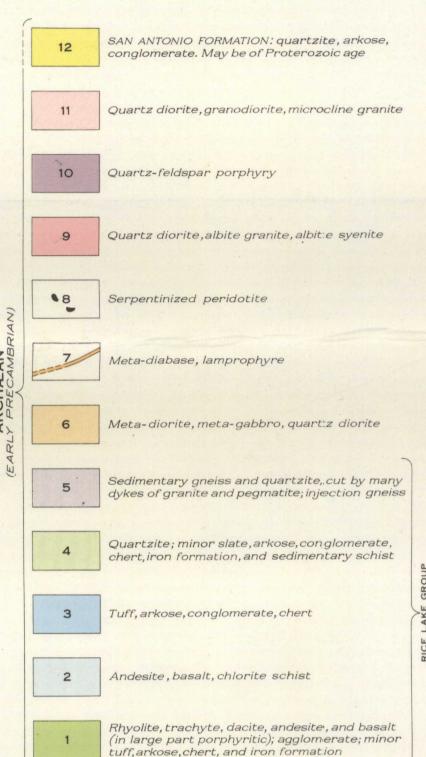
CANADA DEPARTMENT OF MINES AND RESOURCES

MINES AND GEOLOGY BRANCH BUREAU OF GEOLOGY AND TOPOGRAPHY

LEGEND

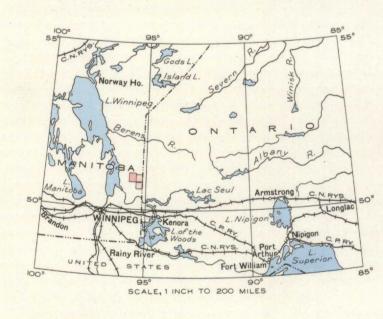


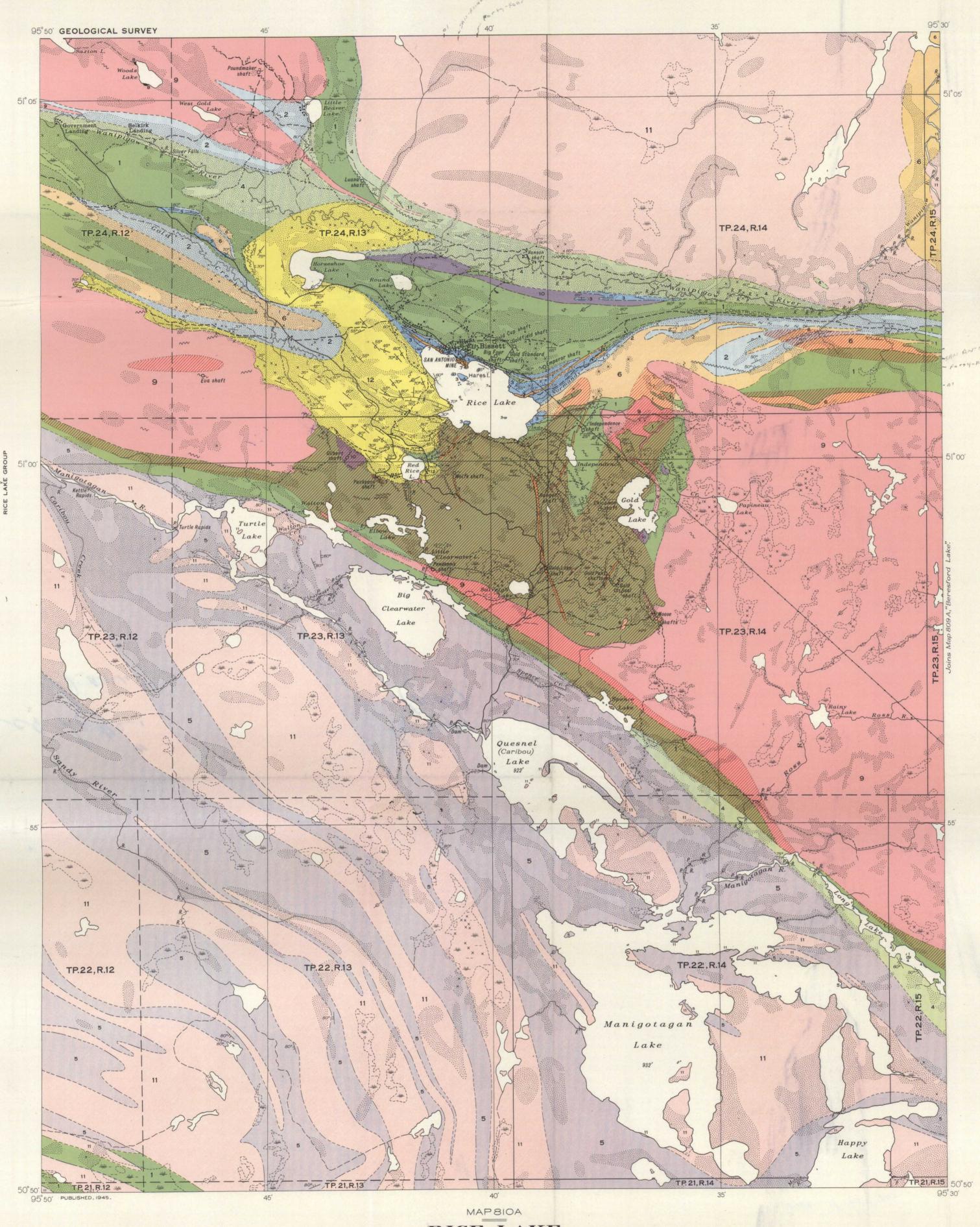
Area within which dykes of feldspar porphyry and quartz-feldspar porphyry are numerous		1111
Drift-covered area		
Small rock outcrop		×
Bedding (inclined, vertical, overturned)	XX	9
Bedding (upper side of bed faces as indicated, direction of dip unknown)		6
Bedding (direction of dip known, upper side of bed unknown)		
Fault or shear zone, generally carrying vein quartz	~~~~	ww.
Synclinal axis		
Glacial striæ		6
Road and buildings		
Road not well travelled		
Winter road or trail		
Portage	<u>P</u>	
Tramway		++++
Transmission line		
Post Office		
School		
Cemetery		+
Shaft		
Township boundary (surveyed)		
Township boundary (unsurveyed)		_
Stream (position approximate)		
Fall and rapid	E	R.
Marsh	(3/4-	- NE

Geology by C. H. Stockwell, 1939, and from published maps of the Geological Survey.

Height in feet above mean sea-level.

Base-map compiled by the Topographical Survey, 1943, from air photographs taken in October 1934, and from published Federal Government maps and from information supplied by the Department of Mines and Natural Resources, Manitoba. Cartography by the Drafting and Reproducing Division, 1944.





DESCRIPTIVE NOTES

Basic to acidic lavas and pyroclastic rocks (1) of the Rice Lake group are interbedded with one another in layers from a few feet to 1,500 feet thick. Much of the lava carries phenocrysts of plagioclase feldspar. Pillow structures are rare. Pyroclastic rocks are about as abundant as the associated flows. Near the contact with the large granitic body (9) east of Gold Lake the volcanic rocks are recrystallized to coarser than normal material and, rarely, are garnetiferous. Interlayered with these grey to black volcanic rocks are dark green lavas (2) that, in part, are to black volcanic rocks are dark green lavas (2) that, in part, are pillowed and amygdaloidal. Sedimentary strata (3) are interbedded with the volcanic rocks (1, 2) of the Rice Lake group. In many places they form bands too thin to map, but their occurrence is indicated by bedding symbols. Pebbles and boulders in the conglomerate consist of rhyolite, porphyritic andesite, quartz-feldspar porphyry, chert, and vein quartz. Much of the tuff is thinly bedded and shows grain variation from coarse, at the bottom of each bed, to fine, at the top. The main body of volcanic and sedimentary rocks (1, 2, and 3) lying north and south of Rice Lake is apparently folded into a broad, easterly-trending anticline with the axis lying 1/2 mile or more south of the lake. Minor synclines occur within the major anticlinal fold. A thick series of sedimentary rocks (4, 5) overlies this main body conformably and outcrops on the limbs of the anticline. The abundant quartzite of the sedimentary rocks (4) is, in places, massive and without bedding structures; in other places it forms beds up to 10 feet thick; and, in still other places, it is thinly laminated. Interbeds of black slate are common. Near the contact with the granitic body in the northeast part of the map-area the slate has been altered to biotite schist, and the quartzite is in places garnetiferous. Conglomerate beds hold boulders and pebbles of granite, trachyte, and porphyritic trachyte. Banded magnetite iron formation was noted at the east end of Horseshoe Lake and strong magnetic attraction in a drift covered area a mile north of the west end of this lake suggests the possible presence of iron formation there. In the southwest part of the map-area, where large sills and stocks of granitic rock are abundant, the sedimentary rocks (4) are partly altered to quartz-plagioclase-biotite-muscovite gneiss cut by numerous dykes, stringers, and lenses of granite and pegmatite and by lit-par-lit injections of these igneous rocks (5). Many beds are garnetiferous

The sedimentary and volcanic rocks of the Rice Lake group are cut by sills and irregular-shaped bodies of meta-diorite, meta-gabbro, and quartz diorite (6), and by apparently somewhat younger dykes and sills of meta-diabase and lamprophyre (7). Some of the lamprophyre dykes are younger than the granitic intrusions (9, 10, 11) and others are older. Very small bodies of serpentinized chromiferous peridotite (8) occur at Woods Lake. The age relation of this peridotite to the other basic intrusions (6, 7) is unknown.

Batholithic and smaller bodies of granitic rocks comprise a sodic type (9, 10) and a potassic type (11). Their age relations are unknown. In the sodic type the feldspar varies from albite to andesine and some of the rock carries phenocrysts of plagioclase or conspicuous eyes of quartz. In the potassic type microcline or orthoclase feldspar is also commonly present. Dykes of pegmatite are common in and near bodies of the potassic type but are rare in and near the sodic type. The latter is, however, the source of large bodies of porphyry (10) and of innumerable smaller porphyry dykes. Areas within which these dykes are abundant are indicated on the map. Minerals noted in the pegmatites include quartz, feldspar, biotite, muscovite, tourmaline, garnet, magnetite, and apatite. was also seen in dykes of aplite and granite.

The beds of the San Antonio formation (12) rest unconformably on rocks of the Rice Lake group and on granitic rocks (9) that cut this group. West of Rice Lake the strata have been folded into a syncline with a gently-dipping south limb and an overturned north limb. South and west of Horseshoe Lake the uppermost beds of this fold face so as, apparently, to underlie adjacent lavas of the Rice Lake group. However, there is a wide zone of intense shearing along the contact and it is probable that the anomalous relationship is due to faulting. Feldspathic quartzite is the most abundant member of the formation. Crossbedding is common. Well rounded pebbles are scattered at random in the quartzite or are concentrated in thin beds. Between some of the quartzite beds are thin partings of finegrained, pale green sediment commonly showing a grain variation from coarse at the bottom to fine at the top. A basal conglomerate from 10 feet, or less, to 80 feet thick occurs along the granite contact on the south limb of the syncline. Rounded boulders and angular blocks of granite derived from the underlying rock are commonly 1 to 10 feet in diameter, the largest noted being 40 feet long and 20 feet wide. These lie in a matrix of smaller boulders and pebbles of a great variety of materials The cement is of schist or arkose, and cracks in the underlying granite are filled with arkose.

Vein quartz has been deposited along shear and fracture zones and faults that either parallel or cross the trend of enclosing rock formations. Minerals noted in the quartz include gold, pyrite, chalcopyrite, pyrrhotite, sphalerite, galena, molybdenite, carbonate, chlorite, bright green chromiferous mica, albite, orthoclase, and tourmaline. Many of the veins in pre-San Antonio rocks are probably genetically related to the sodic intrusions. The origin of those in the San Antonio formation is unknown. Gold production has been almost entirely from the San Antonio mine, on the north shore of Rice Lake; small amounts have come from several other deposits. San Antonio Gold Mines, Limited, reports that gold and a little silver to the value of \$14,488,480 were derived from 1,339,096 tons of ore milled from 1932, when operations commenced, until December 31, 1943, when ore reserves were estimated at 916,471 tons. Production was from a large number of veins arranged in two sets, one striking northeast and the other northwest, in a northdipping sill of meta-diabase. Workings extend to a depth of 2,400 feet.

RICE LAKE EAST OF PRINCIPAL MERIDIAN

MANITOBA

Scale, 63,360 or I Inch to I Mile Approximate magnetic declination, 7°30' East. NOT TO SE TAKEN FROM LIBRARY NE PAS SORTIR DE LA BIBLIOTHEQUE

