



**DESCRIPTIVE NOTES**

The oldest formations in the area, an assemblage of volcanic and sedimentary rocks (1a, 1b), are divided into two groups whose age, and whose relationships to each other have not been established. With the exception of a small outlier near Bass River, these rocks are found only within Cobequid Mountains. The mixed conglomerate (1a) on the south side of the mountains lacks lithological uniformity for considerable distance across its strike. Further north in the mountains, a group of grey shales (1b) shows remarkable uniformity throughout. Whenever these two groups are in contact the shales dip at low angles toward the mixed rocks, and the latter may also dip toward the shales. The contact is, therefore, considered to be a fault, and the stratigraphic relationships of the two groups are not determinable. The shales appear, in general, to be less deformed than the mixed rocks, and are probably younger. The two groups are older than the Pennsylvanian, Riverdale group (11), and their structural relations in the fault block north of Economy and of Bass River seem to indicate that they are pre-Upper Windor. Their age is probably pre-Mississippian and either Devonian or Silurian.

Rocks of the earlier groups (1) are intruded by granite, and by minor acid intrusions associated with it. The granite (3) is usually coarse grained and relatively free of ferro-magmatic minerals, in places being termed alkalic. Where dark minerals are present, hornblende is most common. Biotite is lacking, although biotite granites are found elsewhere in Cobequid Mountains. Aplite (3a) is a pink, sugary textured rock, and is exposed for about a mile on Economy River where it replaces its fine grain throughout. Quartz-feldspar porphyry and quartz porphyry (3b) are very similar to aplite, with addition of relatively large phenocrysts of quartz, and, in places, of orthoclase. The field evidence is not such as to eliminate the possibility of these porphyries being extrusive. The age of the granitic rocks has not been established.

Bordering the main body of granite is a zone of contact rocks (2) in which three types are indicated. The most common type (2b) is that in which the earlier rocks were recrystallized without loss of original structure, and granitic material injected into them. In places (2a) this zone is marked by numerous small bodies of granite cutting the older rocks and becoming more numerous toward the main body of granite. The third type (2c) has resulted from assimilation, by the granitic, of more basic material, and has taken on the composition of rocks varying from quartz diorite to gabbro, and, in one instance, an amphibolite.

The pre-Carboniferous rocks (1, 2, 3) are cut by small dykes and bodies of diabase. These are dark grey to black, with a grain varying with the dimensions of the body from microscopic to about 2 millimetres in size, and may be confused with certain of the contact rocks (2c). The diabase intrusions are either too small or too poorly exposed to be mapped. They are quite common on the northeast branch of East River, on Economy River between the large body of aplite and the granite to the south, and on Bass River for over a mile north of the south face of Cobequid Mountains.

Rocks of the Horton group (4) are fossiliferous, and their age is early Mississippian. Structurally they form part of the so-called Walton anticline, which is bounded on the south by a syncline in Windor and Pennsylvanian rocks occupying the valleys of Sarnett and Tenapee Rivers. Locally the Horton beds are folded into a series of minor anticlines and synclines, one of the latter being deep enough to include some lower Windor strata.

Overlying the Horton rocks with apparent conformity is the Macumber formation (5), 20 to 35 feet thick. It is unfossiliferous, and is included tentatively in the Windor group because of its lime content, no limestone being known in the Horton.

The Pembroke formation (6) overlies the Macumber, and consists of red to red-brown, massive limestone and limestone conglomerate. Pebbles in the latter include angular fragments of the Macumber formation, as well as rounded pebbles of pre-Carboniferous rocks. The Macumber frequently occurs as lenses of erosion, and the possibility that the Macumber formation may represent the top of the Horton.

A bed of anhydrite, locally altered to gypsum, overlies the Pembroke formation, apparently conformably. This bed (7) is about 100 feet thick. Although exposures are few, the extent of the bed has been mapped with considerable accuracy by noting the pitted, Karst topography resulting from solution of the sulphate deposit.

Overlying the anhydrite bed is the Tenapee formation (8), consisting of well-bedded, extremely uniform, bright red shale and sandy shale. The rock is quite soft while wet, but dries to a hard, porous mass. The formation is more than 400 feet thick.

A second bed of anhydrite (9), also locally altered to gypsum, overlies the Tenapee formation. Although Tenapee rocks and gypsum are exposed a few feet apart on the north branch of Tenapee River, and although neither shows any evidence of deformation, there is the possibility that this anhydrite bed is part of the other (7) overlies onto the Tenapee. Both the gypsum and the soft sandy shales used as a lubricant for such a movement, permit it to take place with a minimum of deformation. The rocks overlying this sulphate bed (9) are not exposed, and for this reason it has been necessary to leave a part of the southeast corner of the map unperfected.

Beds of Upper Windor age (10) occur north of Cobequid Bay in a narrow belt protruding through Pennsylvanian rocks. They consist of limestone, and are usually quite well bedded. Fossils obtained from them are of late Upper Windor age.

Pennsylvanian strata of the Riverdale group (11) are exposed in a fault block, protruding through the Triassic rocks, which overlap them on the south, and are downfaulted against them on the north. The Riverdale rocks carry diagnostic fossil shells.

Paralleling the south face of Cobequid Mountains, and separated entirely from the Riverdale group by an intervening belt of Triassic rocks, is a belt of Pennsylvanian sediments whose group designation is somewhat in doubt. The lower, and southern, part of the belt (13) consists of grey beds varying from fine shales to sandstones, the latter occurring in thin beds. Fossils from these beds are of either Cambrian or Pictou age. The upper part of the belt (14) is unfossiliferous, and consists of brownish red, coarse conglomerate and coarse and fine sandstone. As the Pictou group is known elsewhere to be composed of rocks similar to those of the upper part of this belt, and as the coarser rocks overlying the finer shales and sandstones suggest that a break may exist between the two divisions, it is suggested that the lower members (13) are of Cambrian age, and the upper members (14) of Pictou age.

Across most of the map-area, rocks of supposedly Pictou age (14) are down-faulted against the rocks of Cobequid Mountains. Near the west side of the map-area, however, Pennsylvanian rocks (12) rest unconformably on the older rocks of the mountains. They are quite similar lithologically to the supposedly Pictou rocks south of the mountains elsewhere, and like them are unfossiliferous. It is believed, however, that the Cobequid fault here may have left a remnant of the basal Pennsylvanian series on the south side of the mountains, and that this remnant (12) is not probably of Riverdale age.

Unconformably overlying the Pennsylvanian series are rocks of Triassic age, referred to the Annapolis formation (15). They comprise bright red, coarse to fine conglomerate and sandstone, and except near faults dip at low angles. A series of basalt flows of considerable thickness (16) overlies those parts of the Annapolis formation exposed in the area, and may actually be interbedded with it. The rock has everywhere the composition of a basalt, but its texture and appearance vary in different parts of an individual flow.

The rocks of the map-area have all undergone considerable faulting, but, with few exceptions, only those faults of post-Triassic age can be recognized and mapped.

Manganese deposits occur as replacements of limestone, usually that of the Pembroke formation. They also form stringers in rocks of Horton and Riverdale age, and replace the contact in Triassic sandstone and conglomerate.

Iron occurs as lines of specularite along fractures in pre-Carboniferous volcanic and sedimentary rocks (1). Magnetite is known in the Triassic basalt on Gerrish Mountain, as the west border of the map-area, and also in the volcanic and sedimentary rocks on Bass and Economy Rivers.

Gypsum has been quarried from the first or lower anhydrite bed in the Windor group. Limestone is at present being quarried for agricultural purposes from Upper Windor rocks of Upper Economy.