordierite and andalusite. The host volcanic rocks are altered to a gneissic biotite-quartz-feldspar rock resembling a granite gneiss.

Nearly flat-lying beds of Pennsylvanian age (11, 12) unconformably overlie the granite and older rocks along the east boundary of the area, and form small outliers near Tetagouche Lakes, north of Bathurst, and on Clearwater Stream. The Clifton formation is conformable with the Bathurst formation, but is dominantly grey in colour. It contains detrital plant remains and thin coal seams.

The structural pattern of the pre-Silurian rocks may be separated into two distinct units by a line running north-northwesterly from North Little River Lake. Northeast of this line, the dominant structures are two steep anticlines, one south of the Tetagouche Lakes and the other west of the Bathurst granite body, separated by a synclinal trough south of Caliform Lake. These anticlines fold with steep plunges to the southwest and northeast, but are interpreted as plunging regional folds to the northeast. In the Tetagouche region, west of Bathurst these major folds are lost in a series of smaller anticlines and synclines. The acid volcanic rocks in the core of the anticlines are generally schistose. The schistosity parallels the limbs of the folds, but cuts across the axis of the fold south of Caliform Lake. Along the major anticline, north-northwesterly from North Little River Lake, north-northwesterly from the northeast towards the southeast near the axis. The schistosity in places extends in parallel with the bedding. The belts of rock extending from the northeast are deflected by the presence of a major fault zone but it is rather diffuse and cannot be traced continuously across the district. North-northwesterly from the northeast, but these folds do not appear to be counterparts of those on the other side of the structural axis. Horizontal cleavage is more common in this area, and may reflect recumbent folding or thrust faulting. These structures are cut off to the west by large granitic intrusions and by the Devonian rocks that overlie the Ordovician rocks west of Popple Depot. On the north, the Rocky Brook-Millstream thrust separates the complex from the Silurian rocks. There are no intensely folded as the older rocks, but the folds have the same northeast plunges.

In the pre-Silurian rocks, large acid-lead-ore replacement deposits occur in sedimentary beds surrounding or included in the acid volcanic-porphry complex. These deposits are found at different stratigraphic horizons throughout the district. The larger deposits discovered to date, such as the Brunswick Nos. 6 and 12, Heath Steele, Middle River, etc., are in secondary beds and porphyry, and this association has been used successfully as a prospecting guide. In biotite hornfels and quartzite, and it is probable that the porphyry bodies acted as competent members among the incompetent sedimentary beds and that ore-bearing solution rock formed in the channels so formed. Numerous small fracture-filling and replacement deposits containing copper, lead, and zinc are found along and near the locally faulted 'streambreaks' in places remote from any porphyry, and in places in the younger Silurian rocks.

Two distinct mineral associations are found in the replacement deposits - an earlier (1) pyritic-chalcocyanite type, commonly associated with dark green, massive, chlorite wall-rock alteration and a later (2) pyrite-glaucophane or abnormal rust colour in the soil may be only visible surface expression. Iron leached from sulphide minerals has cemented glacial till and stream gravel near certain deposits and may be a guide to prospecting. Low-temperature mineralization is associated for the entire district (see also) and are invaluable for structural studies. The magnetic background of the district is about 1,200 to 1,300 gamma over areas of sedimentary and acid volcanic rocks. Basic volcanic rocks and dykes cause anomalies up to 3,500 gamma. Linear anomalies between 3,800 and 5,000 gamma coincide with magnetic-bearing sedimentary rocks (iron-formation, ferruginous chert), which are generally associated with basic volcanic rocks. Pyritic-chalcocyanite deposits are of economic importance in the district, but cause aeromagnetic anomalies in the order of a few tenths of gamma only, which are generally masked by higher anomalies caused by variations in magnetite content unrelated to economic mineralization. Small aeromagnetic closures when associated with electromagnetic or geochemical anomalies may, however, be indicative of these deposits.

Additional References
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MAP 1-1957
BATHURST-NEWCASTLE AREA
NORTHUMBERLAND, RESTIGOUCHE, AND GLOUCESTER COUNTIES
NEW BRUNSWICK
Scale: One Inch to Two Miles = 1:26,720
Miles 0 2 4



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