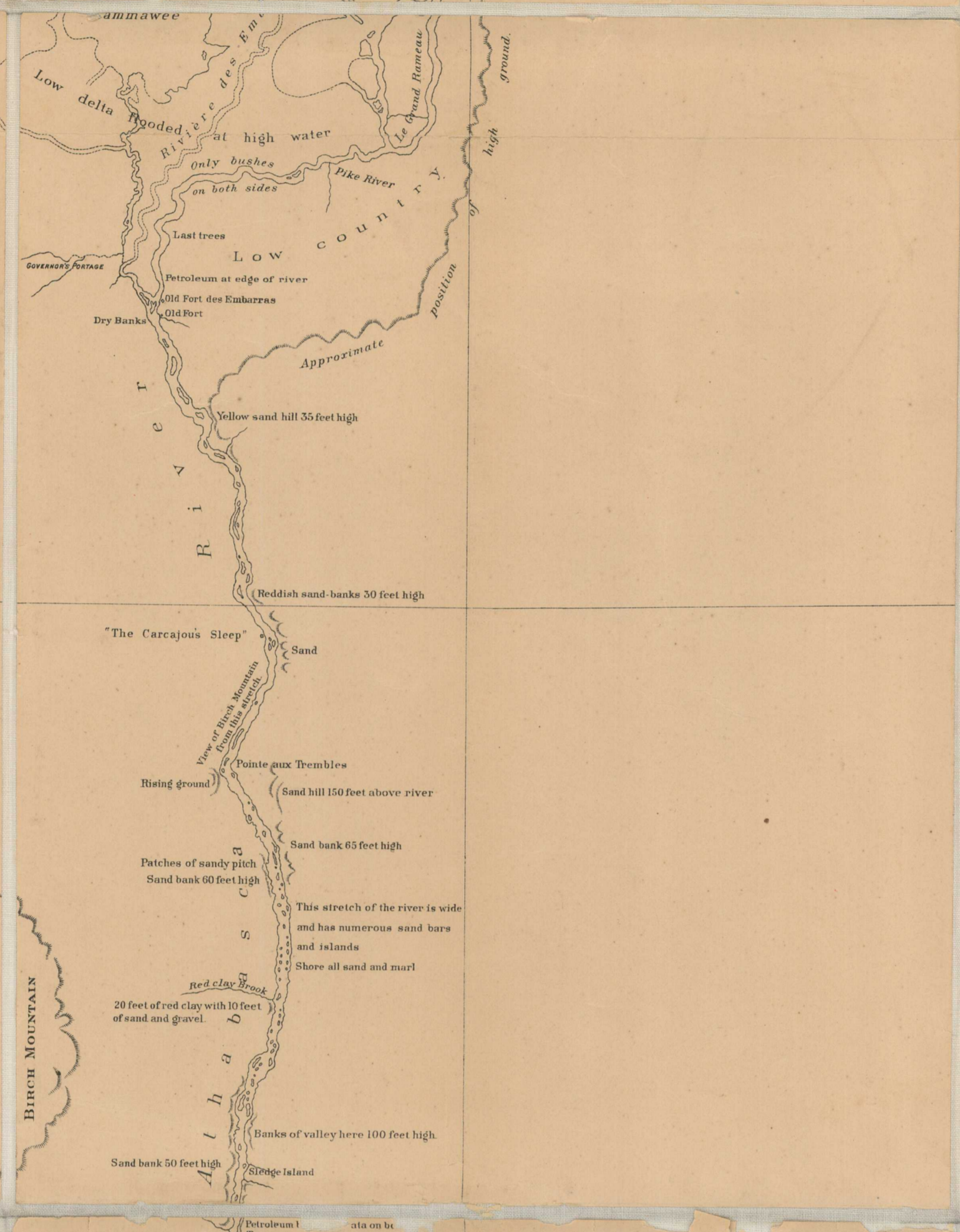


GEOLOGICAL NOTES.

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1. Lake Athabasca occupies a shallow basin excavated in granitic gneiss and sandy shale, of Devonian age.—Page 27, C. C.
2. Small sections of dark chert and sub-colored marls and shales with nodules and thin layers of clay ironstone, lying quite horizontally, occur on the banks of this part of the Athabasca.—Page 7, C. C.
3. From the Athabasca to the Peace River, a distance of about 12 miles, the Athabasca bank between dipping banks from 50 to 150 feet high, consisting of thick shaly Devonian marls having a general horizontal strike, and containing large calcareous sponges and smaller nodules of chert.—Page 7, C. C.
4. From Peace to House River, a distance of 22 miles, the character of the banks of the Athabasca continues the same as in the last section. The rocks consist of soft grey sandstone below, and independent marls above.—Page 9, C. C.
5. For 4 miles in this vicinity, the sandstone forming the lower part of the bank has an almost uniform thickness of 50 feet, and is intersected with some bluish marly bands.—Page 9, C. C.
6. The river, for a few miles above the Grand Rapids, is flanked on both sides by cliffs about 40 feet high, of soft, fine-grained bluish grey sandstone, weathering yellowish grey.—Page 9, C. C.
7. At the Grand Rapids the bed of the river breaks down into a level of sandstone, which is continuous for a distance of 30 miles below, owing to its being more or less thickly underlain with shaly sandstone, differing from the strata in containing some argillaceous matter.—Page 10, C. C.
8. Bank of soft grey bluish sandstone, 50 feet high, with the conglomerate very hard, 20 feet thick at the base, for several miles along the left side of the river.—Page 11, C. C.
9. Banks of soft sandstone, 125 feet high, containing seams of lignite, one of which is 4 feet thick.—Page 12, C. C.
10. Cliffs 170 feet high of homogeneous grey sandstone, parts of which weather into pillars.—Page 12, C. C.
11. Sandstone cliffs with seams of lignite.—Page 12, C. C.
12. Cliffs 200 feet high of soft grey sandstone, weathering yellow, with the conglomerate band at one-third of this height.—Page 13, C. C.
13. Seams of lignite in sandstone cliffs on the right bank, at 3 and 4 miles below Little Buffalo River.—Page 13, C. C.
14. At Burnt Rapids are beds of little high drab-colored sandstone and of calcareous green sandstone containing Cretaceous shells and fragments of siliceous wood.—Page 13, C. C.
15. Cliffs of sandstone 200 feet high, in four terraces.—Page 13, C. C.
16. The bank at Pointe Terre Brule is 200 feet high, in three terraces, the conglomerate band forming the central one.—Page 13, C. C.
17. At the Drowned Rapids, Cretaceous fossils occur in the harder sandstone beds. Here also the black, petroleum-bearing, fine-grained sandstone, first makes its appearance. It becomes abundant and conspicuous hereafter nearly to the delta of the river.—Page 14, C. C.



18. The bank is here about 200 feet high. The top of the conglomerate band, which is 50 feet thick, is 200 feet over the river.—Page 15, C. C.
 19. The petroleum-bearing strata dip in different directions in this neighborhood.—Page 15, C. C.
 20. The Devonian rocks are first seen on the upper side of Crooked Point. They consist of a few feet of bluish grey limestone, mostly of an vertically character, lying horizontally.—Page 15, C. C.
- The following is an estimated section of the Cretaceous rocks along the Athabasca, in descending order from top to bottom:
- | | |
|--|-----|
| Bluish grey marls, mostly indurated, holding thin layers and conical concretions of impure clay ironstone. These occur from the banks from the Delta to the Peace River. | 200 |
| Assorted marls and homogeneous fine-grained grey sandstones, in the banks from Peace River to the foot of Grand Rapids. | 170 |
| Masses of grey sandstone, underlain with concretions which are mostly spherical, and of large size, from 30 to 40 feet in diameter. | 50 |
| Marls, mostly ammoniac, grey, drab, etc. | 120 |
| Fine sandy strata, mostly blackened by petroleum. | 200 |
| | 760 |
21. At the Cascade Rapids, 80 or 90 feet of petroleum-bearing sandy marl are exposed in the bank. This is underlain by about 15 feet of Devonian limestone, over part of which the river passes in low seasons. The limestone is exposed at the foot of the Cretaceous rocks along the edge of the river for many miles below this rapid.—Page 17, C. C.
 22. For 12 miles from The Forks, the right bank of the river is from 100 to 200 feet high, and consists of beds of thick Devonian sandstone composed of fine quartz and saturated with petroleum, which gives it a shaly appearance when freshly exposed, resting on a few feet of Devonian limestone.—Page 16, C. C.
 23. The Devonian limestones which were observed as far down as this point, are generally yellowish-grey, thin-bedded and somewhat earthy, with rough surfaces. They form low ledges and hills along the edge of the river on both sides. The general attitude of the strata is about horizontal, the bedding is, however, seldom quite level for any great distance, but undulates slightly in all directions, and finally disappears under the river.—Page 25, C. C.
 24. Limestone gneiss. At Fort Chipewyan it is strongly bedded and fibrous, the prevailing color being red, and the strike S. 45° W. (true). At the outlet of Lake Hamanowee the strike is S. 35° W. (true).—Page 24, C. C.
 25. Masses of sandy silt in the bed of the Clearwater River at 11, 13 and 17 miles above The Forks.—Page 24, C. C.
 26. At the Cascade Rapids the rock is a hard, yellowish-grey sandstone, with a bluish color on fresh fracture.—Page 25, C. C.
 27. At the Paw Rapids and in the valley of the river above and to the north of it, much porous or spongy grey limestone is exposed. One bed in the vicinity of the rapid is stained with free petroleum. Islands and phases of the limestone stand in the river as the rapid and in the sand which covers the bottom of the valley in the neighborhood. In some places the limestone is cavernous, and all the exposures are much decayed and eroded.—Page 25, C. C.
 28. At the Twin Rapids, the river passes down a narrow high island and points of grey limestone, which is much stained on the surface by the weather, but otherwise appears to be mostly of a massive character. No fossils were observed in the rocks at any of the rapids, but they appear to belong to a part of the Devonian system, somewhat lower down than the fossiliferous beds, which immediately underlie the Cretaceous sandstone on the Athabasca.—Page 25, C. C.
 29. At this place a number of copious springs of mineral water issue from the north bank in a space of about 200 yards. From one of these the saline matter was obtained by evaporation, and found by Mr. Hoffman to consist of potash, soda, magnesia and lime, all in considerable quantity as sulphates, chlorides and carbonates.—Page 25, C. C.
 30. The banks of the Clearwater consist for the most part of a soft, puffy, grey clay. On the Muddy Portage trail, the bank of the river is one and a half miles from the river and 400 feet above its level.



Geological and Natural History Survey of Canada.
Alfred R.C. Selwyn L.L.D., F.R.S., Director.

MAP
OF PART OF THE
ATHABASCA RIVER

To illustrate
DR. BELL'S EXPLORATION.
Report 1882-84, pages 100 to 370c.

Scale 8 miles to 1 inch
Val. Scale 250,000
1884.
Compiled & Drawn by
A.S. COCHRANE, Assistant Topographer.

NOTE.
The Clearwater River and Lake Athabasca are taken from unpublished track surveys made in 1881, by A.S. COCHRANE, the remainder of the map being from track surveys by Dr. R. BELL.